Simulation of boundary layer clouds with double-moment microphysics and microphysics-oriented subgrid-scale modeling

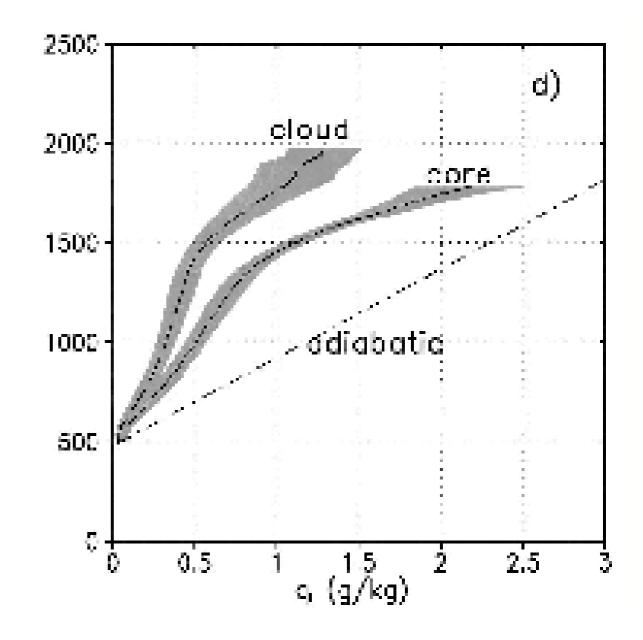
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Motivation

Shallow convective clouds are strongly diluted by entrainment





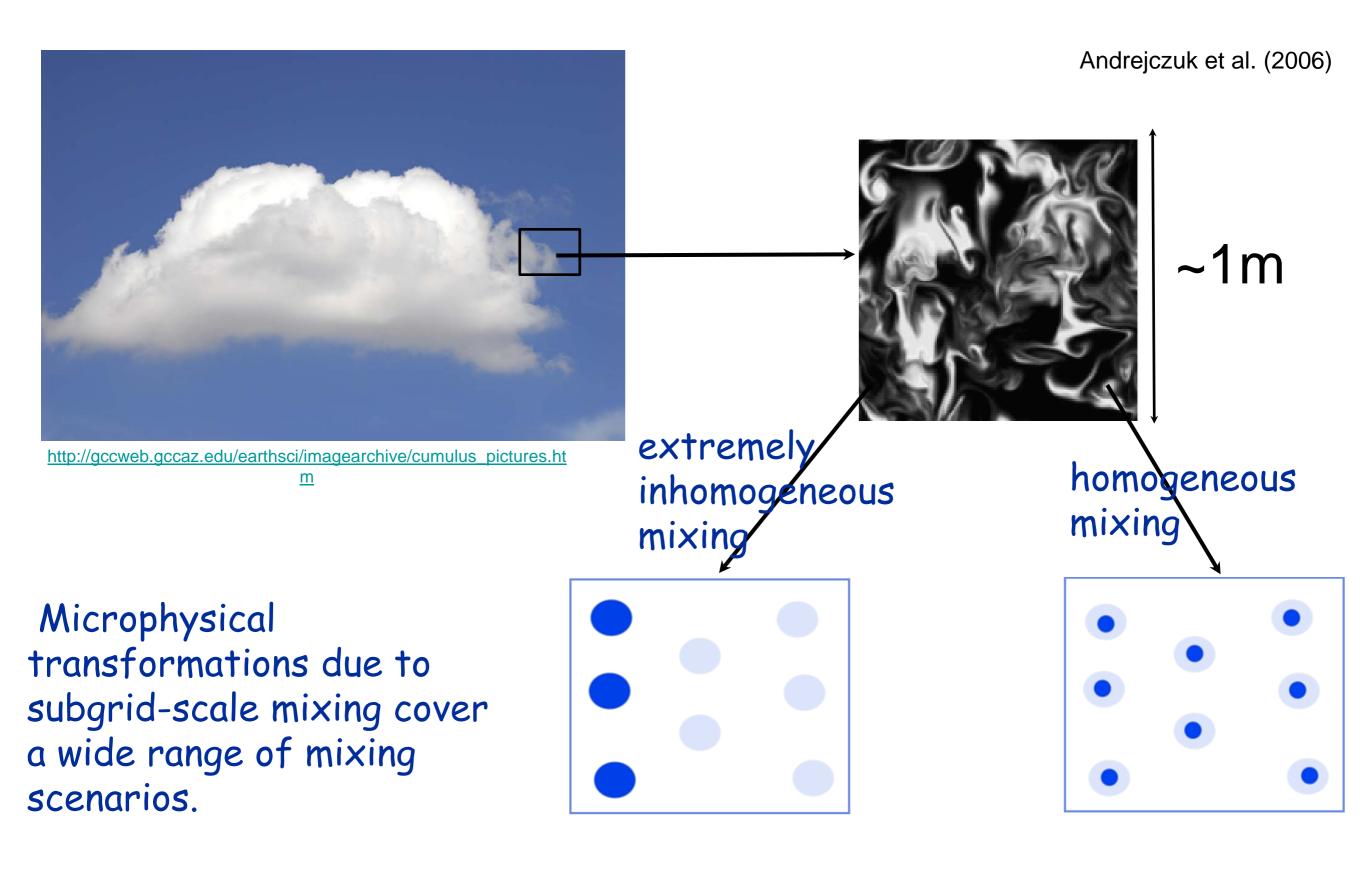
Siebesma et al. JAS 2003

http://greenearthfacts.com/weather/how-much-does-a-cloud-weigh/

QuickTime™ and a H.264 decompressor are needed to see this picture.

from Steve Krueger

Turbulent mixing in clouds



1-moment bulk microphysics

 $q_{\rm l}$ - liquid water mixing ratio, the only information about the cloud water

Cloud water can exist only in the saturated conditions.

In the bulk model, a rate of condensation, C, is calculated assuming the saturation adjustment – C^{sa}

Poor representation of the evaporation due to turbulent mixing!

1 - moment bulk microphysics with the λ -parameterization

λ - spatial scale of the cloudy filament during turbulent mixing

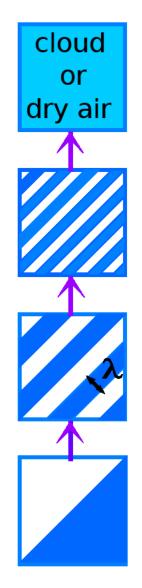
$$\frac{d\lambda}{dt} = -\gamma \epsilon^{\frac{1}{3}} \lambda^{\frac{1}{3}}$$

 Λ - the model gridlength; λ_0 - the homogenization scale (~ 1 cm).

 $\lambda_0 \leq \lambda \leq \Lambda$

 $\gamma \sim 1$ ϵ - the dissipation rate of TKE

Broadwell and Breidenthal (1982); Grabowski (2006)



Evaporation in model with 1-moment microphysics

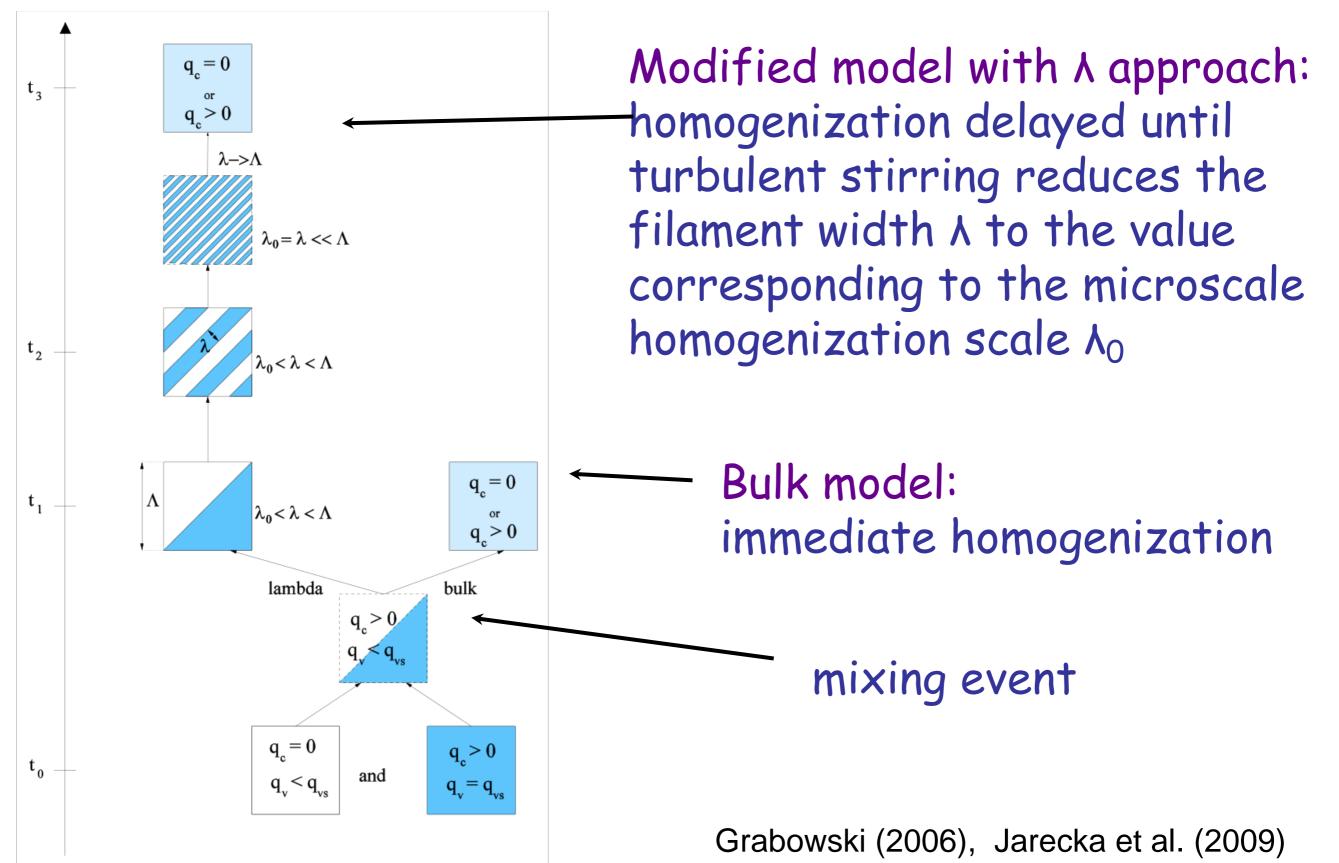
Saturation adjustment is delayed until the gridbox can be assumed homogenized:

 $\lambda = \Lambda \text{ or } \lambda \leq \lambda_0 \quad C = C^{\text{sa}} \quad \text{(saturation adjustment)}$ $\lambda_0 \leq \lambda \leq \Lambda \qquad C = \beta C^{\text{a}} \quad \text{(adiabatic C)}$

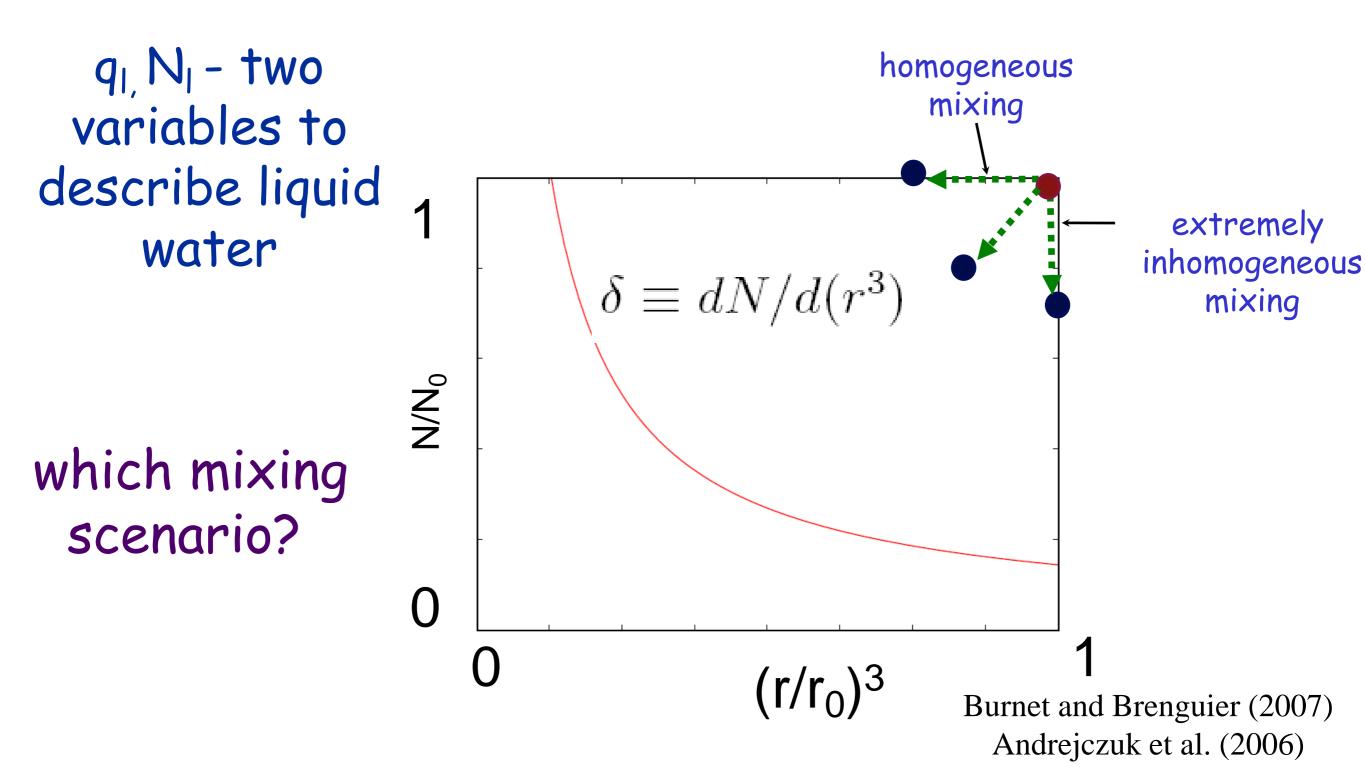
 β - fraction of the gridbox covered by cloudy air

$$C^a = -\frac{dq_{vs}}{dt}$$
 - adiabatic condensation rate

Delay in evaporation in model with 1-moment microphysics



2-moment microphysics and the mixing diagram



2-moment microphysicsmixing scenarios

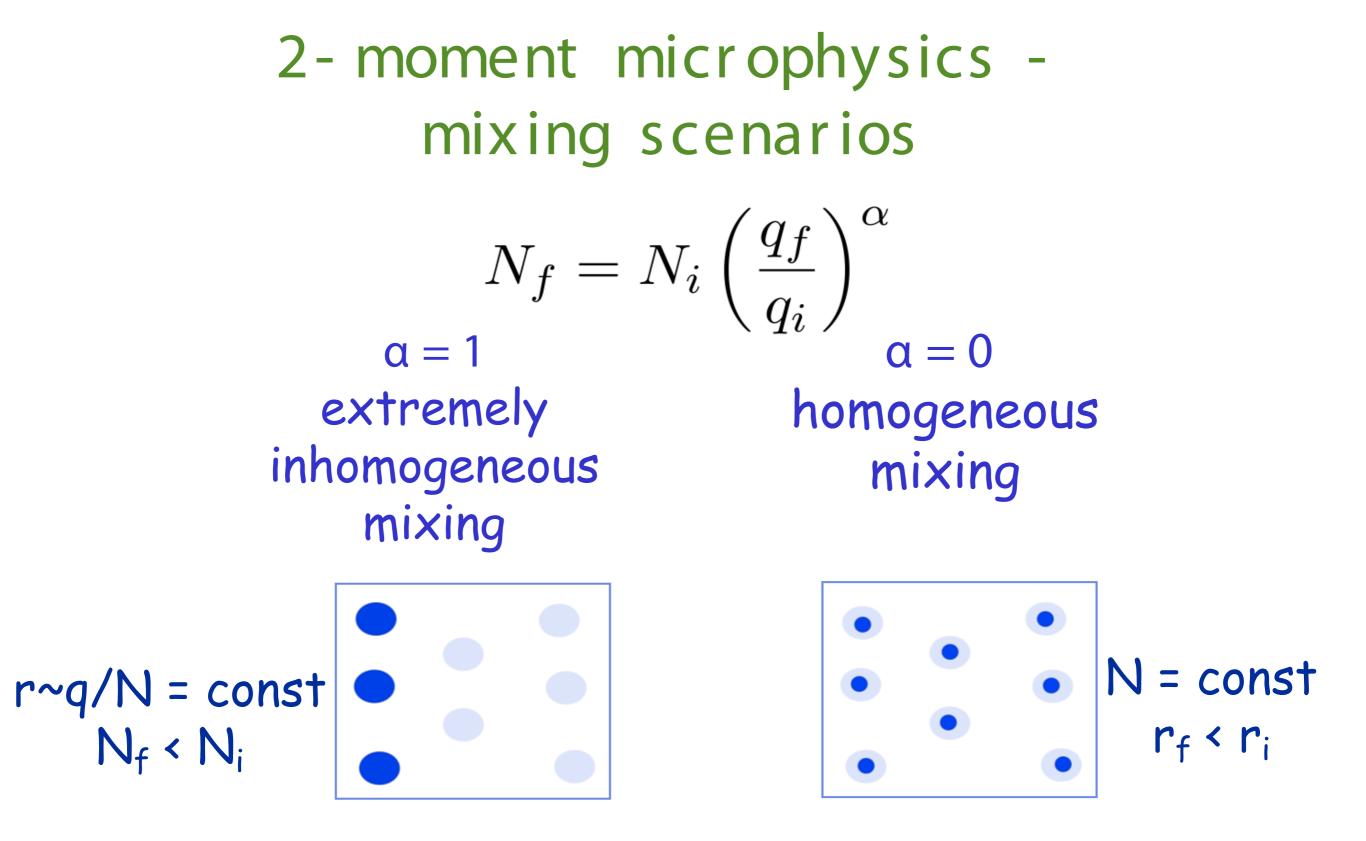
$$N_f = N_i \left(\frac{q_f}{q_i}\right)^{\alpha}$$

Morrison and Grabowski (2008)

 q_i , q_f - initial and final cloud water mixing ratios

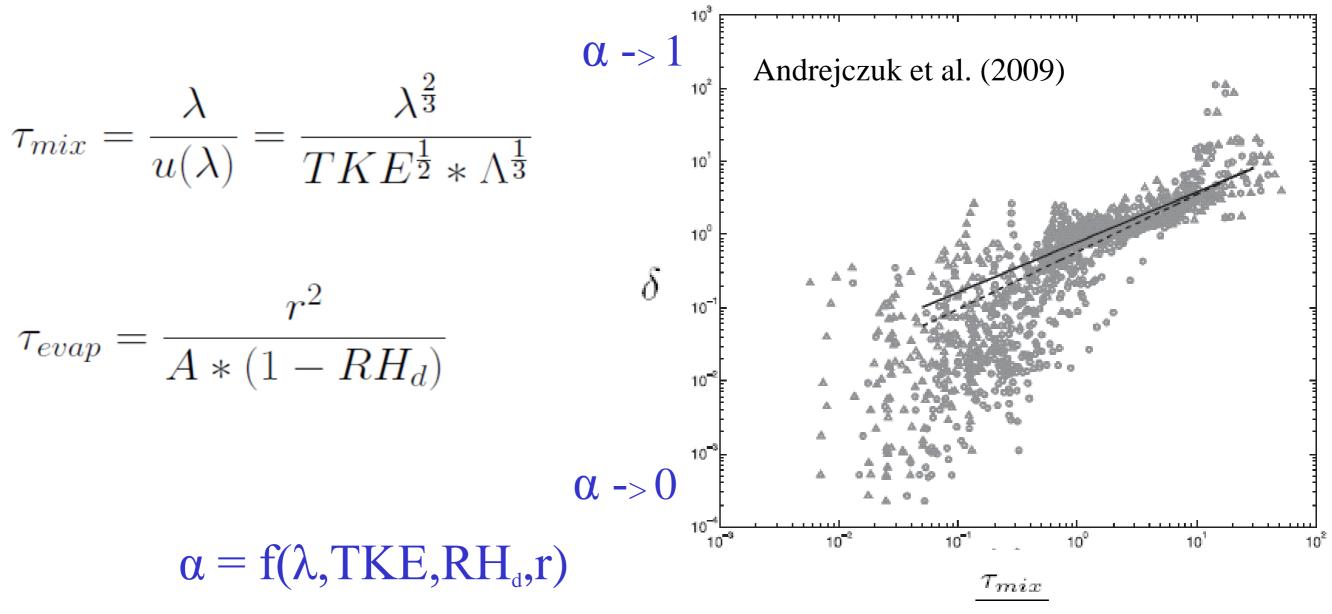
 N_i - droplet concentration after turbulent mixing, the initial value for the microphysical adjustment

 N_{f} , - final (after turbulent mixing and microphysical adjustment) value of the droplet concentration



Previous studies (Slawinska et al. 2010): α=const for entire simulation.

Using DNS results for λ -parameterization: mixing scenarios



 τ_{evap}

We can calculate α locally as a function of these parameters !!

Model and model setup 3D numerical model EULAG

www.mmm.ucar.edu/eulag/

•Eulerian version

Cartesian mesh

Anelastic form

2-moment warm-rain microphysics scheme

(Morrison and Grabowski 2007, 2008)

Simulation setup - BOMEX

•Domain: 6.4km, 6.4km, 3km

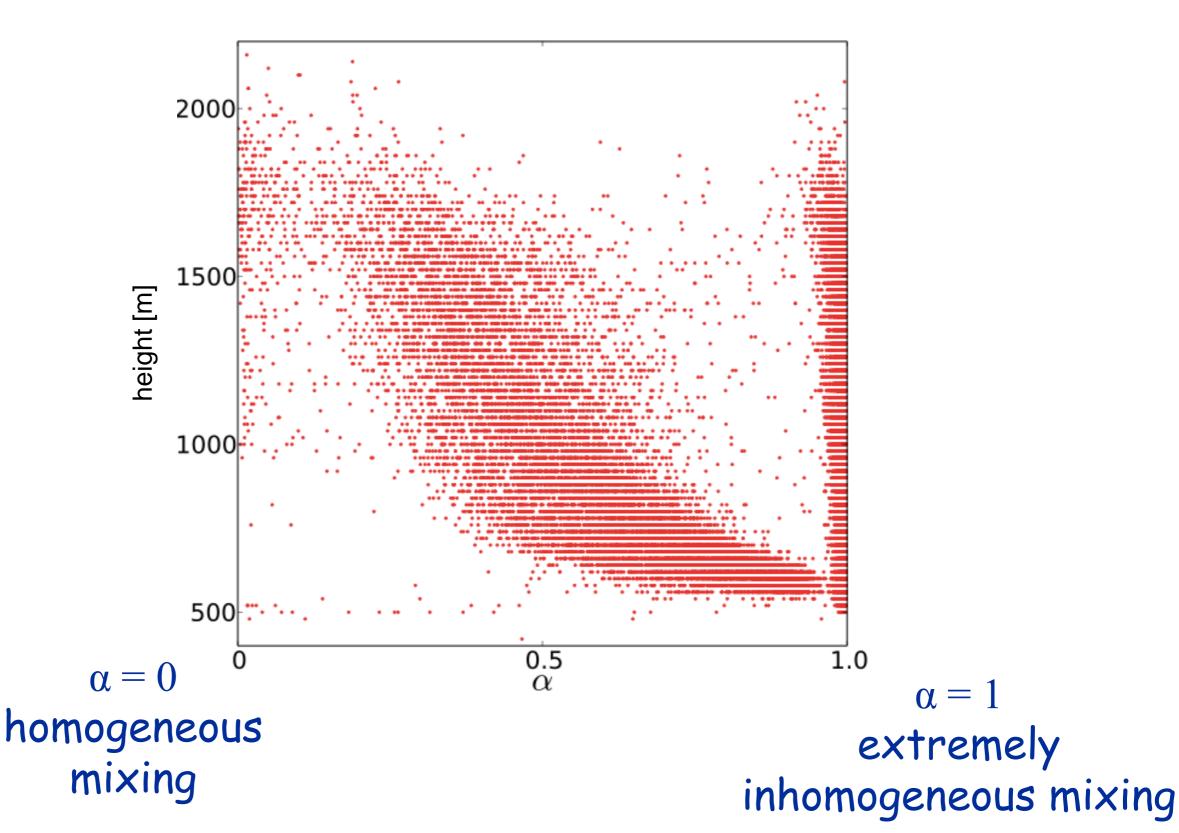
•Grid size: 50m, 50m, 20m

•Time step: 1s

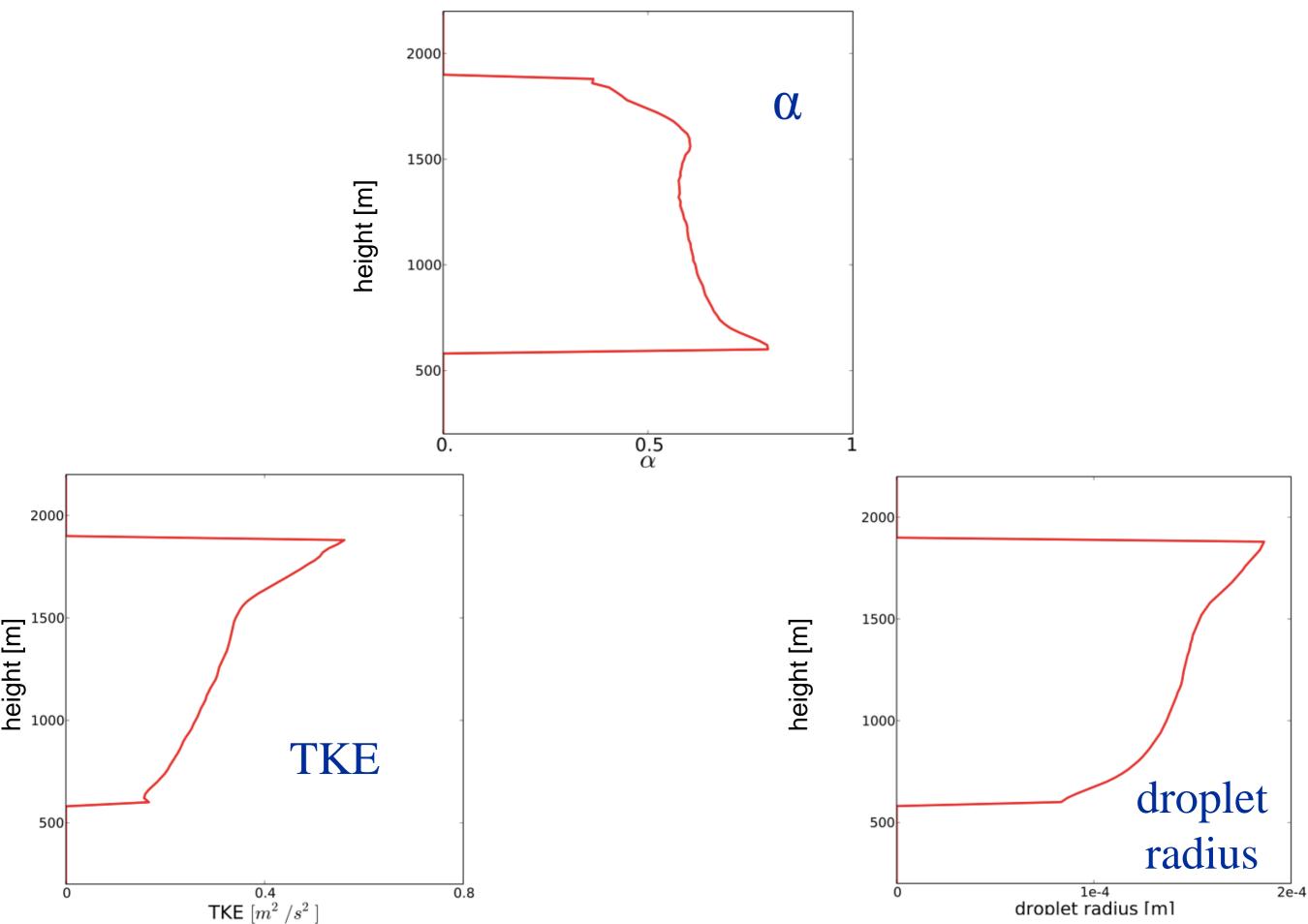
•Initial profiles from Siebesma et al. 2003

Slawinska et al., ICCP 2008

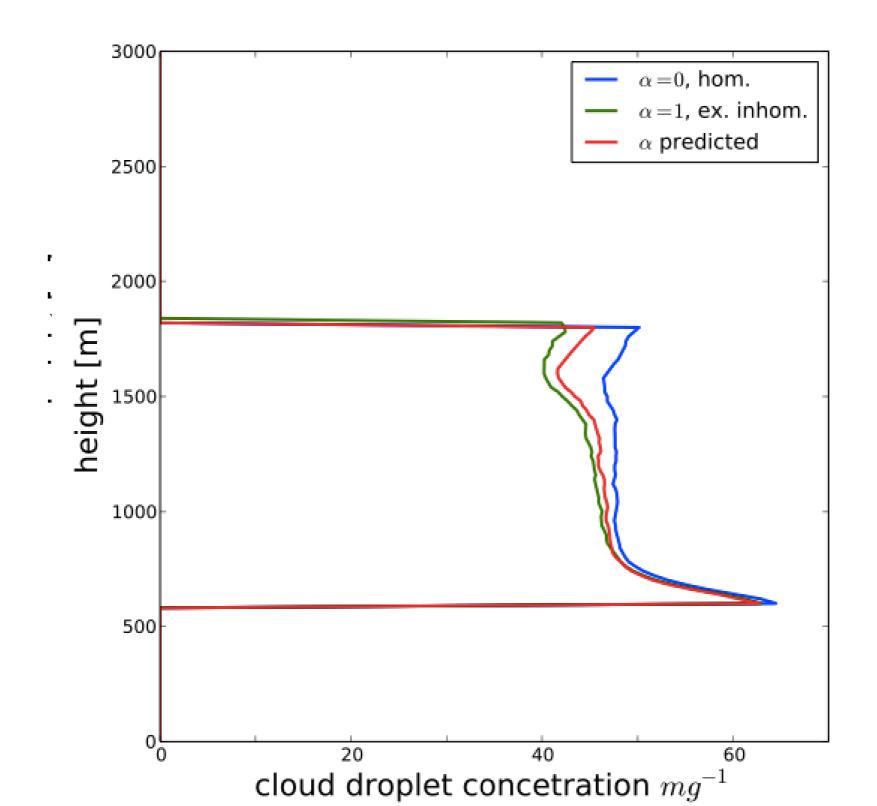
Changes of the parameter a with height



Vertical profiles of α , droplet radius and TKE



Cloud droplet concentration in simulations with various mixing scenarios



S ummary

+ Predicting scale of cloudy filaments λ allows representing in a simple way progress of the turbulent mixing between cloudy air and entrained dry environmental air.

+ Parameter α and the mixing scenario can be predicted as a function of λ , TKE, RH, and droplet radius r.

• In BOMEX simulations, α decreases with height on average, i.e., the mixing becomes more homogeneous. This is consistent with both TKE and droplet radius increasing with height.