Loughborough 2012

The direct modeling of subgrid-scale stresses and its relevance for medium range weather prediction Nils Peter Wedi*

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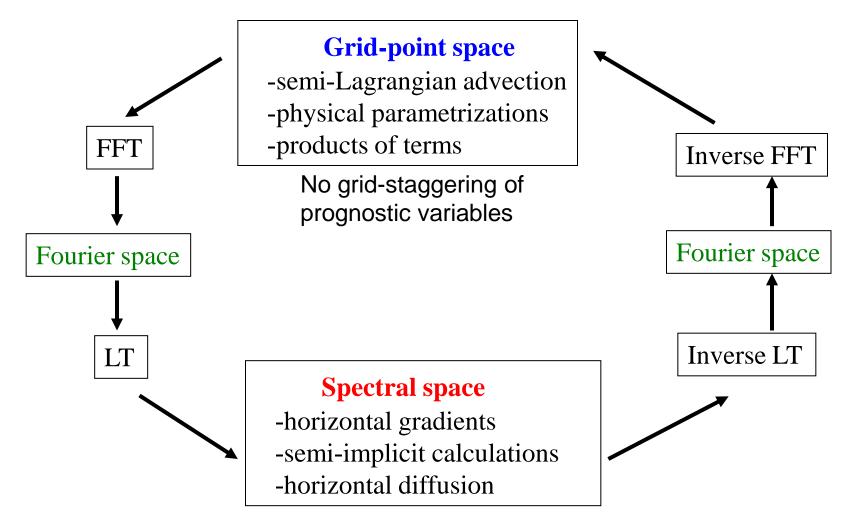


Outline

- The spectral transform method: de-aliasing of spectral computations on the linear grid
- The role of horizontal diffusion
- Non-linear diffusion

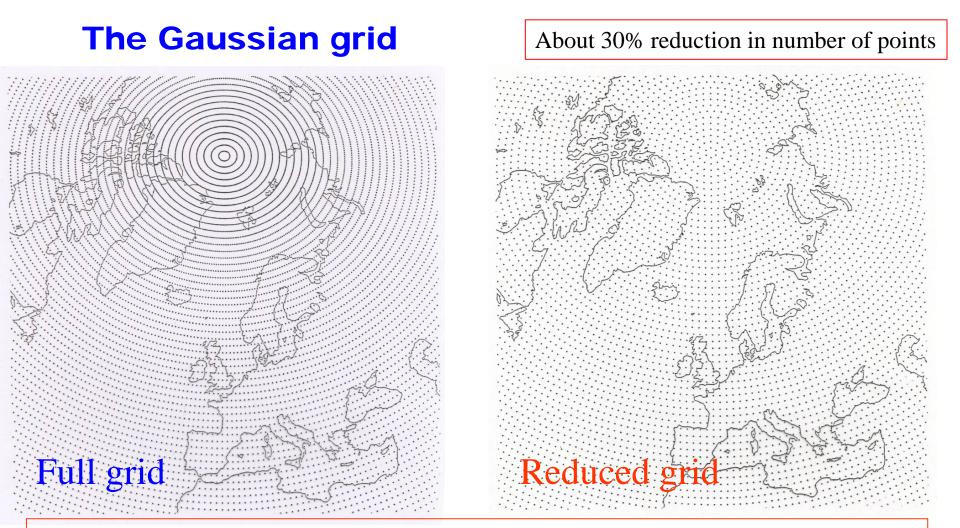


Schematic description of the spectral transform method in the ECMWF IFS model



FFT: Fast Fourier Transform, LT: Legendre Transform





Reduction in the number of Fourier points at high latitudes is possible because the associated Legendre polynomials are very small near the poles for large m.

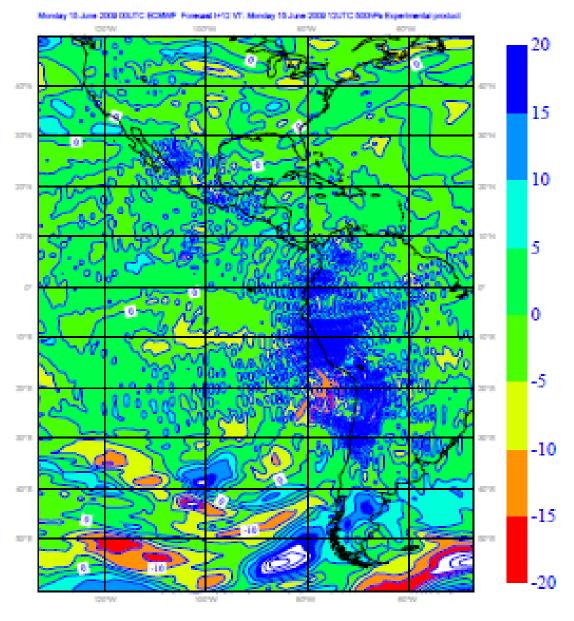


- Aliasing of quadratic terms on the linear grid (2N+1 gridpoints per N waves), where the product of two variables transformed to spectral space cannot be accurately represented with the available number of waves (as quadratic terms would need a 3N+1 ratio).
- Absent outside the tropics in E-W direction due to the design of the reduced grid (obeying a 3N+1 ratio) but present throughout (and all resolutions) in N-S direction.
- By subtracting the difference between a specially filtered and the unfiltered pressure gradient term at every time-step the stationary noise patterns can be removed at a cost of approx.
 5% at T1279 (2 extra transforms).

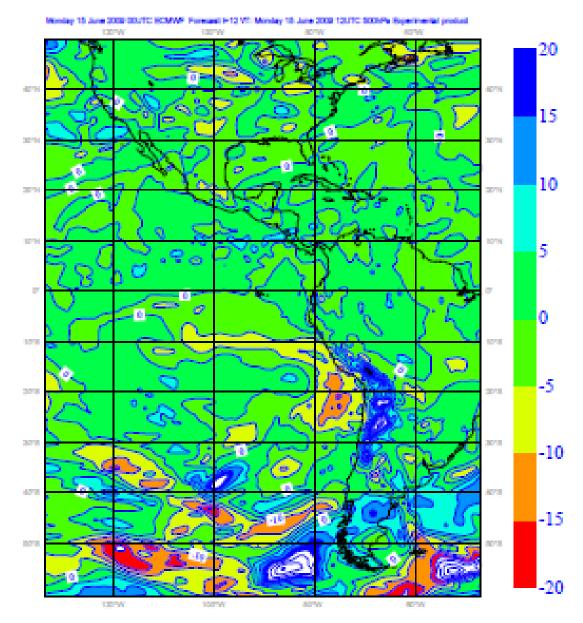


E-W

500hPa adiabatic zonal wind tendencies (T159)

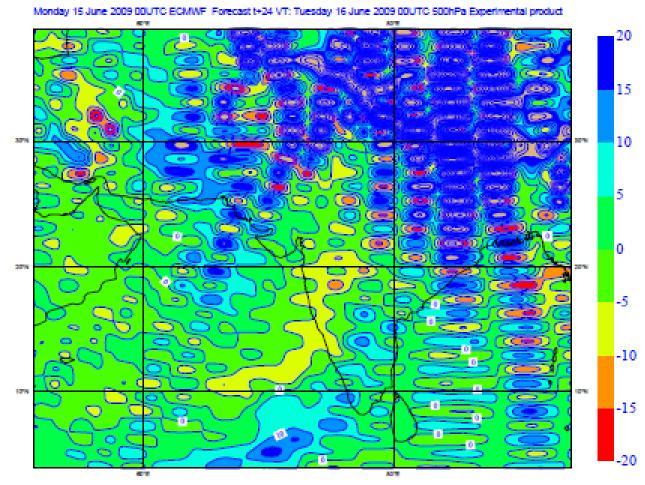




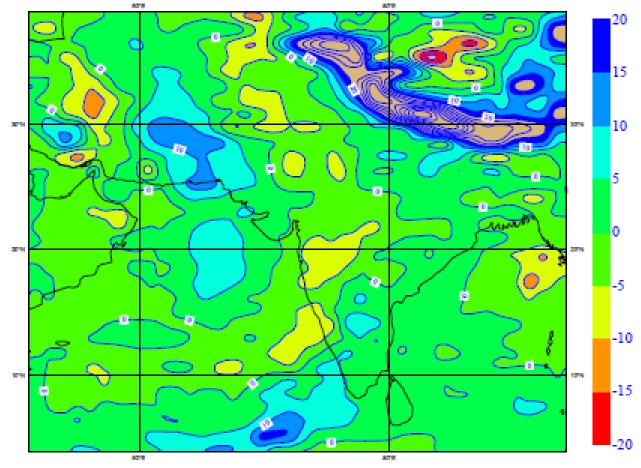


N-S

500hPa adiabatic meridional wind tendencies (T159)



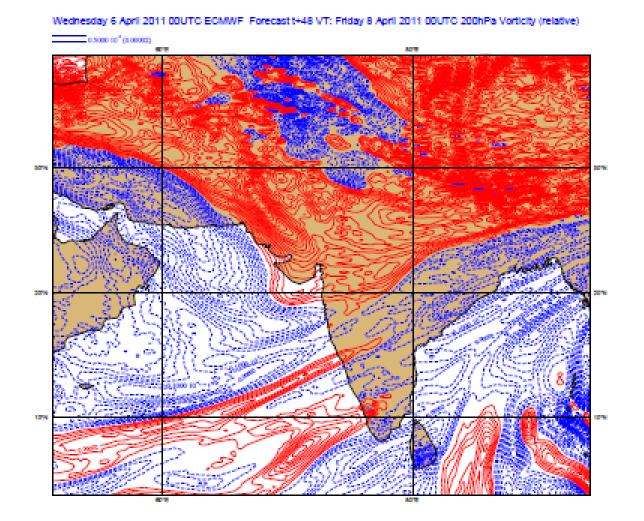




Monday 15 June 2009 00UTC ECMWF. Forecast t+24 VT: Tuesday 16 June 2009 00UTC 500hPa Experimental product.

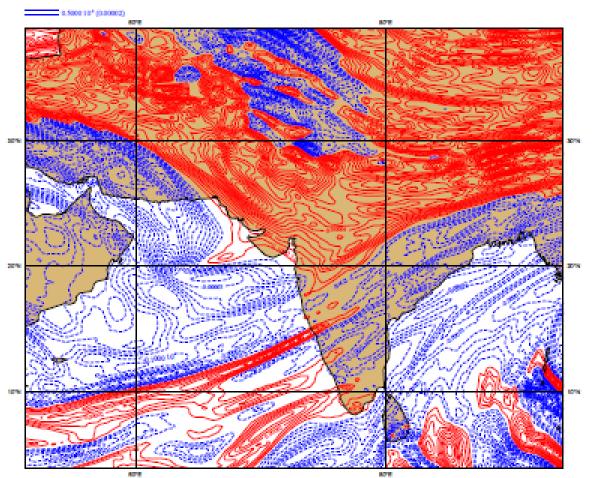


de-aliasing: vorticity noise



ECMWF

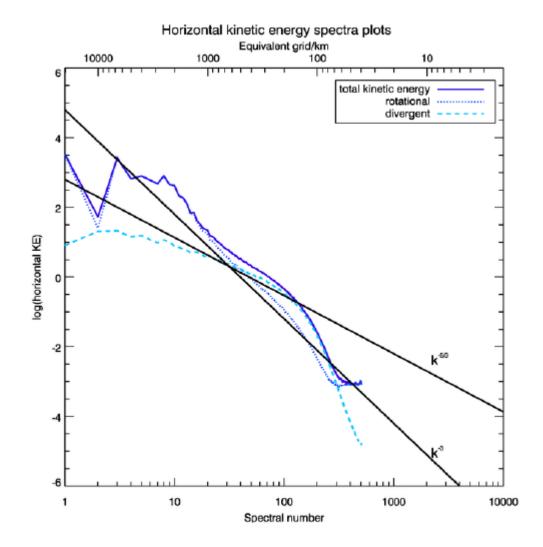
De-aliasing: vorticity noise



Wednesday 6 April 2011 00UTC ECMWF Forecast t+48 VT: Friday 8 April 2011 00UTC 200hPa Vorticity (relative)

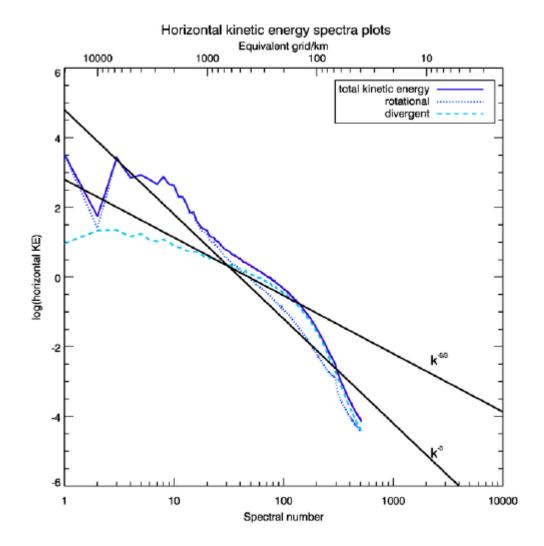


Kinetic Energy Spectra - 100 hPa





Kinetic Energy Spectra - 100 hPa





Horizontal Diffusion

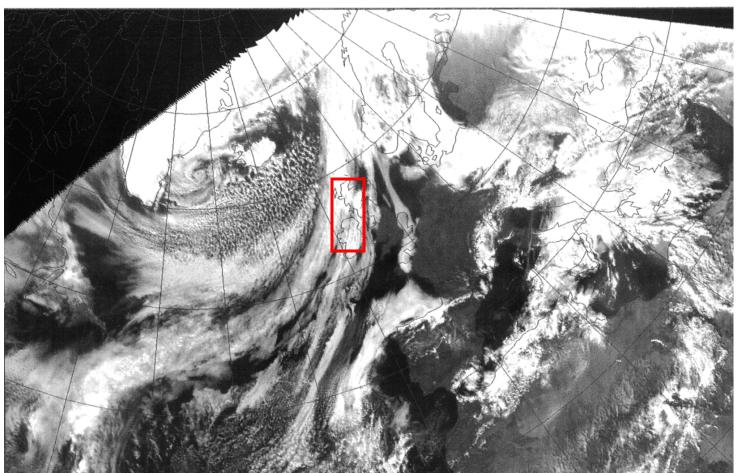
Four reasons for horizontal diffusion:

- Sponge" for vertically propagating gravity waves
- Dampen the accumulation of KE and enstrophy at the smallest resolved scales
- Represent unresolved subgrid-scale mixing
- Make tangent-linear and non-linear evolution more similar (relevant for variational data-assimilation)
- Given the removal of aliasing, horizontal diffusion can now be reduced or alternative schemes applied to represent unresolved subgrid-scale mixing
 - > Non-linear horizontal diffusion



Satellite picture of N. Atlantic

METEOSAT VIS 13 MAR 1995 15:00 UTC

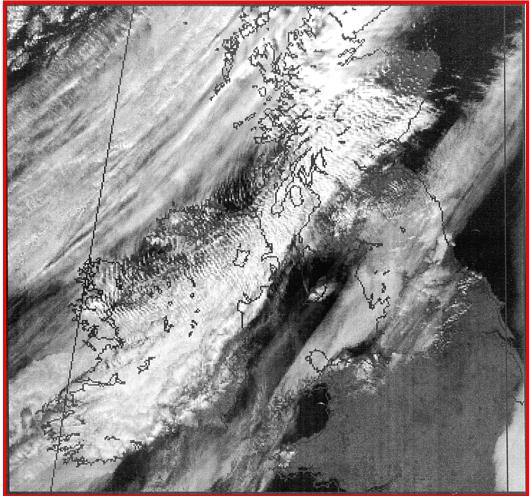


Large-Scale: Frontal systems on scale of N.Atlantic. Cellular convection S. of Iceland. 2d vortex shedding S. of Greenland.



Enlarged picture over U.K.

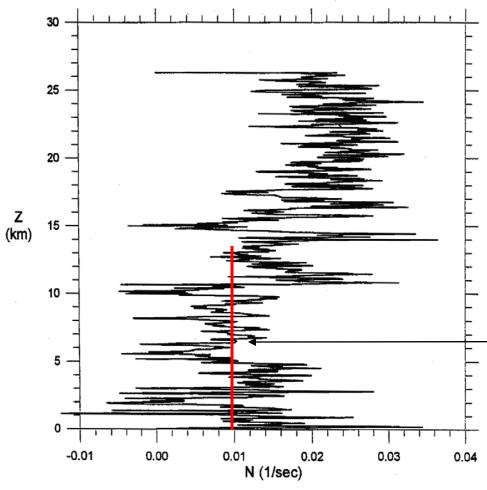
AVHRR VIS 13 MAR 1995 13:15 UTC



More detail ... Gravity-wave train running SW-NE over UK.



Balloon measurements of static stability



$$S = \frac{N^2}{g} = \frac{1}{\theta} \frac{\partial \theta}{\partial z}$$

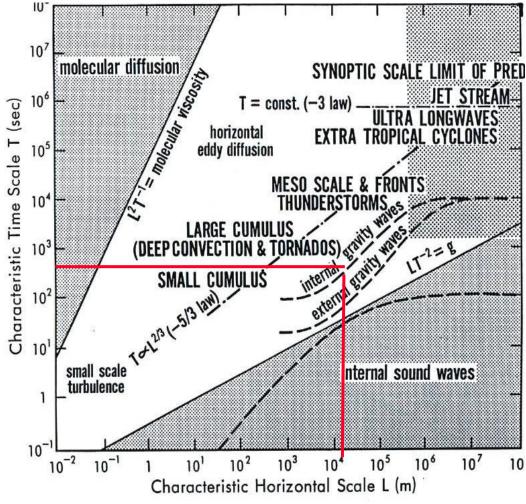
Static stability S Brunt-Väisällä frequency N Potential temperature θ

Idealized reference or initial state often assumed in idealized simulations



3rd EULAG Workshop 2012 Slide 17

Scales of atmospheric phenomena



Practical *averaging* scales do not necessarily correspond to a physical scale separation.

If equations are averaged, there may be strong interactions between resolved and unresolved scales.

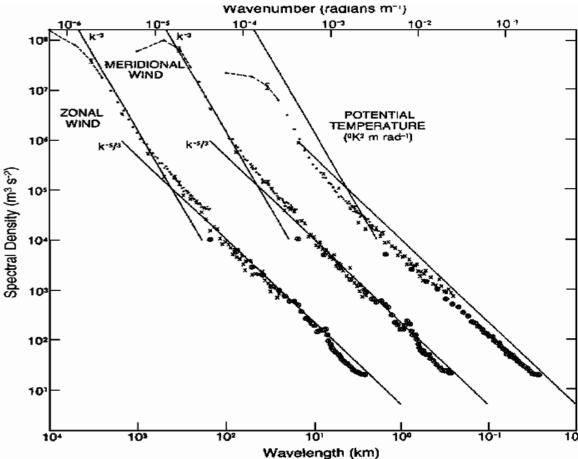


"Averaged" equations

- The equations as used in an operational NWP model represent the evolution of a space-time average of the true solution.
- The sub-grid model represents the effect of the unresolved scales on the averaged flow expressed in terms of the resolved flow variables which represent an averaged state.



(Aircraft) observed spectra of motions in the upper troposphere



Spectral slope near k^{-3} for wavelengths >500km. Near $k^{-5/3}$ for shorter wavelengths.

"It is remarkable that no scientific consensus yet ... whether the energy cascade through the mesoscale range is upscale or downscale. " (Vallgren et al., 2012)

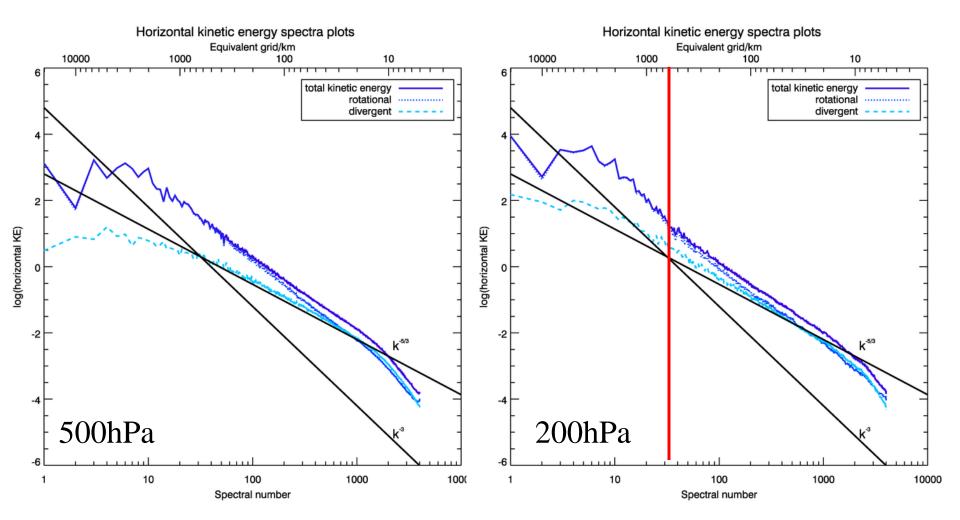


2D vs 3D turbulence

- So far horizontal correlations are simply neglected and only the vertical correlations of subgrid-scale fluctuations are parameterized.
- based on scale arguments accelerations of the larger scale flow caused by eddy motions ...
 - in 2D turbulence decrease as the scale of motion decreases (i.e. as the model resolution increases)
 - in 3D turbulence *increase* as the scale of motion decreases (i.e. as the model resolution *increases*)
- So as long as the resolved flow regime behaves like 2D turbulence the neglect of horizontal eddy fluxes in NWP is okay ... (Tennekes, 1978)

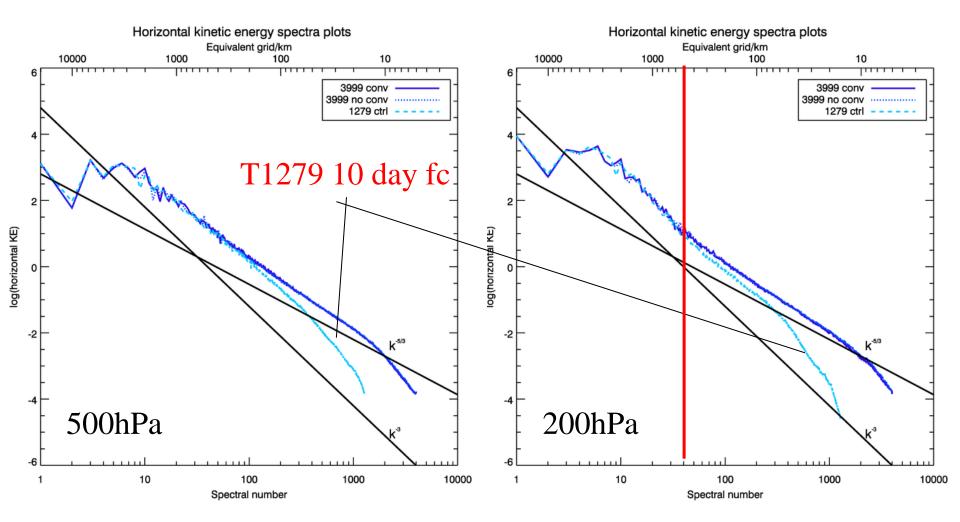


rotational/divergent energy of T3999 10 day fc





Kinetic Energy of T3999 10 day fc





Attempt to apply LES techniques to global NWP

- Take into account horizontal and vertical non-linear interactions of scales that are still represented on the grid.
- Utilize similarity of spectral slope over a wide range of wavenumbers while assuming self-similar behaviour of nonlinear interactions at the smallest resolved scales and at the largest unresolved (subgrid) scales.

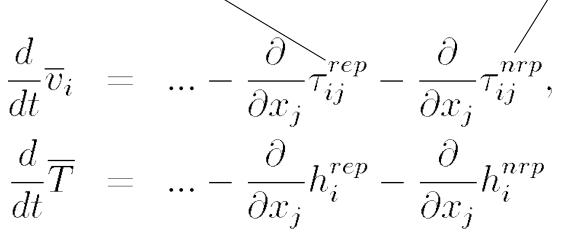


Two parts to the SGS stress tensor

Domaradzki and Adams (2002)

The result of non-linear interactions of scales actually represented on the numerical grid

truly subgrid-scale





Truly subgrid-scale == Model with spectral viscosity in IFS

e.g. Gelb and Gleeson (2001)

$$F_{\zeta_n}^{\ m}(\epsilon,\hat{q}) = -\epsilon\hat{q}^2 \left(\frac{n^2(n+1)^2}{a^4}\right)\zeta_n^m$$

$$\hat{q} = \begin{cases} 0 & n \le n_c ,\\ \exp(-(n-N)^2/(n-n_c)^2) & n_c < n \le N , \end{cases}$$
$$\epsilon = 2a^3/N^3 \text{ and } n_c = N/2,$$

In EULAG ==

ILES of MPDATA?



SGS Stress tensor - similarity model

$$\tau_{ij}^{rep} = \overline{v_i \ v_j} - \overline{v}_i \ \overline{v}_j,$$
$$\tau_{ij}(x) = \int v_i(x') v_j(x') G_{\Delta}(k, l, x - x') dx' - \overline{v}_i(x) \overline{v}_j(x).$$

$$\overline{v}_i(x) = \int v_i(x') G_{\Delta}(k, l, x - x') dx'.$$
De-convolution operation to recover the "true" velocity

$$v_i(x) = \int \overline{v}_i(x') G_{\Delta}^{-1}(k, l, x - x') dx'$$



The Gaussian filter

$$G_{\Delta}(k,l) = \exp\left\{-\frac{\Delta^2 n(n+1)}{24a^2}\right\}$$

e.g.
$$\Delta_1 = 6\Delta\lambda, \quad \Delta_2 = 8\Delta\lambda.$$

Has the same shape in physical and in wave space !



SGS Stress tensor - similarity model

Based on the idea of similar behaviour between the smallest (well) resolved scales and the actual subgrid scales

$$\tau_{ij}^{rep} = C_L L_{ij}$$
 Secondary filter
$$L_{ij} = \widetilde{\overline{v_i} \ \overline{v_j} - \overline{v_i} \ \overline{v_j}},$$

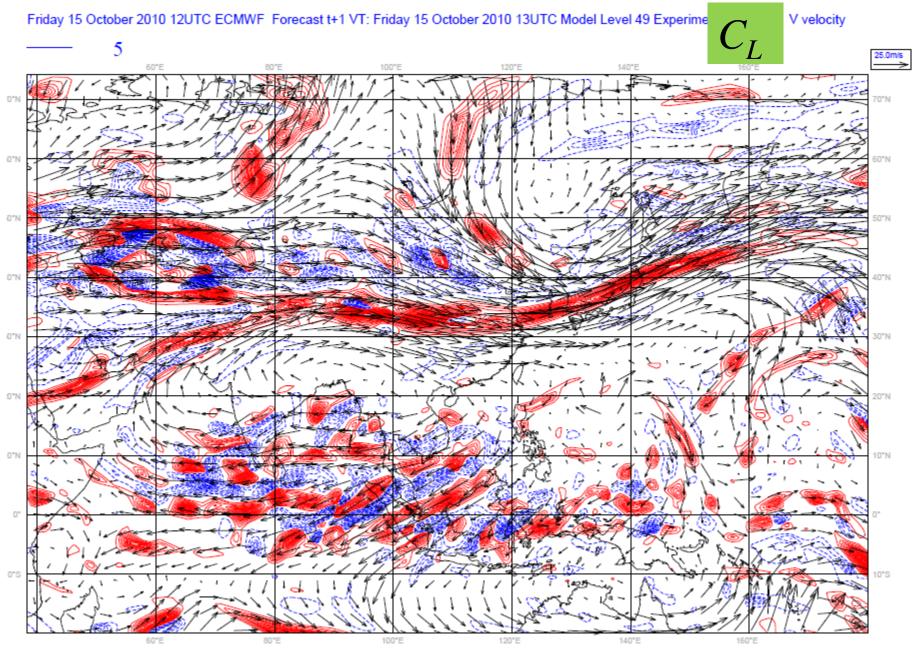
Minimize error norm (Lilly 1991)

$$f(C_L) = ||Q||_F^2 = \sum_{i,j} (\tau_{ij} - C_L L_{ij})^2.$$

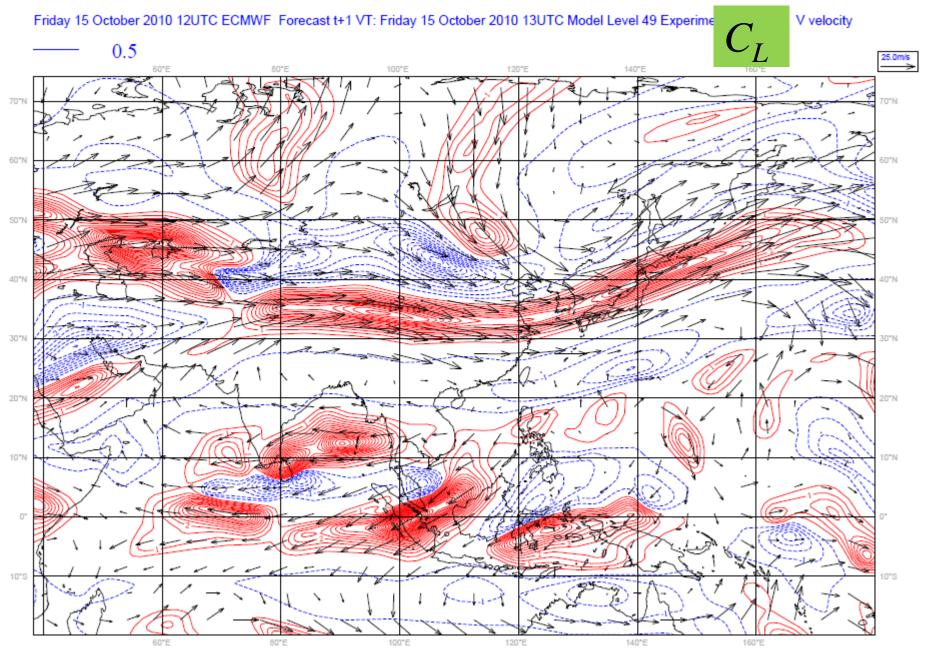
$$C_L = \frac{\sum_{i,j} \tau_{ij} L_{ij}}{\sum_{i,j} L_{ij}^2}.$$

$$\sum_{i,j} \tau_{ij} L_{ij} = \sum_{i} \tau_{ii} L_{ii} + 2 \sum_{i < j} \tau_{ij} L_{ij}.$$
$$\sum_{i,j} L_{ij}^2 = \sum_{i} L_{ii}^2 + 2 \sum_{i < j} L_{ij}^2$$

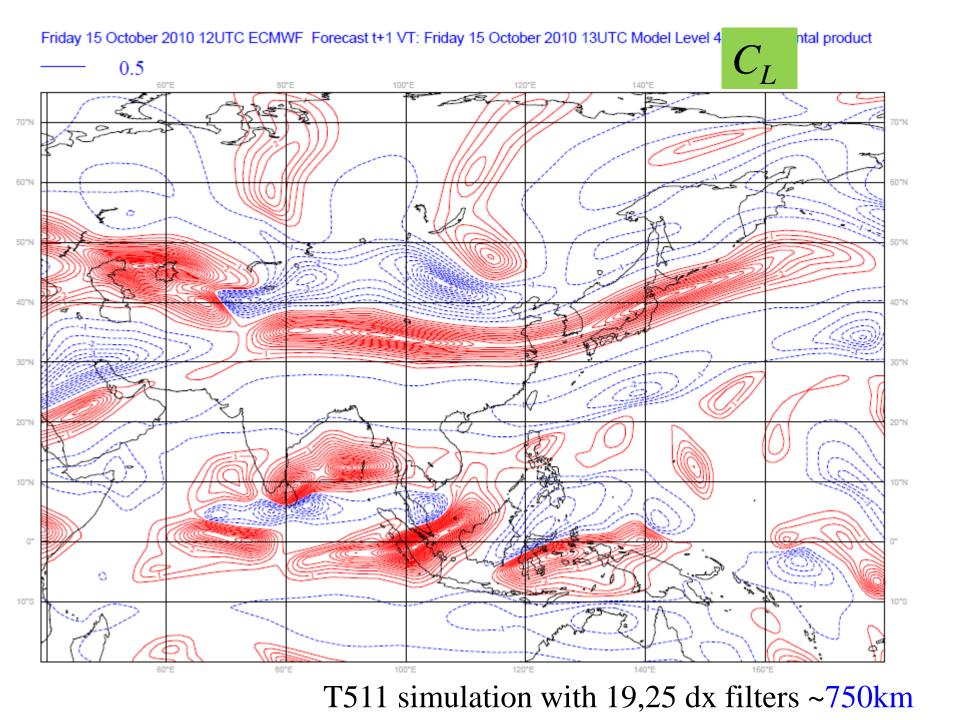




T511 simulation with 6,8 dx filters ~ 240km



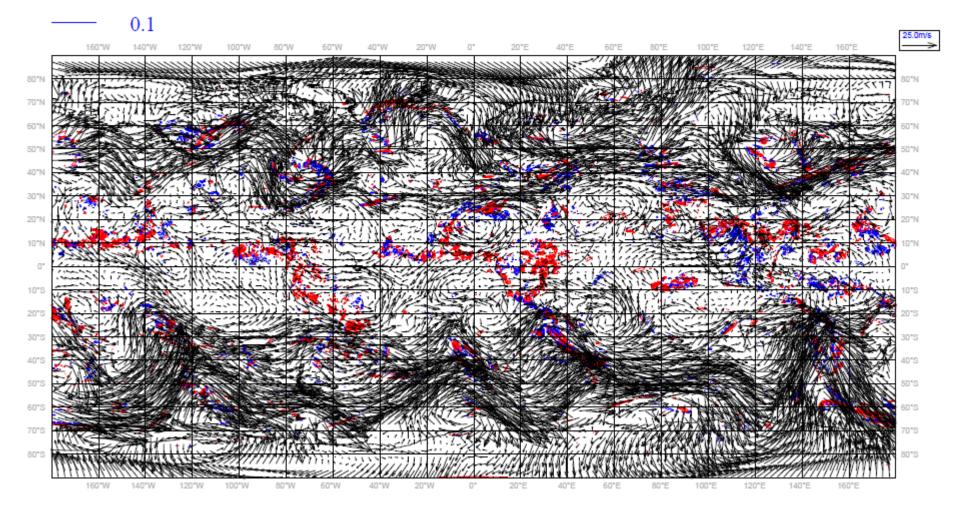
T159 simulation with 6,8 dx filters ~750km



Total physics tendency Δu

T511 simulation

Friday 15 October 2010 12UTC ECMWF Forecast t+4 VT: Friday 15 October 2010 16UTC Model Level 64 Experimental product/ V velocity

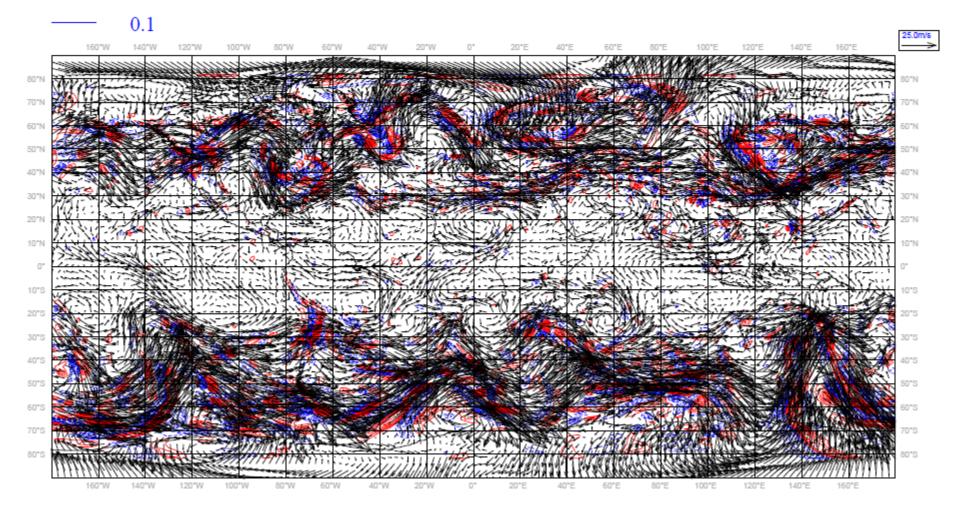




New stress **Au**

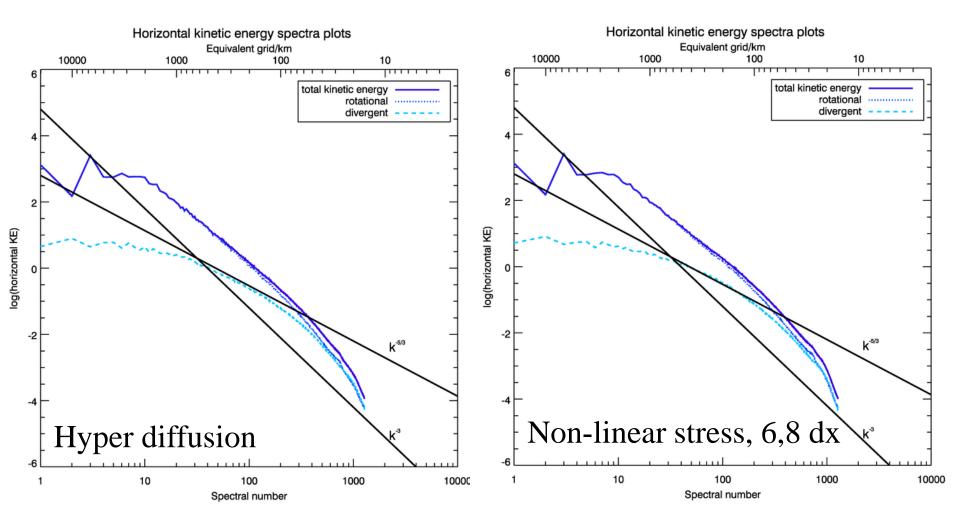
T511 simulation

Friday 15 October 2010 12UTC ECMWF Forecast t+4 VT: Friday 15 October 2010 16UTC Model Level 64 Experimental product/ V velocity





Spectra T1279



But no change in the spectra ...



SGS Stress tensor – similarity model

- The turbulent stress clearly not "random" and complimentary to physical parameterizations
- Similarity constant appears to be a property of the flow, quite independent of the model resolution while it's value is influenced by the chosen physical length-scale of the flow.
- Similarity constant naturally defines (dynamically consistent) *backscatter* (positive and negative C_L) that may be used in generating EPS perturbations.



Summary

- De-aliasing improved mass conservation by 50 %, improved energy spectra and removed unphysical noise "seen" (and compensated in part) by the physical parametrizations.
- Potential for weaker and/or alternative schemes for horizontal (hyper-)diffusion
- Potential for the use of a dynamic similarity model to compute non-linear velocity correlations, in particular in the context of ensemble prediction systems
- More work required to assess the impact of different filter choices on the resulting kinetic energy spectra



Additional slides



The final additional contributions to the IFS equations

$$\frac{du}{dt} = \dots - \left[\nabla_{\eta} \cdot \begin{pmatrix} C_{L}L_{11} \\ C_{L}L_{12} \end{pmatrix} + \frac{1}{RT} \nabla_{\eta} \phi \cdot \pi \frac{\partial}{\partial \pi} \begin{pmatrix} C_{L}L_{11} \\ C_{L}L_{12} \end{pmatrix} - \frac{g}{RT} \pi \frac{\partial}{\partial \pi} C_{L}L_{13} \right],$$

$$\frac{dv}{dt} = \dots - \left[\nabla_{\eta} \cdot \begin{pmatrix} C_{L}L_{12} \\ C_{L}L_{22} \end{pmatrix} + \frac{1}{RT} \nabla_{\eta} \phi \cdot \pi \frac{\partial}{\partial \pi} \begin{pmatrix} C_{L}L_{21} \\ C_{L}L_{22} \end{pmatrix} - \frac{g}{RT} \pi \frac{\partial}{\partial \pi} C_{L}L_{23} \right],$$

$$\frac{dT}{dt} = \dots - \left[\nabla_{\eta} \cdot \begin{pmatrix} C_{LH}H_{1} \\ C_{LH}H_{2} \end{pmatrix} + \frac{1}{RT} \nabla_{\eta} \phi \cdot \pi \frac{\partial}{\partial \pi} \begin{pmatrix} C_{LH}H_{1} \\ C_{LH}H_{2} \end{pmatrix} - \frac{g}{RT} \pi \frac{\partial}{\partial \pi} C_{LH}H_{3} \right]$$



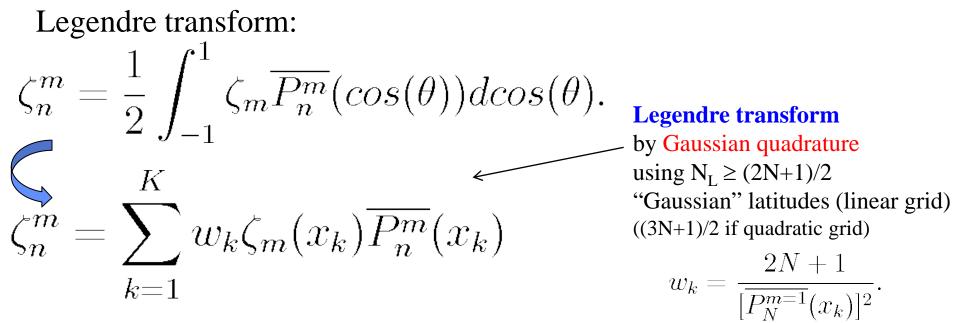
Direct spectral transform

Fourier transform:

$$\zeta_m(\theta) = \frac{1}{2\pi} \int_0^{2\pi} \zeta(\lambda, \theta) e^{-im\lambda} d\lambda$$

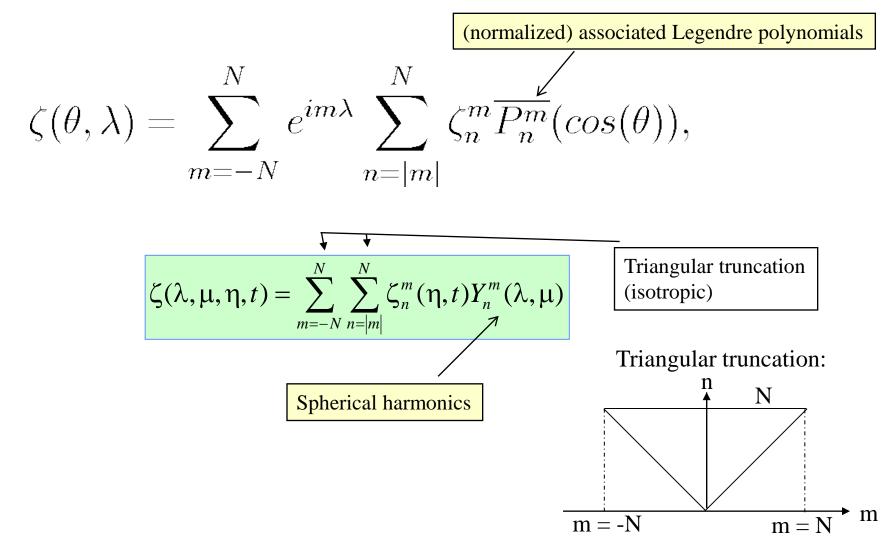
FFT (fast Fourier transform)

using $N_F \ge 2N+1$ points (linear grid) (3N+1 if quadratic grid)





Inverse spectral transform



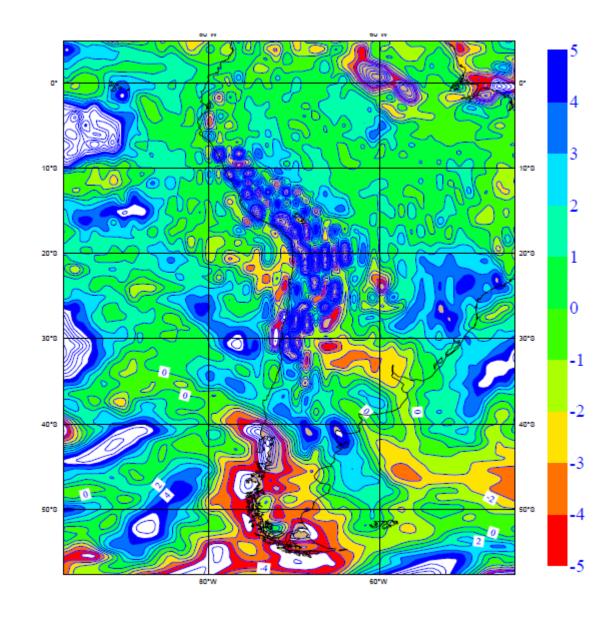


Heat

$$h_{i} = C_{LH}H_{i}.$$
$$H_{i} = \widetilde{\overline{v}_{i}}\overline{\overline{T}} - \widetilde{\overline{v}_{i}}\widetilde{\overline{T}},$$
$$C_{LH} = \frac{\sum_{i}h_{i}H_{i}}{\sum_{i}H_{i}^{2}},$$

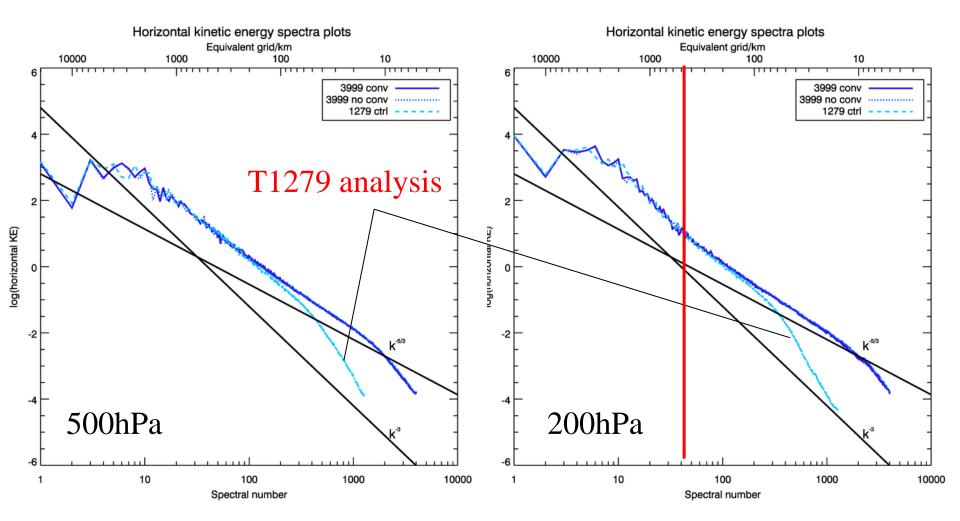


Temperature (Control)





Kinetic Energy of T3999 10 day fc





Temperature variance

Horizontal power spectra of Temperature

Horizontal power spectra of Temperature

