Simulating the morning transition of flow over the Eastern Snake River Plain

Yue Qin, Dan Li Department of Earth and Environment, Boston University

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Effects of topography features

"perfect site" : Flat, uniform and horizontallyhomogeneous





"imperfect site" : Complex terrain, spatiallyinhomogeneous



Free atmosphere Mountain atmosphere Slope atmosphere Valley Atmosphere

Local circulation could complicate the land-atmospheric interaction processes and result in deviation from the MOST-based parameterizations.

Thermally-driven slope wind system

Katabatic case (nighttime)



$$b = \frac{g}{\theta_r} [\theta_2 - \theta_3] < 0$$

Buoyancy forces pushes the air parcel downward.

Anabatic case (daytime)



$$b = \frac{g}{\theta_r} [\theta_2 - \theta_1] > 0$$

Buoyancy forces pushes the air parcel upward.

Why does the diurnal slope wind system matter?

- The distribution of air temperature
- Air quality forecasts
- Local weather forecasting
- Wind farms
- Agriculture
- Low clouds/fogs



Challenges & Objectives

- The nocturnal boundary layer is hard to parameterize.
- One needs to prevent the "run-away" cooling in the model.
- The assumptions under MOST are violated.
- High-resolution land surface data are required.

We aim to explore the performance of numerical weather models in simulating the morning transition of flow over the terrain with a mild slope.

- How well can numerical models capture the local flow and temperature features under the influence of slope winds?
- What are the key factors that affect the simulation of the slope winds?

Study area: Eastern Snake River Plain



Geography features:

- Mid-latitude
- Bordered by high-rise mountains
- Northeast-southwest oriented
- Very mild slope
- Mostly covered by sagebrush

Climatology:

- Thermally-driven drainage flow during the night
- Westerly and channel winds during the day

Local observing network

INL Mesonet system



Eddy covariance at GRI



Clawson, K. L. et al. (2018) NOAA Technical Memorandum OAR ARL

WRF configuration

- D03 horizontal resolution: 1 km by 1 km
- Lowest model level: 10 m



- Baseline simulation: 2020-10-04, 7h spin-up
 - NARR
 - RRTM radiation scheme
 - ACM2 PBL scheme
 - Pleim-Xiu LSM
 - Pleim-Xiu surface layer scheme
- Sensitivity test:
 - YSU, MYJ
 - Noah LSM
 - Revised MM5, Janjic Eta
 - Spin-up, grid spacing, initial soil conditions, FNL bc

Land surface dataset: topography (STRM_30m), land use (NLCD~270m), GVF (ASTER_15m)

Evaluation of the wind field

Before sunrise



Underestimation of the drainage flow (northeast and northwest)

Evaluation of the wind field

After sunrise



Earlier morning transition (from northeast to southwest)

Evaluation of the 2-m temperature



Air temperature increases with height, namely, positive vertical temperature gradient.

The hot bias accumulates throughout the night.

The hot bias is the largest at the lowest level and gradually collapses when it goes up.

Connection between temperature and wind biases







OBS: downslope wind sustained

WRF: downslope wind weakened/ceased 11

Assumptions of the causes of "hot bias"





• To much turbulent mixing from above to the surface.

$$H = \overline{w'\theta'} = -\rho C_P K_h \frac{\partial \bar{\theta}}{\partial z}$$

- Under-resolving of local topography
 - Cold pool effect
 - Gravity waves

Summary

- WRF underestimates the drainage flow and significantly overestimate the nearsurface temperature at night over the study area.
- There exists a feedback loop between the underestimation of cold air drainage effect and the hot bias at night, which in turns leads to an earlier reversal of wind direction in the morning transition.
- The parameterization of the eddy diffusivity in stably-stratified boundary layer generates over-mixing between the surface and atmosphere above. (needing further examination)

Any suggestions and remarks are very welcome!



Radiation & Fluxes



2-m temperature



Richardson Number



Sensible heat



Wind speed and wind direction







Decrease eddy diffusivity



