

14.2

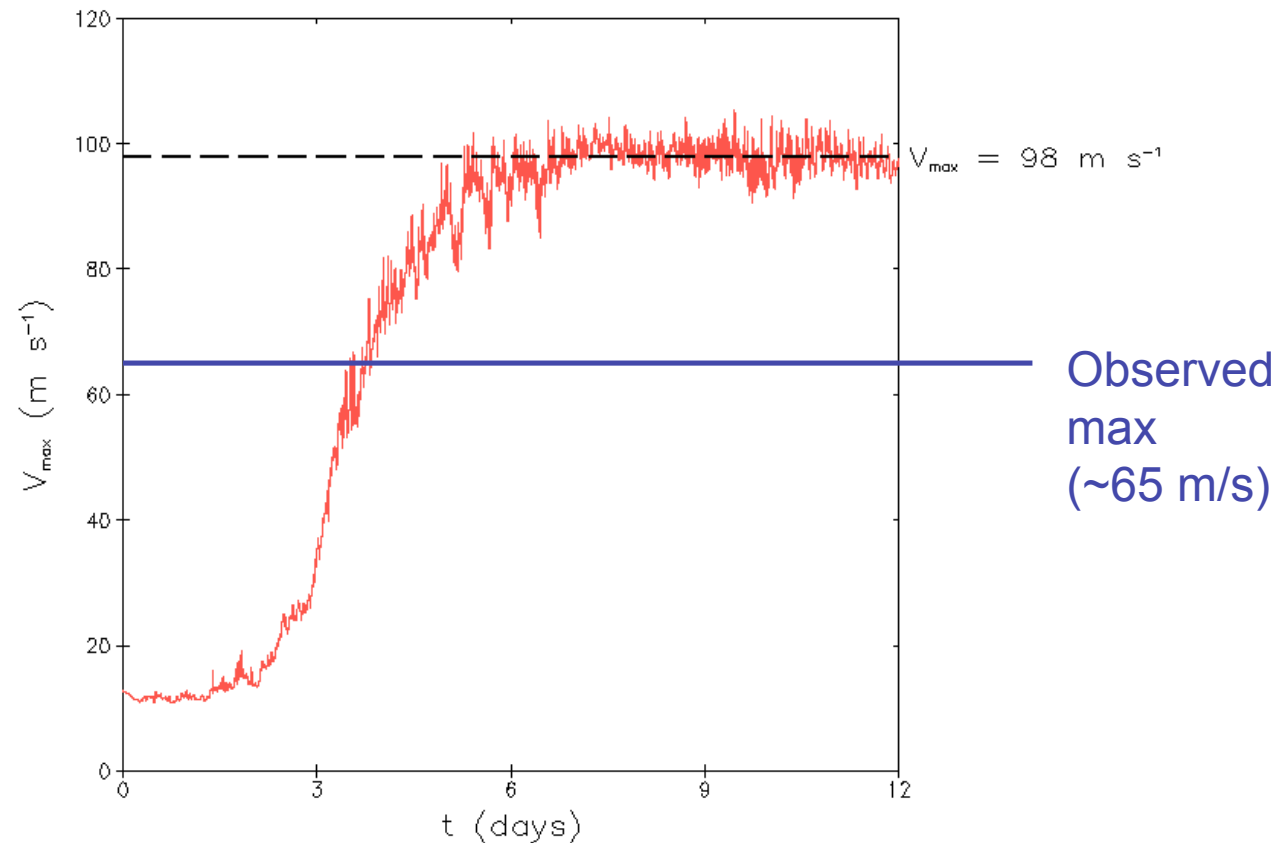
The effects of small-scale turbulence on maximum hurricane intensity

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- Axisymmetric numerical model (2d: r-z)
 - Nonhydrostatic, primitive equations, time-dependent
 - See Bryan and Rotunno (2009b, MWR) for details
 - $T_s = 26^\circ\text{C}$, same setup as Rotunno-Emanuel (1987) and Persing-Montgomery (2003)

Time series of max. azimuthal velocity (m s^{-1}):



Model components investigated:

[see Bryan and Rotunno (2009b, MWR) for details]

- Resolution
- Numerics
- Initial vortex
- Governing equations
 - unique mass/energy conservation in this model
- Microphysics
 - liquid / ice processes
 - single-moment / double-moment
 - fall velocity of condensate
- Surface exchange coefficients (C_E / C_D)
- Turbulence

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- **Turbulence**

Turbulence eddy viscosities:

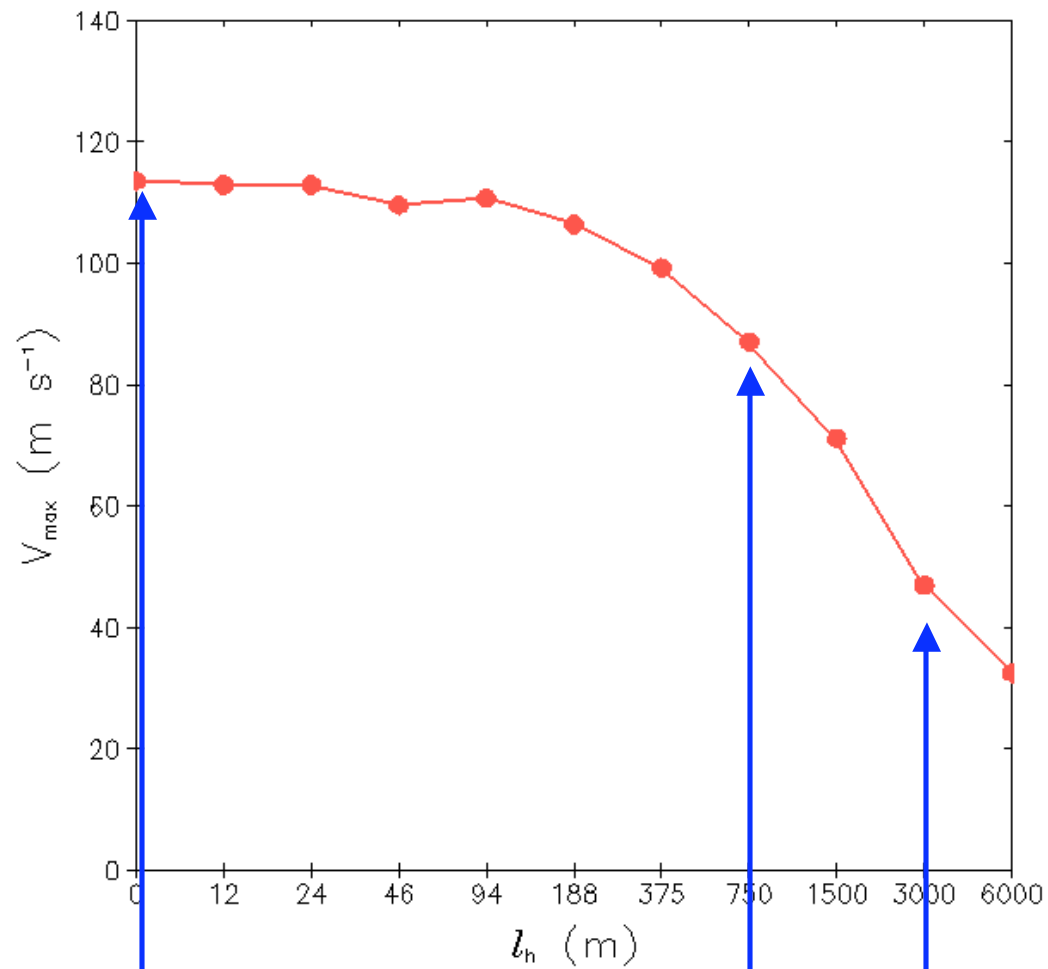
$$\nu_h = l_h^2 S_h,$$

$$\nu_v = l_v^2 (S_v^2 - N_m^2)^{1/2}.$$

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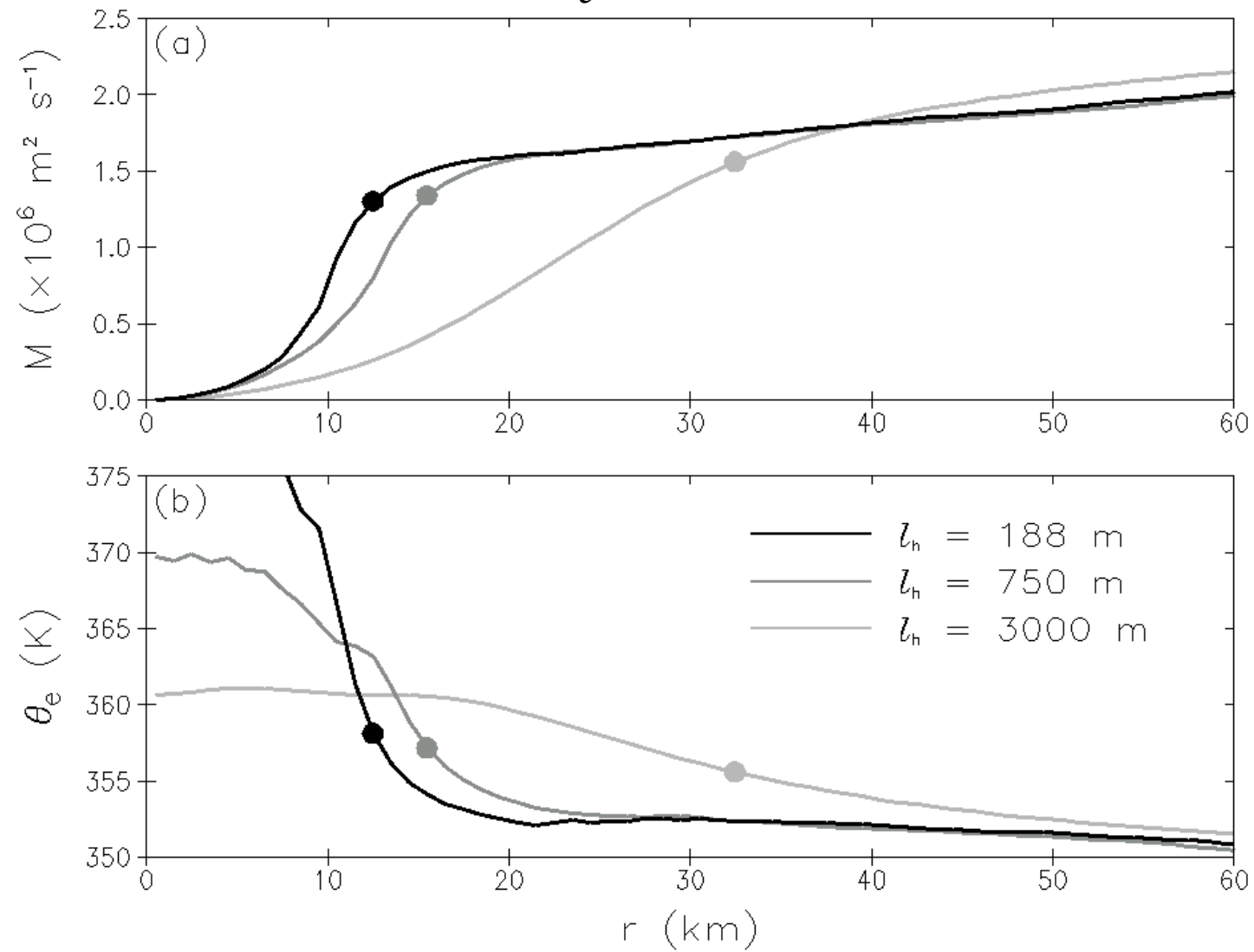
HOS06

PM03

RE87

Bryan and Rotunno (2009b, MWR)

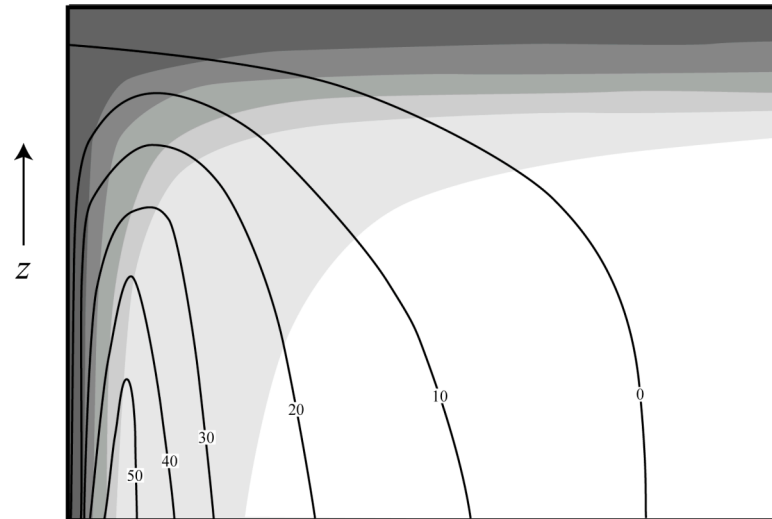
M and θ_e at $z = 1.1$ km



Thermal-wind balance

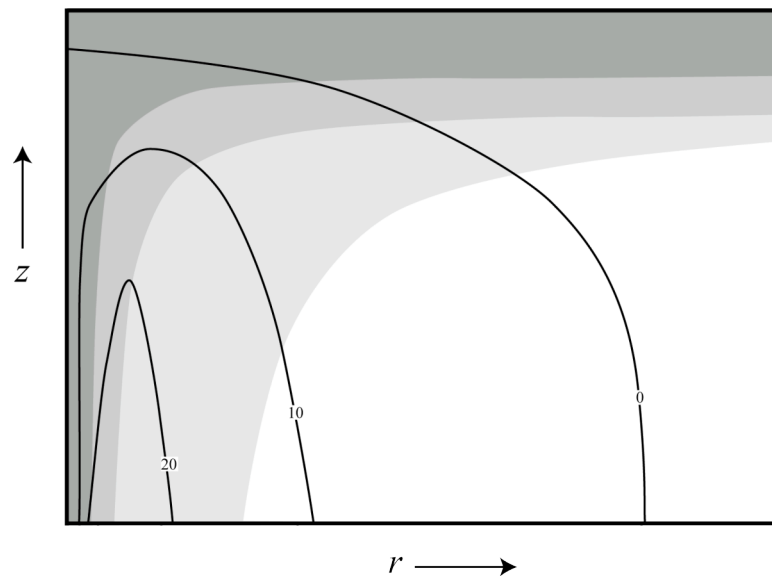
- vertical wind shear is proportional to the horizontal temperature (entropy) gradient

A strong vortex:



shading = entropy

A weak vortex:

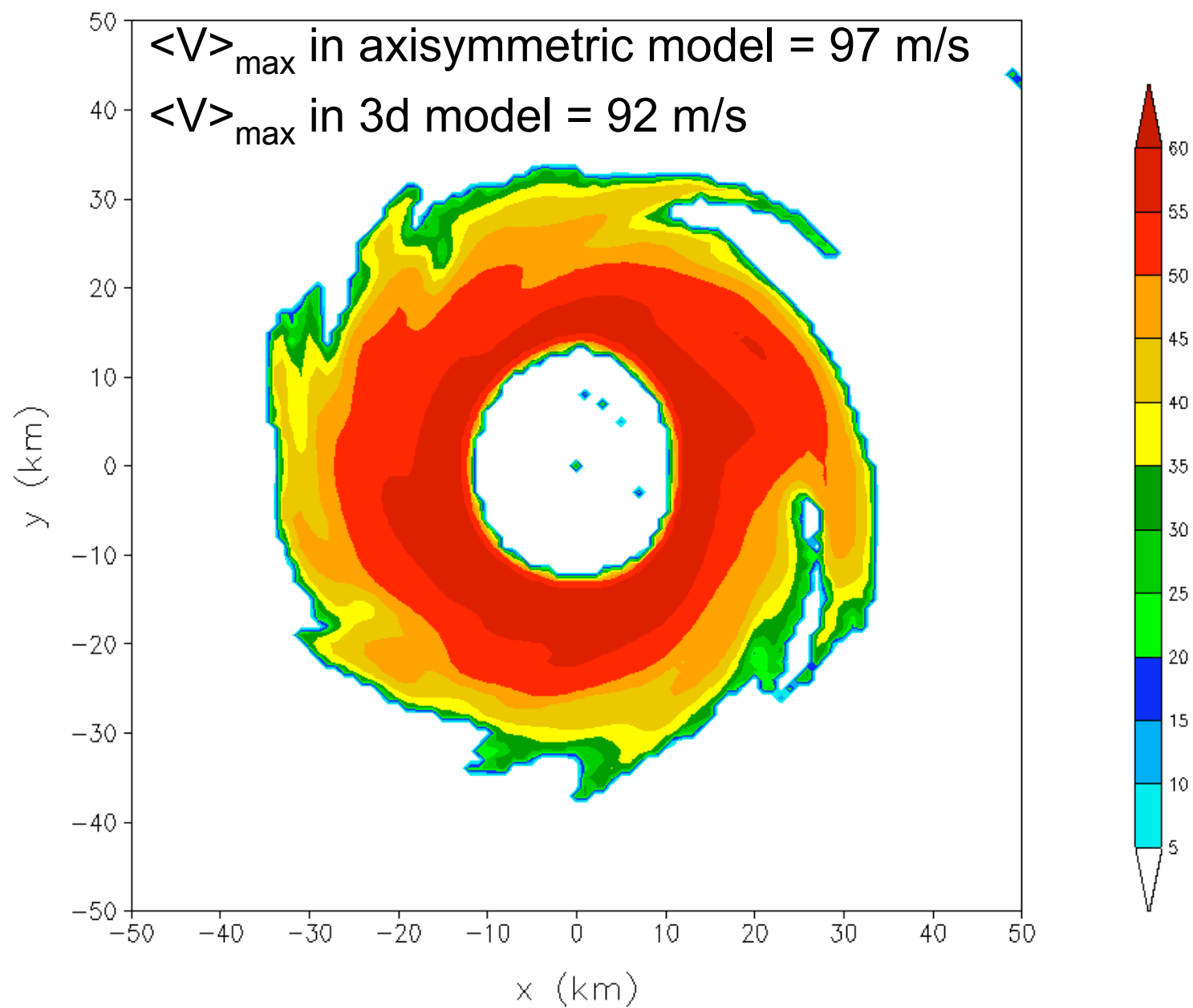


contours = v (m s^{-1})

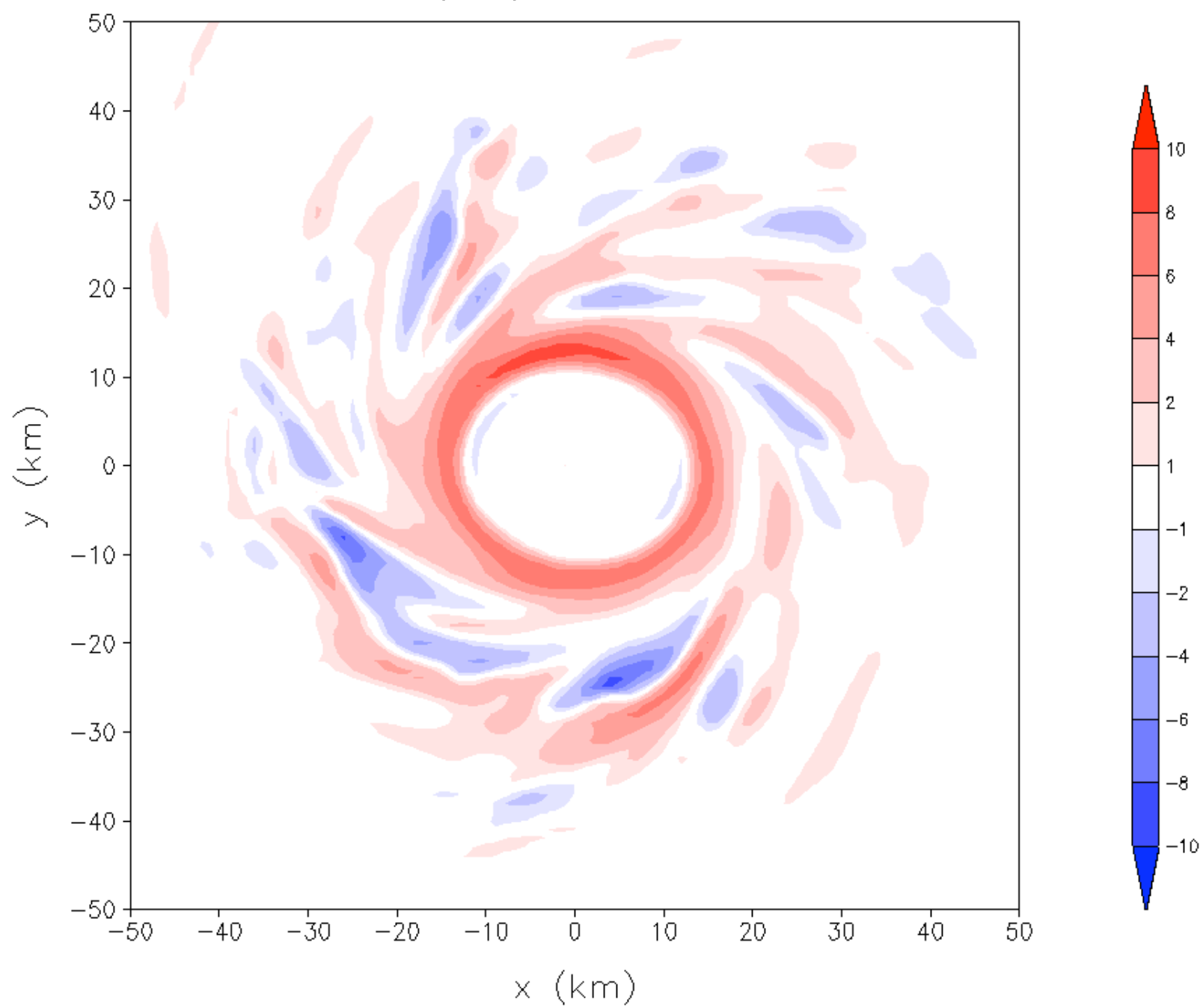
Outstanding question:

- What happens in 3d numerical simulations?
- Setup:
 - Same initial conditions, physics, etc
 - 3d geometry (Cartesian grid)
 - $\Delta x = \Delta y = 1 \text{ km}$, $\Delta z = 250 \text{ m}$
(same as axisymmetric simulations)

Surface reflectivity

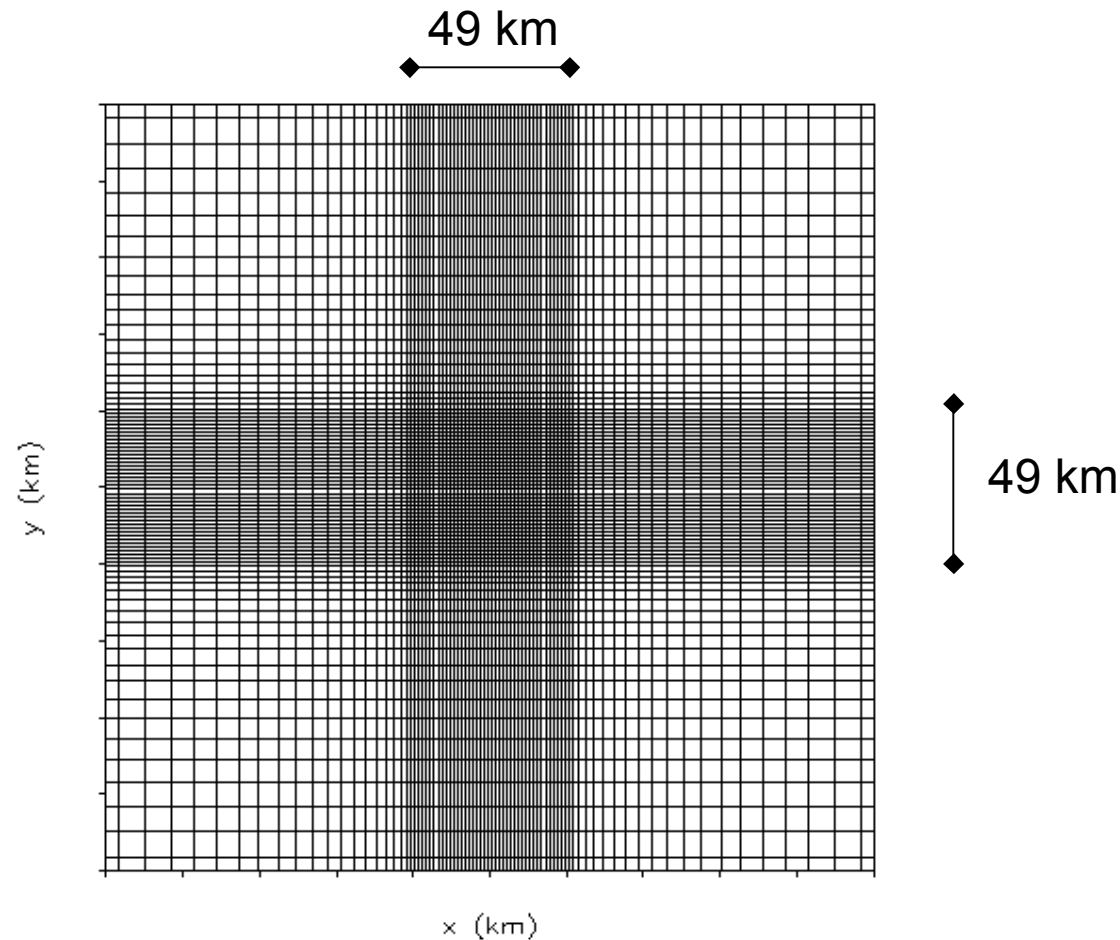


w (m/s) at $z = 1$ km

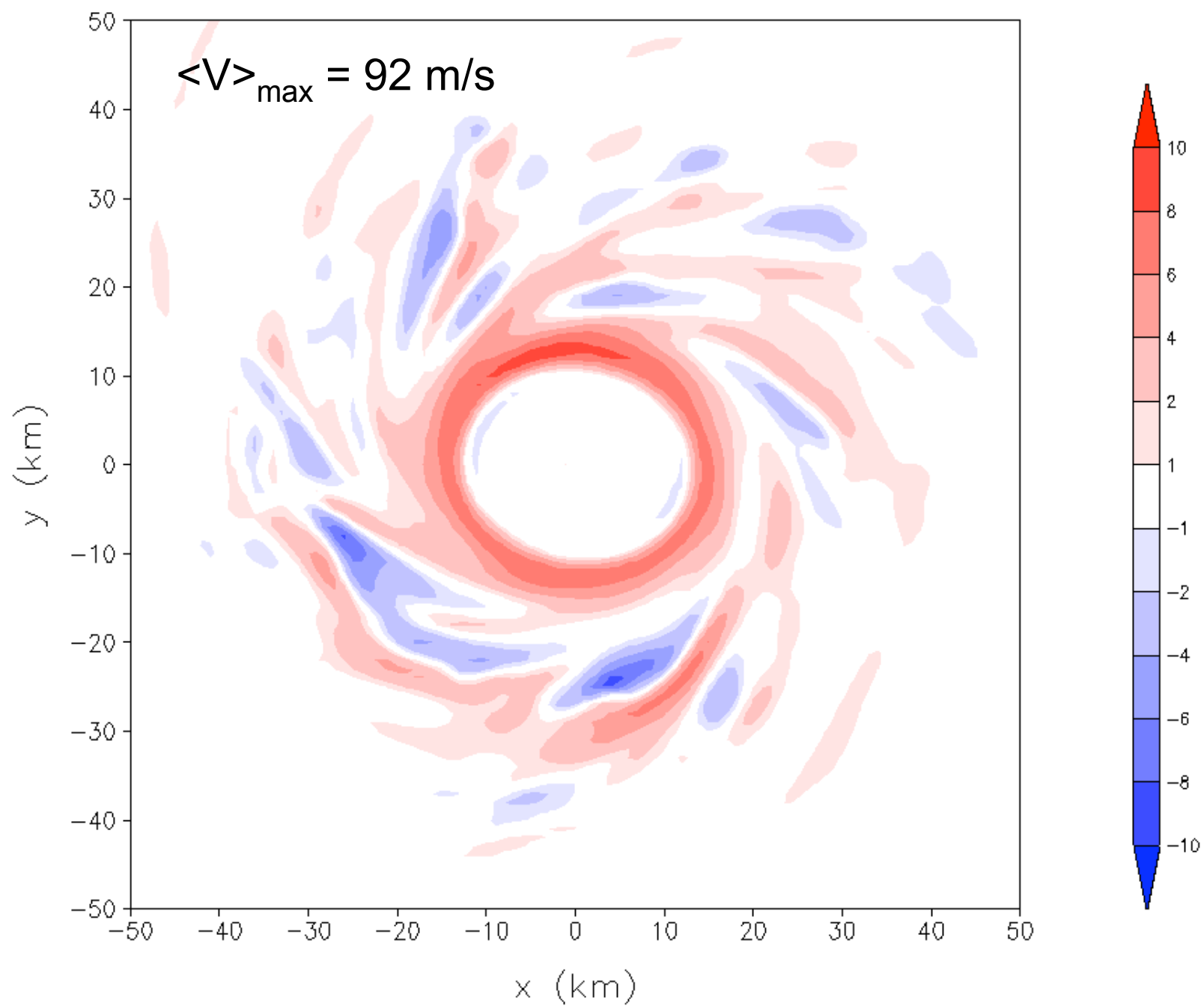


A very high-resolution simulation

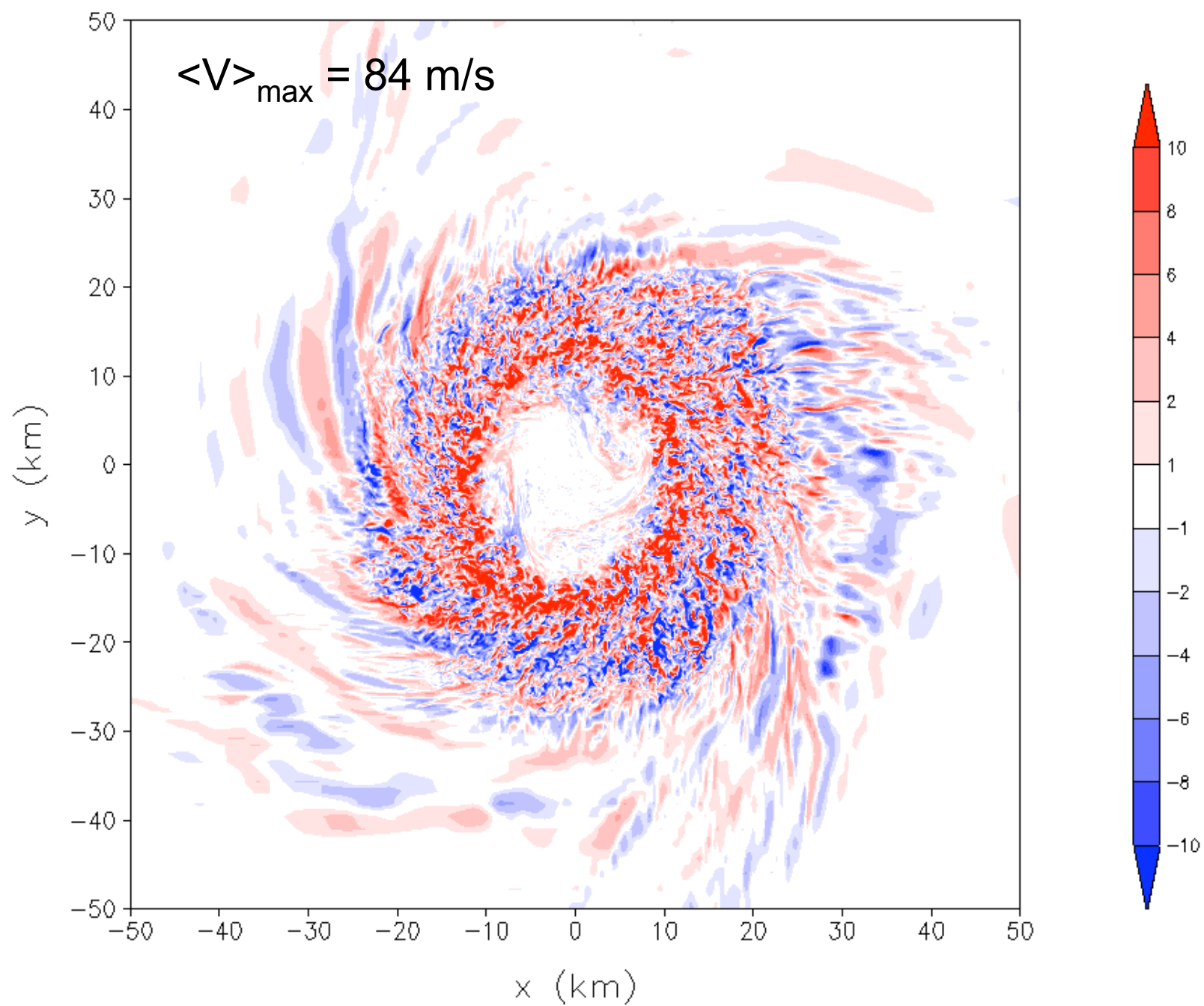
- Motivated by Rotunno et al. (2009, BAMS, in press)
- In center: $\Delta x = \Delta y = \Delta z = 62.5$ m
- (stretched structured grid)
- Initialized from 1-km simulation



w (m/s) at $z = 1$ km: $\Delta = 1000$ m

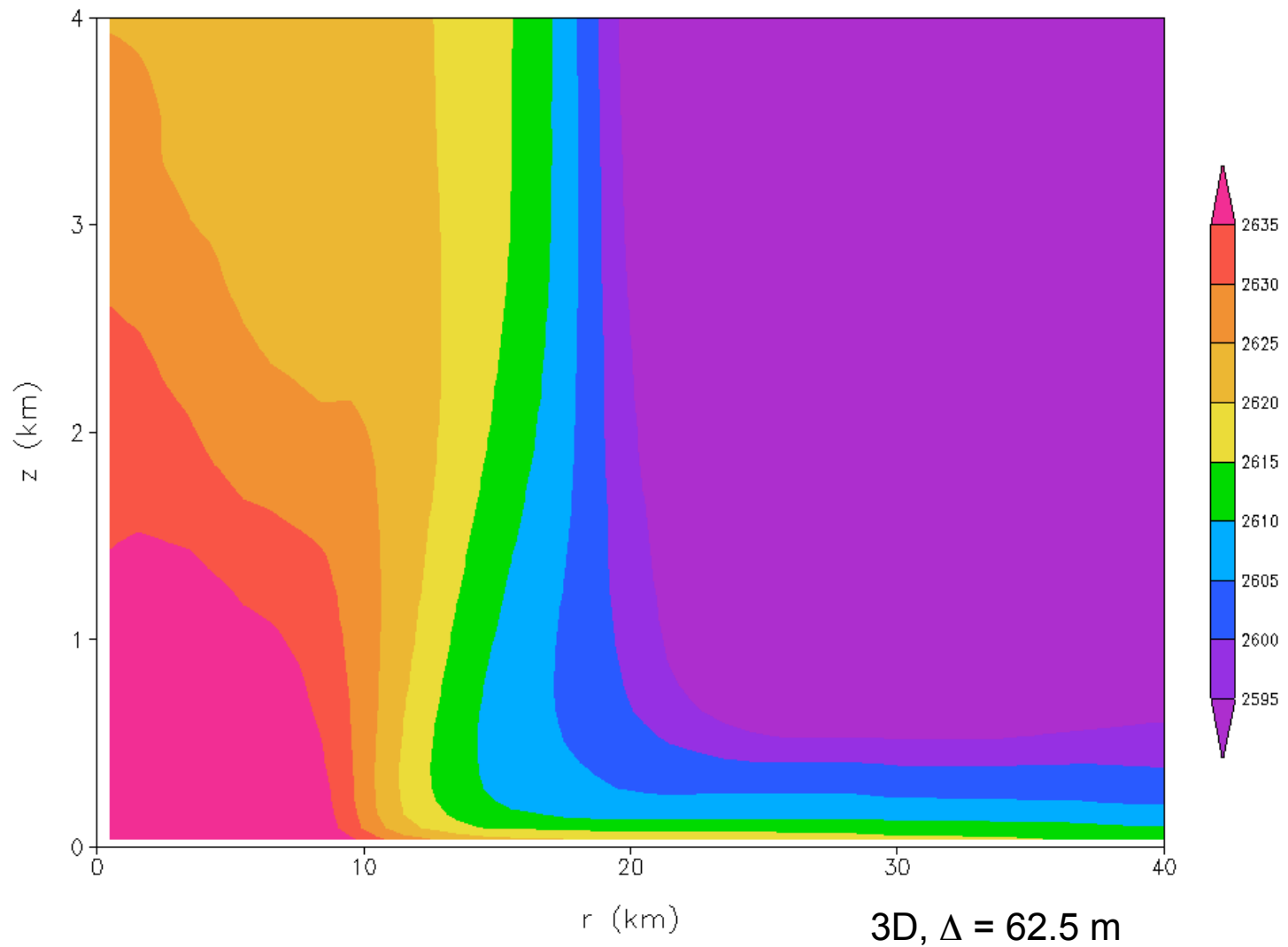


w (m/s) at $z = 1$ km: $\Delta = 62.5$ m



Let $\alpha(r, \phi, z) = \langle \alpha \rangle(r, z) + \alpha'(r, \phi, z)$

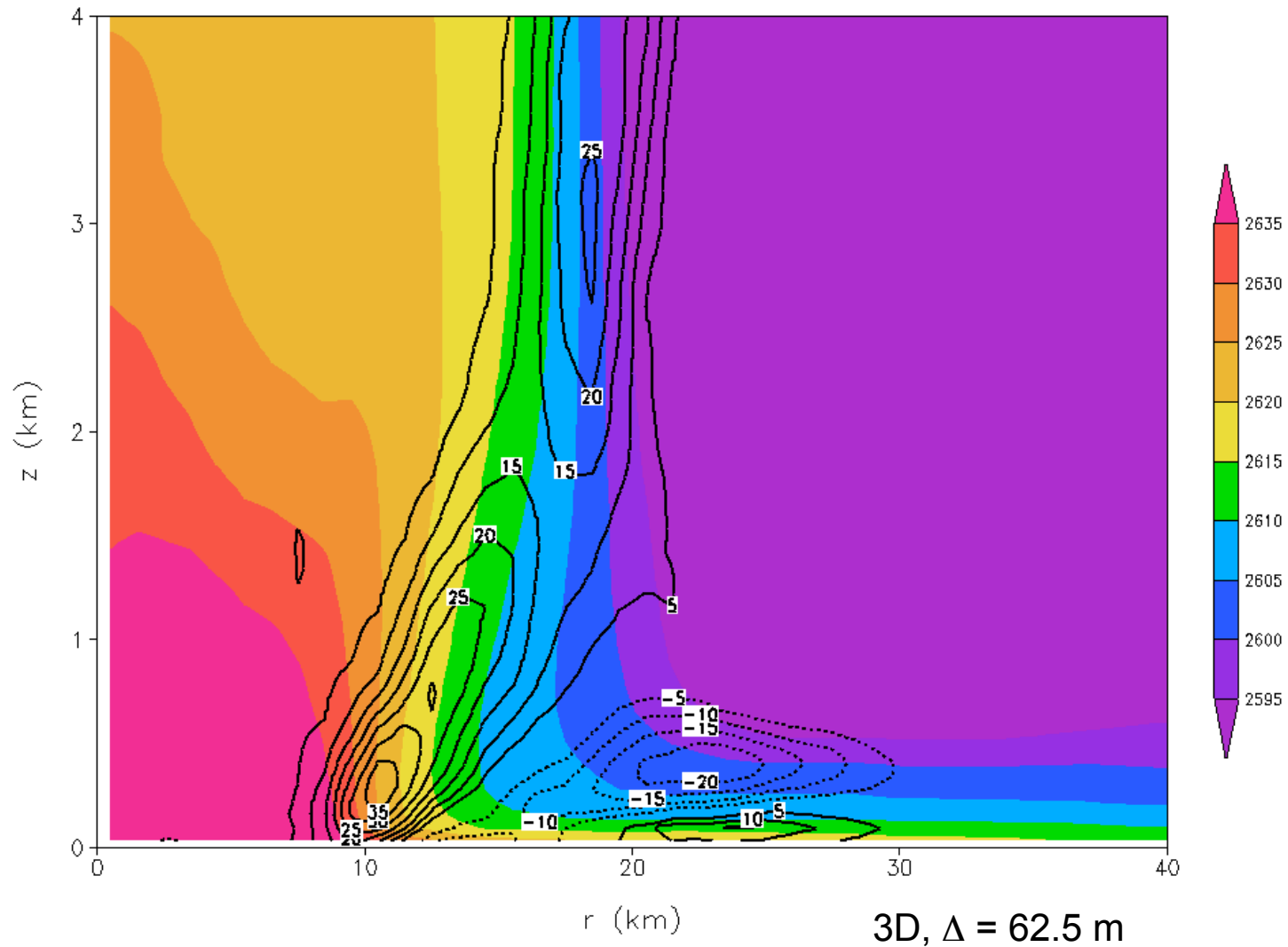
$\langle s \rangle$ (shaded)



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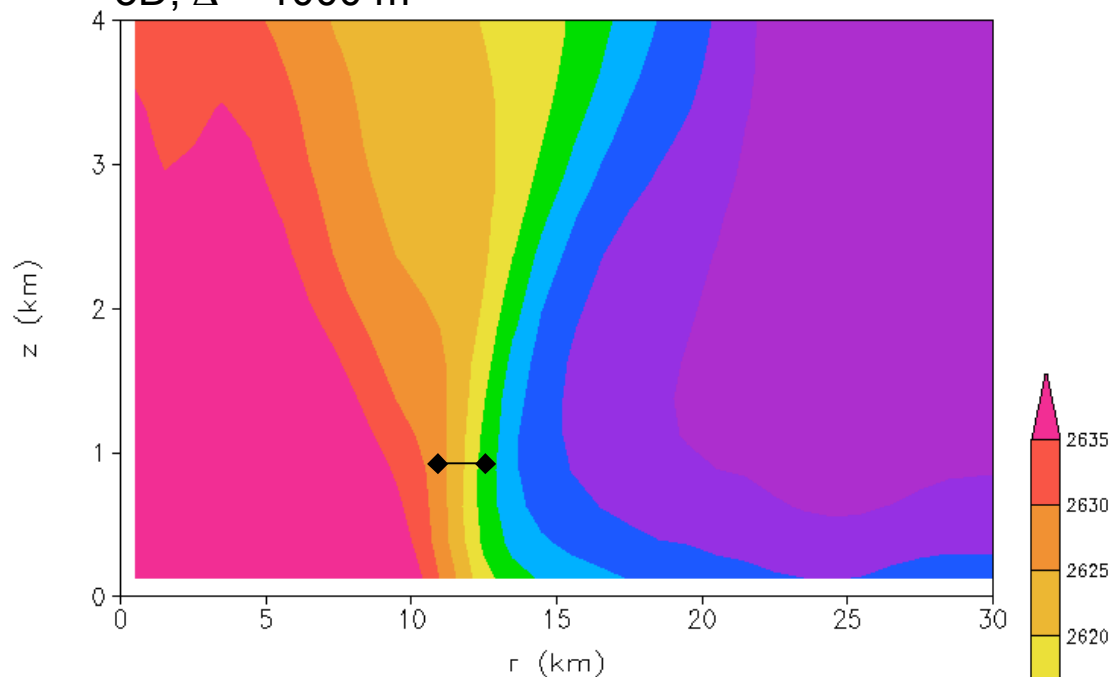
and turbulent flux of s in radial direction, $\langle u's' \rangle$ (contours)



total moist entropy, $\langle s \rangle$:

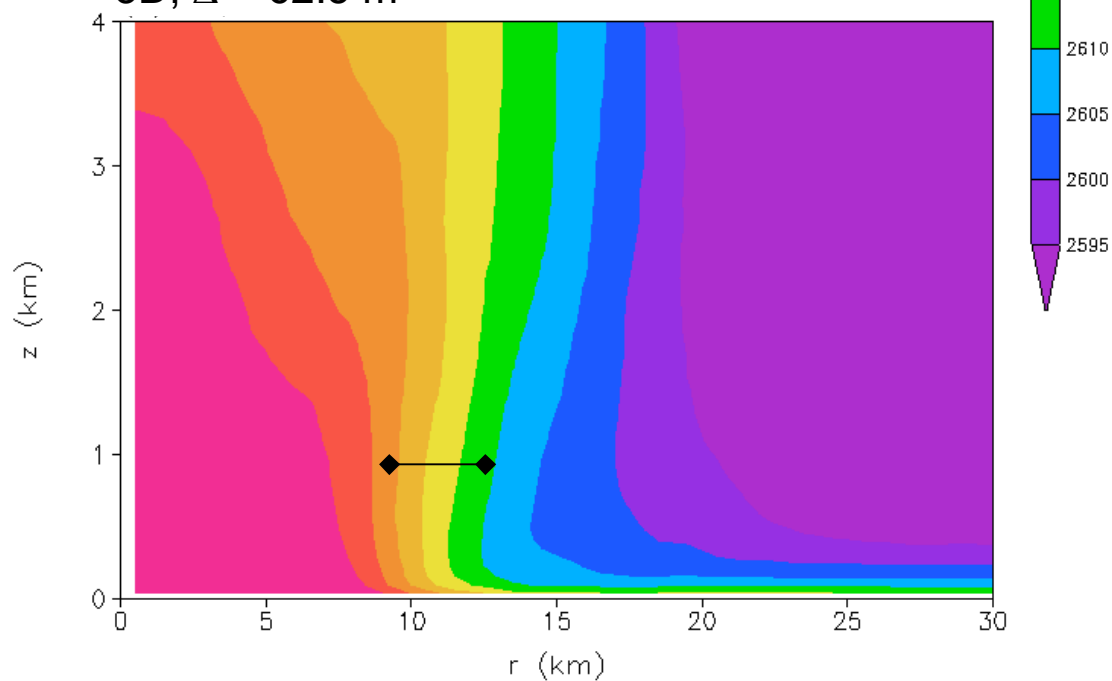
$$\frac{\partial \langle s \rangle}{\partial r} = -8.8 \times 10^{-3} \text{ m s}^{-2} \text{ K}^{-1}$$

3D, $\Delta = 1000 \text{ m}$



$$\frac{\partial \langle s \rangle}{\partial r} = -4.3 \times 10^{-3} \text{ m s}^{-2} \text{ K}^{-1}$$

3D, $\Delta = 62.5 \text{ m}$



Hurricane Isabel (2003):

[from Montgomery et al. (2006, BAMS)]

$$\frac{\partial \langle s \rangle}{\partial r} = -1.7 \times 10^{-3} \text{ m s}^{-2} \text{ K}^{-1}$$

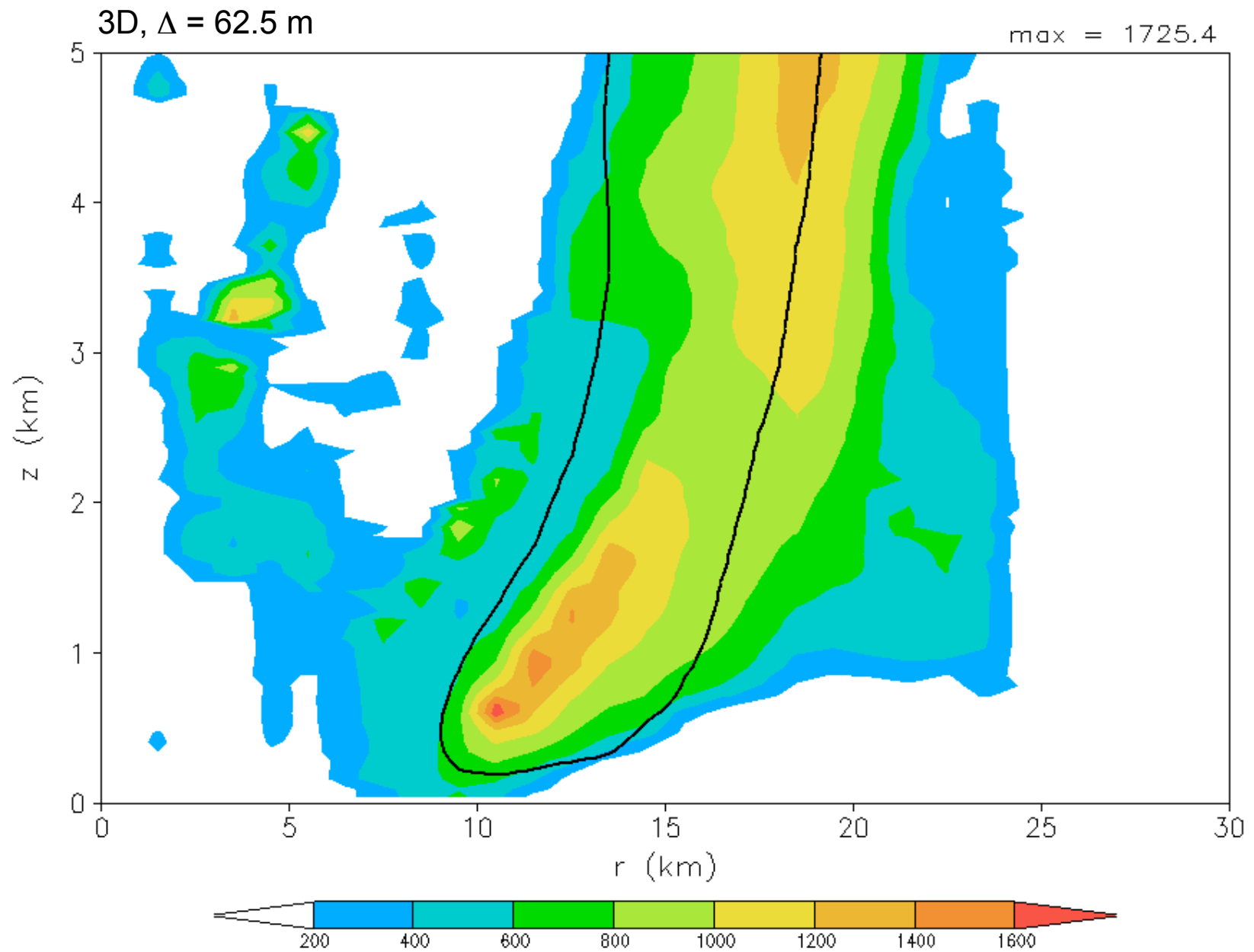
turbulence length scale:

$$\langle u' s' \rangle = -K \frac{\partial \langle s \rangle}{\partial r}$$

where:

$$K = l_h^2 \left[2 \left(\frac{\partial \langle u \rangle}{\partial r} \right)^2 + 2 \left(\frac{\langle u \rangle}{r} \right)^2 + \left(\frac{\partial \langle v \rangle}{\partial r} - \frac{\langle v \rangle}{r} \right)^2 \right]^{1/2}$$

turbulence length scale (l_h):



Summary

- Turbulence in the eyewall of hurricanes reduces hurricane intensity
- Very high resolution ($\Delta < 100$ m) in three dimensions is required to simulate turbulent processes (see also Rotunno et al. 2009, BAMS, in press)

... otherwise, turbulent processes must be *parameterized* (even with $\Delta \approx 1$ km)

... $l_h \approx 1000$ m

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