

Mountain wave trapping on the tropopause inversion layer revealed by idealized numerical simulations and evaluated by mid-latitude radiosonde measurements

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**DFG Research Group: MSGwaves
Multiscale Dynamics of Gravity
Waves) 2014-2020**

**BMBF Research Initiative: ROMIC (Role
of the Middle atmosphere In Climate)
2014 -2017**



Knowledge for Tomorrow

Reference:
PhD-Thesis S. Gisinger
LMU 2018



Gravity waves?



“While we cannot see atmospheric gravity waves, we can see the effects the waves have on the atmosphere.”
(Nappo, 2002)



(a) rotor clouds, (b) lenticularis clouds and (d) small scale wave structures in cirrus clouds, (c) polar stratospheric clouds, (d) noctilucent clouds

Why to care about gravity waves?

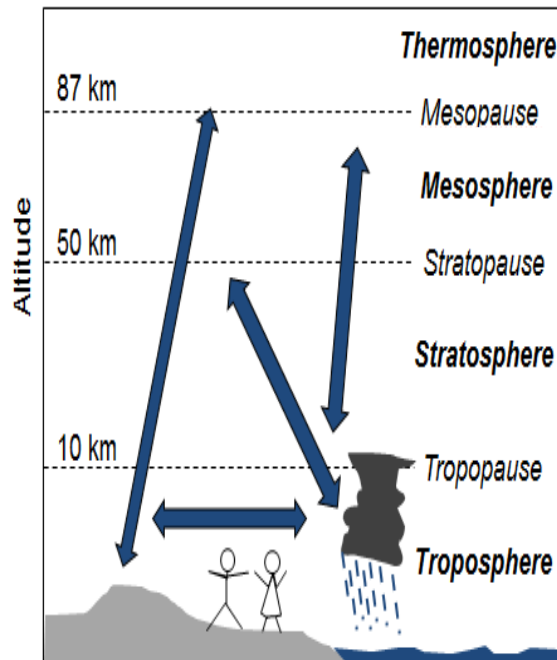
- hazard to aviation (**clear air turbulence**)
- horizontal and vertical distribution of momentum in the atmosphere (**coupling**)
- **global circulation** (without gravity waves atmospheric circulation models do not get the correct circulation patterns, mesospheric wind reversal)
- horizontal spatial scales from tens to hundreds of kilometers → **parameterizations** needed in weather and climate models
- models have deficits assigned to **missing gravity wave drag**, e.g., southern hemispheric cold pole bias

before



Photo: Ken Meiris DC8 über Evergreen, 9.12.1992

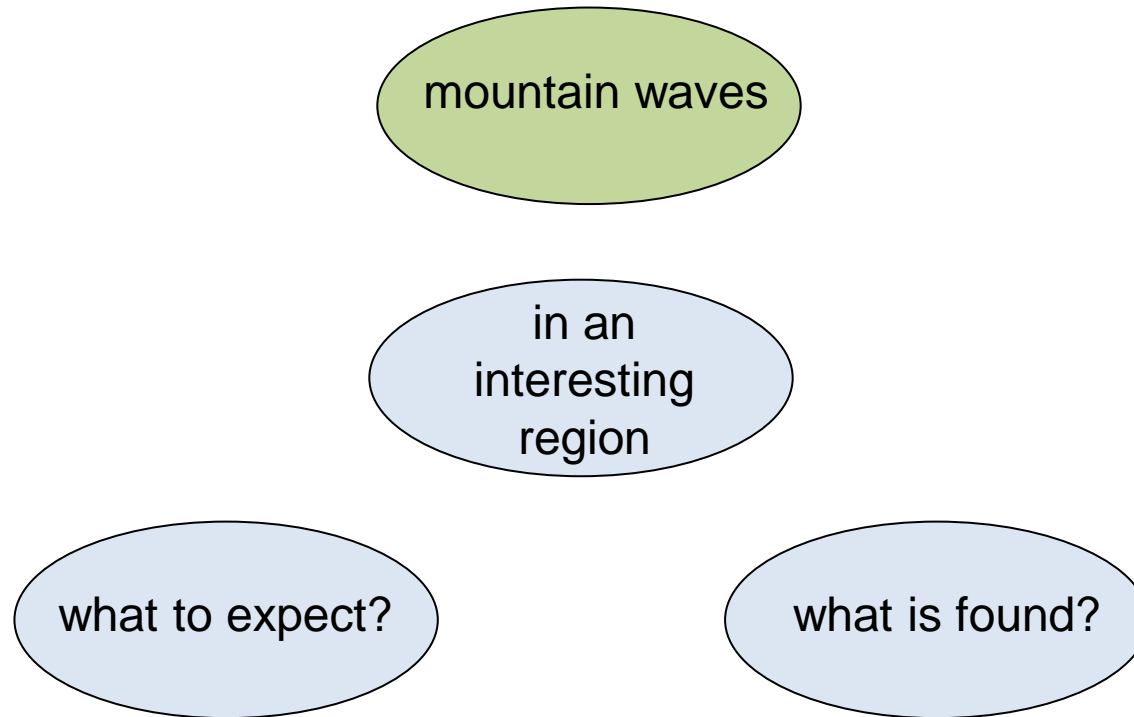
after



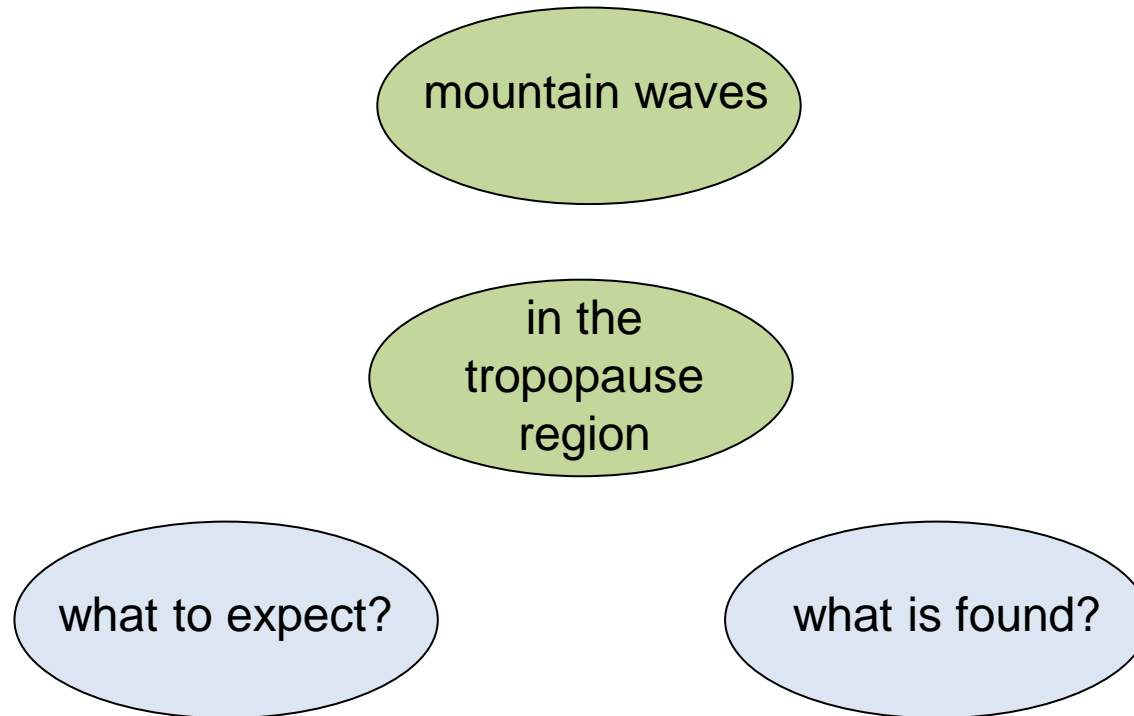
Strategy to increase the knowledge about gravity waves and their behaviour in the real atmosphere



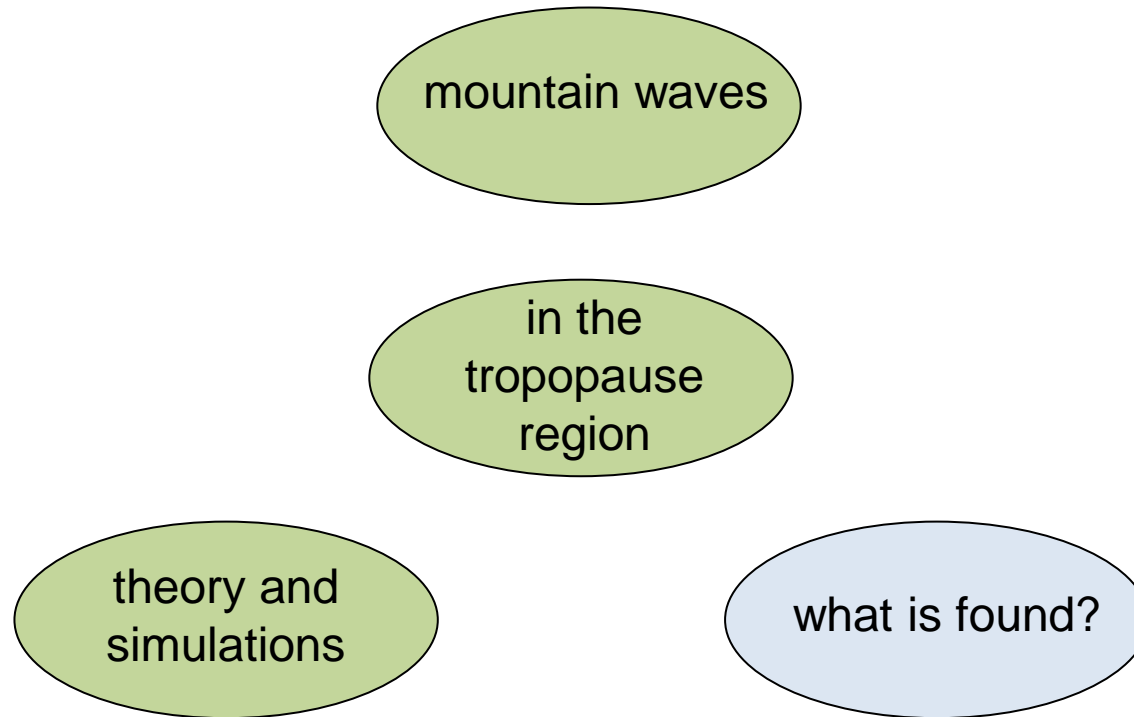
Strategy to increase the knowledge about gravity waves and their behaviour in the real atmosphere



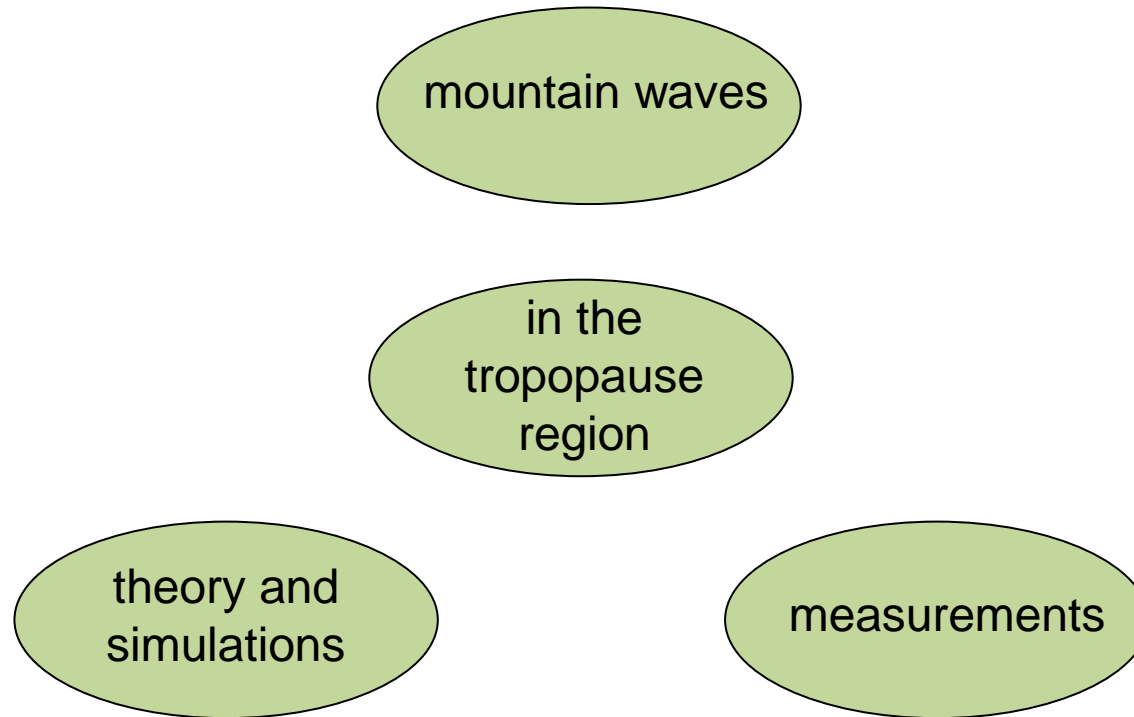
Strategy to increase the knowledge about gravity waves and their behaviour in the real atmosphere



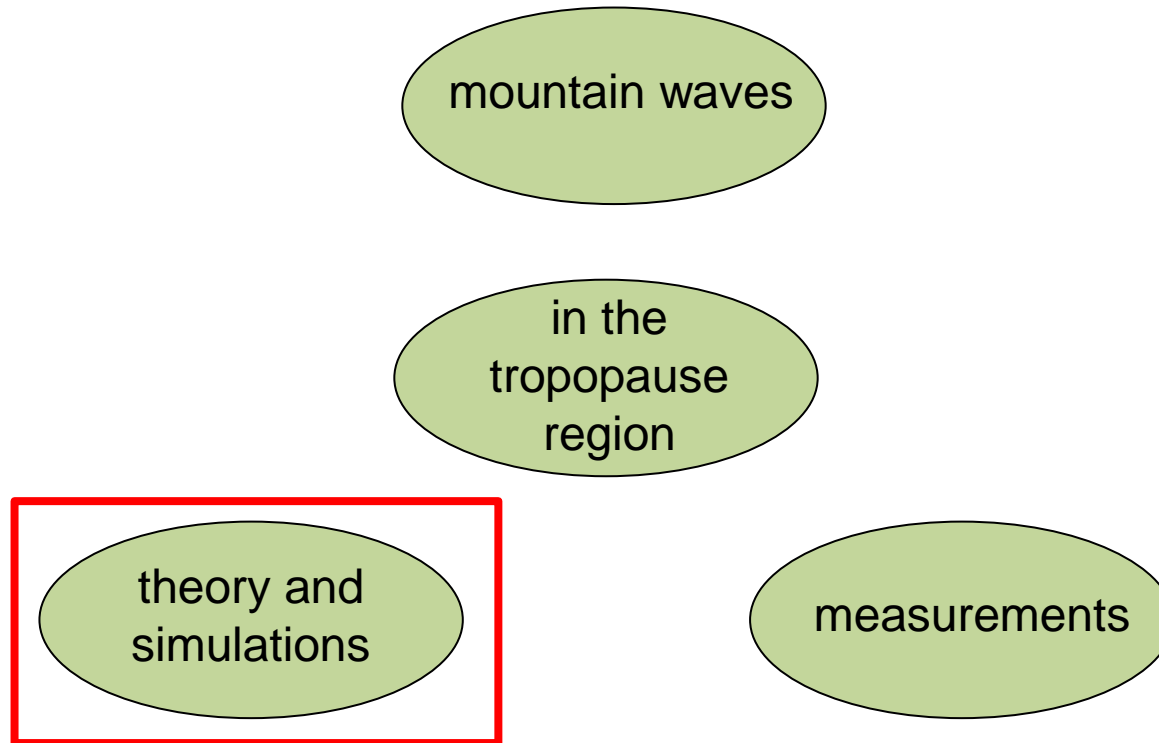
Strategy to increase the knowledge about gravity waves and their behaviour in the real atmosphere



Strategy to increase the knowledge about gravity waves and their behaviour in the real atmosphere

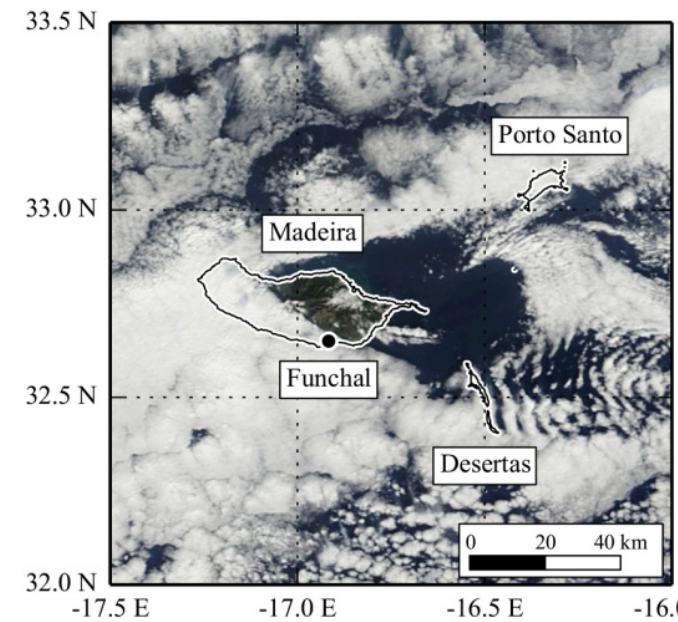
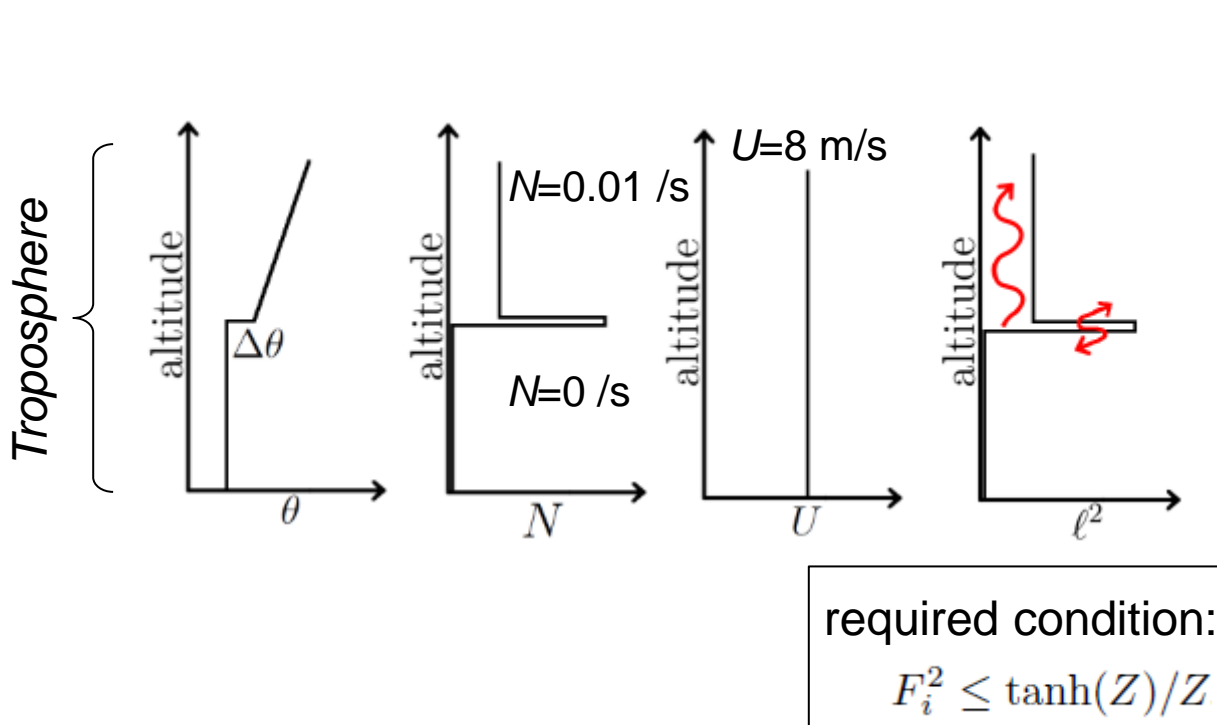


Strategy to increase the knowledge about gravity waves and their behaviour in the real atmosphere



Mountain wave propagation: trapping on an inversion

- an **inversion in the troposphere** (e.g., at the top of the boundary layer) can be a waveguide for trapped waves (**interfacial waves**) which propagate downwind of the mountain (Vosper 2004, Sachsperger et al. 2015)



Sachsperger et al. 2015

occurrence of interfacial waves depends on **strength of inversion $\Delta\theta$**

$$F_i = \frac{U}{\sqrt{gz_i \Delta\theta/\theta}}$$

$$Z = \frac{Nz_i}{U}$$

Mountain wave propagation: trapping on an inversion

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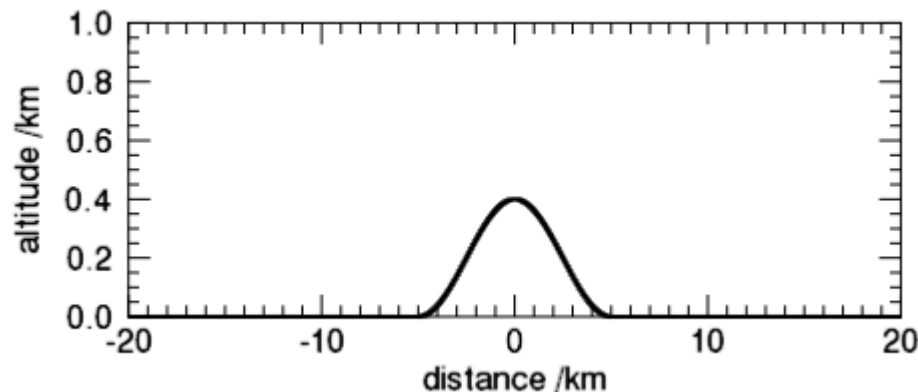
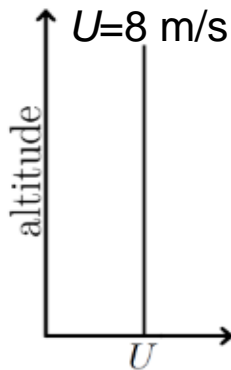
EULAG Setup

- 2D with 1032 x 2000 grid points
- $\Delta x = 100$ m and $\Delta z = 10$ m
- Δt is set to 1 s
- incompressible Boussinesq approximation
- inviscid
- TKE subgrid-scale model

- idealized ridge:
$$h(x) = \begin{cases} h_0[1 + \cos(Kx)]/2 & \text{for } |x| \leq \pi/K \\ 0 & \text{for } |x| > \pi/K, \end{cases}$$
- $U(z)=\text{const.}=8$ m/s

$$K = 2\pi/L$$

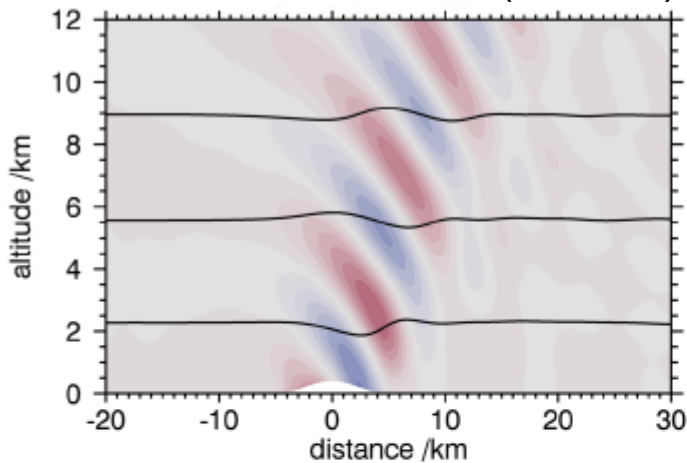
h_0 is set to 400 m and width L to 10 km or 5 km



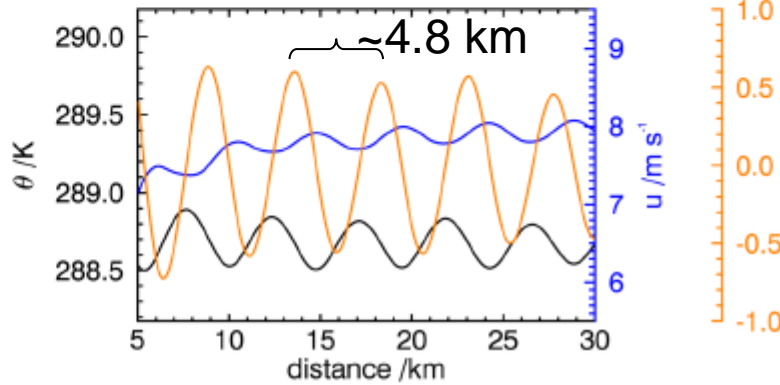
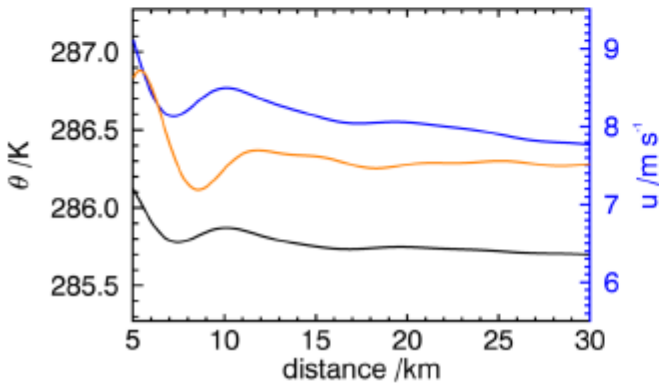
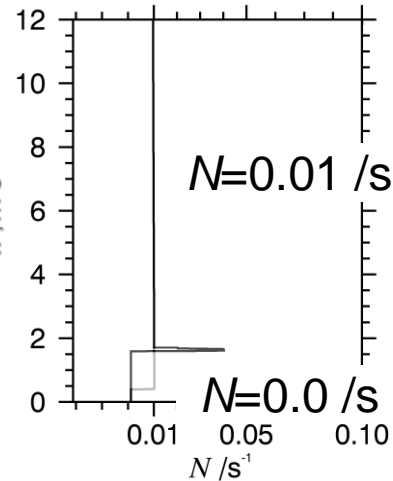
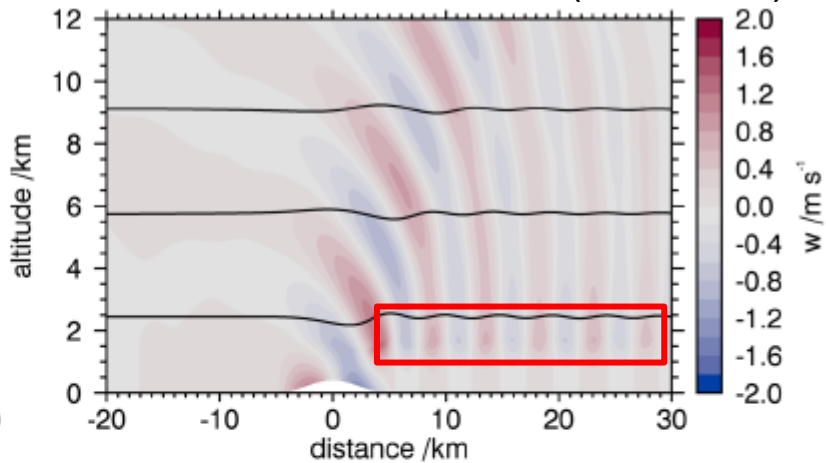
Mountain wave propagation: trapping on an inversion

- an **inversion in the troposphere** (e.g., at the top of the boundary layer) can be a waveguide for trapped waves (**interfacial waves**) which propagate downwind of the mountain (Vosper 2004, Sachsperger et al. 2015)

no inversion ($\Delta\theta=0$ K)



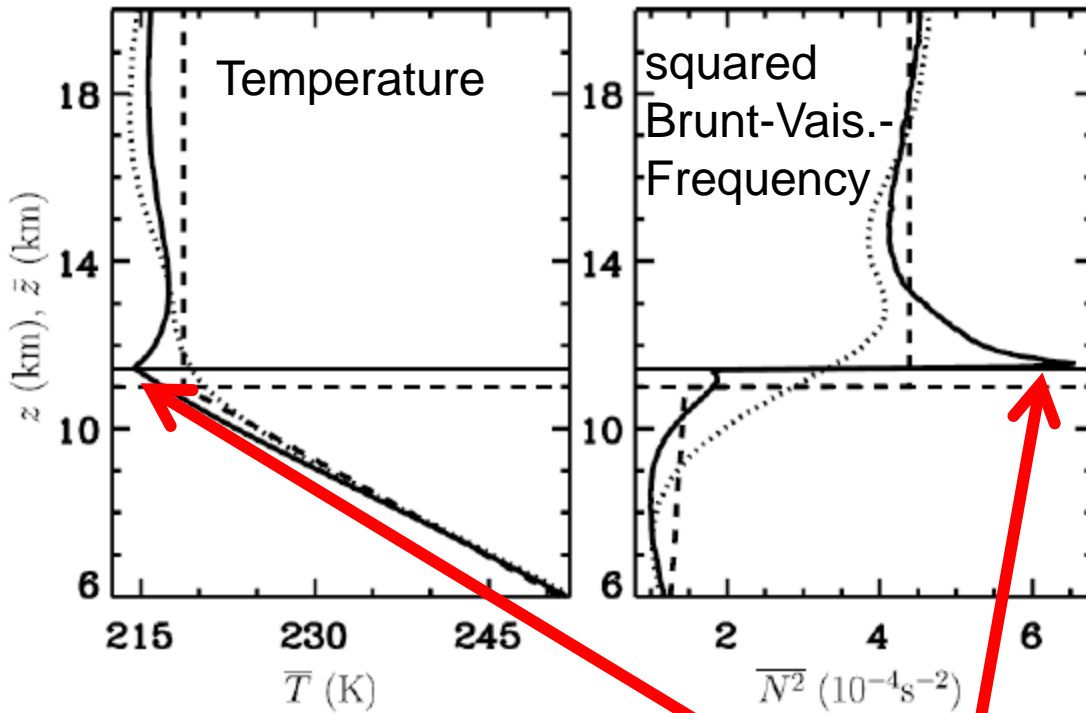
inversion at 1.6 km ($\Delta\theta=3.3$ K)



potential
temperature,
horizontal and
vertical wind
at 2 km

EULAG
Large Eddy Simulation

Mountain wave propagation: trapping on the TIL



Differences between tropospheric inversion and TIL:

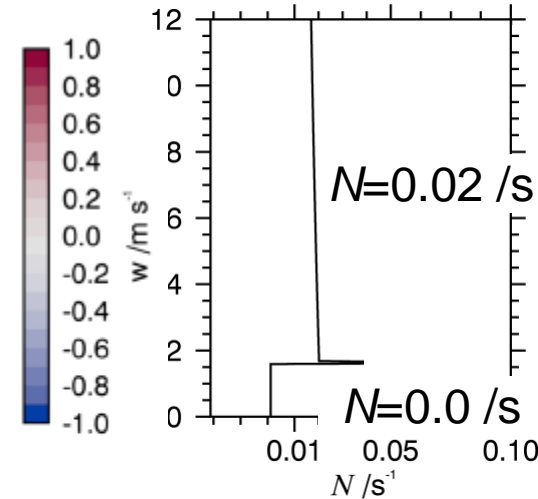
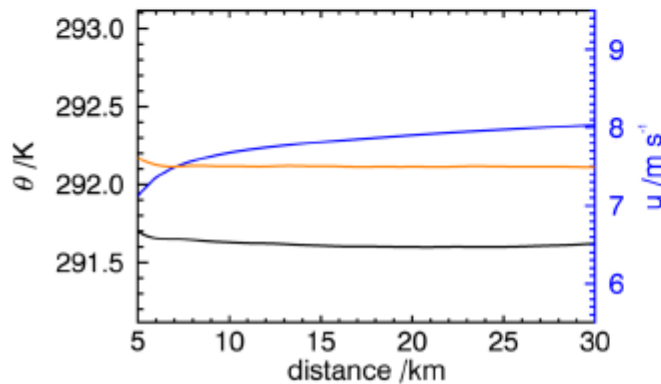
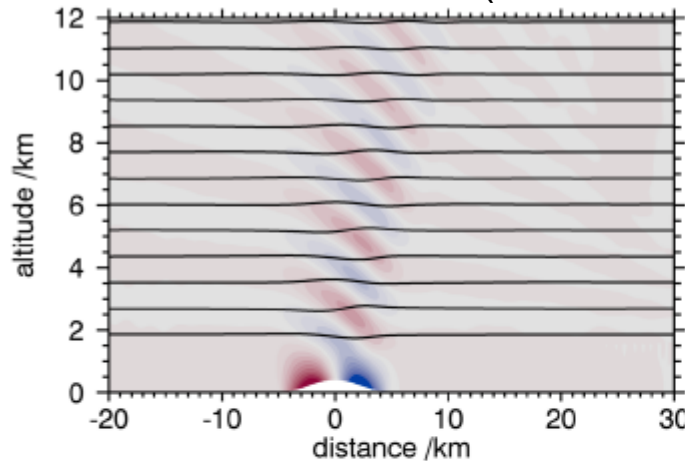
- N above inversion is 2 times larger
- inversion is at **higher altitude**
- **stable stratification below TIL** in the troposphere

Tropopause Inversion Layer (TIL)

Mountain wave propagation: trapping on the TIL

- Is concept of trapping on an inversion in the troposphere (Vosper 2004, Sachsperger et al. 2015) applicable for the **TIL**? ($N=0.02$ /s)

inversion at 1.6 km ($\Delta\theta=3.3$ K)



potential
temperature,
horizontal and
vertical wind
at 2 km

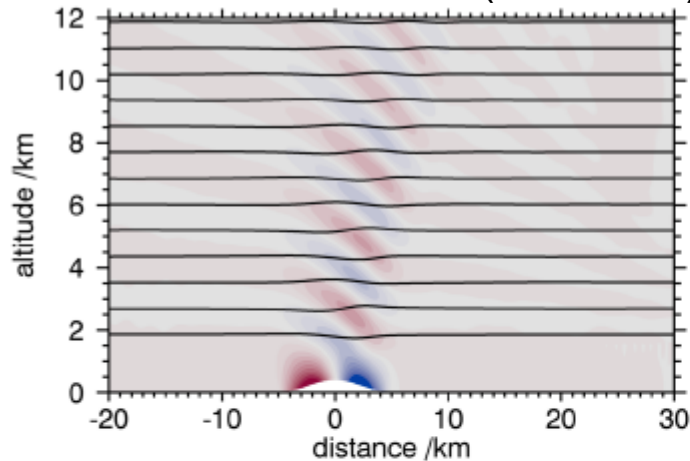
EULAG
Large Eddy Simulation

Mountain wave propagation: trapping on the TIL

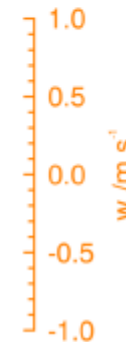
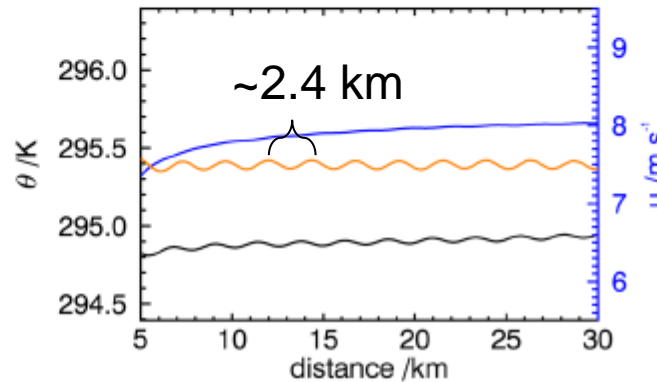
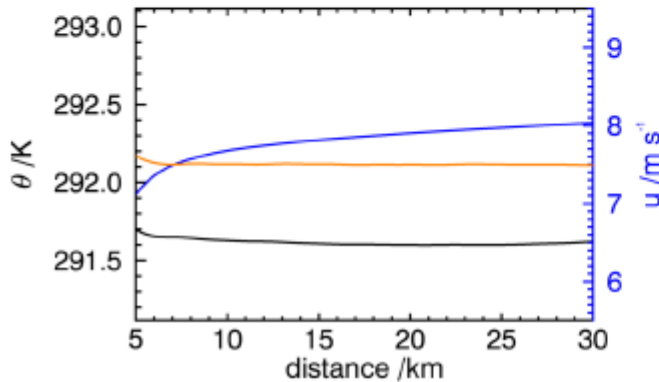
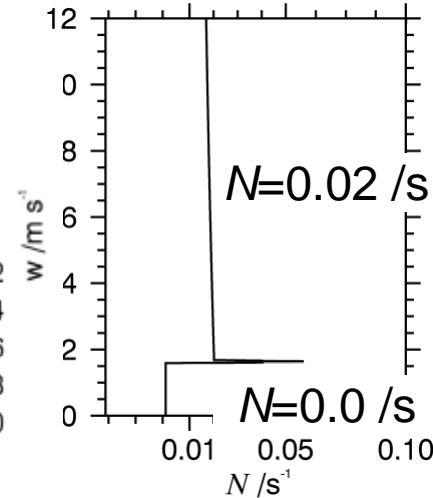
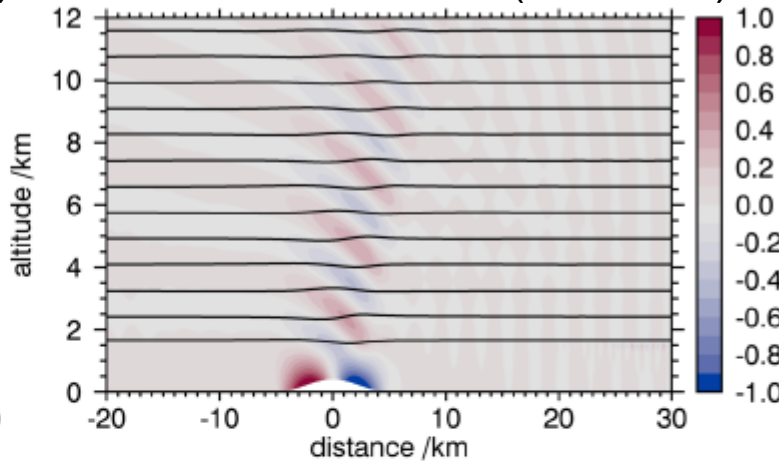
- Is concept of trapping on an inversion in the troposphere (Vosper 2004, Sachsperger et al. 2015) applicable for the **TIL**? ($N=0.02$ /s)

→ inversion must be 2 times **stronger** for trapping

inversion at 1.6 km ($\Delta\theta=3.3$ K)



inversion at 1.6 km ($\Delta\theta=6.6$ K)



potential
temperature,
horizontal and
vertical wind
at 2 km

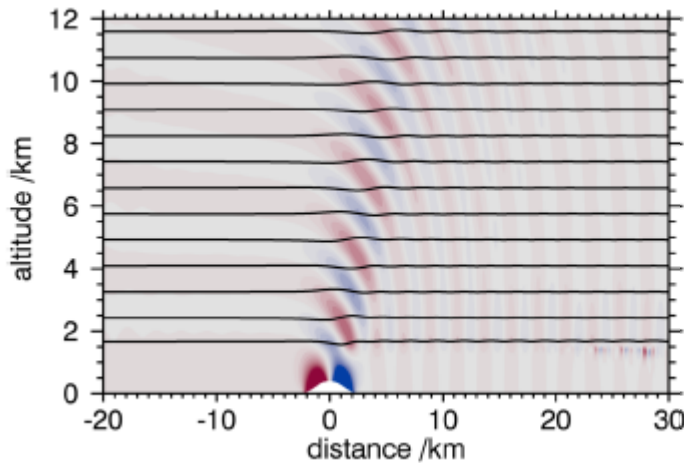
EULAG
Large Eddy Simulation

Mountain wave propagation: trapping on the TIL

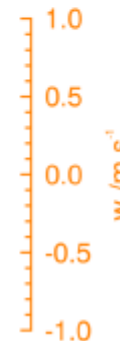
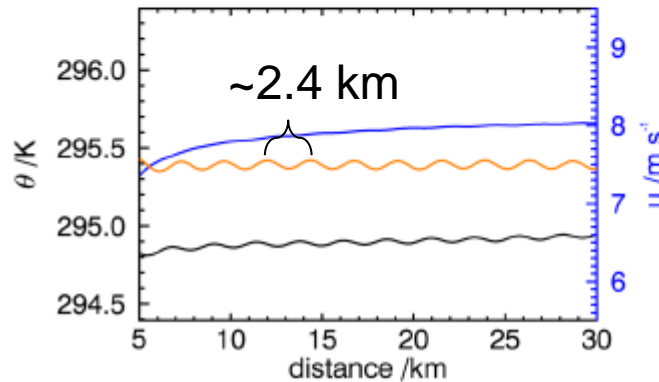
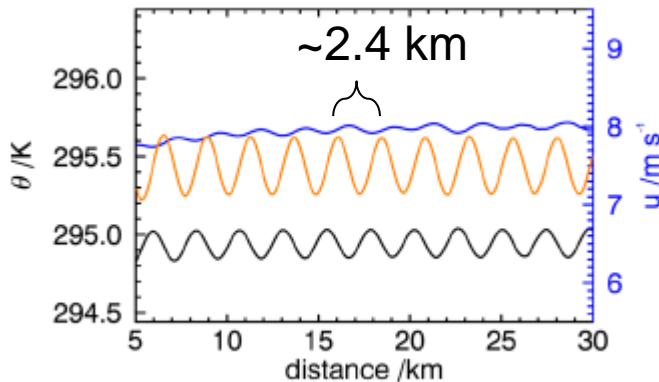
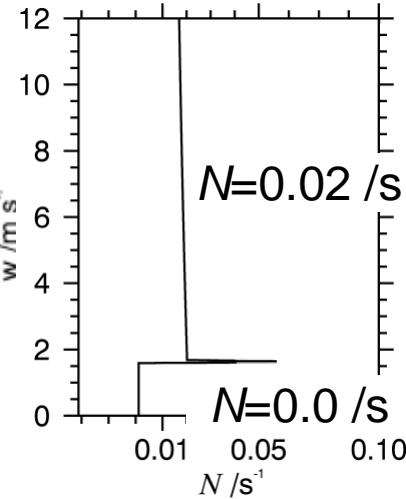
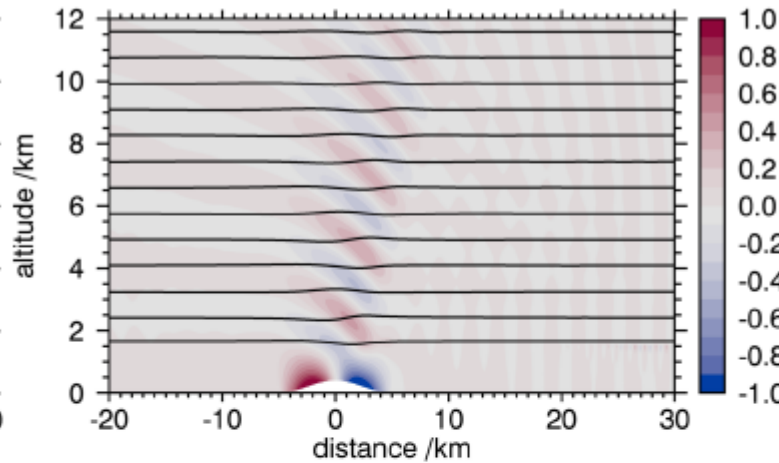
- Is concept of trapping on an inversion in the troposphere (Vosper 2004, Sachsperger et al. 2015) applicable for the **TIL**? ($N=0.02$ /s)

inversion at 1.6 km ($\Delta\theta=6.6$ K)

mountain width 5 km



mountain width 10 km

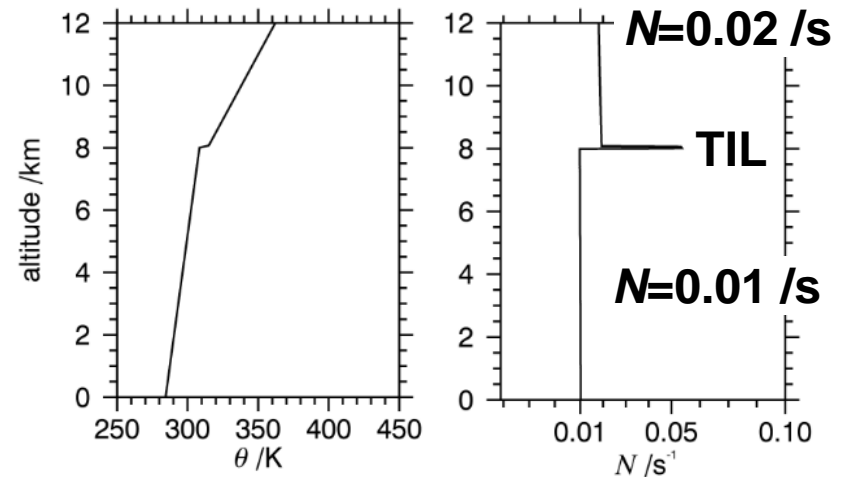


potential
temperature,
horizontal and
vertical wind
at 2 km

EULAG
Large Eddy Simulation

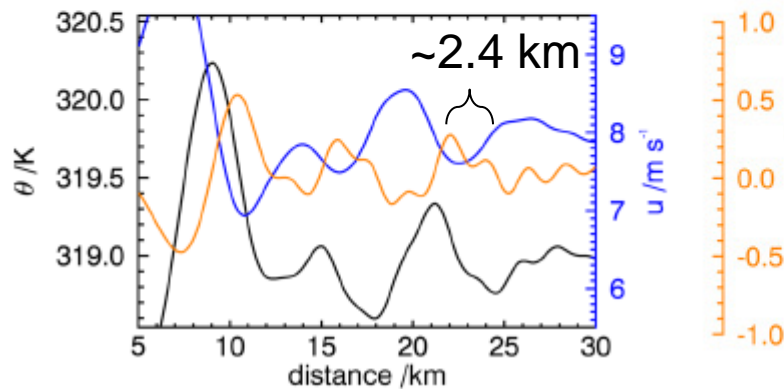
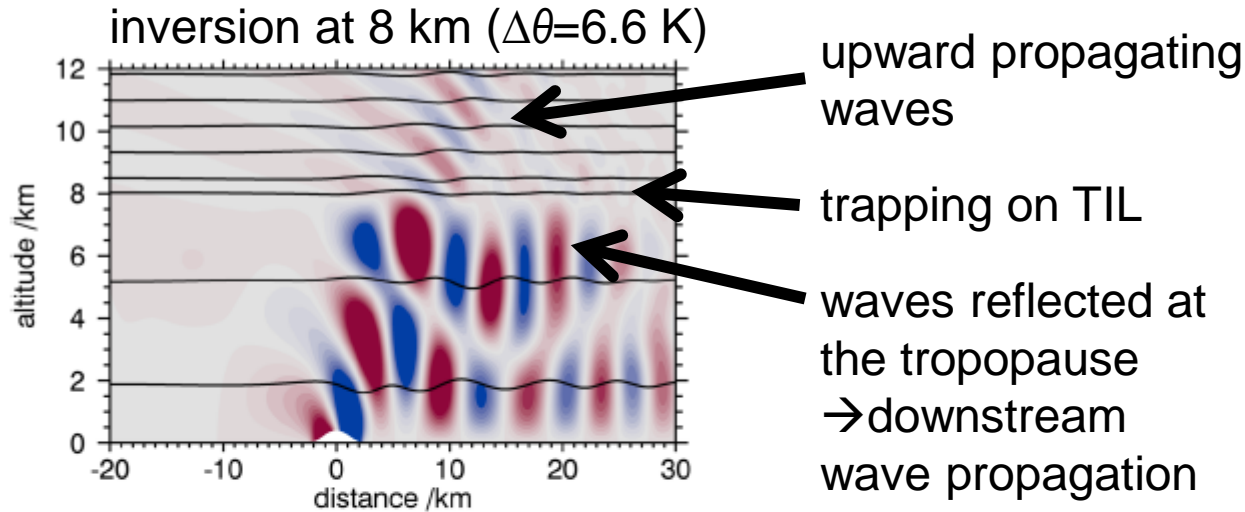
Mountain wave propagation: trapping on the TIL

- Is concept of trapping on an inversion in the troposphere (Vosper 2004, Sachsperger et al. 2015) applicable for the **TIL**? ($N=0.02$ /s, $z_T= 8$ km)

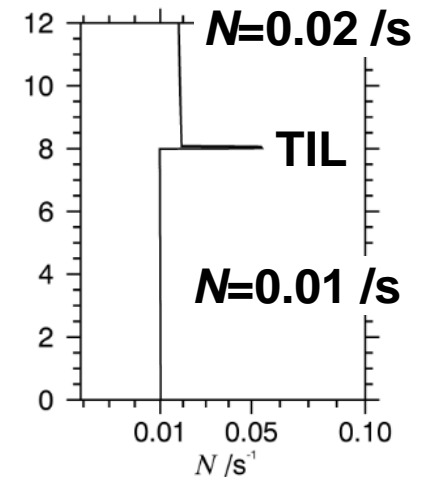


Mountain wave propagation: trapping on the TIL

- Is concept of trapping on an inversion in the troposphere (Vosper 2004, Sachsperger et al. 2015) applicable for the **TIL**? ($N=0.02 \text{ /s}$, $z_T=8 \text{ km}$)



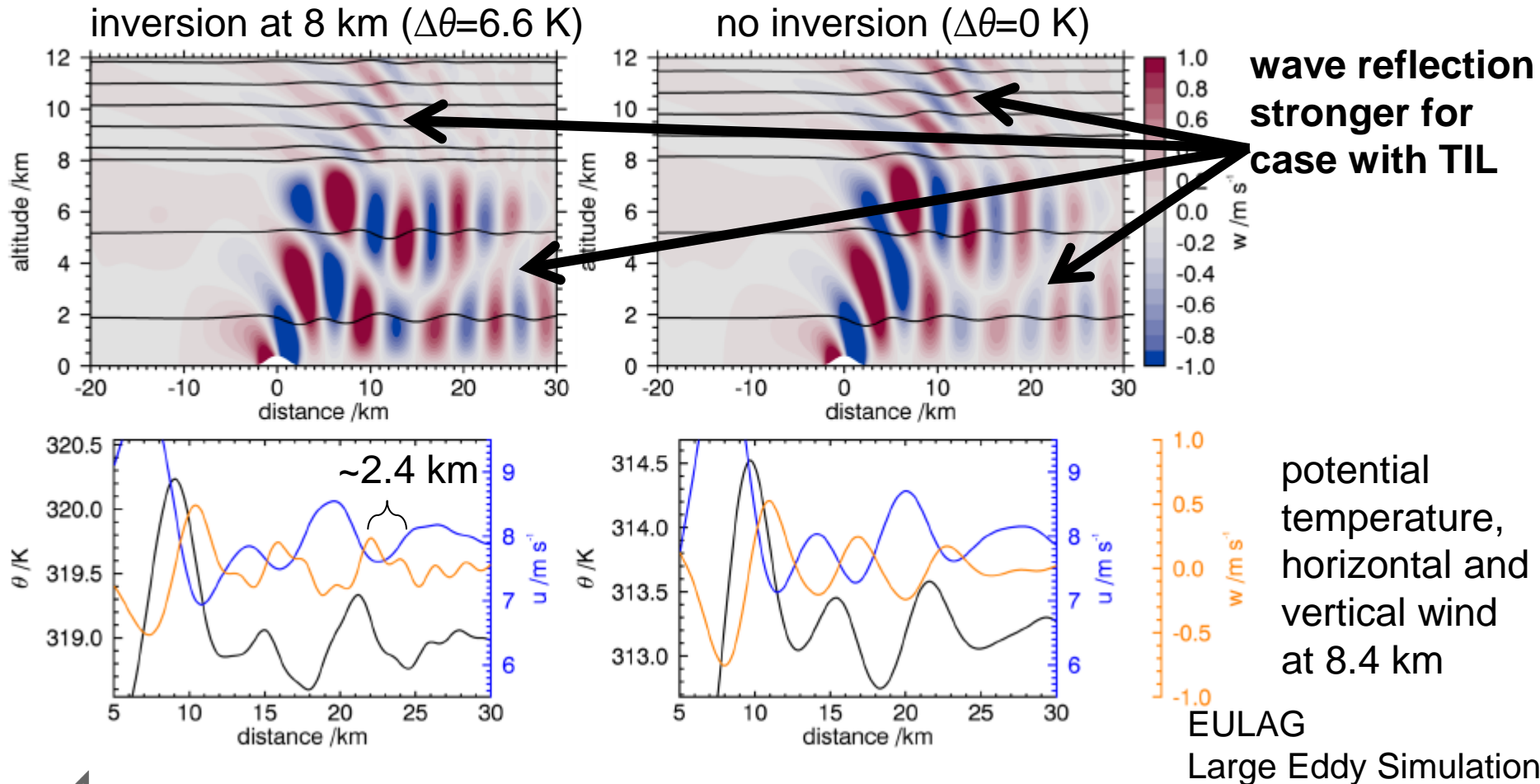
potential
temperature,
horizontal and
vertical wind
at 8.4 km



EULAG
Large Eddy Simulation

Mountain wave propagation: trapping on the TIL

- Is concept of trapping on an inversion in the troposphere (Vosper 2004, Sachsperger et al. 2015) applicable for the **TIL**? ($N=0.02$ /s, $z_T=8$ km)



Mountain wave propagation: role of the tropopause

Concept of **trapping on an inversion** in the troposphere (Vosper 2004, Sachsperger et al. 2015) is applicable for the **TIL**

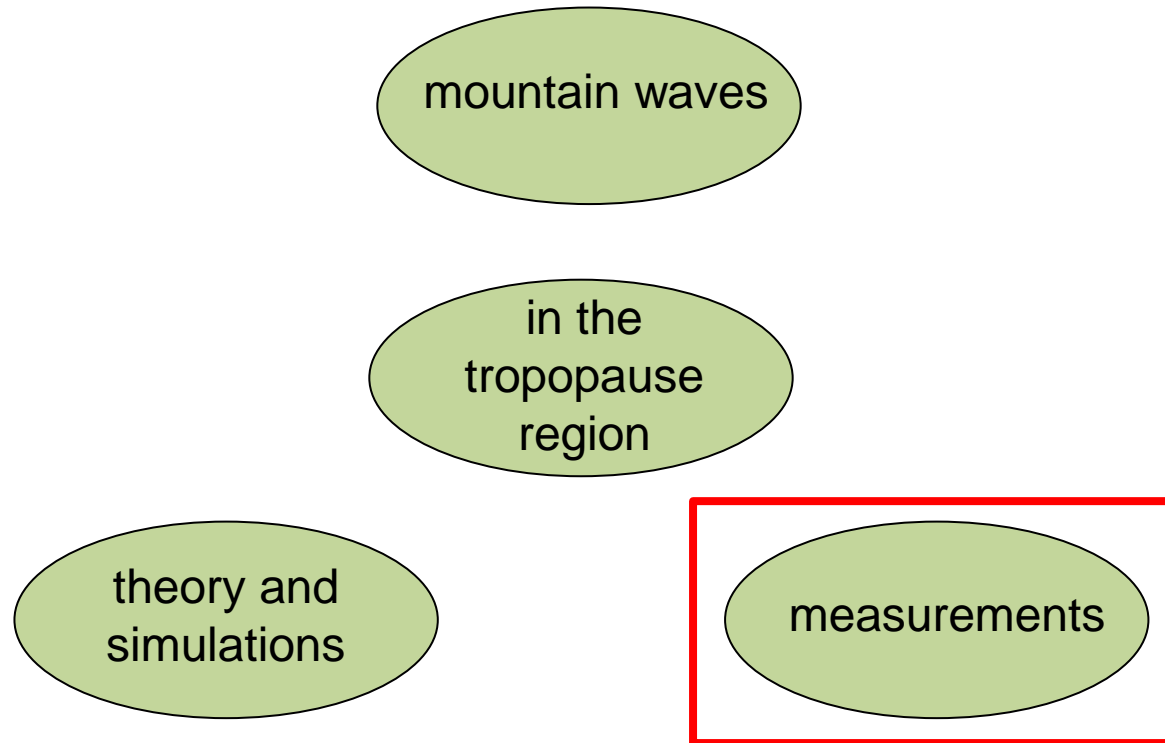
- **Inversion must be stronger than for the troposphere**
- Inversion must be even stronger when wind speed is larger
- **Horizontal wavelength decreases with increasing stability above the inversion (amplitudes depend on the generating terrain)**
- **tropopause causes wave reflection** and waves propagate downstream in the troposphere even though they are not evanescent in the stratosphere (**not classical Scorer trapping**)
- **TIL** causes stronger wave reflection than just the tropopause without inversion



the **amount of mountain wave energy** downstream in the troposphere and in the stratosphere **should vary** in dependence of the **trapping on the TIL and strength of reflection** → **strength of the TIL**



Strategy to increase the knowledge about gravity waves and their behaviour in the real atmosphere

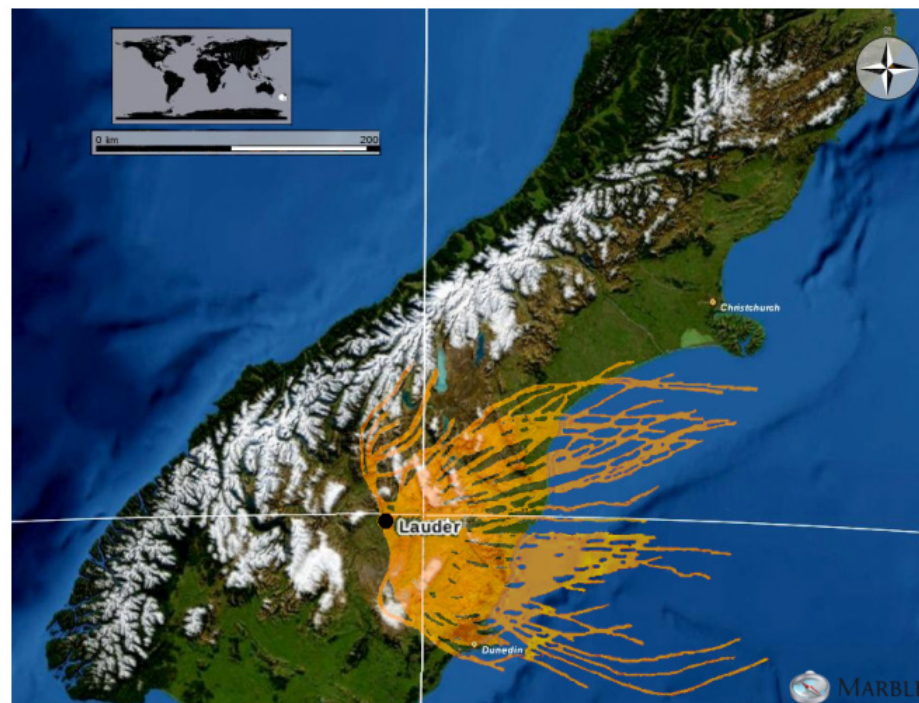


Measurements: DEEPWAVE radiosondes (Lauder)



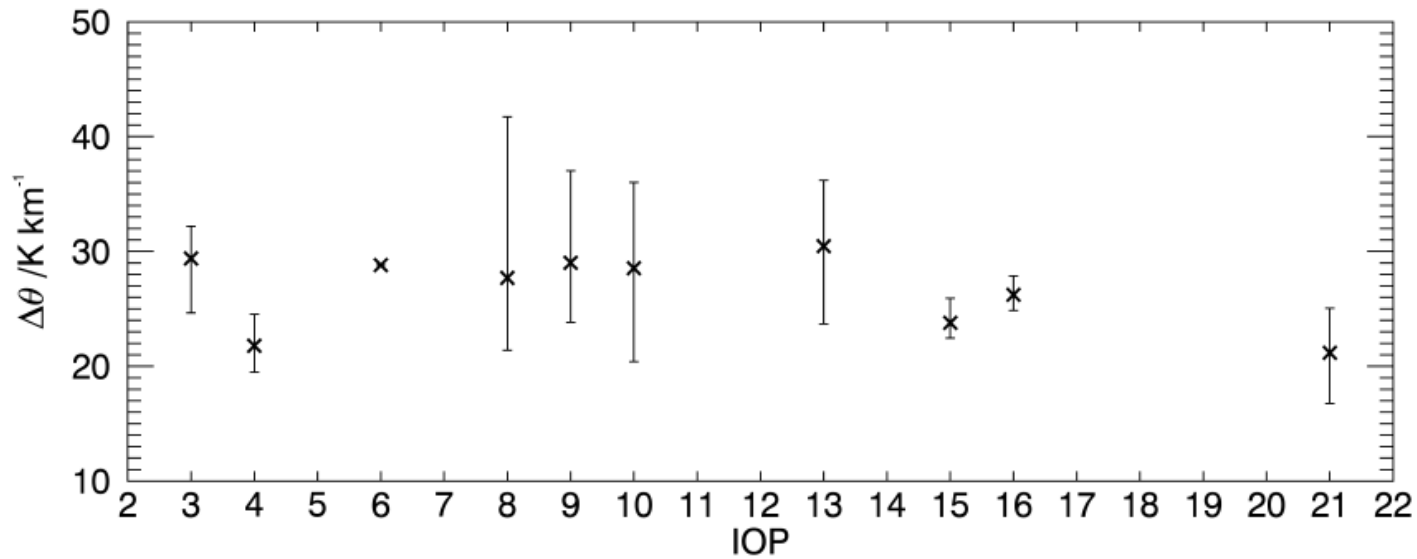
LAUDER soundings

98 soundings during
14 intensive observation
periods (**IOPs**)



45° S

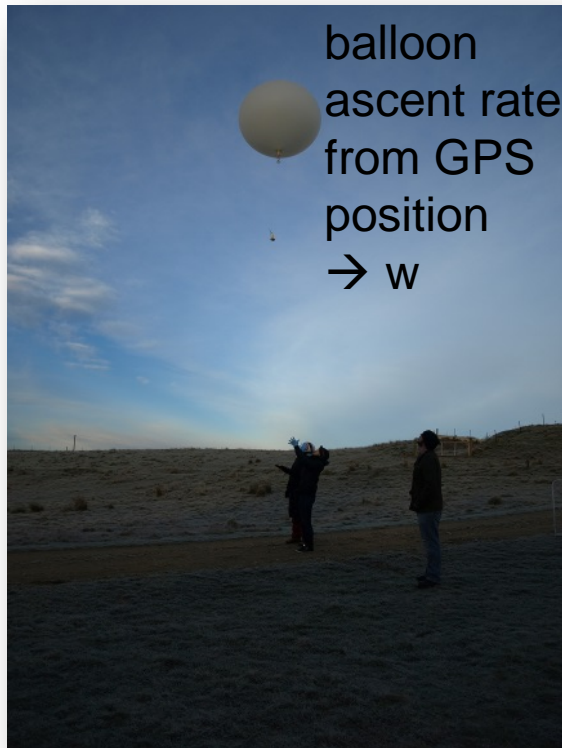
Strength of the TIL from DEEPWAVE soundings



**strength of the TIL of around 25-30 K
suggests trapping on the TIL and enhanced
mountain wave reflection during DEEPWAVE**



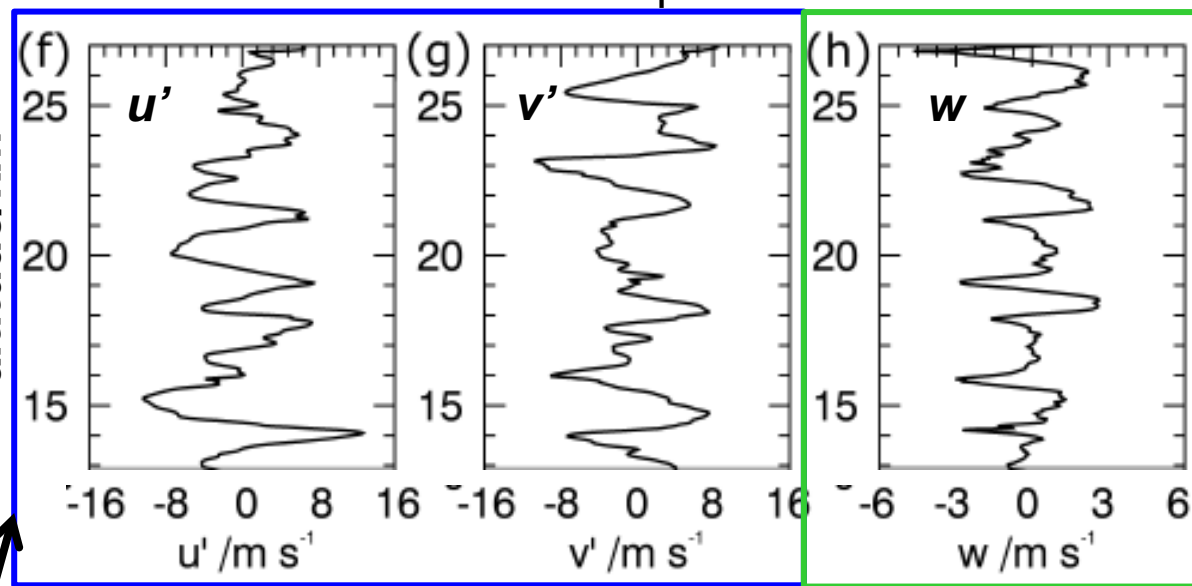
Analysis of radiosonde data



perturbation profiles

measured profile minus background fit
(2nd-polynomial-fit with
additional 5-km running mean)

altitude/km



Assumption: perturbations are caused by GWs

→ measurements of **hz. wind** emphasize
low frequency waves (inertia-GWs)

→ measurements of **vert. wind** emphasize
medium to high frequency waves

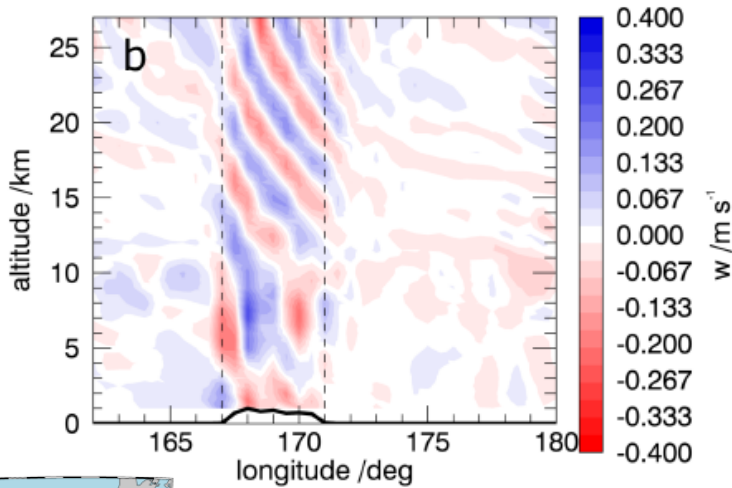
(Lane et al 2000, Lane et al 2003, Geller and Gong 2010)

Medium to high frequency waves: mountain waves?

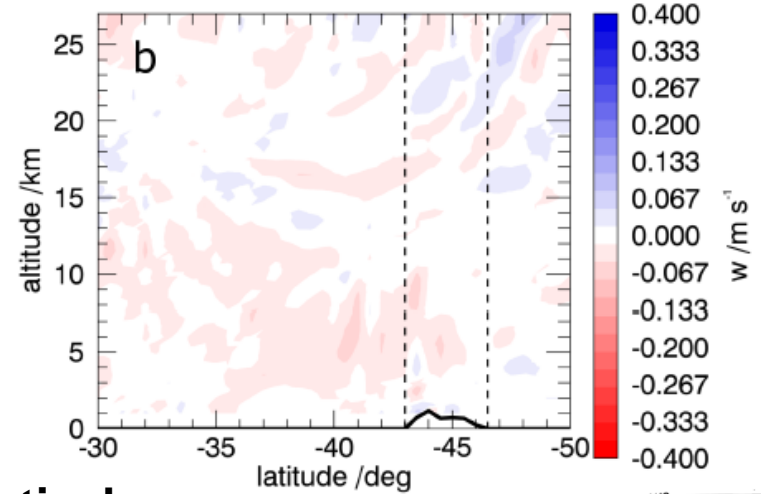
Approach: comparison of a mountain wave case to a non-mountain wave case

mountain wave case (IOP 4)

non-mountain wave case (IOP 15)



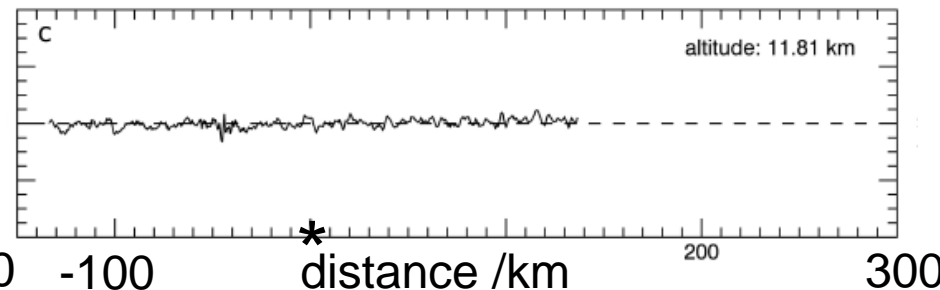
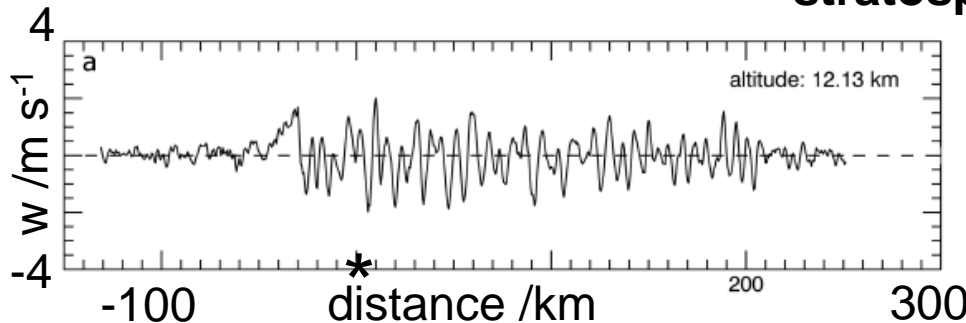
ECMWF
vertical
velocity



**Aircraft vertical
velocity
in lower
stratosphere**

NCAR/NSF GV

DLR Falcon



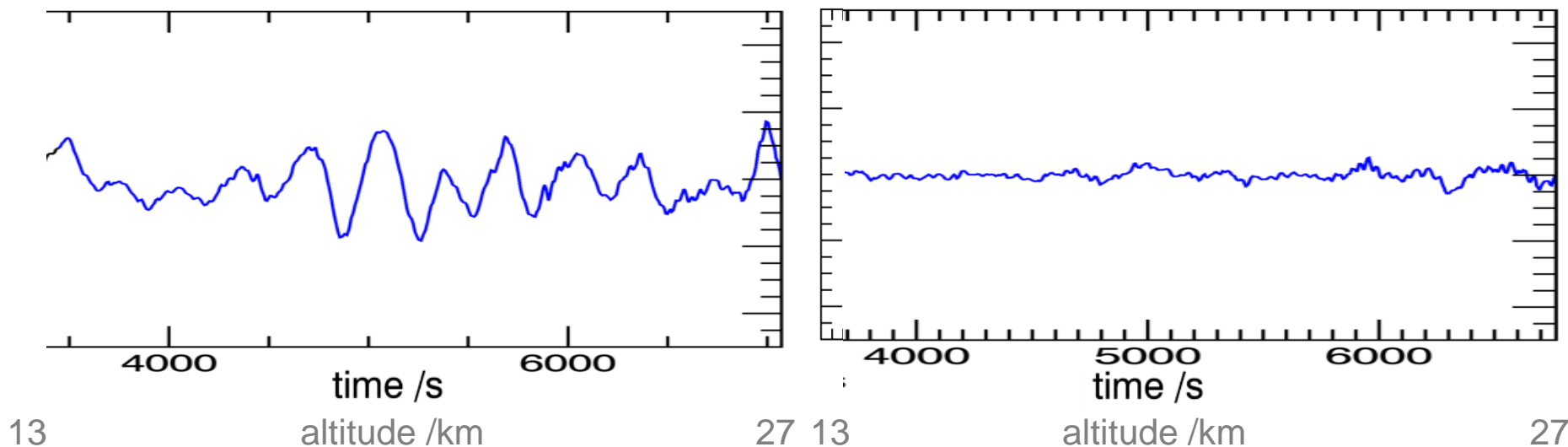
Medium to high frequency waves: mountain waves?

Approach: comparison of a mountain wave case to a non-mountain wave case

mountain wave case (IOP 4)

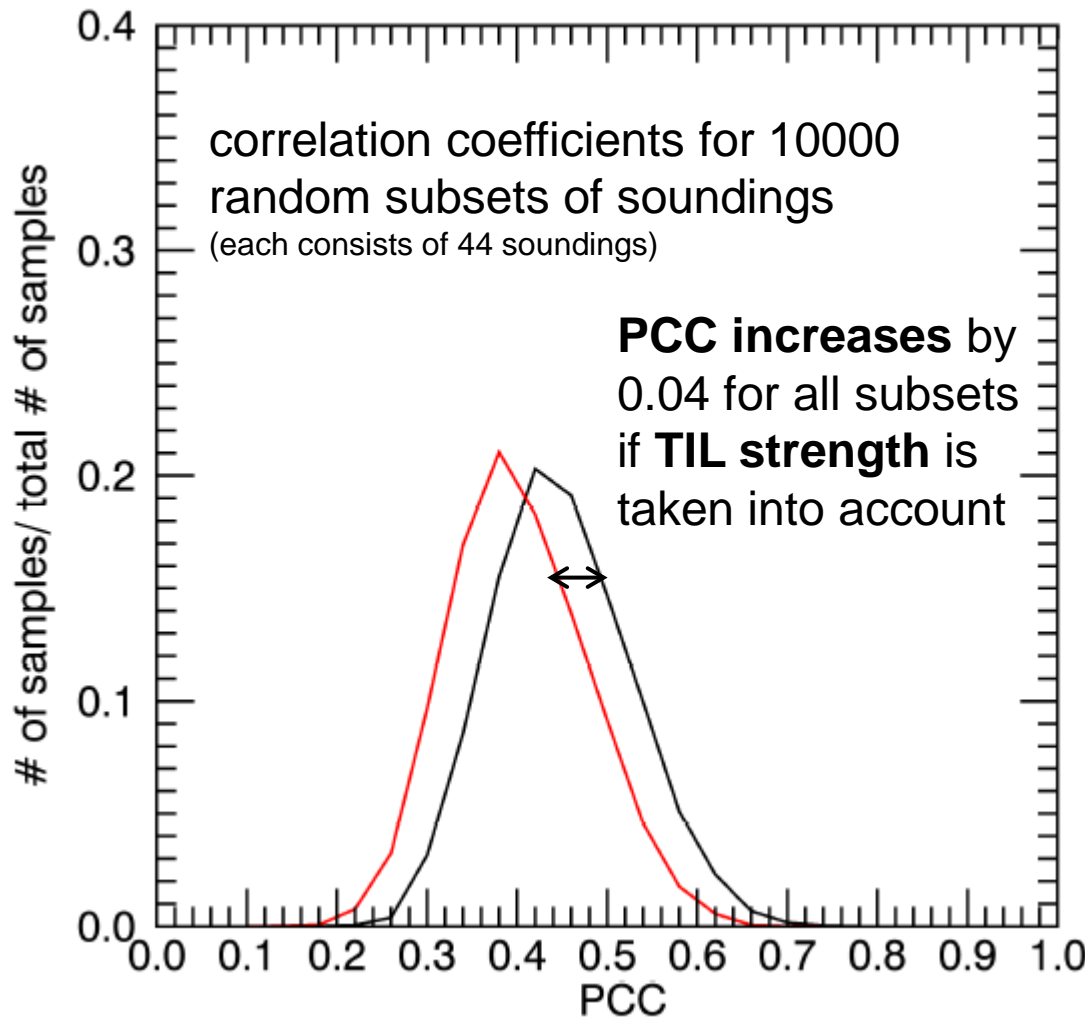
non-mountain wave case (IOP 15)

radiosondes
ascent rate perturbation (w) in the
stratosphere



ascent rate perturbations (w) of
soundings are caused by
mountain waves

Correlation between MW forcing, strength of TIL and stratospheric MW activity ($=1/2w^2$) in DEEPWAVE soundings

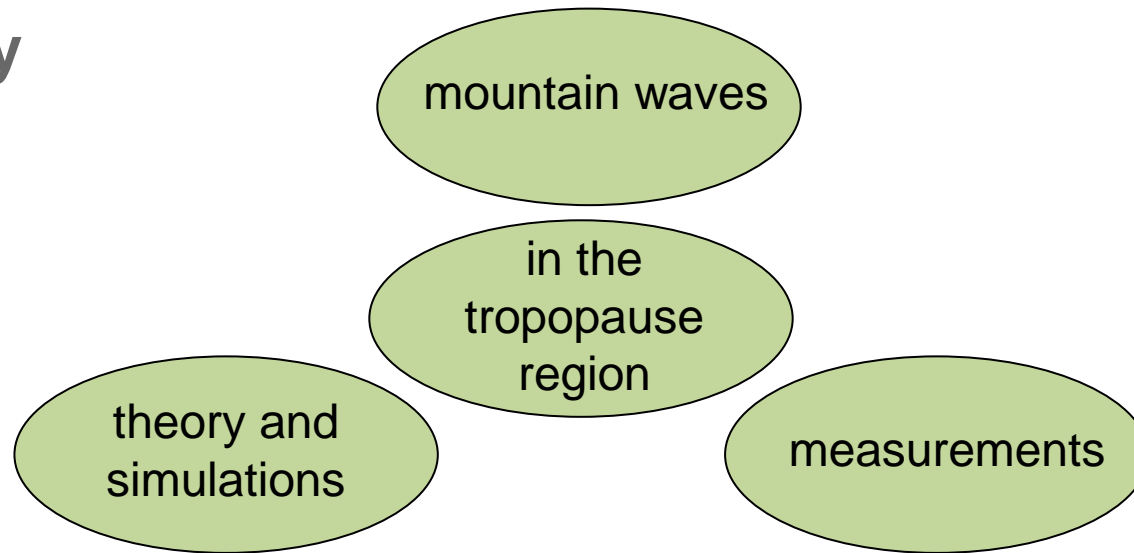


Pearson Correlation Coefficient (linear correlation):

- low level MW forcing and MW activity
→ **0.396**
- low level MW and **TIL strength** and MW activity (multiple linear regression)
→ **0.434**

TIL strength and associated processes (trapping, reflection) influence the observed **stratospheric MW activity**

Summary



Concept of **trapping on an inversion** in the troposphere (Vosper 2004, Sachsperger et al. 2015) is applicable for the **TIL**

the **amount of mountain wave energy** downstream in the troposphere and in the stratosphere **should vary** in dependence of **the trapping on the TIL and strength of reflection** → **strength of the TIL**

strength of the TIL of around 25-30 K **suggests trapping on the TIL and enhanced mountain wave reflection** during DEEPWAVE

ascent rate perturbations (w) of soundings are caused by **mountain waves**

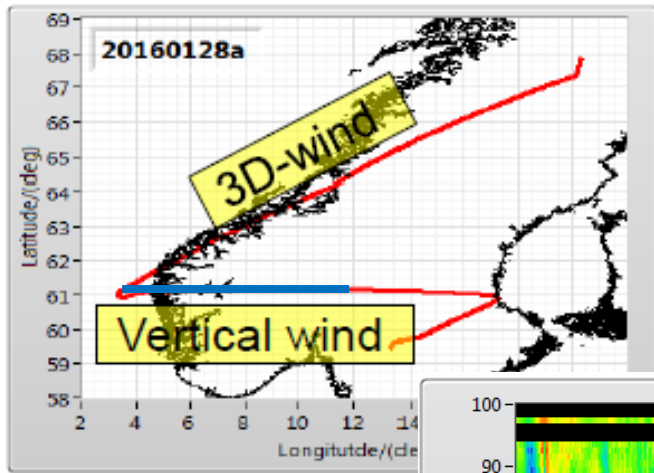
TIL strength and associated processes (trapping, reflection) influence the observed **stratospheric MW activity**



Outlook

measurements

GW-LCYCLE 2 wind lidar (B. Witschas)



- horizontal wavelength from wavelet analysis
below TP spectral max $\lambda_h > 10$ km, above
TP spectral max $\lambda_h < 10$ km

vertical wind

Falcon in-situ

