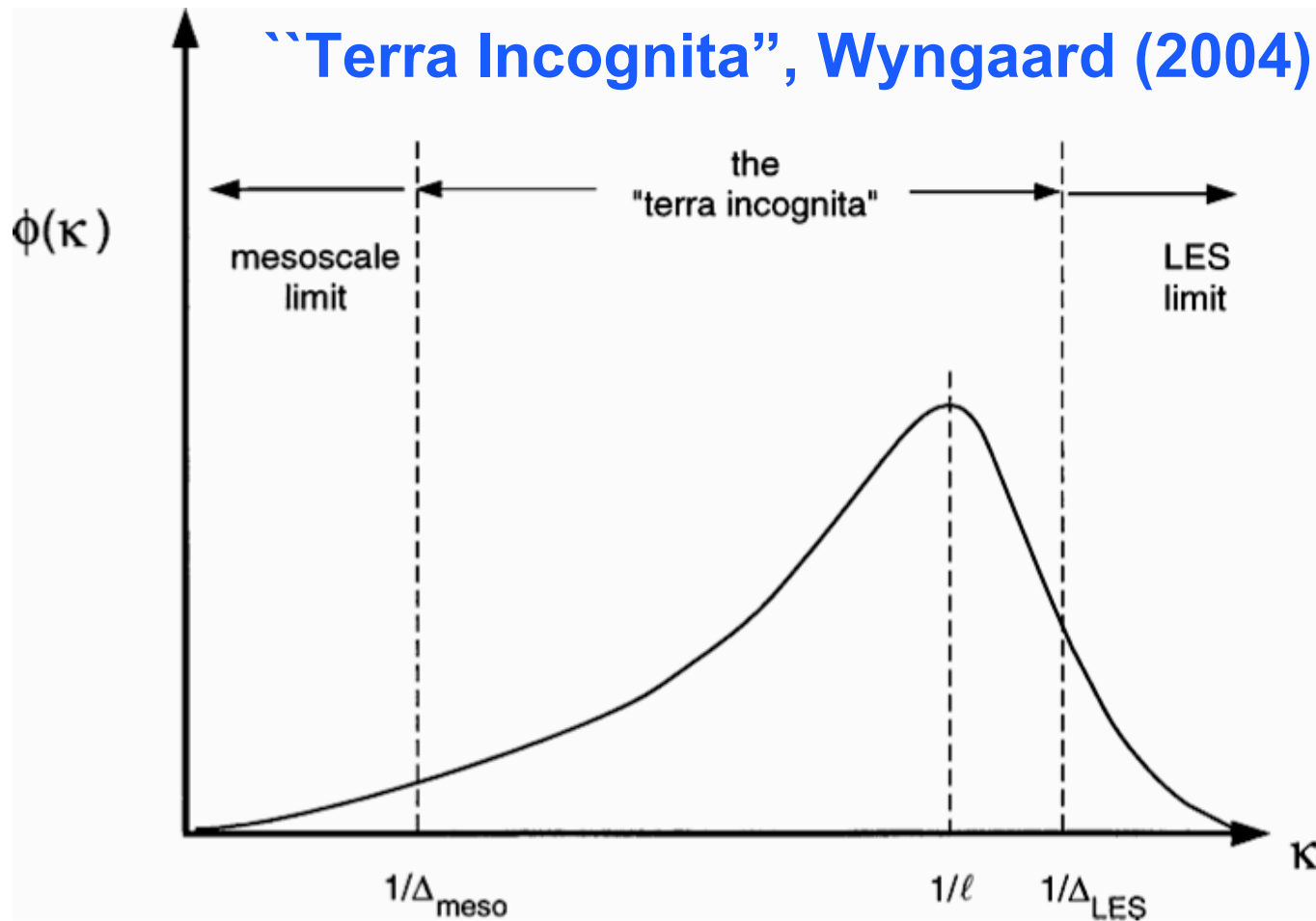


ATMOSPHERIC BOUNDARY LAYER SIMULATION AND SUBGRID-SCALE DYNAMICS

1. An Introduction to Outdoor LES
2. Observations and SGS Model Equations
3. LES Applications

Peter P. Sullivan
National Center for Atmospheric Research

“Terra Incognita”, Wyngaard (2004)



*Where is your
“LES” ?*

FIG. 1. A schematic of the turbulence spectrum $\phi(\kappa)$ in the horizontal plane as a function of the horizontal wavenumber magnitude κ . Its peak is at $\kappa \sim 1/l$, with l the length scale of the energetic eddies; Δ is the scale of the smoothing filter. In the mesoscale limit (left), $\Delta_{\text{meso}} \gg l$ and none of the turbulence is resolved. In the LES limit (right), $\Delta_{\text{LES}} \ll l$ and the energy-containing turbulence is resolved.

LARGE-EDDY SIMULATIONS AND OBSERVATIONS

Why LES of the PBL:

- Outdoor 4-D measurements are challenging
- Unsteady nature of the atmosphere and ocean
- Systematic investigation of the parameter space
- *Advances in parallel computing*



Gulf of Tehuantepec U ~[20-25] m/s

Validating/Improving LES with observations:

- Test the output
- Test the input subgrid-scale parameterizations

APPROACHES TO SUBGRID SCALE MODELING

SGS parameterizations are physical models for small scale processes and filters that stabilize the numerical scheme.

- Traditional approaches to building SGS models
 - Lean on theory, *e.g.*, analysis by Lilly and others
 - Filter DNS databases to obtain resolved and SGS pieces
- The literature is vast, especially in physics/engineering
 - Simple algebraic models *e.g.*, Smagorinsky
 - Stochastic approaches [*Mason and Thomson*(1992)]
 - One equation TKE models [*Deardorff*(1980)]
 - Hybrid TKE/RANS [*Sullivan et al.*(1994)]
 - Algebraic Reynolds stress models [*Schmidt and Schumann*(1989)]
 - Dynamic models [*Germano et al.*(1991)]
 - Nonlinear mixed models [*Meneveau and Katz*(2000)]
 - Deconvolution methods [*Stolz et al.*(2001)]
 - Velocity estimation models [*Dubrulle et al.*(2002)]
 - Algorithmic models [*Domaradzki and Horiuti*(2001)]
 - Transport equations [*Wyngaard*(2004)]

***Can we use targeted observations to
provide insight as to the nature of SGS
motions in high Re PBLs?***

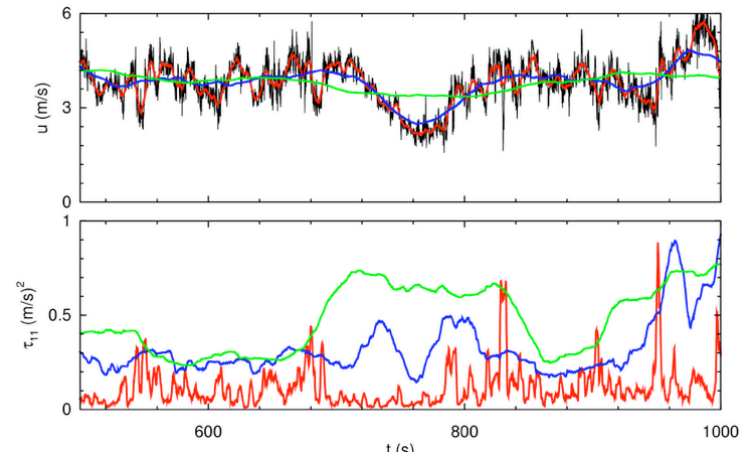
HIGH REYNOLDS NUMBER OBSERVATIONS AND LES

- **SINGLE-POINT MEASUREMENTS**

- Cannot be used directly to improve LES

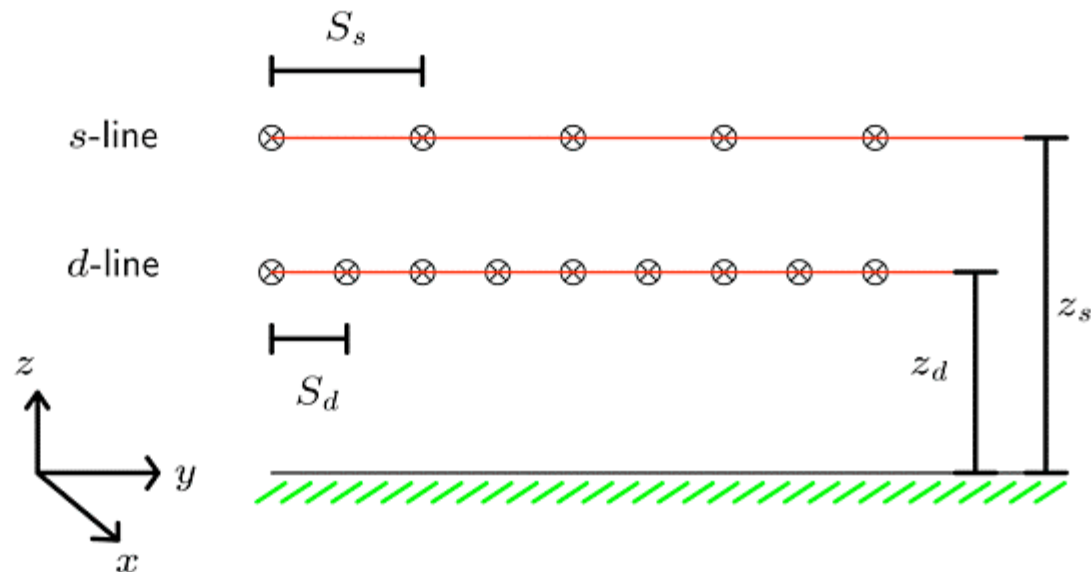
- **MULTI-POINT MEASUREMENTS**

- Span a range of filter widths, *e.g.*, $\mathcal{O}(\text{m})$ to $\mathcal{O}(100\text{m})$
- Ideally 3-D, time varying “volume” of turbulence and scalars in canonical flows with shear, stratification, near boundaries, ...
- Horizontal Array Turbulence Study field campaigns, HATS (2000), OHATS (2004), CHATS (2007), AHATS (2008)



HORIZONTAL ARRAY TURBULENCE STUDY (HATS)

- Field campaign to measure $\tau_{ij}, \tau_{i\theta}$ over a wide range of stratification Horst *et al.* (2002)
- Based on the horizontal array technique Tong *et al.* (1998), (1999) and Porté-Agel *et al.* (2001)
- 4 different sonic arrays, $-2 < z/L < 2$



RATIONALE FOR EXPERIMENTAL DESIGN

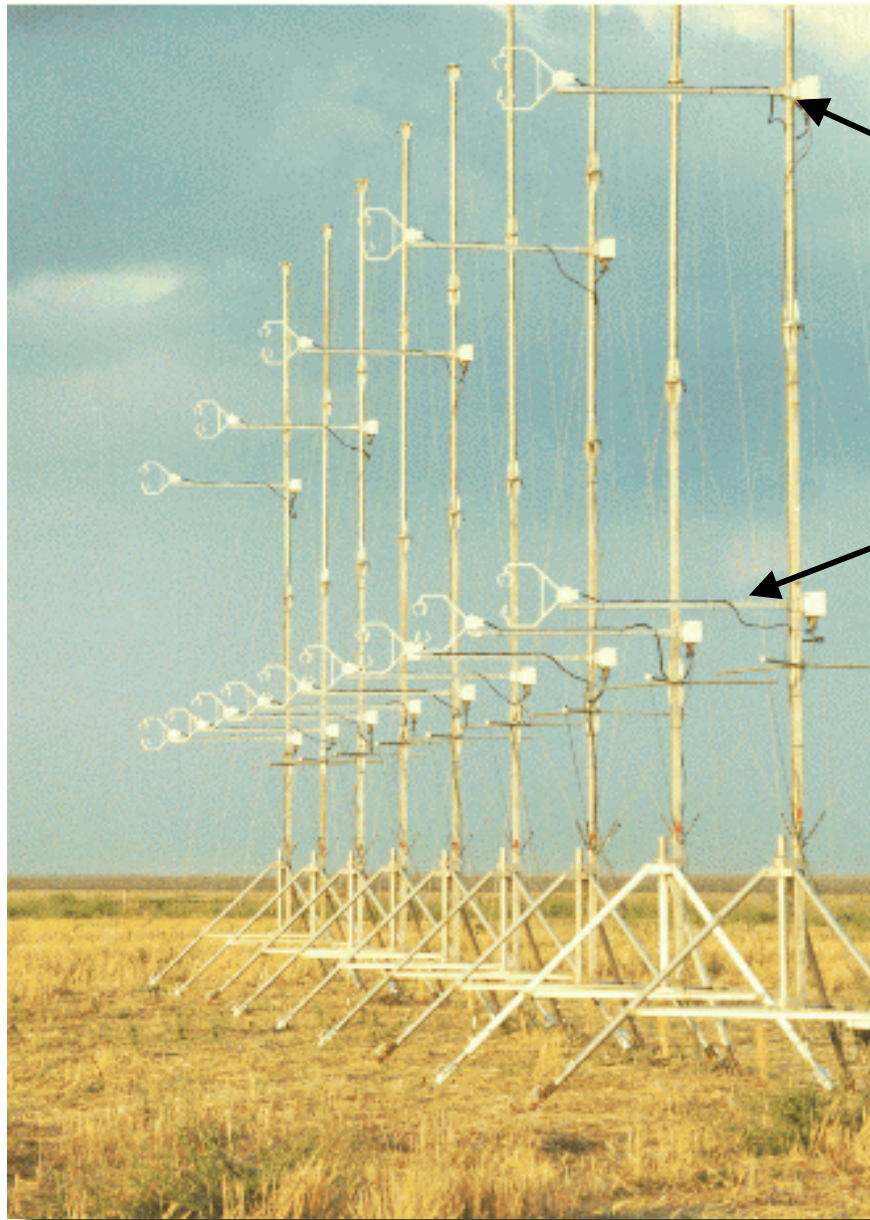
- Allows *spatial* filtering of flow field and decomposition into resolved and subfilter scale velocities $(\overline{U}_i, u_i)$:

$$U_i = \overline{U}_i + u_i \equiv \int U(x'_j) G(x_i, x'_j) dx'_j + u_i$$

- Allows construction of SFS fluxes:

$$\mathcal{T}_{ij} = \overline{U_i U_j} - \overline{U}_i \overline{U}_j$$

- Allows measurement of resolved gradients $\partial \overline{U}_i / \partial x$, $\partial \overline{U}_i / \partial y$ and $\partial \overline{U}_i / \partial z$
- Allows expansion of SFS fluxes $\mathcal{T}_{ij}$ into Leonard, Cross, and Reynolds terms which requires *double* spatial filtering, *e.g.*, $\overline{\overline{U}_i u_j}$



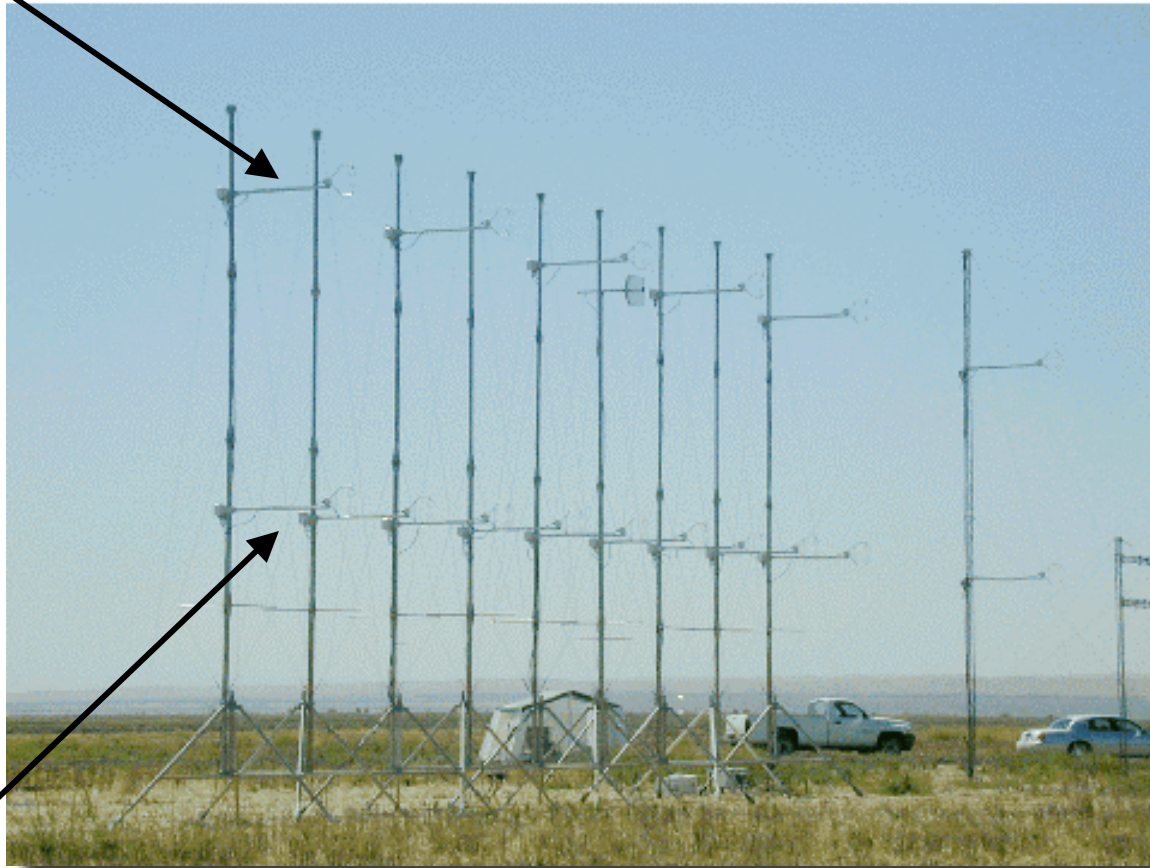
$z_s = 6.90\text{m}$, $dy_s = 6.70\text{m}$

$z_d = 3.45\text{m}$, $dy_d = 3.35\text{m}$

ARRAY-1

ARRAY-2

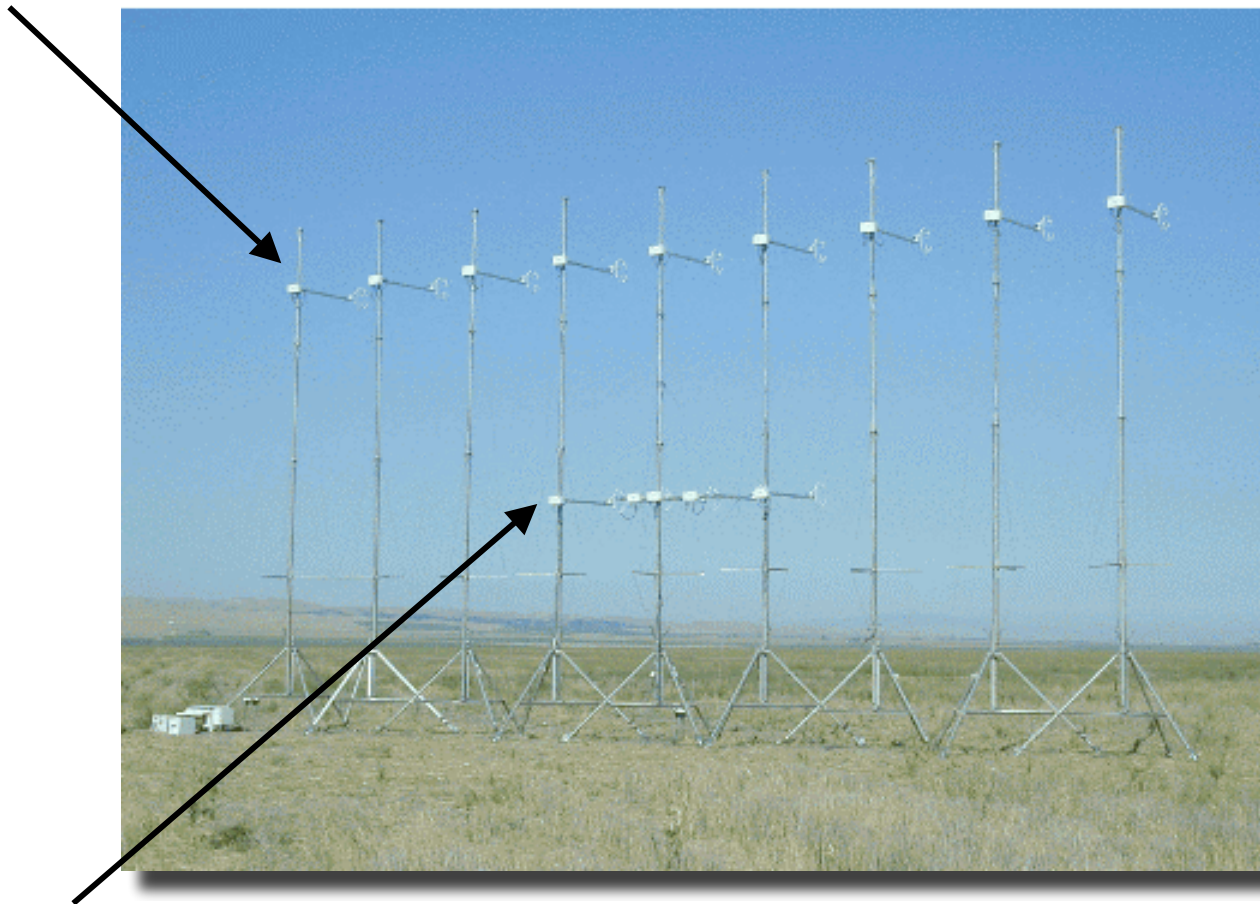
$z_s = 8.66\text{m}$, $dy_s = 4.33\text{m}$



$z_d = 4.33\text{m}$, $dy_d = 2.17\text{m}$

ARRAY-3

$z_d = 8.66\text{m}$, $dy_d = 2.17\text{m}$



$z_s = 4.33\text{m}$, $dy_s = 1.08\text{m}$

ARRAY-4

$$z_s = 5.15\text{m}$$

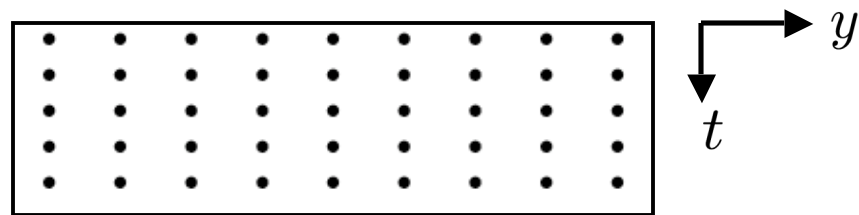
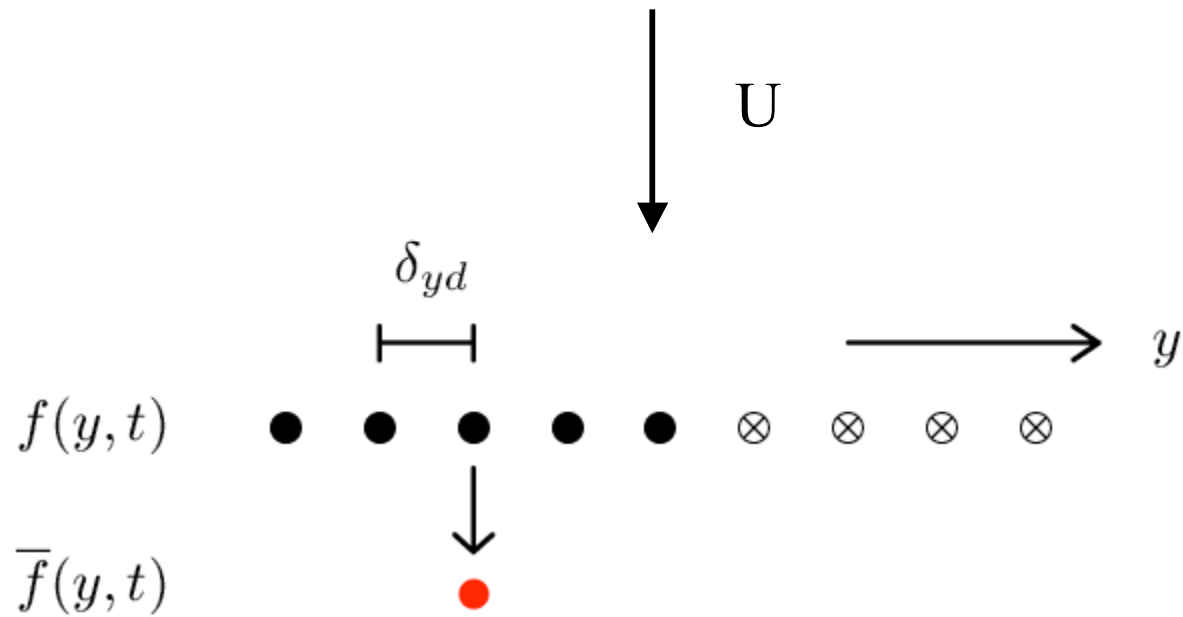
$$dy_s = 0.63\text{m}$$

$$z_d = 4.15\text{m}$$

$$dy_d = 0.50\text{m}$$

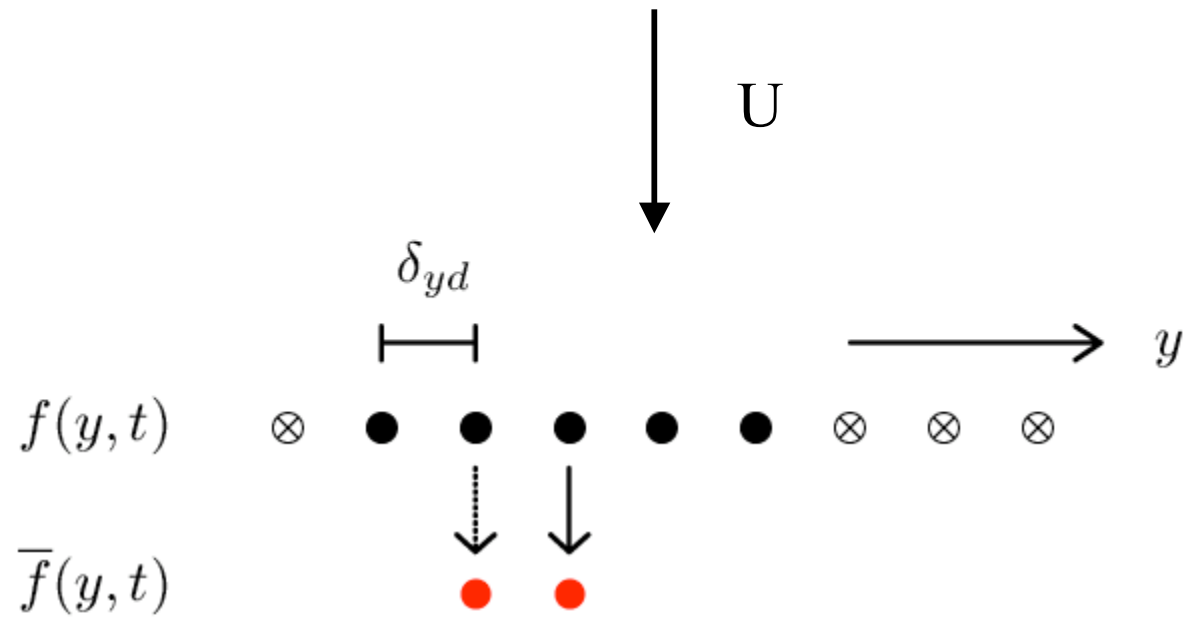


AN EXAMPLE OF LATERAL (Y) FILTERING

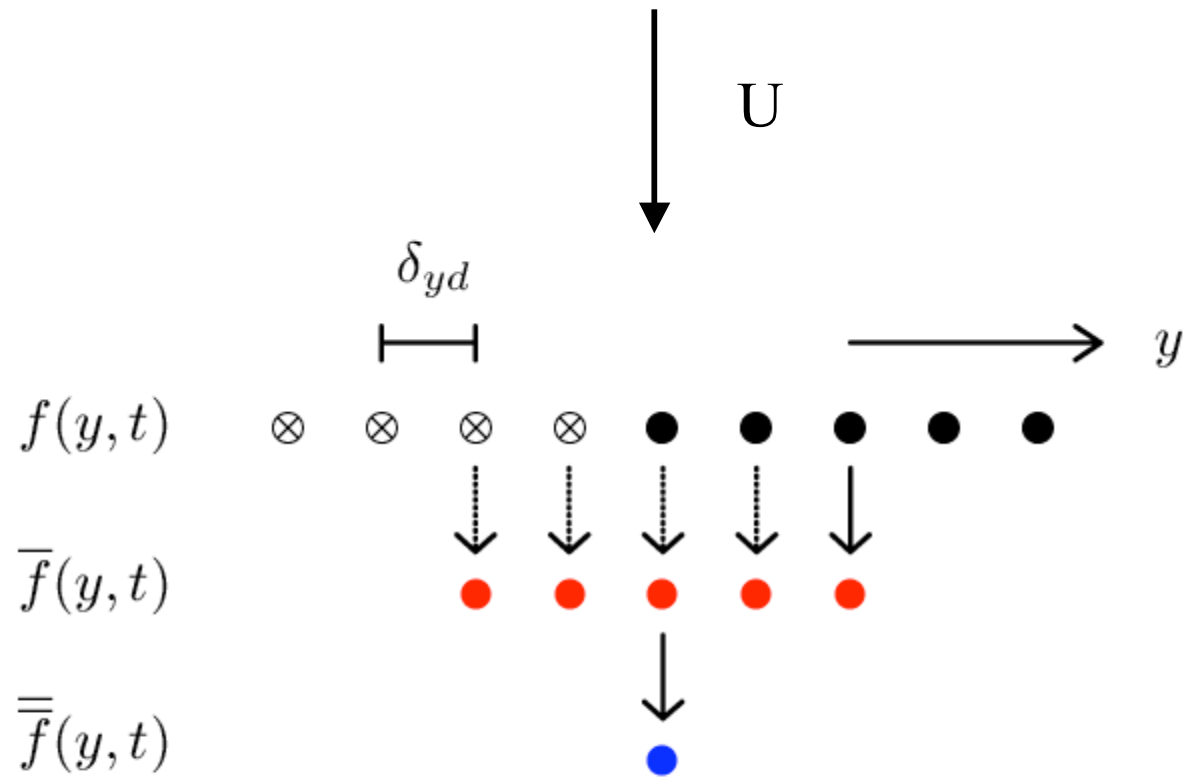


“2D plane of turbulence”

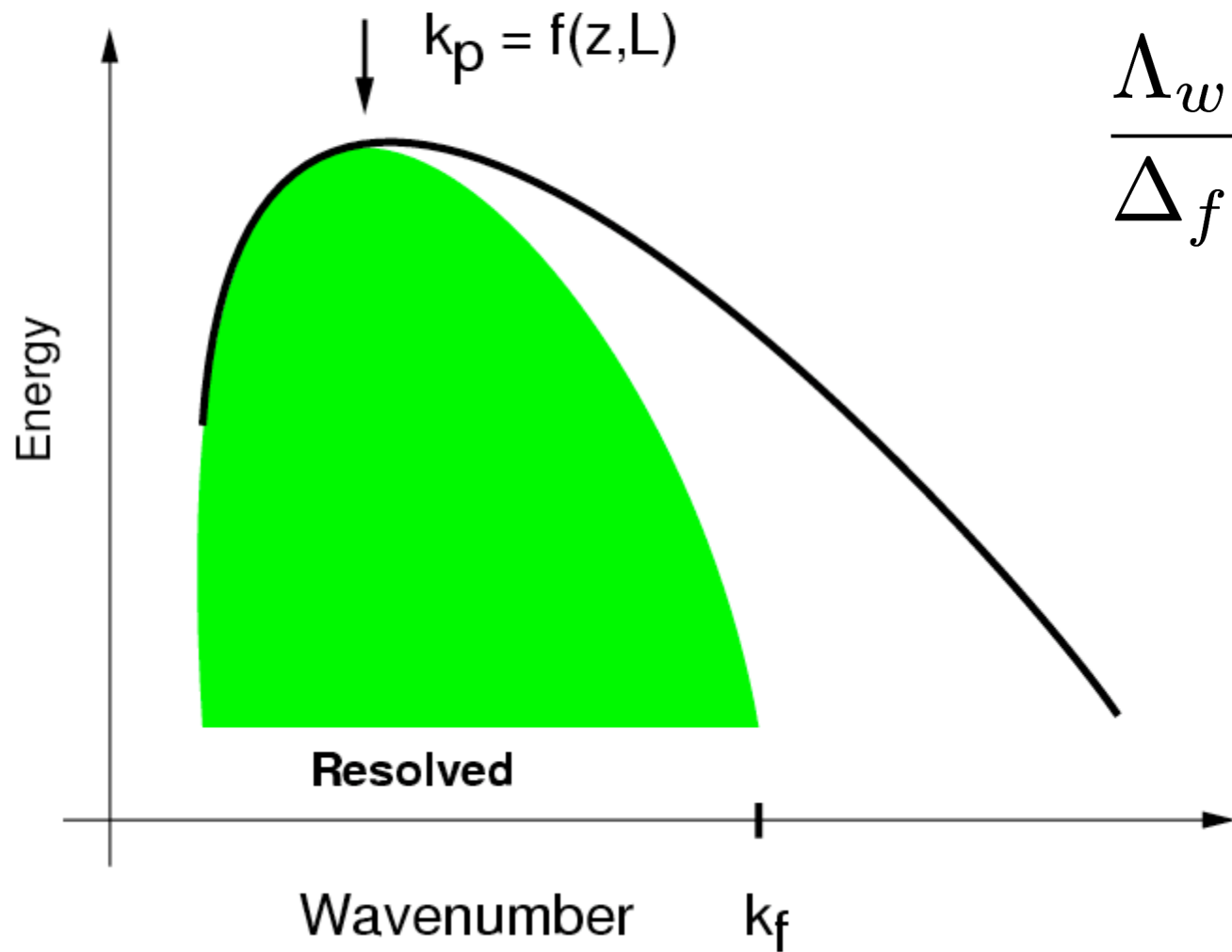
AN EXAMPLE OF LATERAL (Y) FILTERING



AN EXAMPLE OF LATERAL (Y) FILTERING

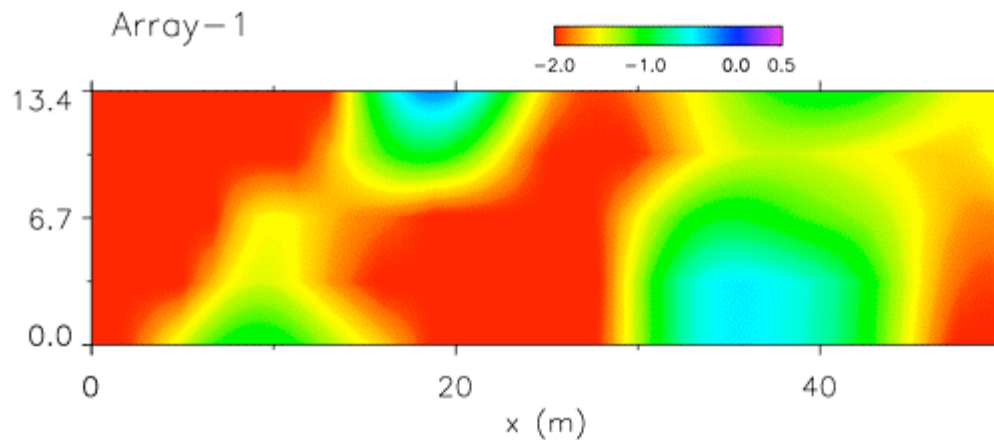
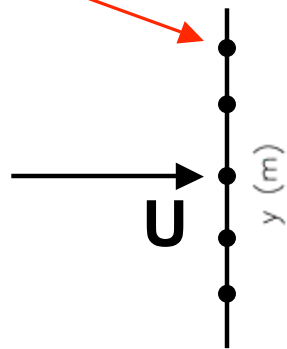


SPECTRAL PEAK AND FILTER CUTOFF WAVENUMBERS



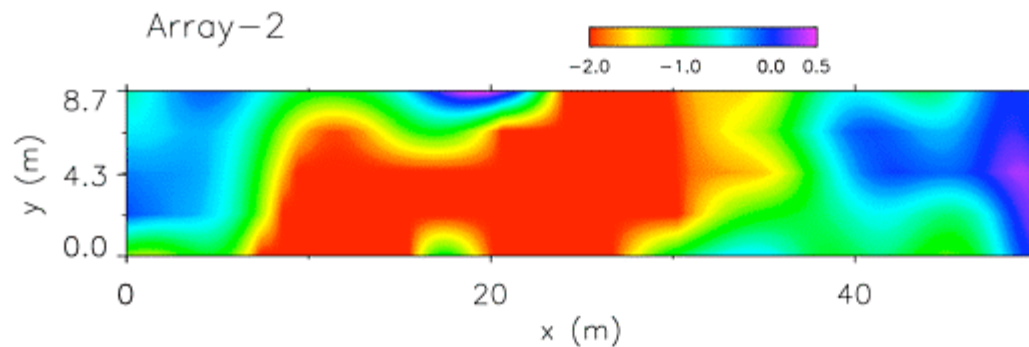
SFS Flux $\tau_{13} = \overline{U_1 U_3} - \overline{U_1} \overline{U_3}$ for Varying Filter Widths

Sonic array

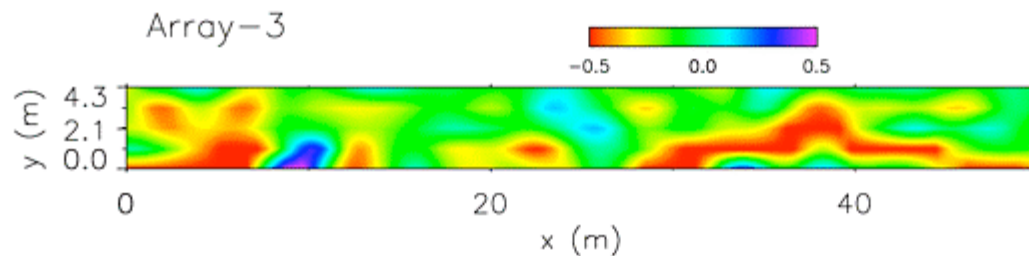


$$\frac{\Lambda_w}{\Delta_f} = 0.58$$

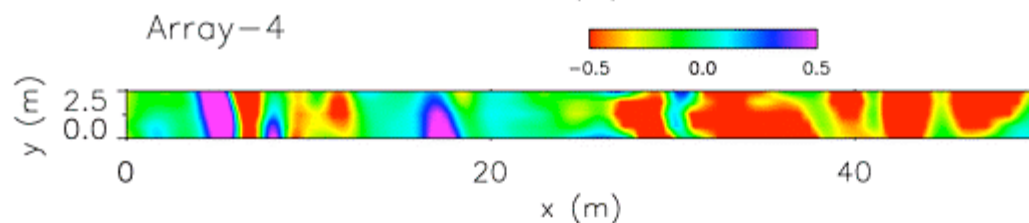
top view



$$1.18$$

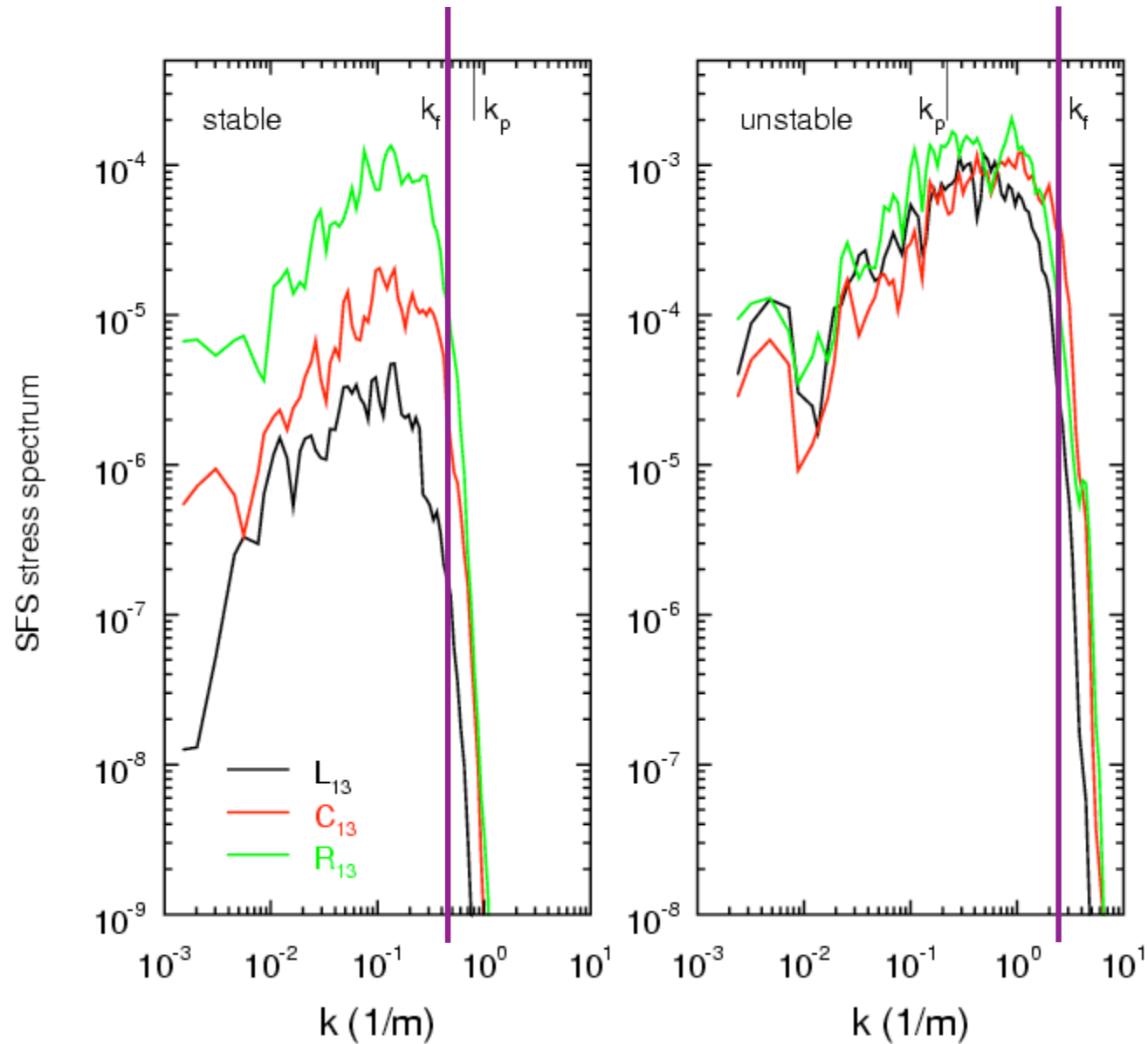


$$5.00$$

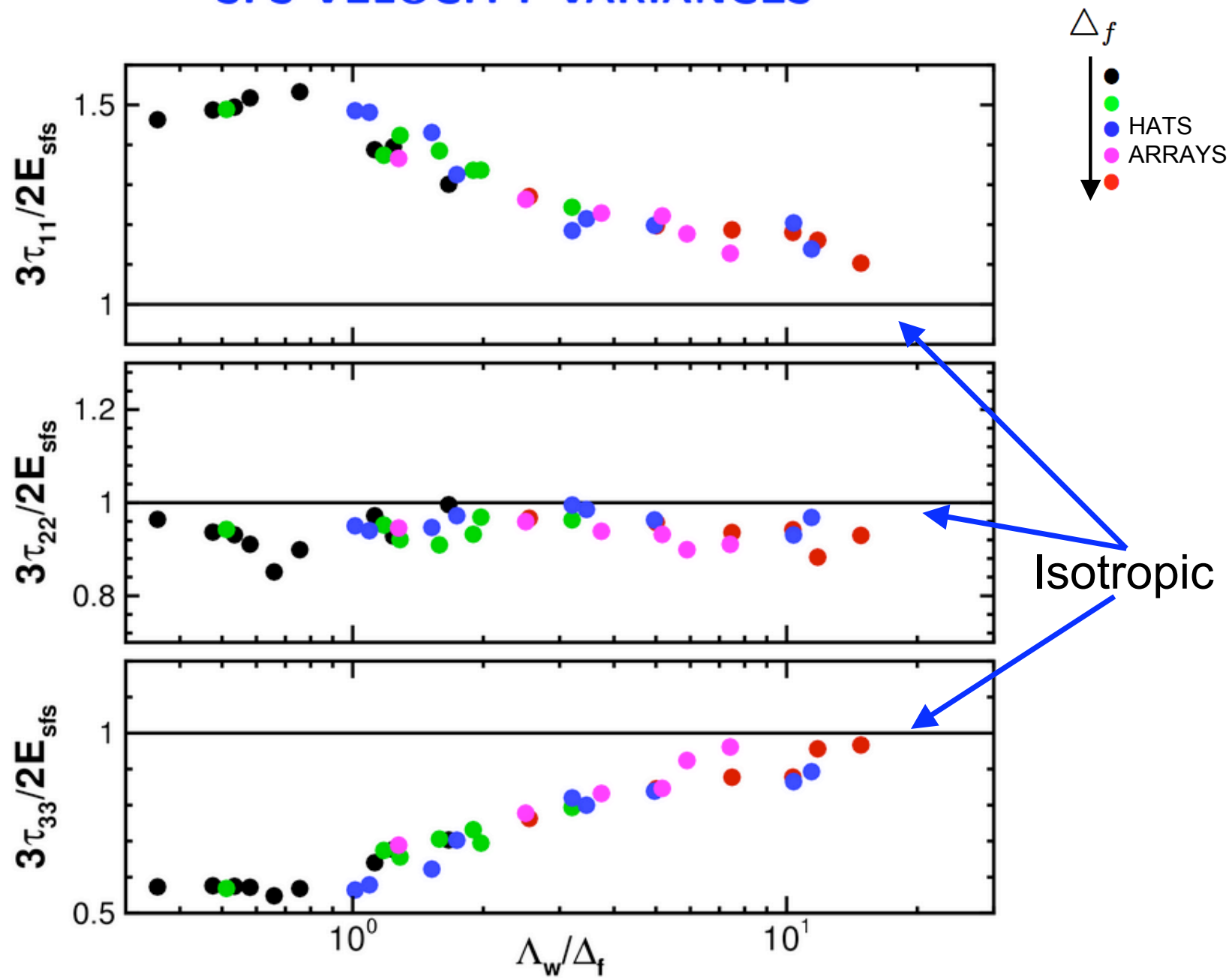


$$11.4$$

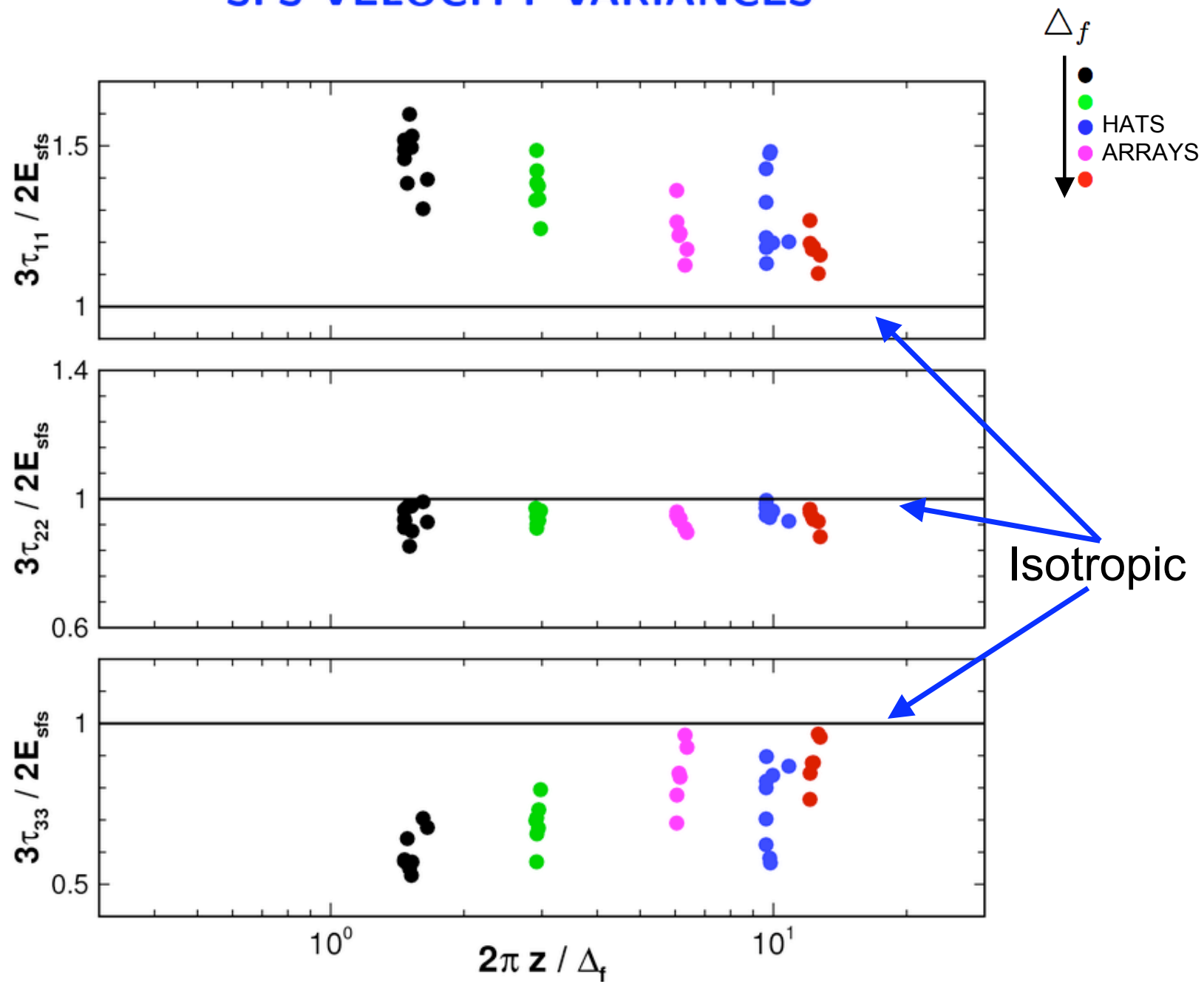
Spectra of Leonard, Cross, Reynolds Terms for (1,3) Component



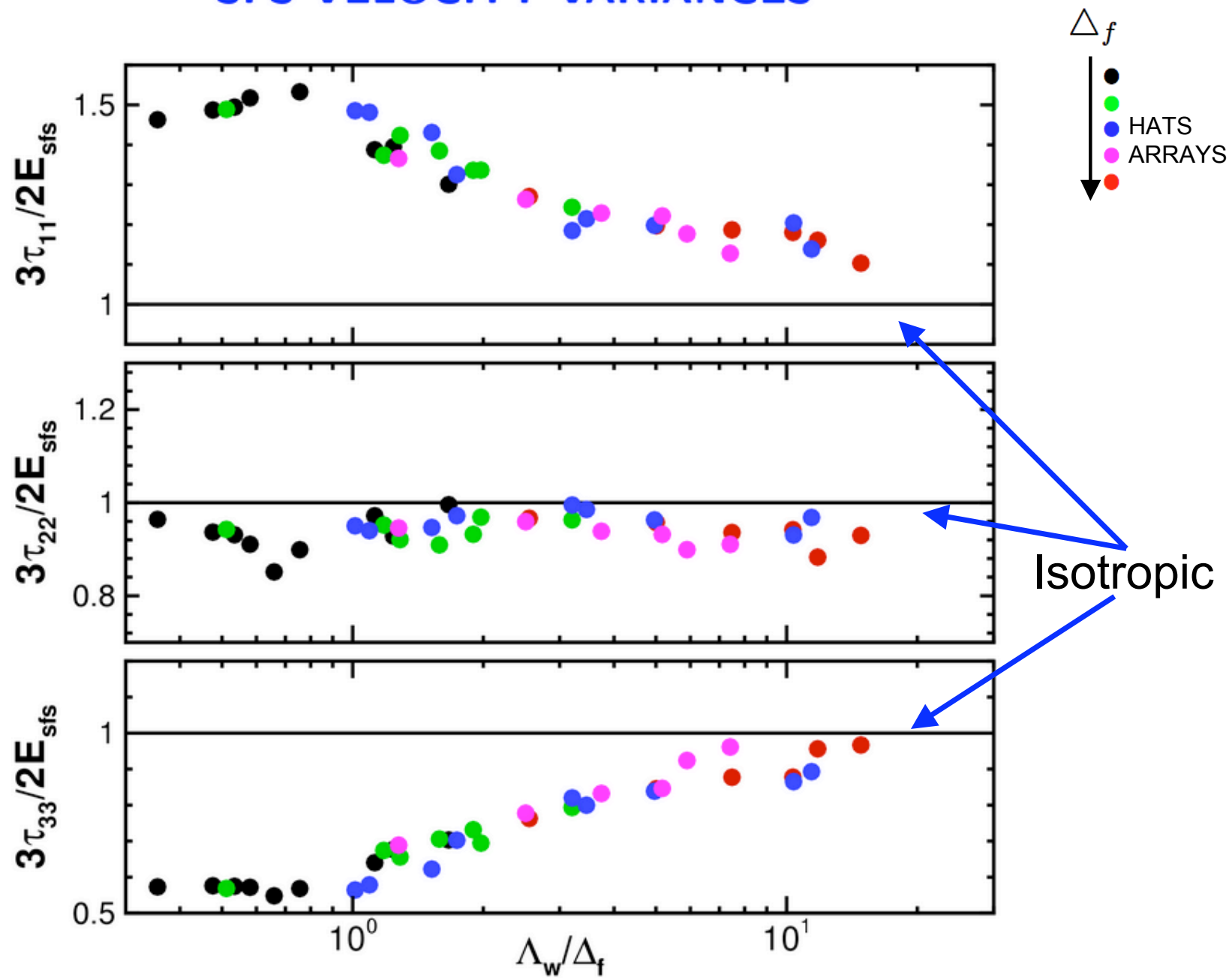
SFS VELOCITY VARIANCES



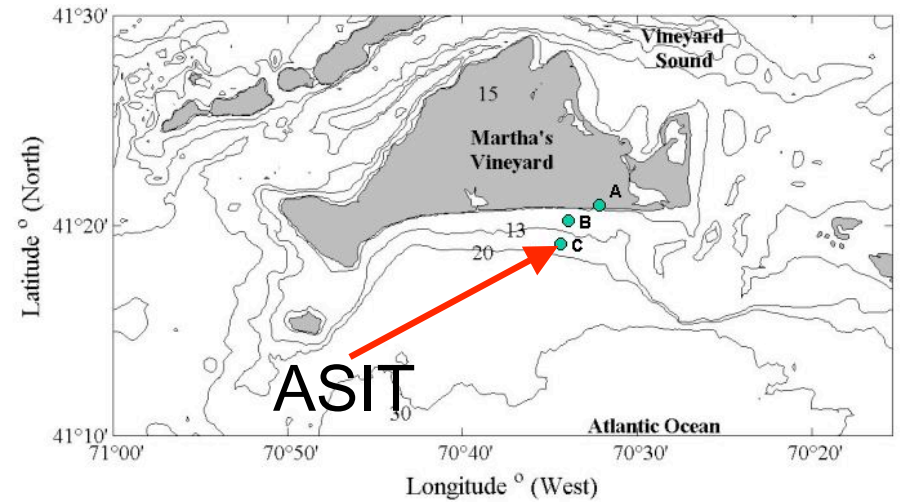
SFS VELOCITY VARIANCES



SFS VELOCITY VARIANCES



OHATS FIELD CAMPAIGN

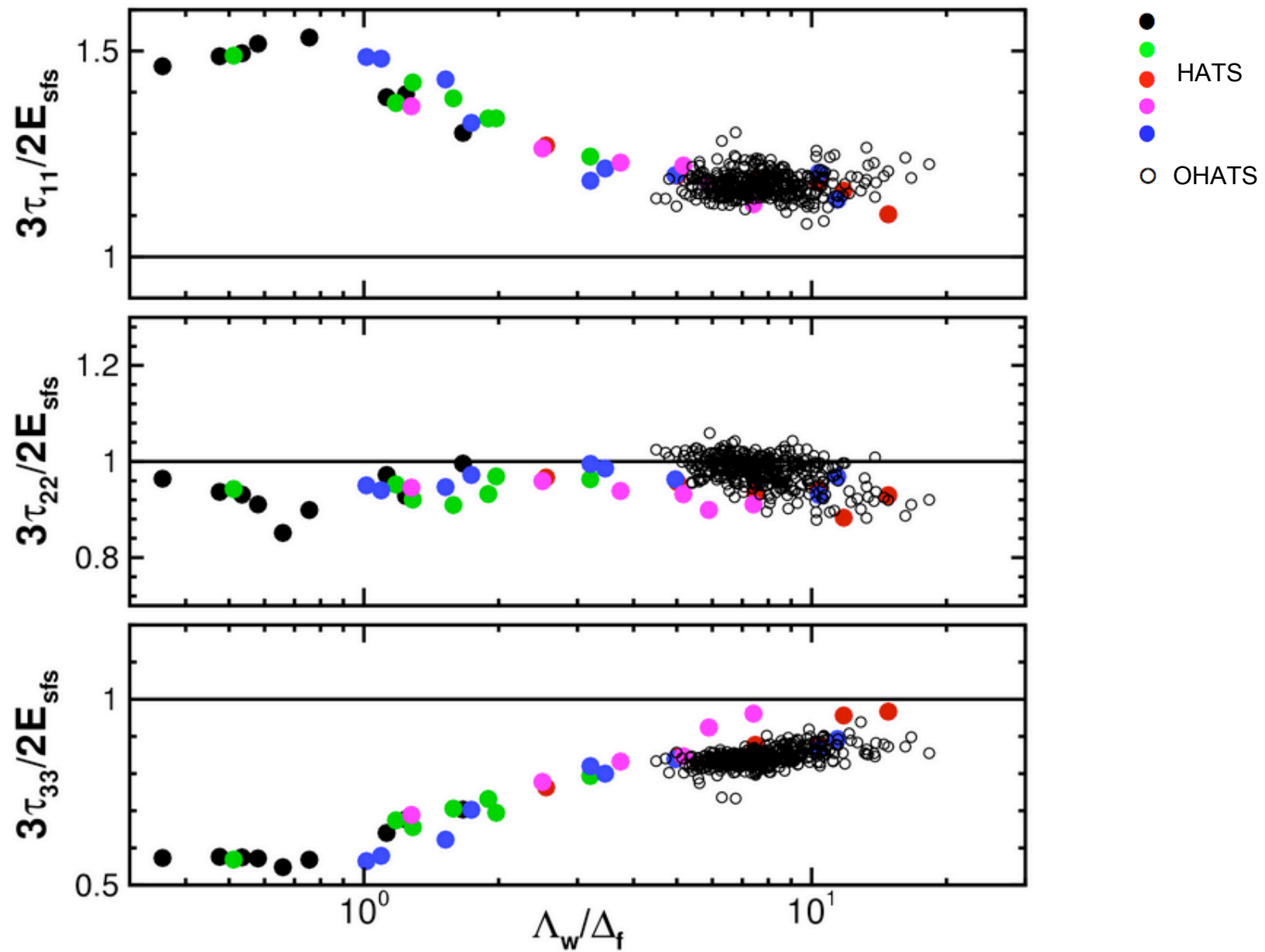


Laser altimeters

18 CSATS

275 hours ``12 days of data''
analyzed

SFS VELOCITY VARIANCES



RATE EQUATIONS FOR SUBGRID DEVIATORIC STRESS

- What are the parent equations for the Smagorinsky model?

RATE EQUATIONS FOR SUBGRID DEVIATORIC STRESS

The SGS stress is

$$\tau_{ij} = \overline{u_i u_j} - \bar{u}_i \bar{u}_j$$

To get the “rate equation” for SGS τ_{ij}

$$\frac{\partial \tau_{ij}}{\partial t} = \left[\overline{u_j \frac{\partial u_i}{\partial t}} - \bar{u}_j \frac{\partial \bar{u}_i}{\partial t} \right]$$

Substitution steps:

$$u_j \frac{\partial u_i}{\partial t} = \textcircled{u_j \mathcal{R}_i}$$

$$\bar{u}_j \frac{\partial \bar{u}_i}{\partial t} = \textcircled{\bar{u}_j \bar{\mathcal{R}}_i}$$

The difference is now τ_{ij} is the deviatoric stress, *i.e.*, $-2/3e\delta_{ij}$

Considerable Algebra !

RATE EQUATIONS FOR SUBGRID DEVIATORIC STRESS

- **What are the parent equations for the Smagorinsky model?**
 - Lilly (1967), Deardorff (1973), Wyngaard (2004), Hatlee & Wyngaard (2007)

$$\begin{aligned}
 \frac{D\tau_{ij}}{Dt} = & \frac{2}{3}e \left(\frac{\partial \bar{u}_i}{\partial x_j} + \frac{\partial \bar{u}_j}{\partial x_i} \right) \quad \text{Isotropic production} \\
 & - \left[\tau_{ik} \frac{\partial \bar{u}_j}{\partial x_k} + \tau_{jk} \frac{\partial \bar{u}_i}{\partial x_k} - \frac{1}{3} \delta_{ij} \tau_{kl} \left(\frac{\partial \bar{u}_k}{\partial x_l} + \frac{\partial \bar{u}_l}{\partial x_k} \right) \right] \\
 & - \frac{1}{\rho} \left[\overline{p \left(\frac{\partial u_i}{\partial x_j} + \frac{\partial u_j}{\partial x_i} \right)} - \bar{p} \left(\frac{\partial \bar{u}_i}{\partial x_j} + \frac{\partial \bar{u}_j}{\partial x_i} \right) \right] \\
 & + \text{transport} + \text{buoyancy production}
 \end{aligned}$$

Pressure destruction

Anisotropic deviatoric production

RATE EQUATIONS FOR SUBGRID DEVIATORIC STRESS

- What are the parent equations for the Smagorinsky model?
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 & + \text{transport} + \text{buoyancy production}
 \end{aligned}$$

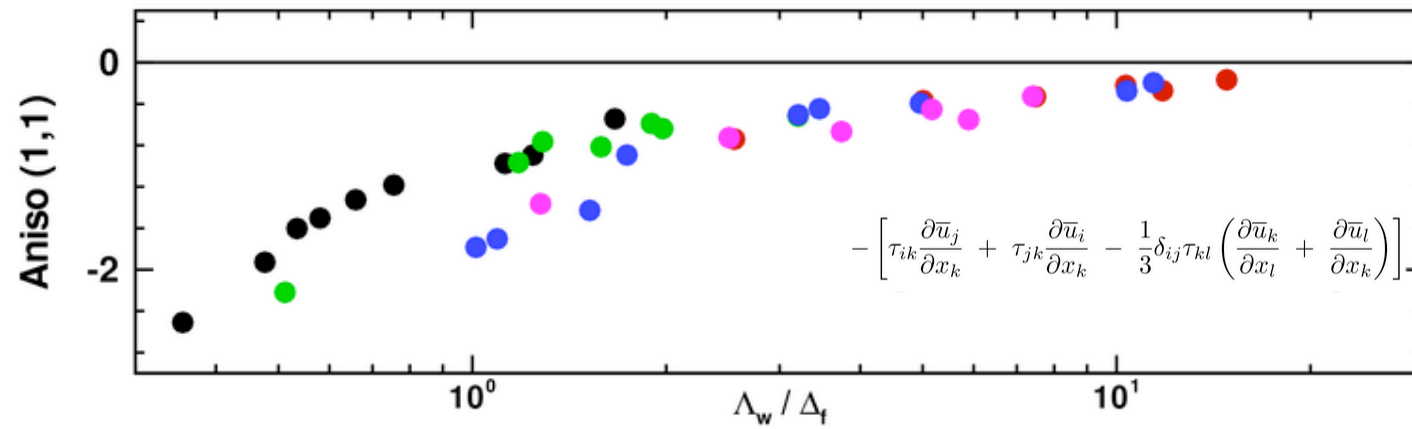
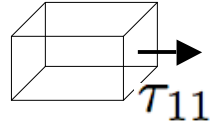
Rotta model

$$\frac{\tau_{ij}}{T} = \frac{2}{3}e \left(\frac{\partial \bar{u}_i}{\partial x_j} + \frac{\partial \bar{u}_j}{\partial x_i} \right)$$

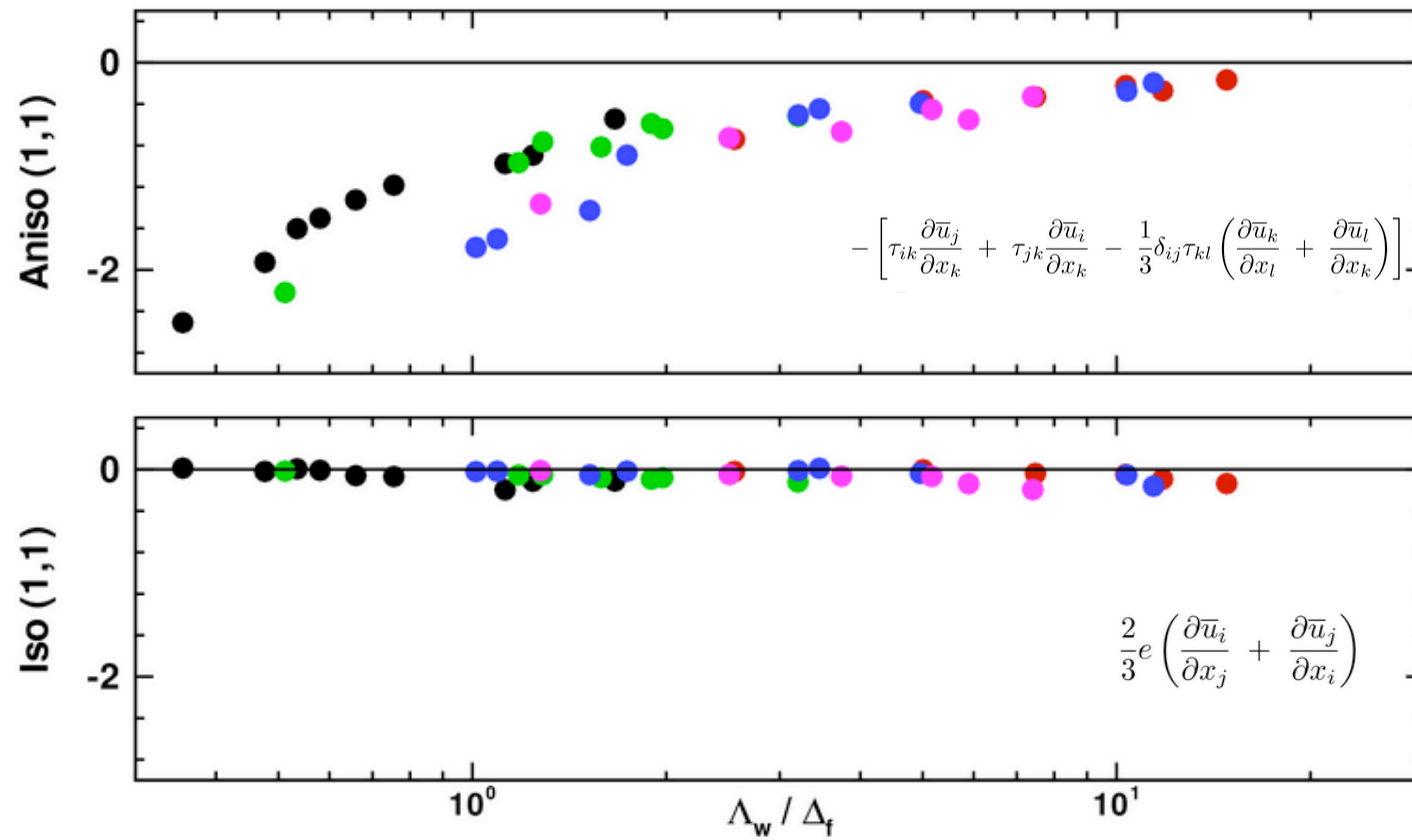
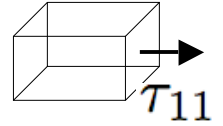
$$T = c \frac{\Delta_f}{\sqrt{e}}$$

Time scale

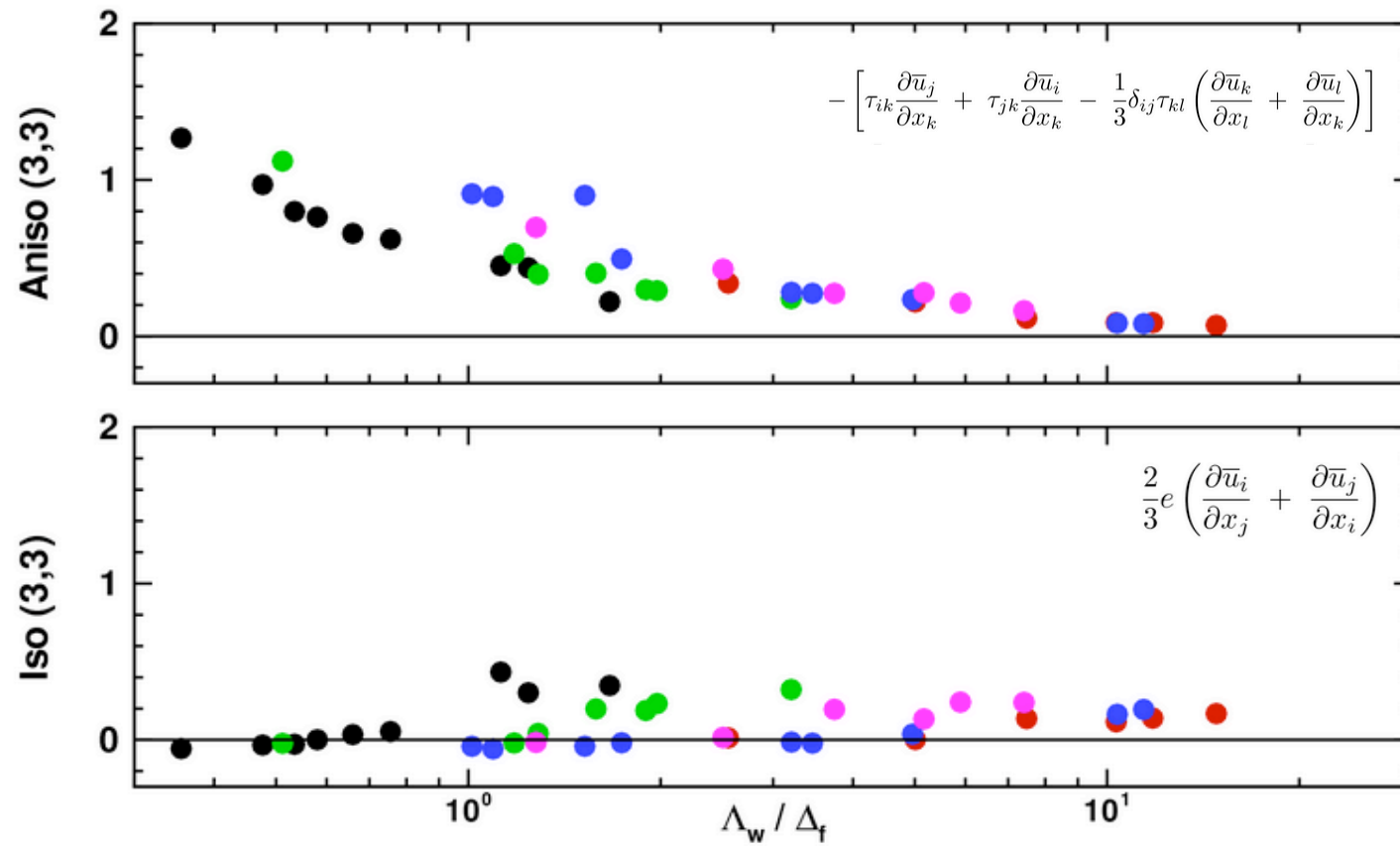
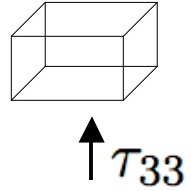
PRODUCTION OF SUBFILTER SCALE FLUX τ_{11}



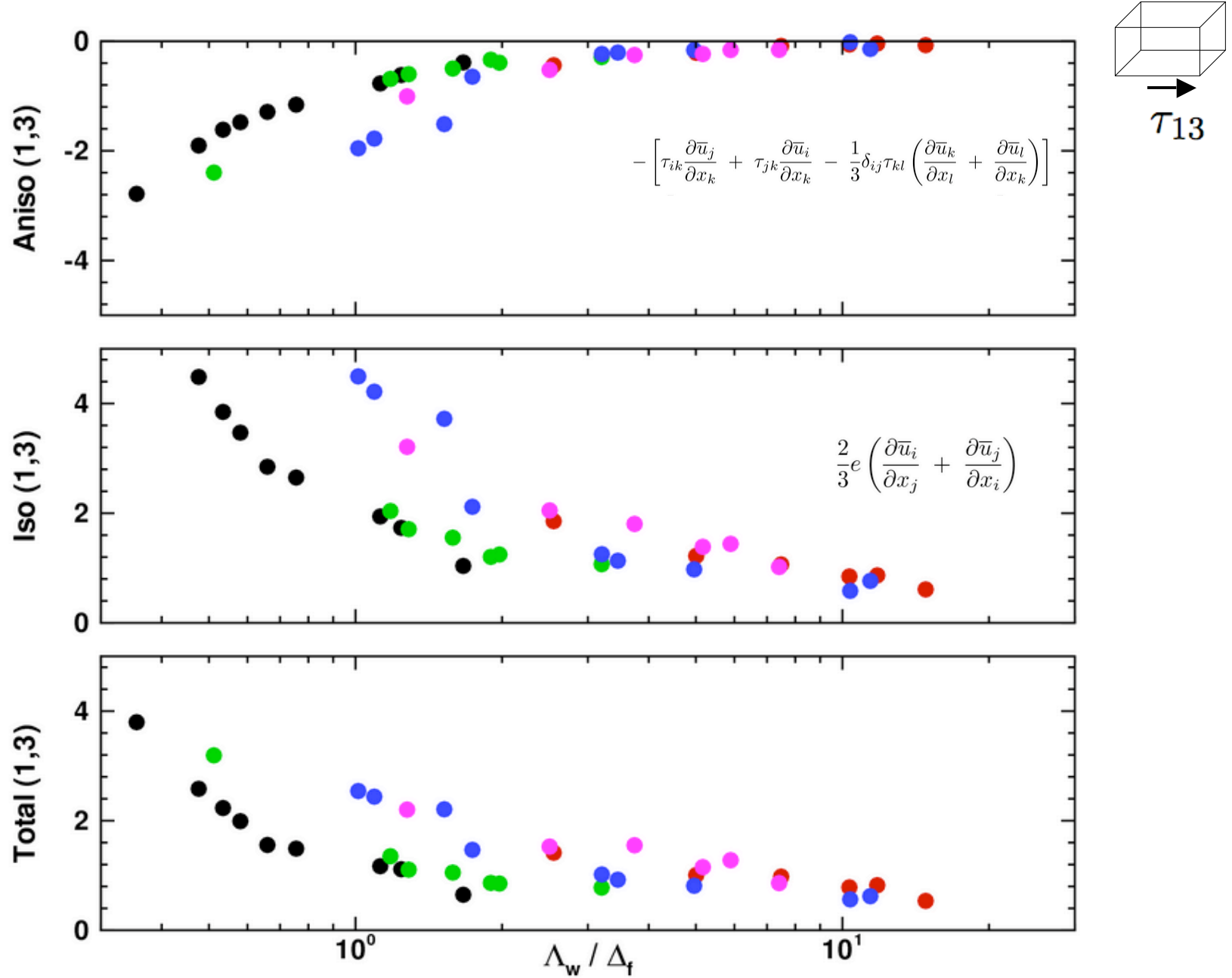
PRODUCTION OF SUBFILTER SCALE FLUX τ_{11}



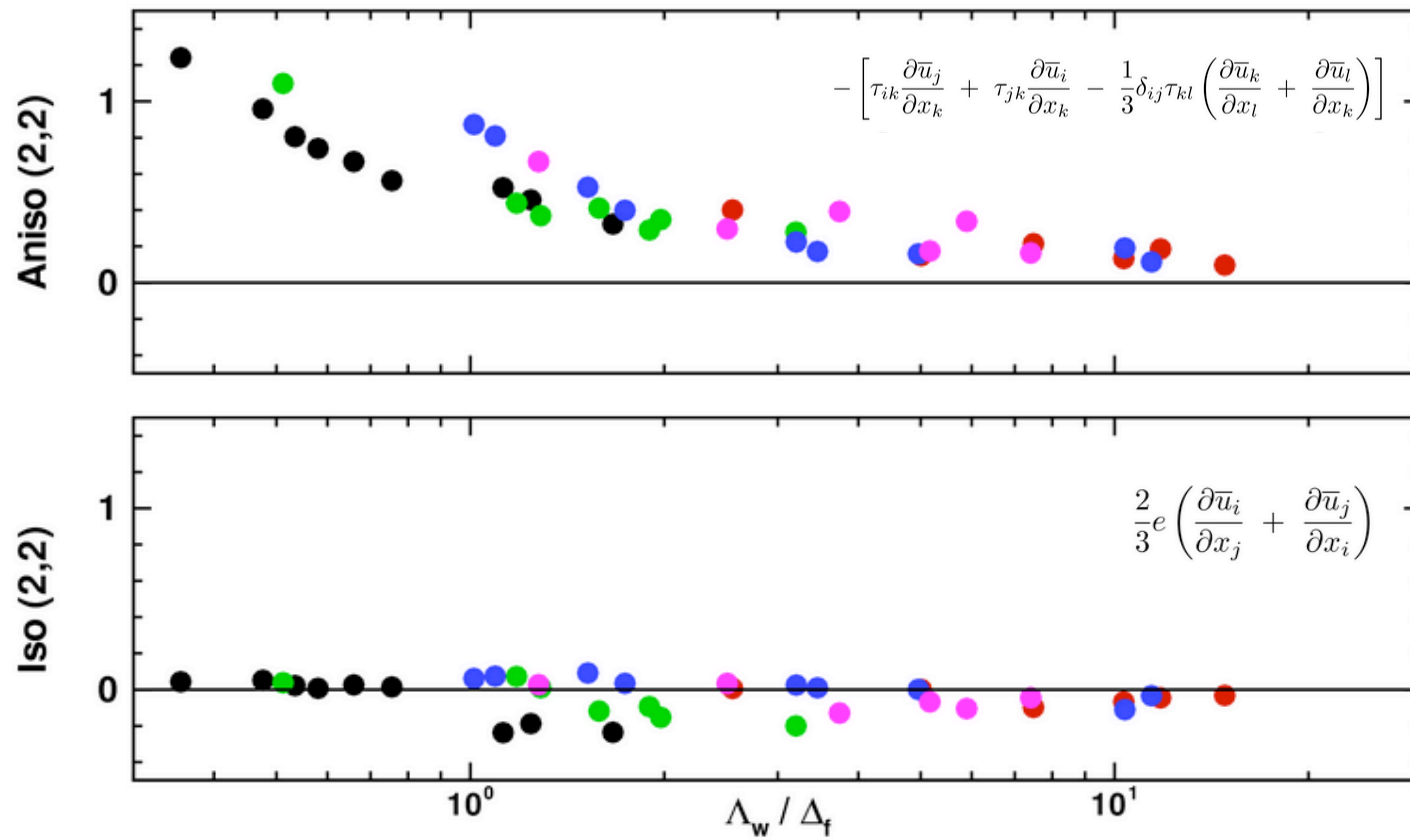
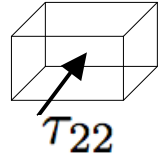
PRODUCTION OF SUBFILTER SCALE FLUX τ_{33}



PRODUCTION OF SUBFILTER SCALE FLUX τ_{13}



PRODUCTION OF SUBFILTER SCALE FLUX τ_{22}



VARIATION OF DEVIATORIC STRESS IN LIMIT $\Lambda_w/\Delta_f \rightarrow 0$

$\langle \tau_{11} \rangle = T \left(-2 \langle \tau_{13} \rangle \frac{\partial U}{\partial z} + \frac{2}{3} \epsilon \right)$	$\langle \tau_{11} \rangle = 0$
$\langle \tau_{22} \rangle = T \left(\frac{2}{3} \epsilon \right)$	$\langle \tau_{22} \rangle = 0$
$\langle \tau_{33} \rangle = T \left(\frac{2}{3} \epsilon \right)$	$\langle \tau_{33} \rangle = 0$
$\langle \tau_{13} \rangle = T \left(\frac{2}{3} e \frac{\partial U}{\partial z} - \langle \tau_{33} \rangle \frac{\partial U}{\partial z} \right)$	$\langle \tau_{13} \rangle = T \left(\frac{2}{3} e \frac{\partial U}{\partial z} \right)$

Steady-state rate equations

Smagorinsky model

WHAT ABOUT SCALARS?

RATE EQUATIONS FOR SUBGRID SCALAR FLUX

- What are the parent equations for subgrid-scale scalar flux?

$$f_i = \overline{u_i c} - \bar{u}_i \bar{c}$$

$$\begin{aligned} \frac{Df_i}{Dt} = & -\frac{2}{3}e \frac{\partial \bar{c}}{\partial x_i} \quad \leftarrow \text{Isotropic production} \\ & -f_j \frac{\partial \bar{u}_i}{\partial x_j} + \tau_{ij} \frac{\partial \bar{c}}{\partial x_j} \\ & + \frac{1}{\rho} \left(\overline{p \frac{\partial c}{\partial x_i}} - \bar{p} \frac{\partial \bar{c}}{\partial x_i} \right) \\ & + \text{transport} + \text{buoyancy} \end{aligned}$$

Pressure destruction

Anisotropic production

RATE EQUATIONS FOR SUBGRID SCALAR FLUX

- What are the parent equations for subgrid-scale scalar flux?

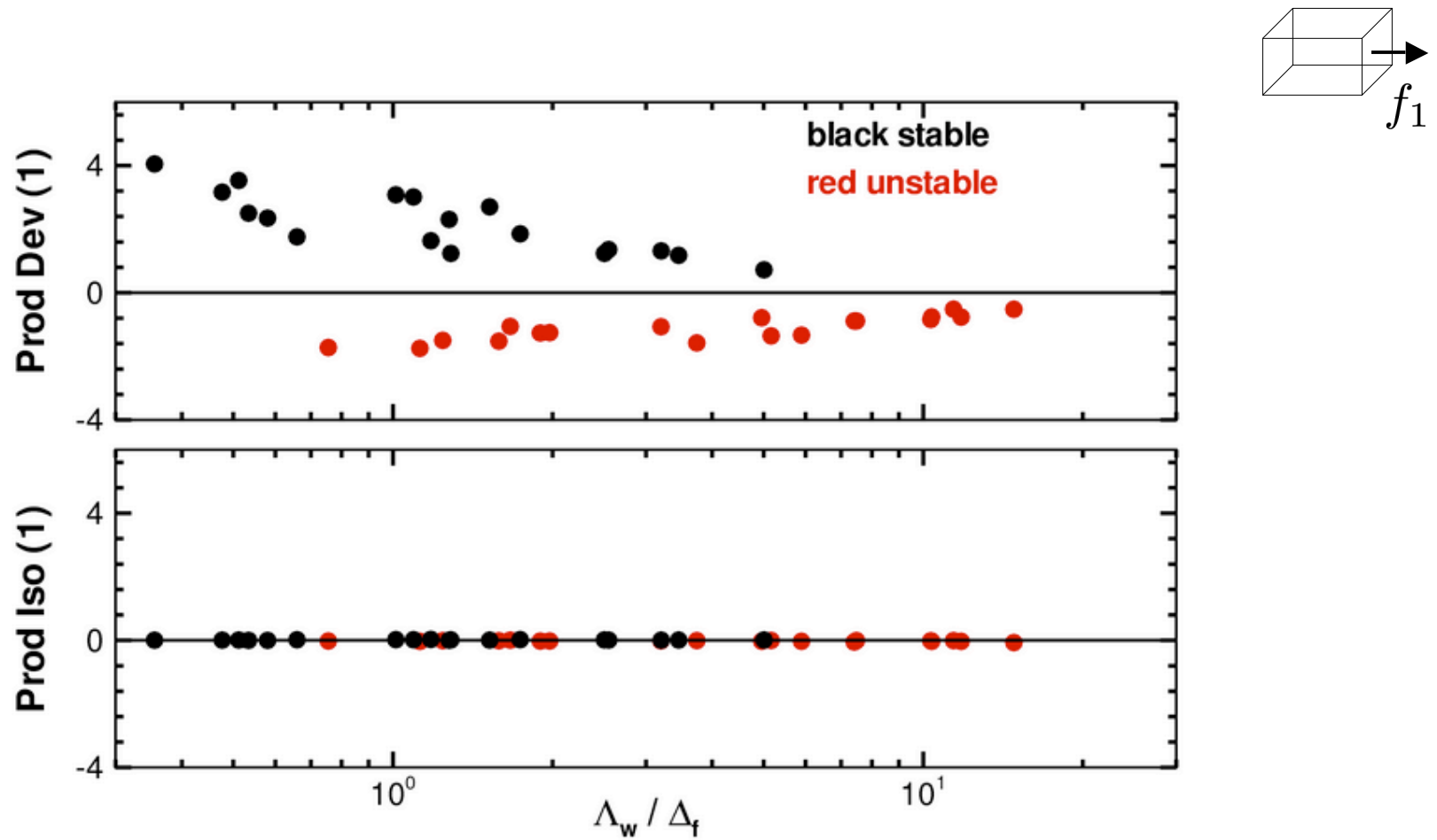
$$f_i = \overline{u_i c} - \bar{u}_i \bar{c}$$

$$\frac{Df_i}{Dt} = -\frac{2}{3}e \frac{\partial \bar{c}}{\partial x_i} - f_j \frac{\partial \bar{u}_i}{\partial x_j} + \tau_{ij} \frac{\partial \bar{c}}{\partial x_j} + \frac{1}{\rho} \left(\overline{p \frac{\partial c}{\partial x_i}} - \bar{p} \frac{\partial \bar{c}}{\partial x_i} \right) + \text{transport} + \text{buoyancy}$$

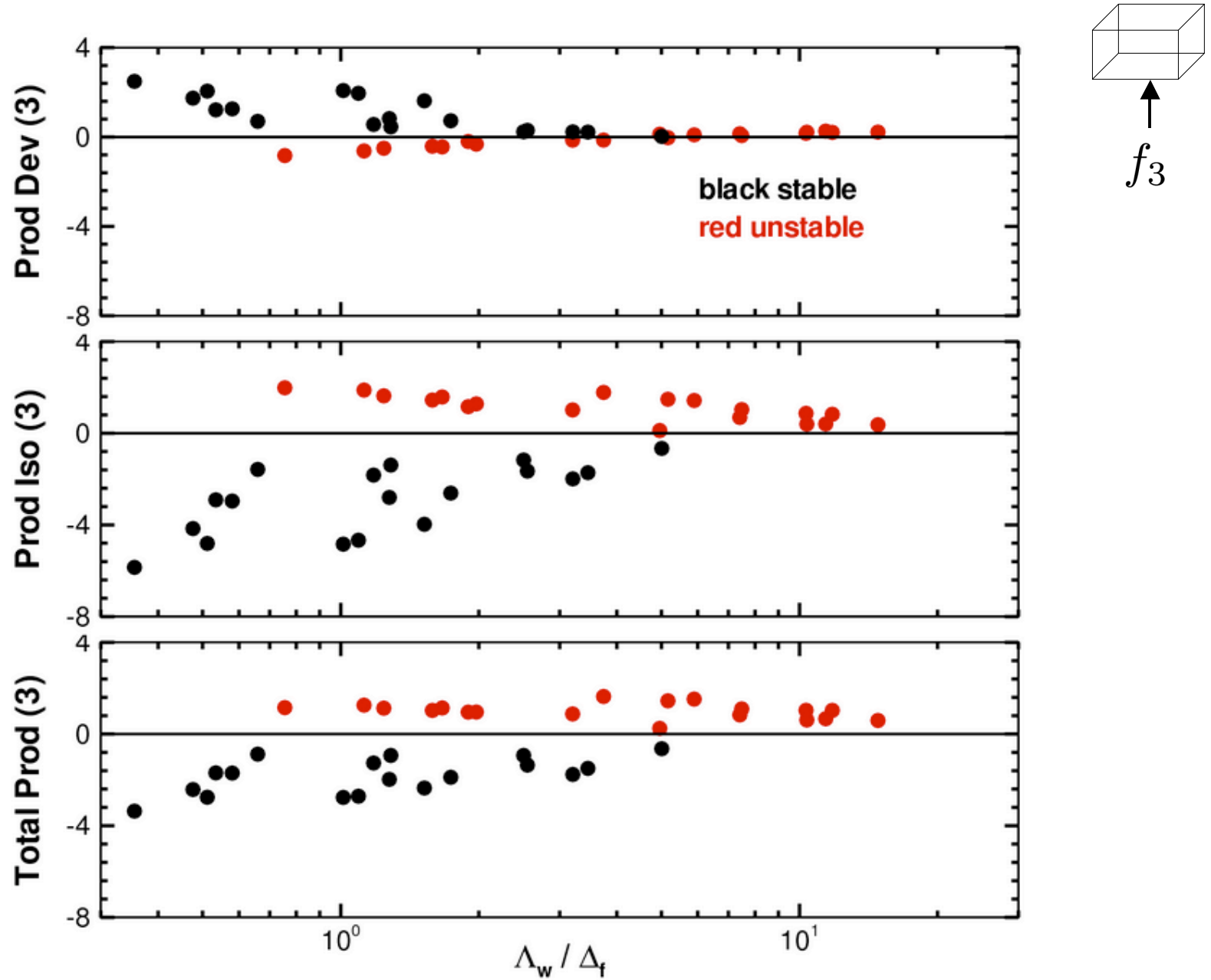
Eddy viscosity model

$$f_i = -\nu_h \frac{\partial \bar{c}}{\partial x_i} \quad \nu_h = \frac{2c_h \Delta_f \sqrt{e}}{3}$$

PRODUCTION OF SUBFILTER SCALE SCALAR FLUX f_1



PRODUCTION OF SUBFILTER SCALE SCALAR FLUX f_3

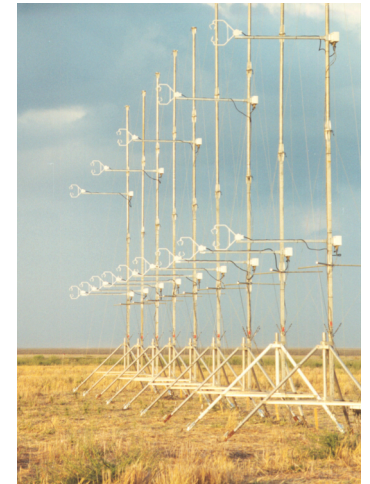
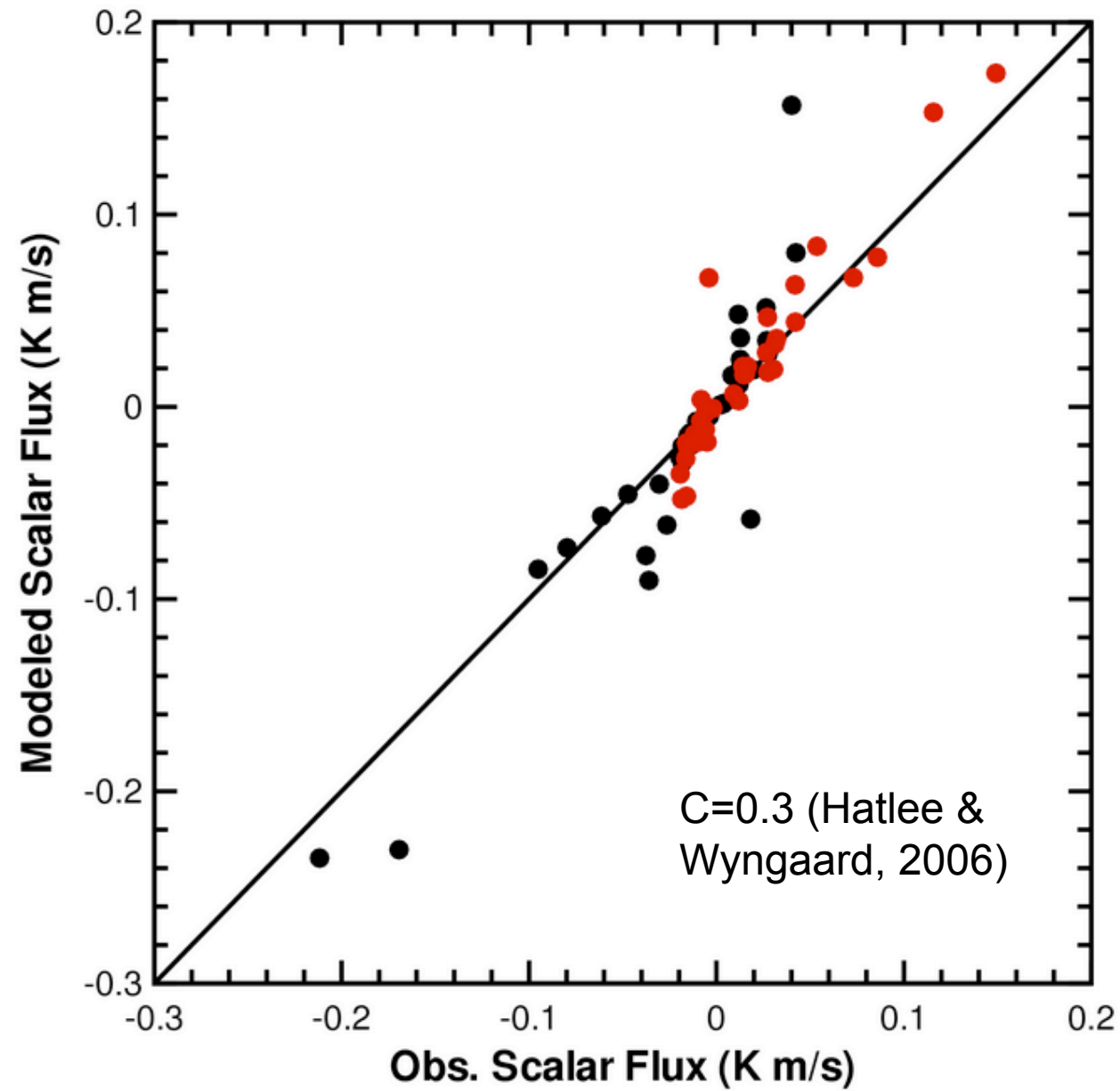


SUBGRID-SCALE SCALAR FLUX

Comments:

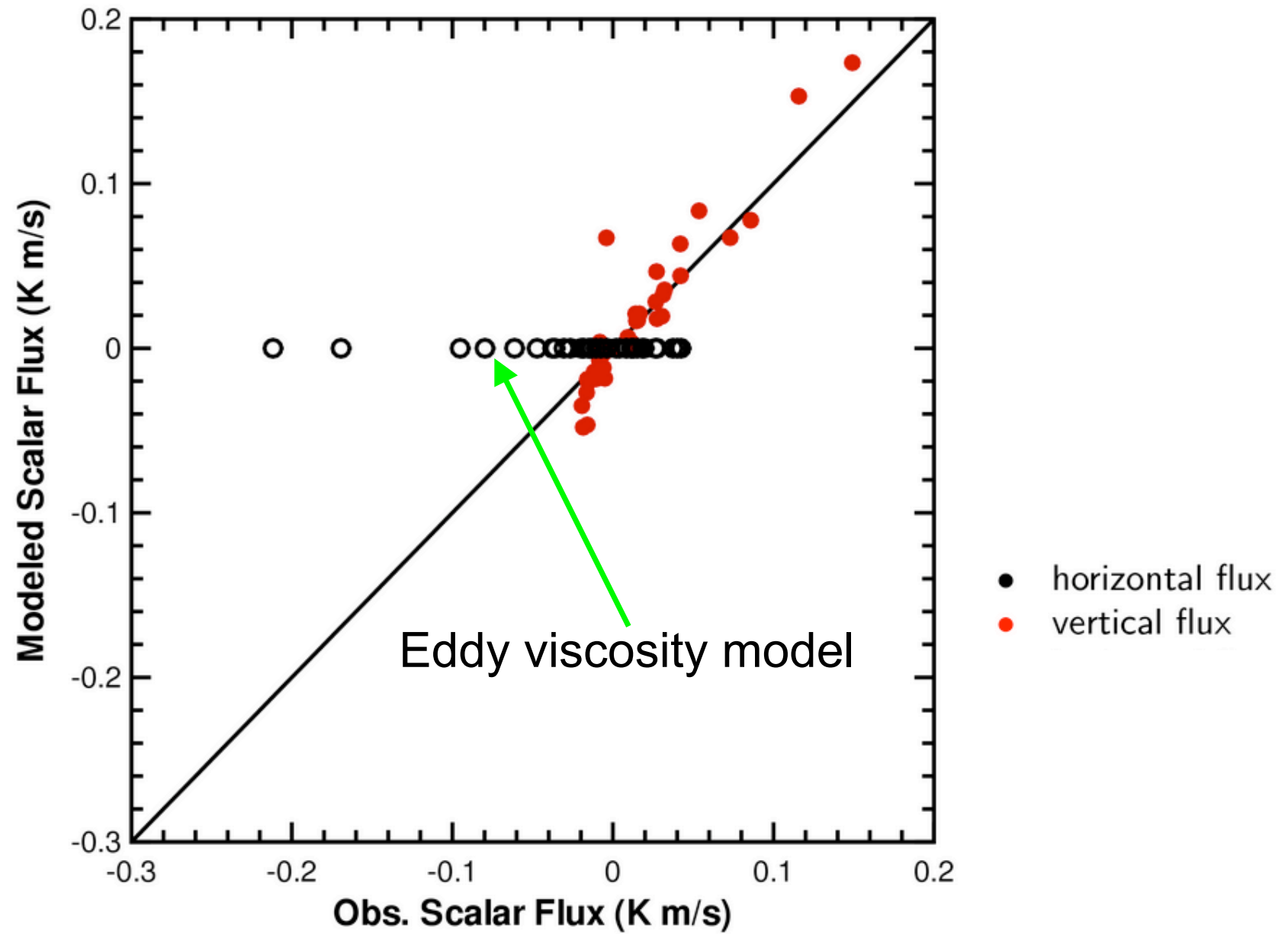
- Net horizontal scalar flux $f_1 = \langle \overline{uc} - \overline{u} \overline{c} \rangle \neq 0$ even for horizontally homogeneous PBLs, *i.e.*, $\frac{\partial}{\partial x} \langle C \rangle = 0$
- Tilting of vertical flux by vertical shear is important
$$f_1 \sim -f_3 \frac{\partial \overline{u}}{\partial z} T$$
- No eddy viscosity model can capture anisotropic production

SFS SCALAR FLUXES IN HATS



- horizontal flux
- vertical flux

SFS SCALAR FLUXES IN HATS



TESTING OTHER TERMS

SUBFILTER-SCALE PRESSURE DESTRUCTION

$$\begin{aligned}
 -\frac{1}{\rho} \left[\overline{p \left(\frac{\partial u_i}{\partial x_j} + \frac{\partial u_j}{\partial x_i} \right)} - \bar{p} \left(\frac{\partial \bar{u}_i}{\partial x_j} + \frac{\partial \bar{u}_j}{\partial x_i} \right) \right] &= -\frac{\tau_{ij} \sqrt{e}}{C_m \Delta_f} \\
 +\frac{1}{\rho} \left(\overline{p \frac{\partial c}{\partial x_i}} - \bar{p} \frac{\partial \bar{c}}{\partial x_i} \right) &= -\frac{f_i \sqrt{e}}{C_s \Delta_f}
 \end{aligned}$$

Momentum

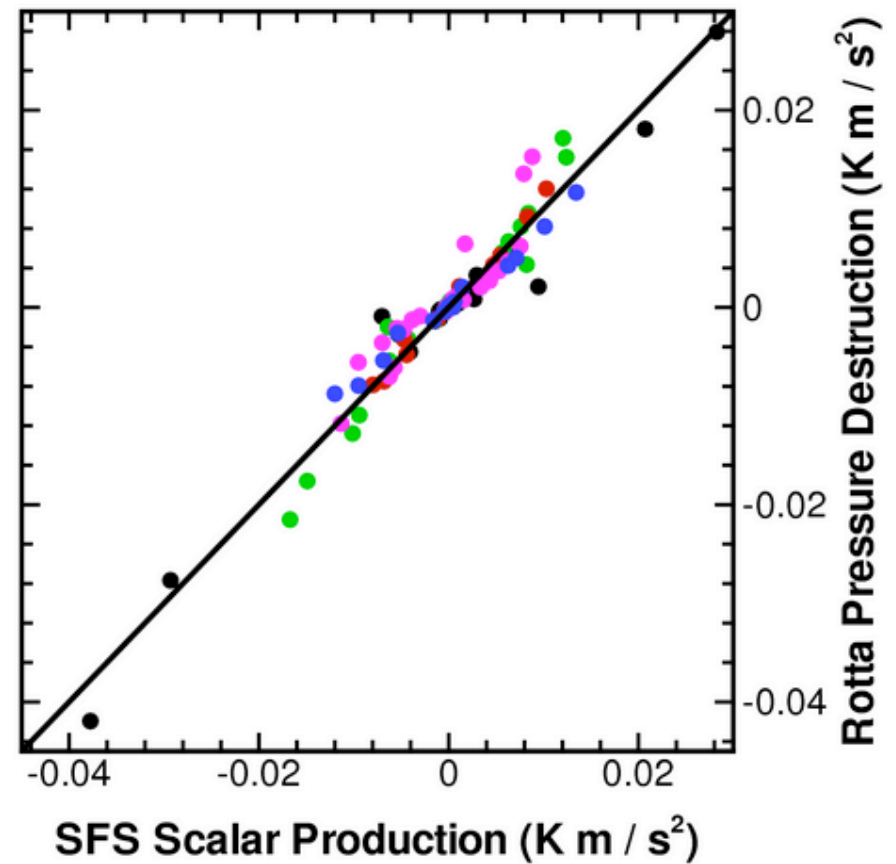
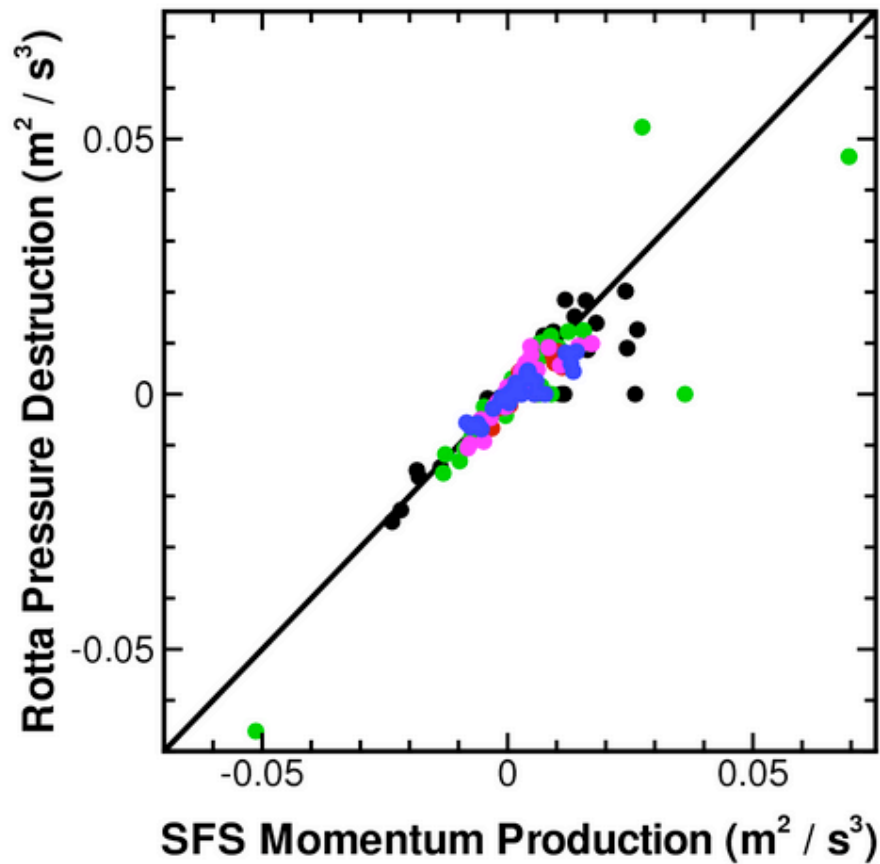
 Scalar




CHATS PRESSURE SENSOR (Steven Oncley)

AHATS (2008) "HORIZONTAL ARRAY" OF PRESSURE SENSORS

VALIDATION OF ROTTA MODEL FOR MOMENTUM AND SCALARS



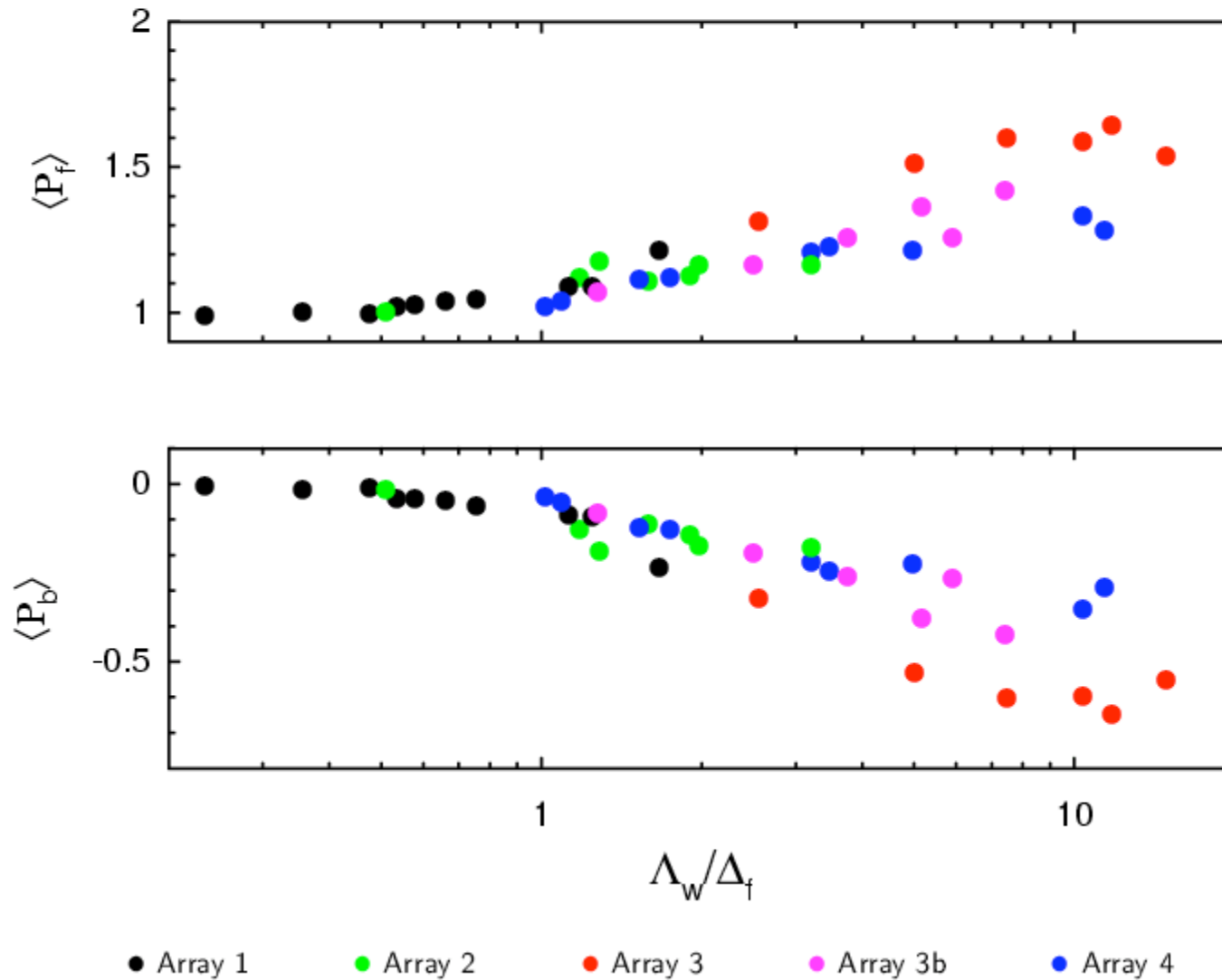
Production $\approx$ Destruction

Decomposition of SFS Production into Forwardscatter and Backscatter

$$P = -\mathcal{T}_{ij}S_{ij} \quad P_f = (P + |P|)/2 \quad P_b = (P - |P|)/2$$

Decomposition of SFS Production into Forwardscatter and Backscatter

$$P = -\mathcal{T}_{ij}S_{ij} \quad P_f = (P + |P|)/2 \quad P_b = (P - |P|)/2$$



SUMMARY

- LES is being applied to a richer set of boundary layer flows because of advances in parallel computing
- Subgrid-scale parameterizations in LES need to be validated/improved for geophysical applications
- Multi-point measurements from the HATS field campaigns compliment our ability to compute
 - Evaluation of subgrid scale models with high Re data
 - Rate equations provide insight into SGS dynamics
 - Importance of anisotropic production for stress and scalar especially for $\Lambda_w/\Delta_f \sim \mathcal{O}(1)$ or less
 - Data highlights the shortcomings of an eddy viscosity approach

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