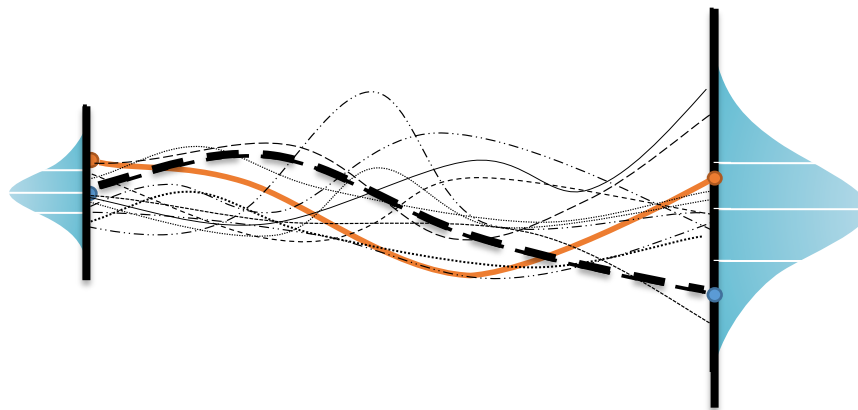
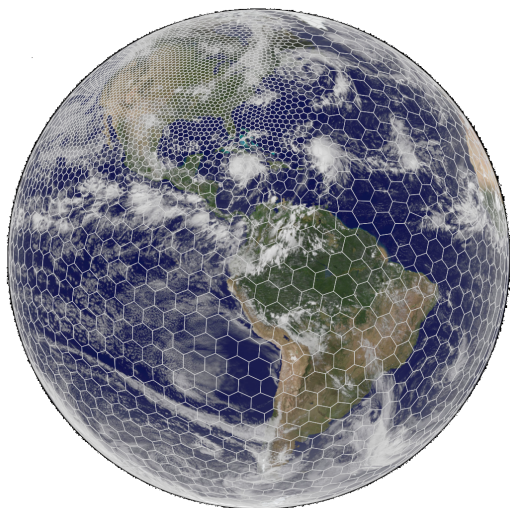


The global Ensemble Kalman filter (EnKF) analysis for the Model for Prediction Across Scales (MPAS) on the variable-resolution meshes

Soyoung Ha, Chris Snyder, Bill Skamarock

*Mesoscale and Microscale Meteorology (MMM) Laboratory
National Center for Atmospheric Research*



Contributions from Jeff Anderson, Nancy Collins (IMAGE),
Joe Klemp, Michael Duda, and Laura Fowler (MMM)

- **Forecast model: Model for Prediction Across Scales (MPAS)-Atmosphere**
 - Unstructured meshes on C-staggering, a terrain-following height coordinate
 - Global, nonhydrostatic simulations w/ horizontally variable resolution grids
 - ‘mesoscale_reference’ and ‘convection-permitting’ physics suites
 - A **cycling** capability
 - &restart
 - config_do_DAcycling = true
 - config_do_restart = true
 - **Incremental Analysis Updates (IAU)** available from V5
 - &IAU
 - config_IAU_option = ‘on’
 - config_IAU_window_length_s = 21600.
 - Current release V5.1: <http://mpas-dev.github.io/>
- **Analysis system: Data Assimilation Research Testbed (DART)**
 - Ensemble Kalman filter (EnKF) analysis
 - Observation operators are built on **native** MPAS meshes
 - => Analysis on variable-resolution meshes
 - Analysis variables in DART \approx prognostic variables in MPAS
 - The latest release named “Manhattan”
<http://www.image.ucar.edu/DARes/DART>

Challenges in MPAS data assimilation

1. Unstructured meshes

=> Develop a new observation operator **H**

- Barycentric interpolation using a triangular mesh
- All locations are considered on the cartesian planes

=> Hard to estimate **P^b** (esp., on variable-resolution meshes);

Estimate **P^bH^T** and **HP^bH^T** from ensemble forecasts on native meshes.

Standard Kalman filter for a linear, unbiased analysis:

$$\mathbf{x}^a = \mathbf{x}^b + \mathbf{K}(\mathbf{y}^o - \mathbf{H}(\mathbf{x}^b))$$

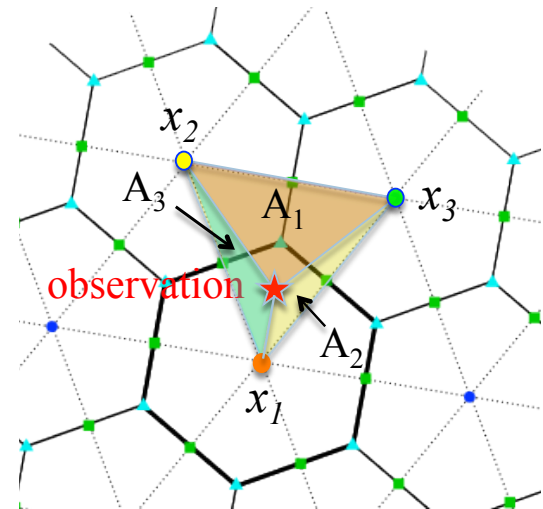
$$\mathbf{K} = \mathbf{P}^b \mathbf{H}^T (\mathbf{H} \mathbf{P}^b \mathbf{H}^T + \mathbf{R})^{-1}$$

\mathbf{x}^a : analysis state; \mathbf{x}^b : background (or prior)

\mathbf{y}^o : observations; $\mathbf{H}(\mathbf{x}^b) = \mathbf{y}^b$: prior obs

\mathbf{P}^b : background error covariance

\mathbf{R} : observation error covariance

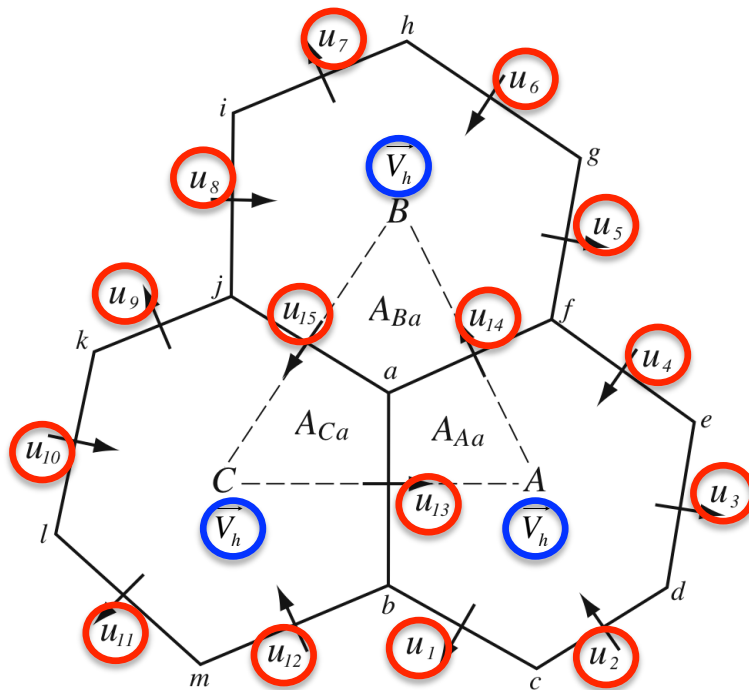


$$y_{obs}^b = (A_1 x_1 + A_2 x_2 + A_3 x_3) / (A_1 + A_2 + A_3)$$

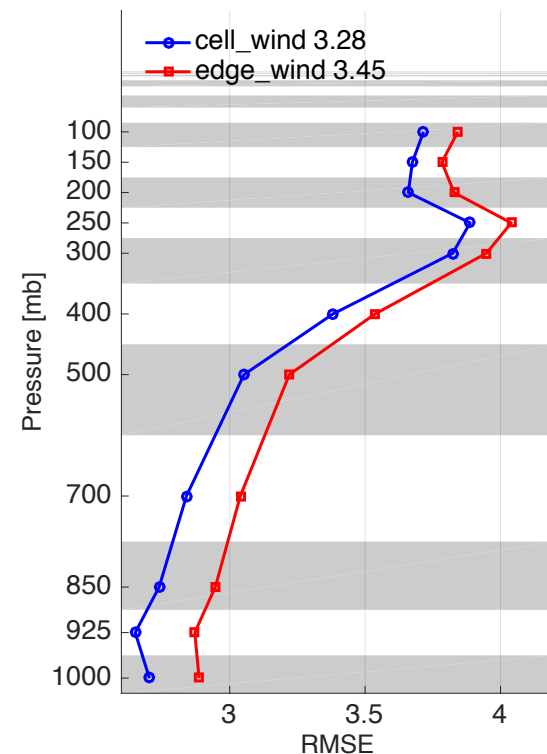
Challenges in MPAS data assimilation

2. Treatment of horizontal wind => **new H for winds**

- A prognostic wind variable is normal velocity (u) at cell edges
- Zonal and meridional winds (\vec{V}_h) are reconstructed at cell centers using Radial Basis Functions (RBFs); used for physics



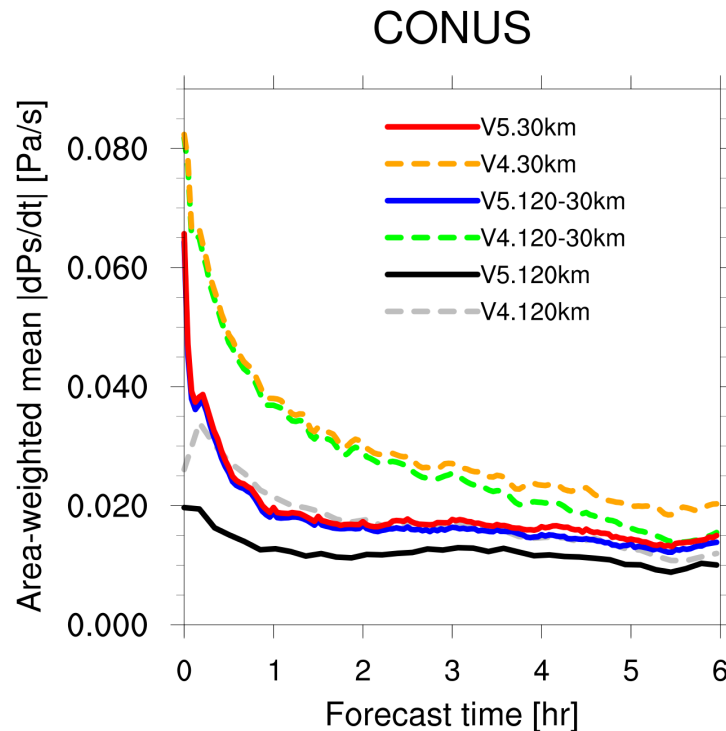
6-h forecast, global RMS error (m/s) against
RADIOSONDE_U_WIND, June 2012



Challenges in MPAS data assimilation

3. Noise Control

- In the global analysis, it is critical to effectively suppress high-frequency oscillations for the numerical stability and the quality of subsequent forecasts.
- MPAS V5.1 introduces a new, scale-aware 3-D divergence damping for acoustic waves.

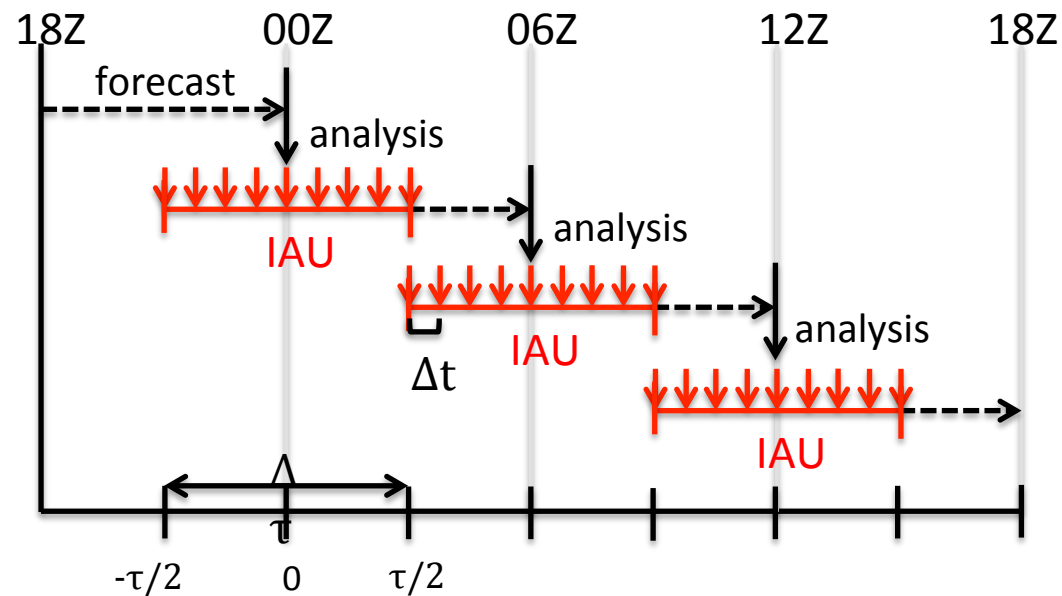


- MPAS forecasts initialized from the NCEP FNL analysis at 2012-06-11_12:00:00 UTC
- Area-weighted global mean surface pressure tendencies at each time step to measure high-frequency noise
- Version 5 reduces the noise from Version 4 in all different meshes

Challenges in MPAS data assimilation

3. Noise Control - IAU

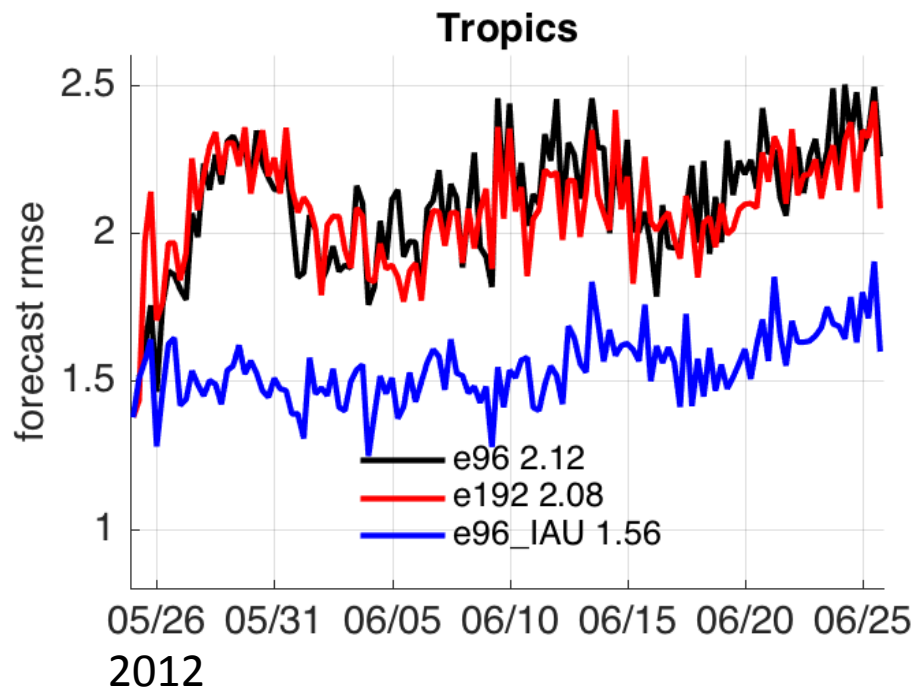
- The analysis quality is often degraded by spurious waves generated from dynamical imbalances arising from analysis increments.
- During cycling, the analysis is usually used as an initial condition for the following forecast. => Sudden localized changes can shock the model.
- The IAU incorporates analysis increments (rather than full fields) into the model tendency in a gradual fashion. (Bloom et al. 1996)



Challenges in MPAS data assimilation

3. Noise Control (cont'd)

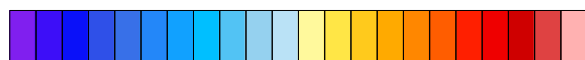
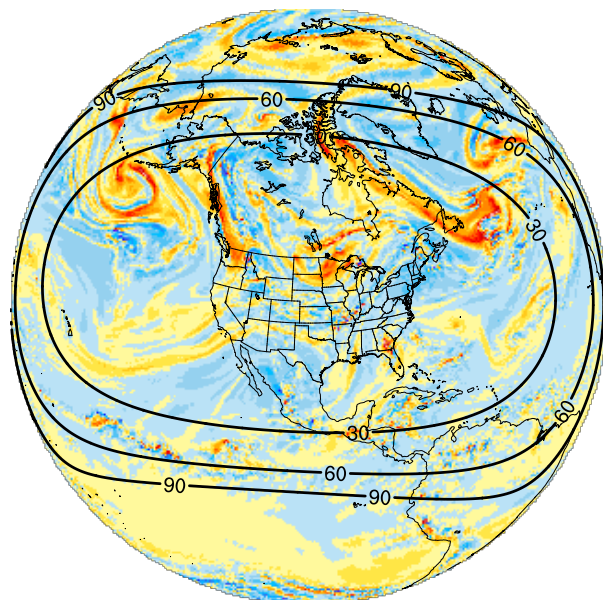
- Cycling w/ a 96-member ensemble (“e96”)
- Cycling w/ a 96-member ensemble and IAU (“e96_IAU”)
- Cycling w/ a 192-member ensemble (“e192”)



- 6-h ensemble mean forecast rms fit to surface altimeter observations during the cycles
- The IAU improved the EnKF analysis, especially over the tropics.
- The use of IAU was more effective than double the ensemble size.

Variable-resolution DA with MPAS/DART

36h forecast valid at 2012-05-29_12:00:00

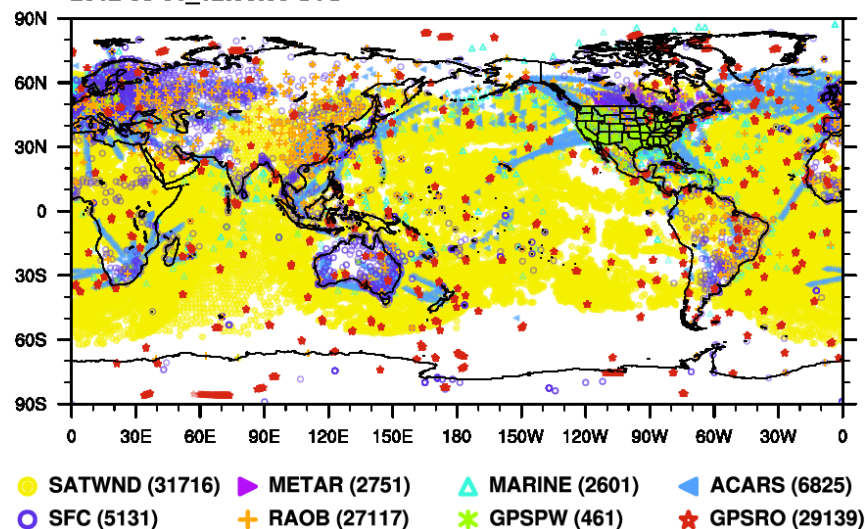


-30 -24 -18 -12 -6 0 6 12 18 24 30

500hPa Vorticity [$\times 10^5$]

Grid resolution in 120-30 km mesh (named “x4”), contouring every 30 km.

2012-06-01_12:00:00 UTC

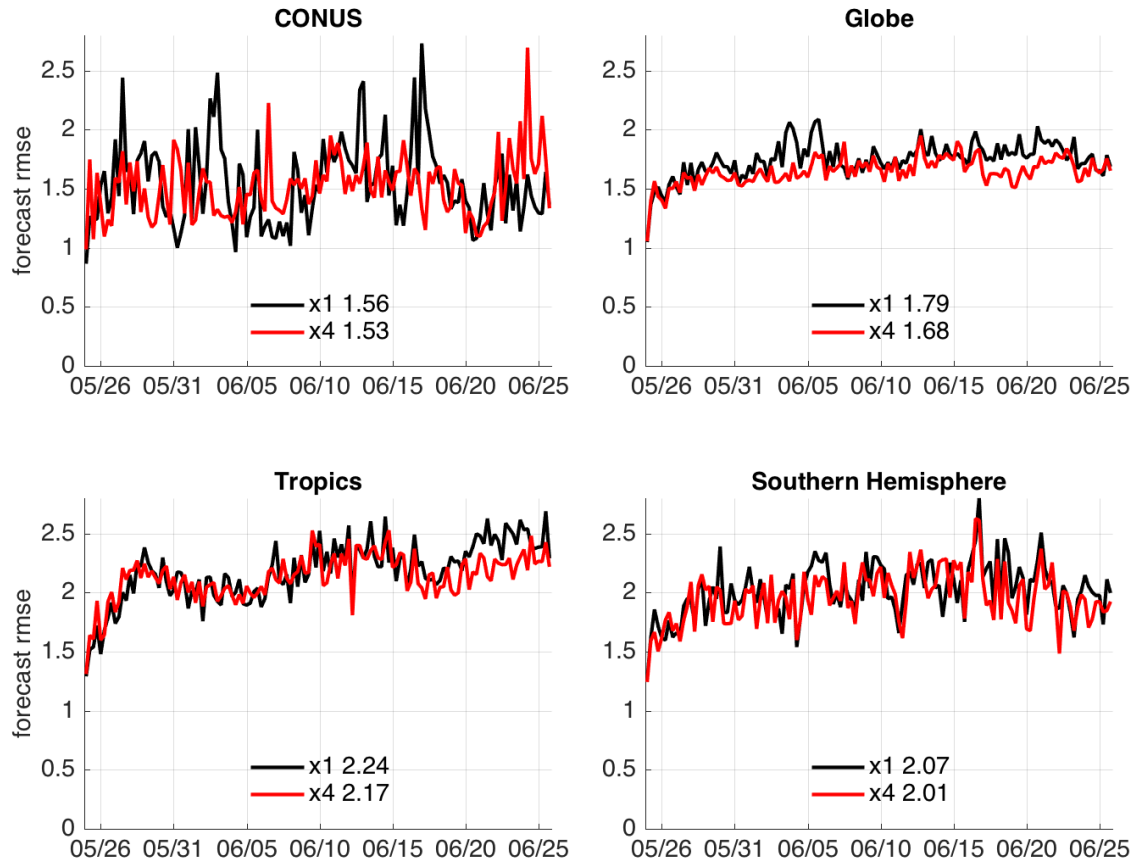


- A 96-member global ensemble
- A variable-resolution mesh (120-30 km) is used in *both* analysis and forecast in the cycling experiment.
- Compare the variable-resolution (x4) to a coarse uniform (x1) mesh w/ 120-km resolution

Ha et al. (Submitted to *Mon. Wea. Rev.*)

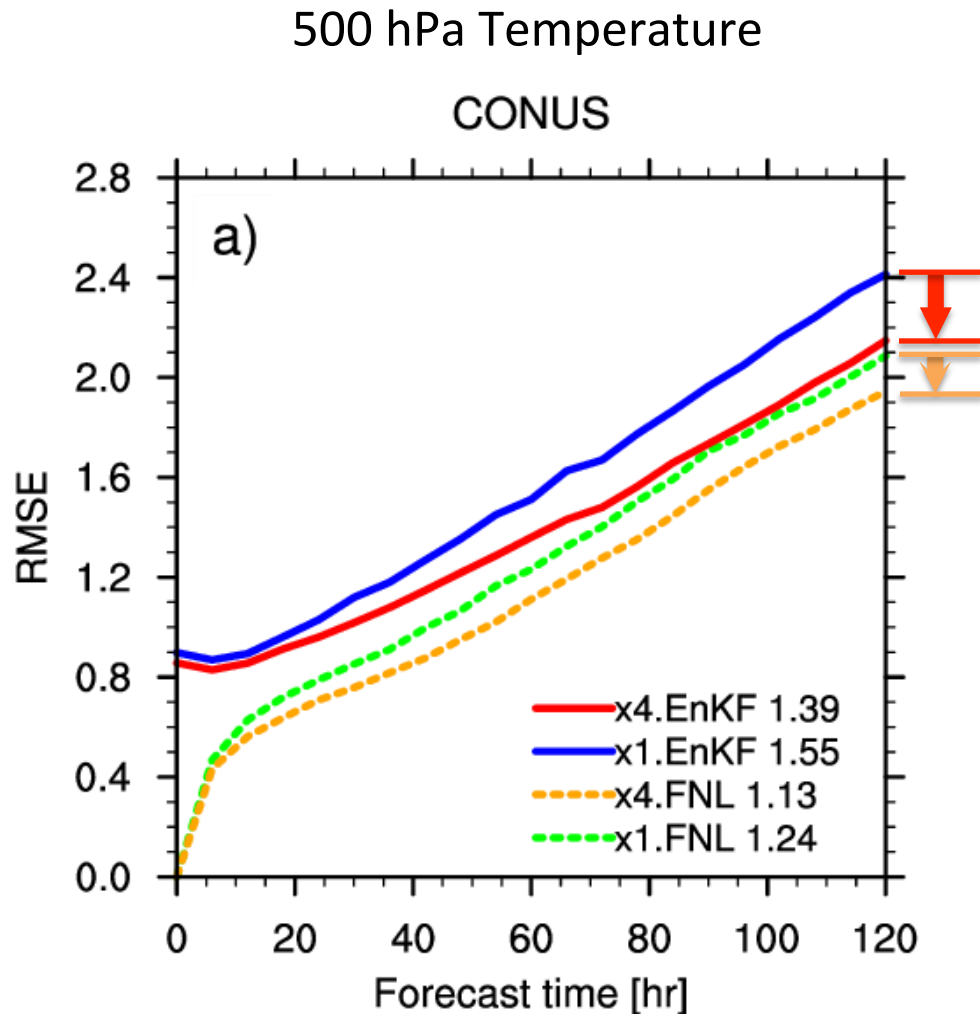
6-h ensemble mean forecast error

LAND_SFC_ALTIMETER @ 1 surface



- Both meshes are reliable throughout the cycling period.
- They are mostly comparable to each other in 6-h prior forecast.

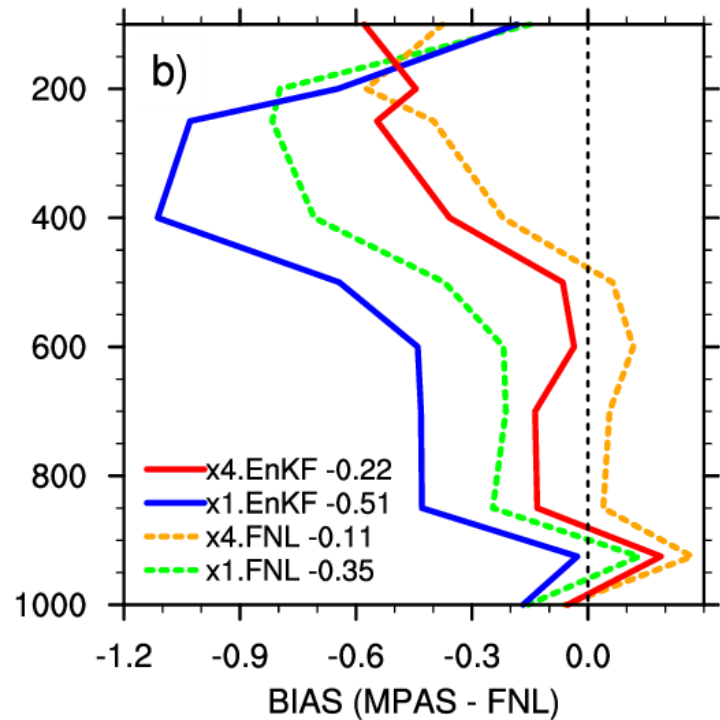
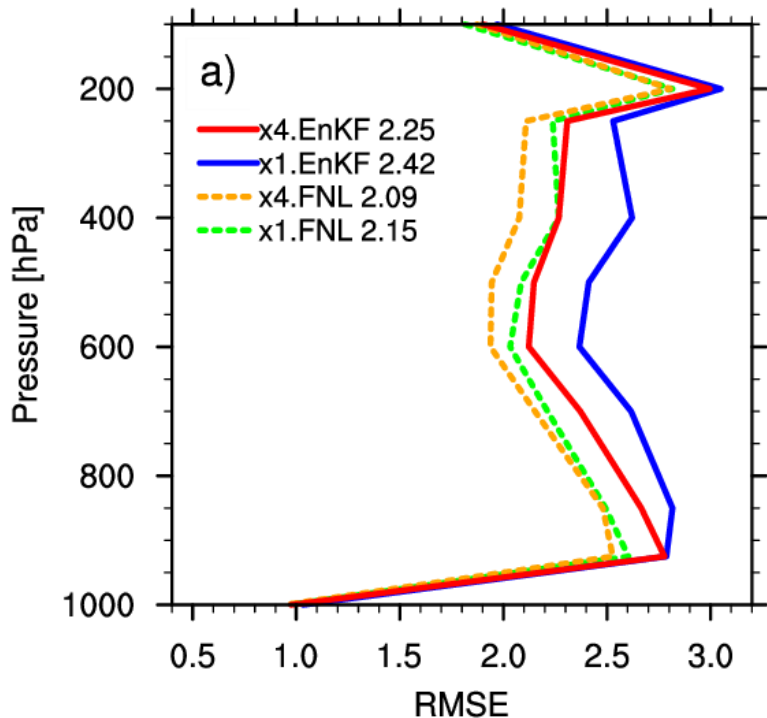
5-day MPAS forecast from the EnKF mean analysis



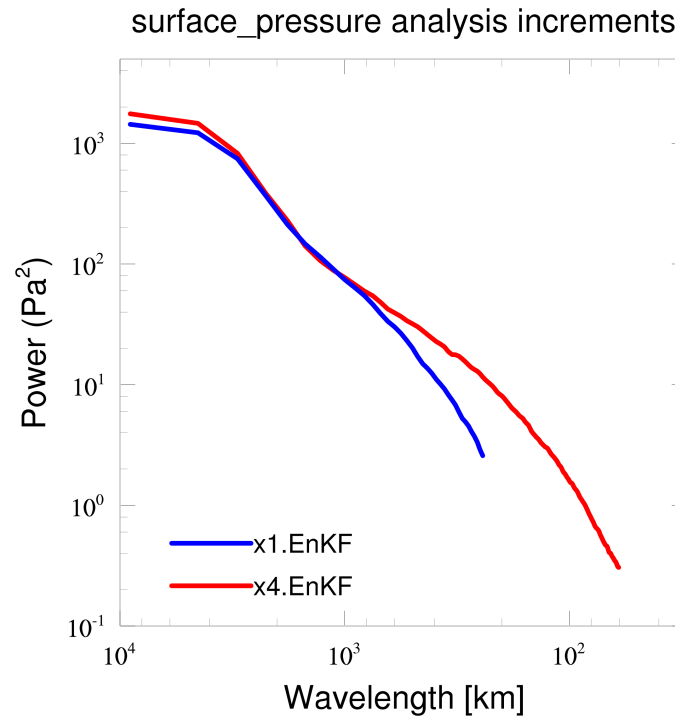
Verification w.r.t. NCEP FNL analysis: The rms error was computed from May 28 to June 25, 2012, twice daily, every other day.

- Benefits of high-resolution model forecasts are well shown in cold-start runs. (orange vs. green)
- The use of variable-mesh in the EnKF analysis improves the benefit nearly as twice as large that in cold-start runs. (red vs. blue)

5-day MPAS forecast from the EnKF mean analysis (cont'd)



Power spectra of ensemble analysis increments (CONUS)

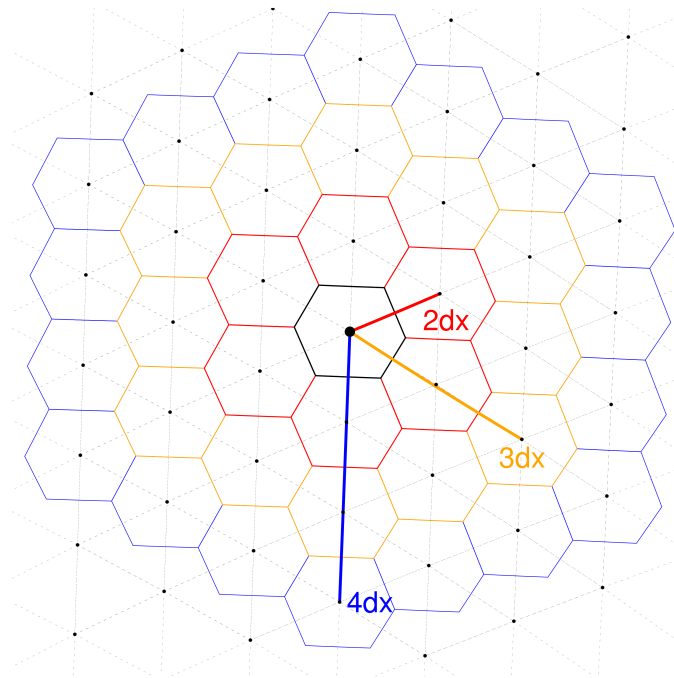


- Both meshes show almost the same power at synoptic scale
- Variable mesh (x4) has much more power at mesoscale range ($< 1,000$ km)

Precipitation forecast

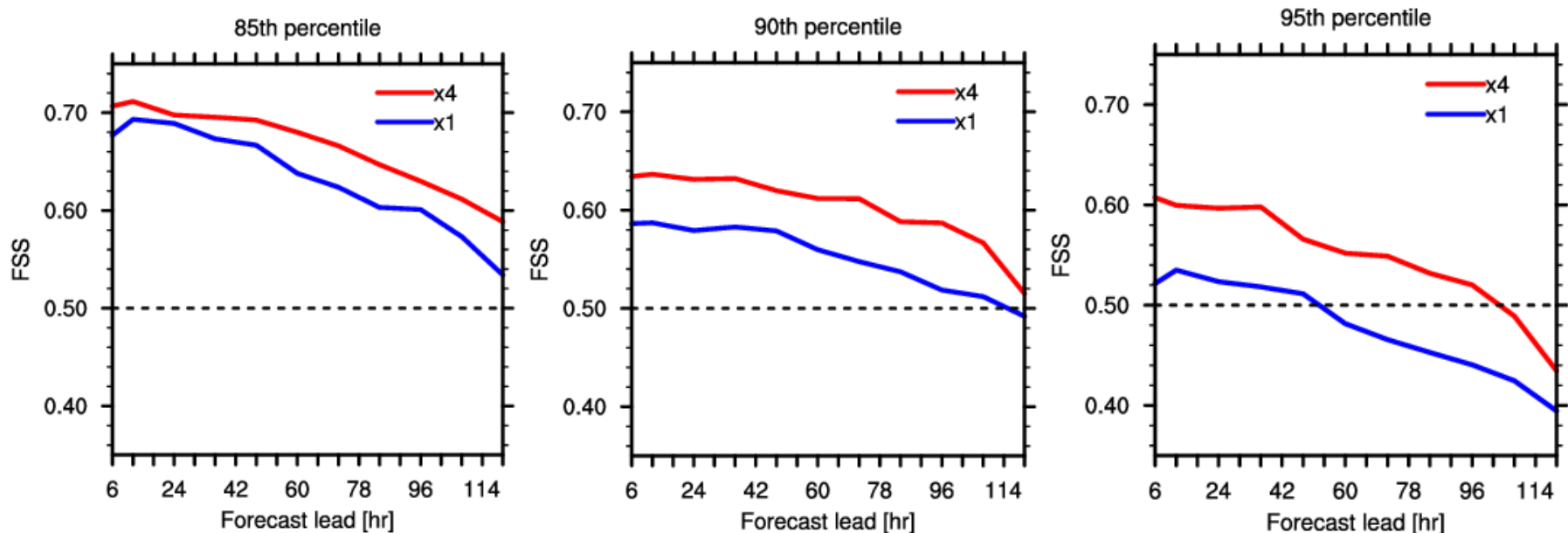
Fractions Skill Scores (FSS) over the MPAS unstructured mesh

- FSS is computed to find the smallest scale over which the model is skillful.
- Both model and observations are projected onto the same verification cells.
- The fraction of occurrences of specified rainfall accumulations is computed over a range of spatial scales (ndx).
- Verification area: CONUS (w/ $dx = 30\text{km}$ in the variable-resolution mesh)



FSS: x1 vs. x4

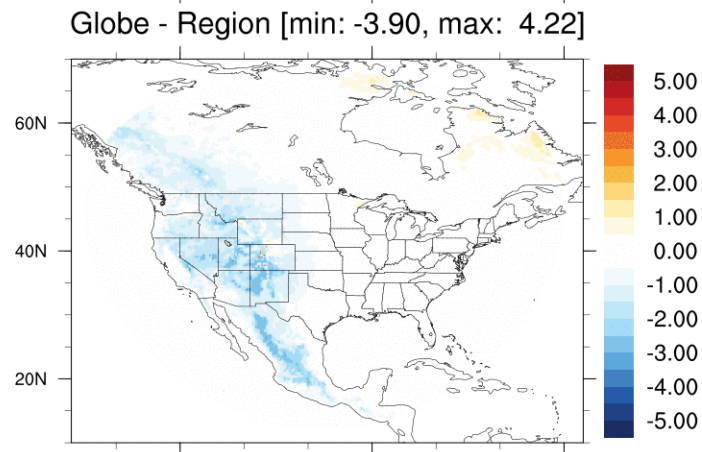
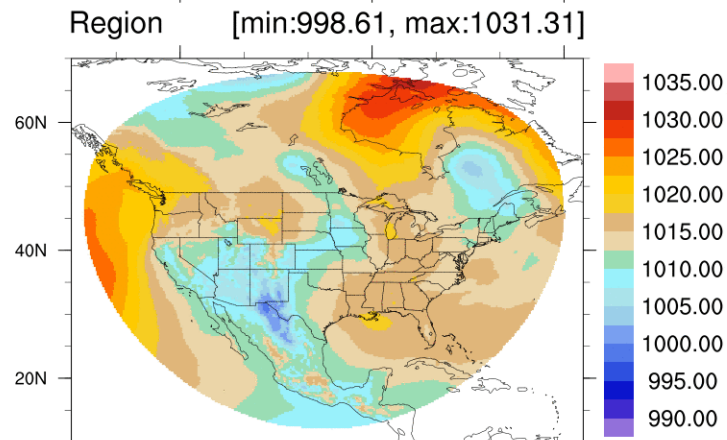
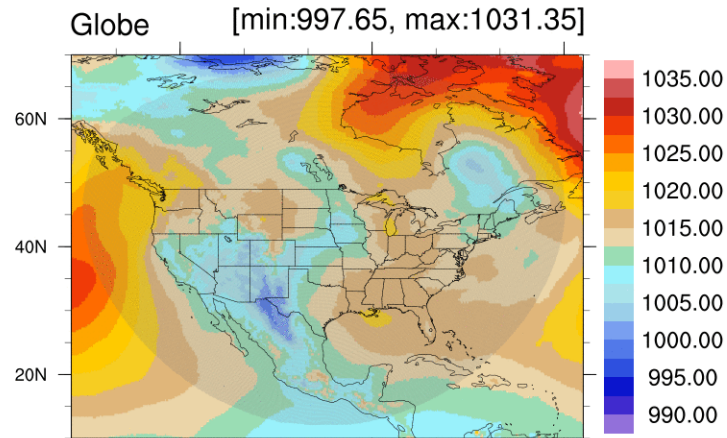
- Observations: NCEP Stage IV data (in 4-km resolution) projected onto the x4 mesh over the CONUS domain
- 5-day forecasts from the coarse uniform (x1) mesh are also projected onto the x4 mesh.
- FSS is compared within the radius of 600 km (e.g. 5dx for “x1” mesh) for three different percentile thresholds.



Summary

- ❑ Both analyses and forecasts are conducted using the *native* MPAS unstructured mesh => applicable to variable-resolution meshes as well as quasi-uniform meshes.
- ❑ Assimilating real observations, the global ensemble analysis/forecast system is reliable throughout a month-long cycling period regardless of the mesh configuration.
- ❑ MPAS forecasts on the variable mesh is benefited from high-resolution simulations in the local refinement area.
- ❑ The EnKF analysis on the variable mesh further improves MPAS forecasts, suppressing forecast error growth.

FCST 000H at 2017-05-09_00 in mslp [hPa]



Future plan

- Conduct cycled MPAS forecasts initialized with EnKF analyses using regional and global MPAS
- Comparison between WRF and regional-MPAS forecasts using the same LBCs from global-MPAS forecasts in the ensemble cycling context
- High-resolution ($O(1\text{km})$) cycled forecasts using variable-resolution meshes in the regional MPAS

Recap on Data Assimilation: Kalman Filter Algorithm

- Analysis step

Assuming that observation and background errors are unbiased, normally distributed, and uncorrelated, the Best Linear Unbiased Estimate (BLUE) is analytically defined as

$$\mathbf{x}^a = \mathbf{x}^b + \mathbf{K}(\mathbf{y}^o - \mathbf{H}(\mathbf{x}^b))$$
$$\mathbf{K} = \mathbf{P}^b \mathbf{H}^T (\mathbf{H} \mathbf{P}^b \mathbf{H}^T + \mathbf{R})^{-1}$$

- Forecast step

$$\mathbf{x}^b = \mathbf{M}(\mathbf{x}^a) + \boldsymbol{\varepsilon}^b$$

$$\mathbf{P}^b = \mathbf{M} \mathbf{P}^a \mathbf{M}^T + \mathbf{Q}$$

where $E(\boldsymbol{\varepsilon}^b) = 0$ and $\mathbf{Q} = E[\boldsymbol{\varepsilon}^b (\boldsymbol{\varepsilon}^b)^T]$

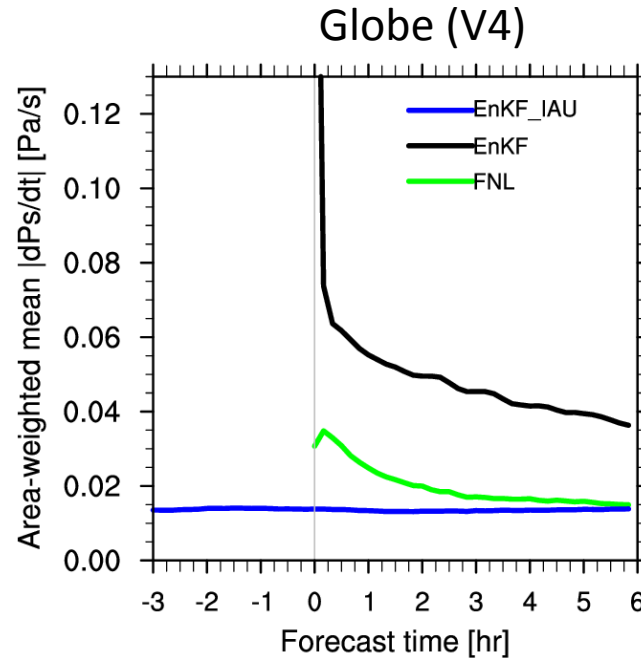
But \mathbf{P}^b is too large and the true state is unknown. In an Ensemble Kalman Filter, ensemble-based approximations are made as below.

$$\mathbf{P}^b \mathbf{H}^T \equiv \frac{1}{N_e - 1} \sum_{e=1}^{N_e} (\mathbf{x}_e^b - \overline{\mathbf{x}}^b) (\mathbf{H} \mathbf{x}_e^b - \overline{\mathbf{H} \mathbf{x}}^b)^T$$

$$\mathbf{H} \mathbf{P}^b \mathbf{H}^T \equiv \frac{1}{N_e - 1} \sum_{e=1}^{N_e} (\mathbf{H} \mathbf{x}_e^b - \overline{\mathbf{H} \mathbf{x}}^b) (\mathbf{H} \mathbf{x}_e^b - \overline{\mathbf{H} \mathbf{x}}^b)^T$$

$$\text{where } \overline{\mathbf{x}}^b = \frac{1}{N_e} \sum_{e=1}^{N_e} \mathbf{x}_e^b \text{ and } \overline{\mathbf{H} \mathbf{x}}^b = \frac{1}{N_e} \sum_{e=1}^{N_e} \mathbf{H} \mathbf{x}_e^b$$

IAU



MPAS forecast from different analyses valid at
2012-06-11_12:00:00 UTC (120-km uniform mesh)