The WRF NMM Core

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Talk modified and presented by
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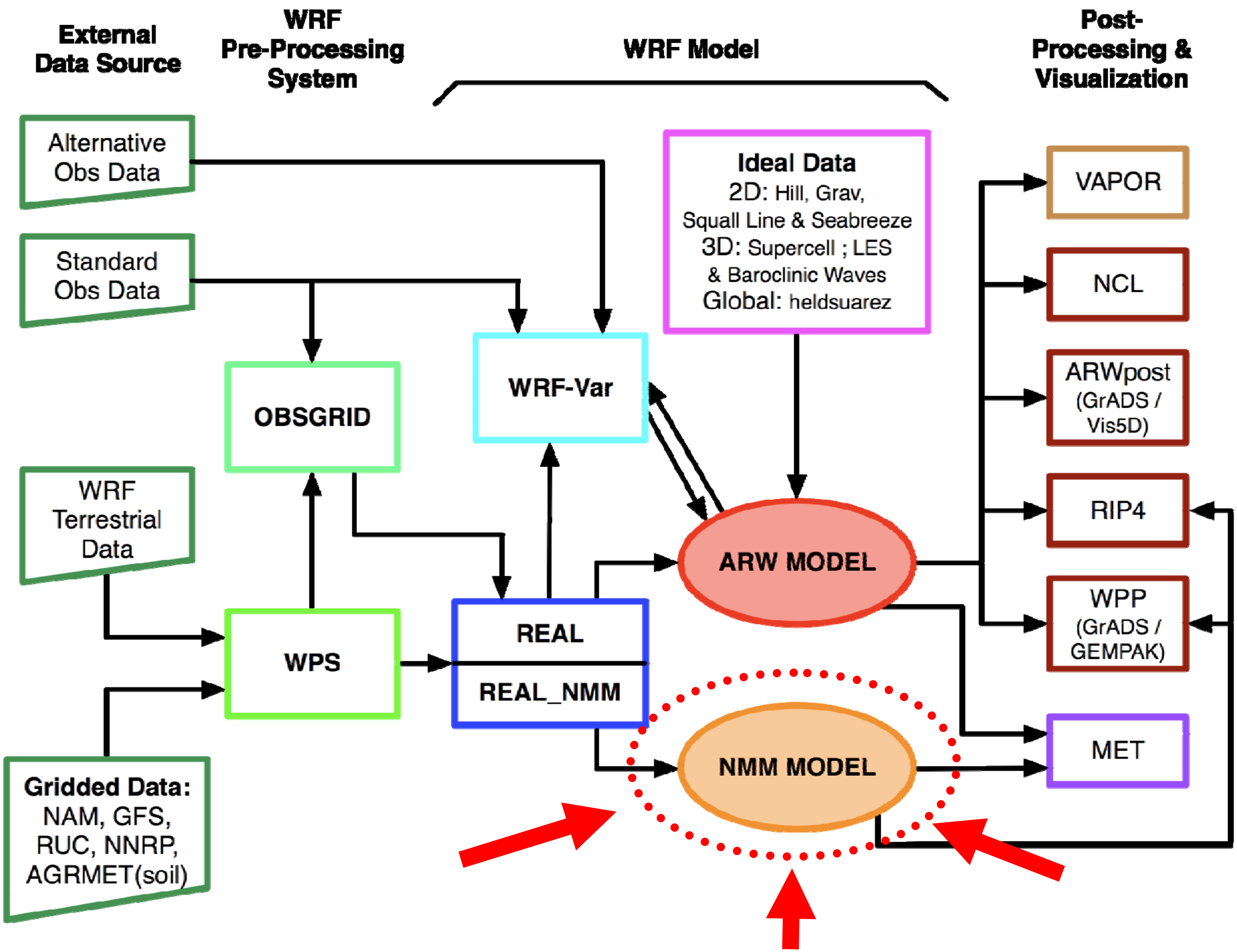
WRF Modeling System Flow Chart



NMM Dynamic Solver

- Basic Principles
- Equations / Variables
- Model Integration
- Horizontal Grid
- Spatial Discretization
- Vertical Grid
- Boundary Conditions
- Dissipative Processes
- Summary

Basic Principles

- Use full compressible equations split into hydrostatic and nonhydrostatic contributions
  - Easy comparison of hydro and nonhydro solutions
  - Reduced computational effort at lower resolutions
- Use modeling principles proven in NWP and regional climate applications
- Use methods that minimize the generation of small-scale noise
- Robust, computationally efficient

Mass Based Vertical Coordinate

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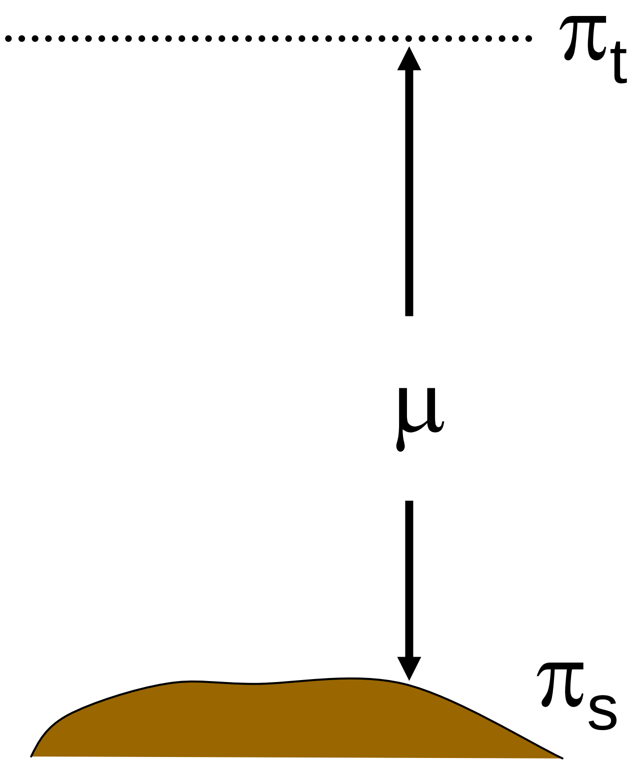
To simplify discussion of the model equations, consider the sigma coordinate to represent a vertical coordinate based on hydrostatic pressure (π):

$$\mu = \pi_s - \pi_t$$

$$\sigma = \frac{\pi - \pi_t}{\mu}$$



WRF-NMM dynamical equations in inviscid, adiabatic,* sigma form (Janjic et al., 2001, MWR)

Analogous to a hydrostatic system, except for *p* and ε, where p is the total (nonhydrostatic) pressure and ε is defined below.

Momentum eqn.

$$\frac{\partial \mathbf{v}}{\partial t} = -\mathbf{v}\cdot\nabla_\sigma \mathbf{v} - \dot{\sigma}\frac{\partial \mathbf{v}}{\partial \sigma} - (1+\varepsilon)\nabla_\sigma \Phi - \alpha\nabla_\sigma p + f\mathbf{k}\times\mathbf{v}$$

Thermodynamic eqn.

$$\frac{\partial T}{\partial t} = -\mathbf{v}\cdot\nabla_\sigma T - \dot{\sigma}\frac{\partial T}{\partial \sigma} + \frac{\alpha}{c_p}\left[\frac{\partial p}{\partial t} + \mathbf{v}\cdot\nabla_\sigma p + \dot{\sigma}\frac{\partial p}{\partial \sigma}\right]$$

Continuity eqn.

$$\frac{\partial \mu}{\partial t} + \nabla_\sigma\cdot(\mu\mathbf{v}) + \frac{\partial(\mu\dot{\sigma})}{\partial \sigma} = 0$$

$$\varepsilon \equiv \frac{1}{g}\frac{dw}{dt}$$

$$\alpha = RT/p$$

*

Hypsometric eqn.

$$\frac{\partial \Phi}{\partial \sigma} = -\mu \frac{RT}{p}$$

Nonhydro var. definition (restated)

$$\varepsilon \equiv \frac{1}{g}\frac{dw}{dt}$$

3rd eqn of motion

$$\frac{\partial p}{\partial \pi} = 1 + \varepsilon$$

ε generally is quite small. Even a large vertical acceleration of 20 m/s in 1000 s produces ε of only ~0.002, and nonhydrostatic pressure deviations of ~200 Pa.

Nonhydrostatic continuity eqn.

$$w = \frac{1}{g}\frac{d\Phi}{dt} = \frac{1}{g}\left(\frac{\partial \Phi}{\partial t} + \mathbf{v} \bullet \nabla_{\sigma}\Phi + \dot{\sigma}\frac{\partial \Phi}{\partial \sigma}\right)$$

*

Properties of system

- Φ, w, ε are not independent, no independent prognostic equation for w!

- ε <<1 in meso and large scale atmospheric flows.

- Impact of nonhydrostatic dynamics becomes detectable at resolutions <10km, important at 1km.

Vertical boundary conditions for model equations*

Top: $\dot{\sigma} = 0$, $p - \pi = 0$

Surface: $\dot{\sigma} = 0$, $\dfrac{\partial (p - \pi)}{\partial \sigma} = 0$

WRF-NMM predictive variables

- Mass variables:
  - **PD** – hydrostatic pressure depth (time and space varying component) (Pa)
  - **PINT** – nonhydrostatic pressure (Pa)
  - **T** – sensible temperature (K)
  - **Q** – specific humidity (kg/kg)
  - **CWM** – total cloud water condensate (kg/kg)
  - **Q2** – 2 * turbulent kinetic energy (m^2/s^2)
- Wind variables:
  - **U,V** – wind components (m/s)

Model Integration

General Philosophy

- Explicit time differencing preferred where possible, as allows for better phase speeds and more transparent coding:
  - horizontal advection of u, v, T
  - passive substance advection of q, cloud water, TKE
- Implicit time differencing for very fast processes that would require a restrictively short time step for numerical stability:
  - vertical advection of u, v, T and vertically propagating sound waves

*

Model Integration

Horizontal advection of u, v, T

2nd order Adams-Bashforth:

$$\frac{y^{\tau+1} - y^{\tau}}{\Delta t} = \frac{3}{2} f(y^{\tau}) - \frac{1}{2} f(y^{\tau-1})$$

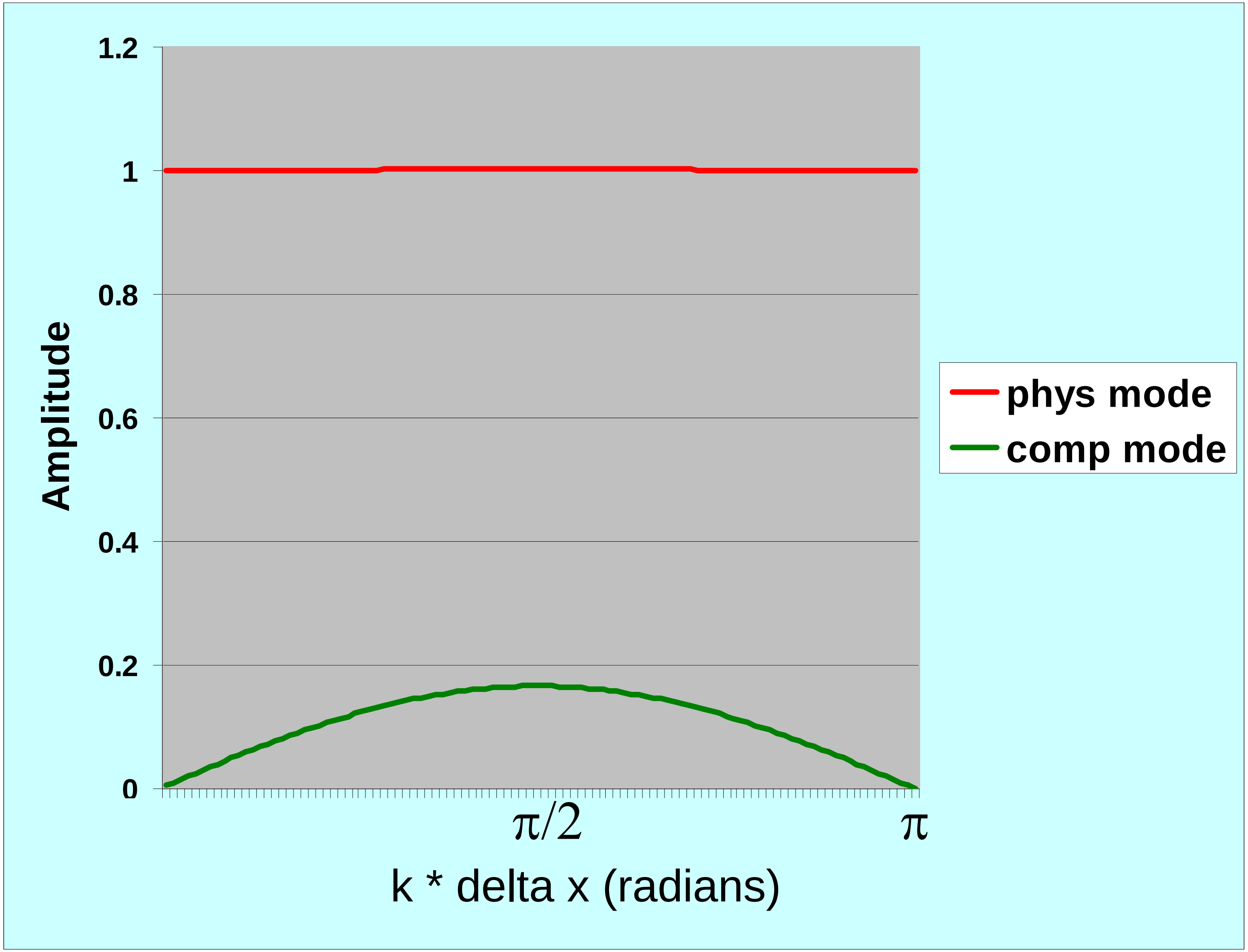
Stability/Amplification:

A-B has a weak linear instability (amplification) which can be tolerated in practice or stabilized by a slight off-centering as is done in the WRF-NMM.

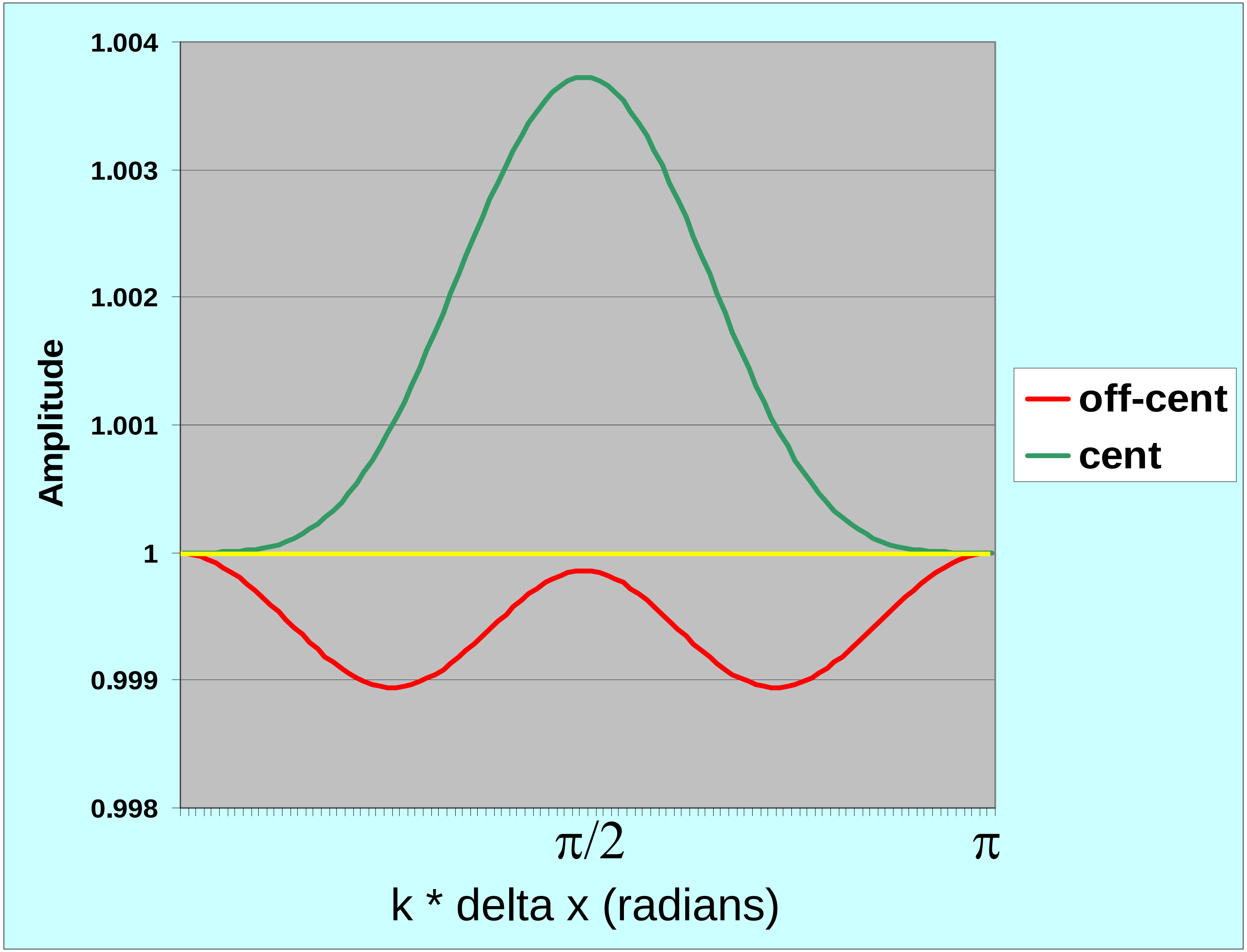
$$\frac{y^{\tau+1} - y^{\tau}}{\Delta t} = 1.533 f(y^{\tau}) - 0.533 f(y^{\tau-1})$$

Adams-Bashforth amplification factor derived from *

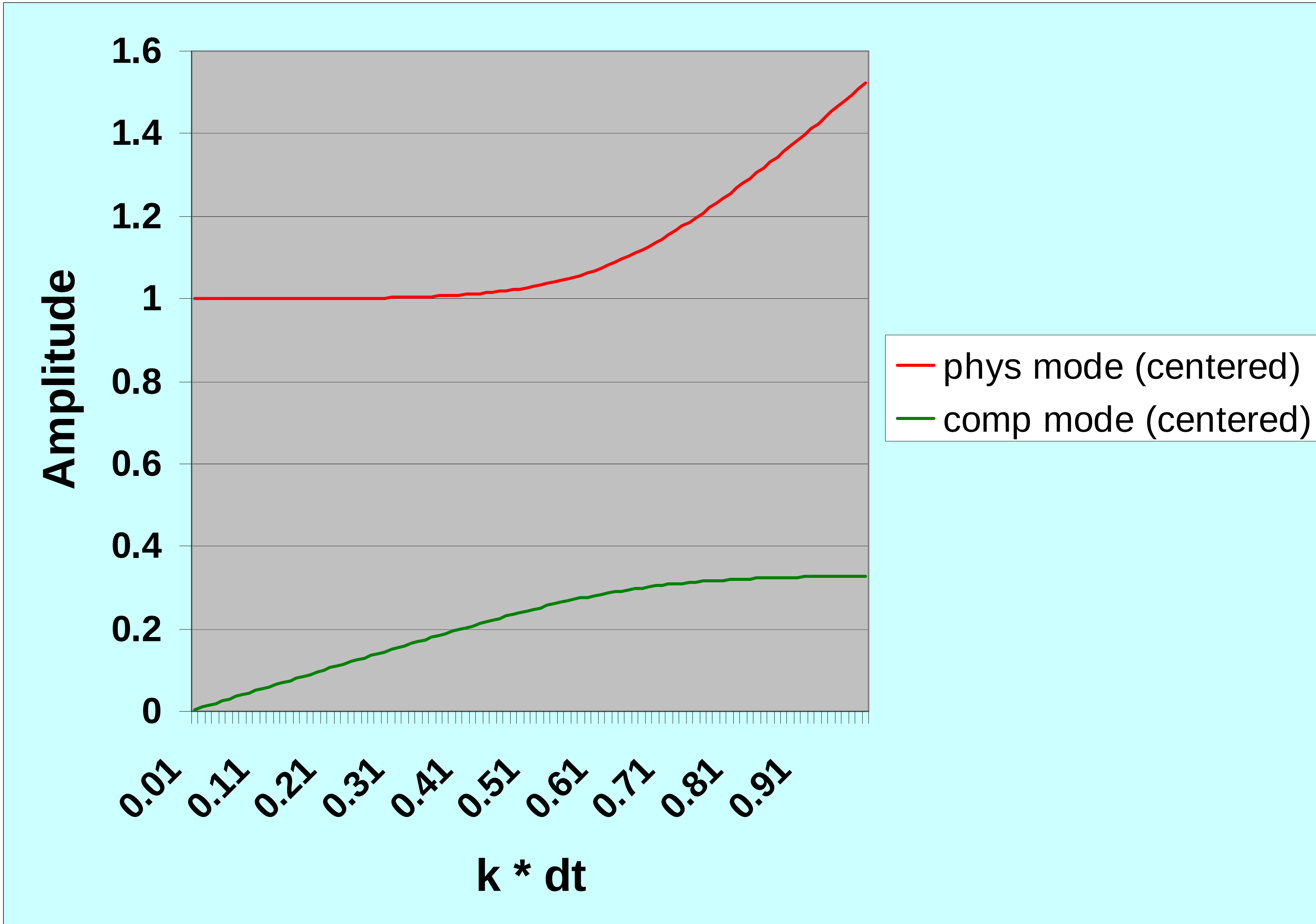
$$\frac{\Delta u}{\Delta t} + c\frac{\Delta u}{\Delta x} = 0$$ with cΔt/Δx = 0.33



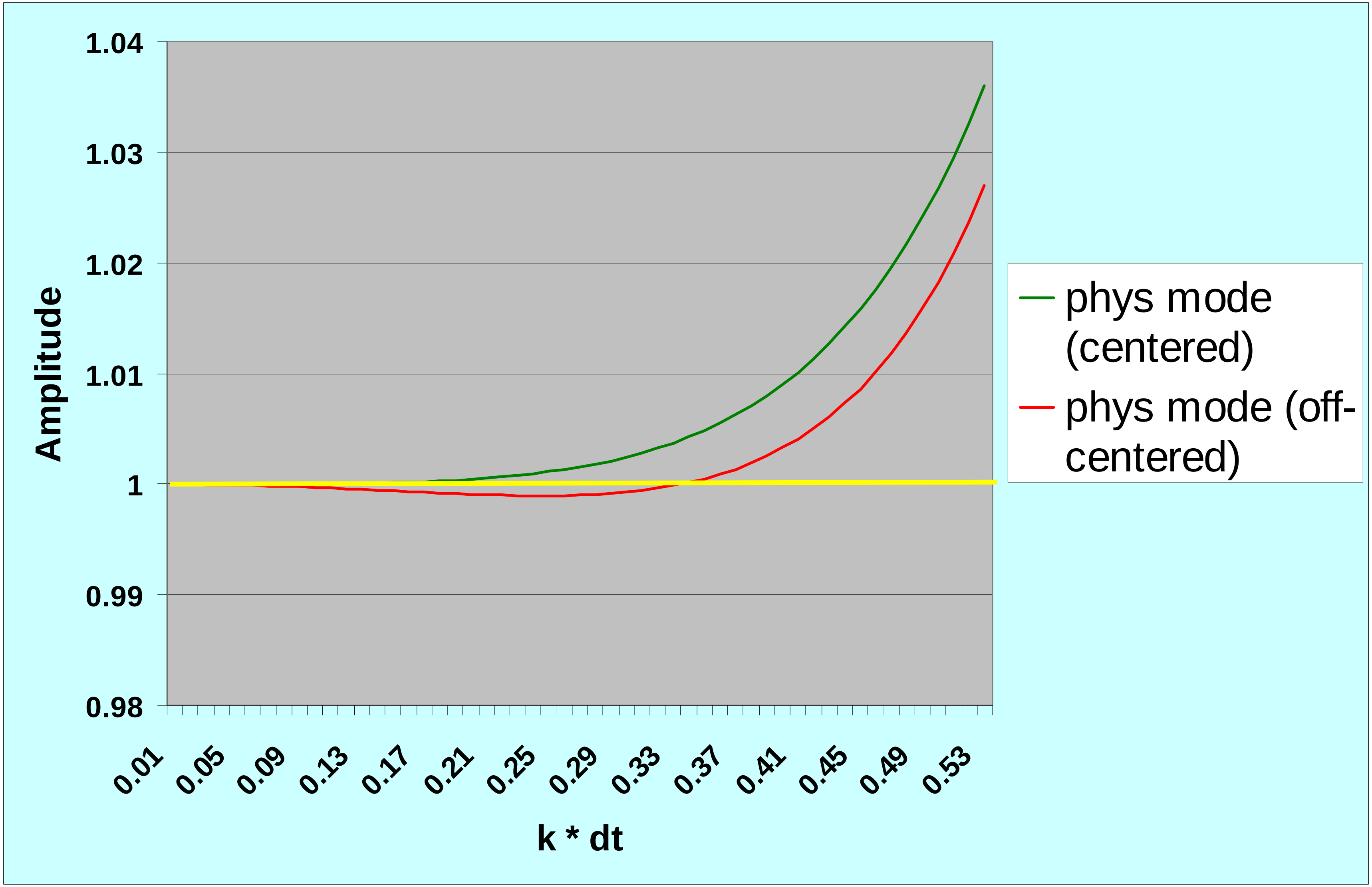
Adams-Bashforth amplification factor, Impact of off-centering



Adams-Bashforth amplification factor derived from oscillation equation ($d\psi/dt=ik\psi$)



Adams-Bashforth amplification factor, Impact of off-centering



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Model Integration

Vertical advection of u, v, & T

Crank-Nicolson (w/ off centering in time):

$$\frac{y^{\tau+1} - y^{\tau}}{\Delta t} = \frac{1}{2}[1.1 f(y^{\tau+1}) + 0.9 f(y^{\tau})]$$

Stability:

An implicit method, it is absolutely stable numerically. Short time steps still needed for *accuracy.*

Model Integration

Advection of TKE (Q2) and moisture (Q, CWM)

- Traditionally has taken an approach similar to the Janjic (1997) scheme used in Eta model:
  - Starts with an initial upstream advection step
  - anti-diffusion/anti-filtering step to reduce dispersiveness
  - conservation enforced after each anti-filtering step – maintain global sum of advected quantity, and prevent generation of new extrema.

Model Integration

Advection of TKE (Q2) and moisture (Q, CWM)

- Included in WRFV3.2 is a new "Eulerian" advection option for the NMM:
  - Improved conservation of advected species, and the method is more consistent with remainder of the NMM dynamics.
  - The old advection (previous slide) is used by default, but this new option can be invoked in the model namelist by adding:

```
&dynamics
 euler_adv        = .true.,
 idtadt           = 2, (time step intvl for advecting met tracers)
```

  - Forecast impact not always positive, but better on average. Largest positive benefit is in reducing warm season high bias in precipitation at finer scales.

Model Integration

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Fast adjustment processes - gravity wave propagation

Forward-Backward (Ames, 1968; Janjic and Wiin-Nielsen, 1977; Janjic 1979): Mass field computed from a forward time difference, while the velocity field comes from a backward time difference.

In a shallow water equation sense:

$$\frac{\partial u}{\partial t} = -g\frac{\partial h}{\partial x}, \frac{\partial h}{\partial t} = -H\frac{\partial u}{\partial x}$$

$$h^{\tau+1} = h^{\tau} - \Delta t\, H\frac{\partial u^{\tau}}{\partial x}$$

$$u^{\tau+1} = u^{\tau} - \Delta t\, g\frac{\partial h^{\tau+1}}{\partial x}$$

Mass field forcing to update wind from τ+1 time

*

Model Integration

Vertically propagating sound waves

In case of linearized equations, equivalent to implicit solution of:

$$\frac{\partial^2 p'}{\partial t^2} \rightarrow \frac{p'^{\tau+1} - 2p'^{\tau} + p'^{\tau-1}}{\Delta t^2} = \frac{c_p}{c_v} R T_0 \frac{\partial^2 p'^{\tau+1}}{\partial {z_0}^2}$$

Where p' is the perturbation pressure from a hydrostatic basic state, and τ represents the time level. (Janjic et al., 2001; Janjic, 2003). *Embedded into full equations, not actually used in the model in this form.*

Model Integration

- Sequence of events within a solve_nmm loop (ignoring physics):

(0.6%) ▪ PDTE – integrates mass flux divergence, computes vertical velocity and updates hydrostatic pressure.

(26.4%) ▪ ADVE – horizontal and vertical advection of T, u, v, Coriolis and curvature terms applied.

(1.2%) ▪ VTOA – updates nonhydrostatic pressure, applies $\omega\alpha$ term to thermodynamic equation

(8.6%) ▪ VADZ/HADZ – vertical/horizontal advection of height. w=dz/dt updated.

(10.6%) ▪ EPS – vertical and horizontal advection of dz/dt, vertical sound wave treatment.

(relative % of dynamics time spent in these subroutines)

Model Integration

- Sequence of events within a solve_nmm loop (cont):

(19.5%) ▪ VAD2/HAD2 (every other step) – vertical/horizontal advection of q, CWM, TKE

(11.8%) ▪ HDIFF – horizontal diffusion

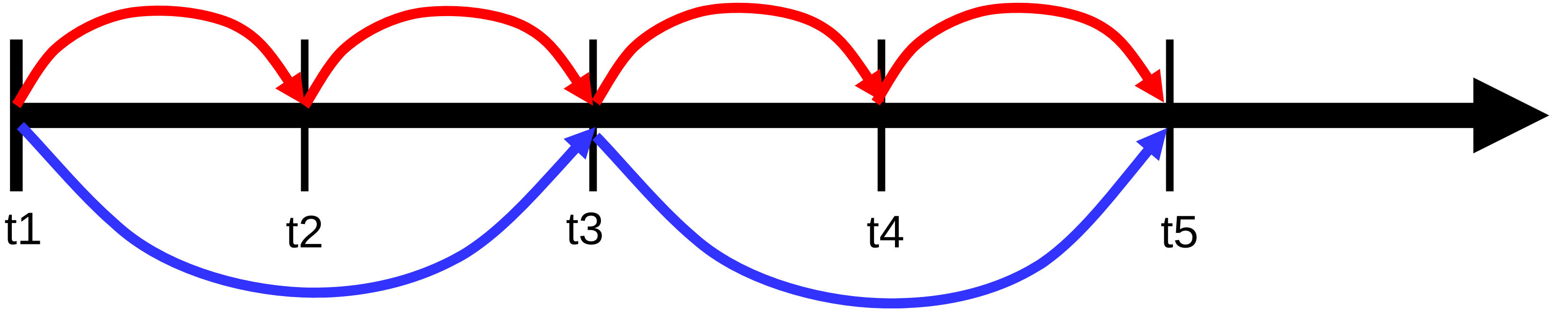
(1.2%) ▪ BOCOH – boundary update at mass points

(17.5%) ▪ PFDHT – calculates PGF, updates winds due to PGF, computes divergence.

(2.3%) ▪ DDAMP – divergence damping

(0.3%) ▪ BOCOV – boundary update at wind points

All dynamical processes every fundamental time step, except….



…passive substance advection, every other time step

Model time step “dt” specified in model namelist.input is for the fundamental time step.

Generally about 2.25x** the horizontal grid spacing (km), or 350x the namelist.input “dy” value (degrees lat).

** runs w/o parameterized convection may benefit from limiting the time step to about 1.9-2.0x the grid spacing.

Now we'll take a look at two items specific to the WRF-NMM horizontal grid:

- Rotated latitude-longitude map projection (only projection used with the WRF-NMM)

- The Arakawa E-grid stagger

Rotated Latitude-Longitude

- Rotates the earth’s latitude/longitude grid such that the intersection of the equator and prime meridian is at the center of the model domain.
- The rotation minimizes the convergence of meridians over the domain, and maintains a more uniform earth-relative grid spacing than exists for a regular lat-lon grid.

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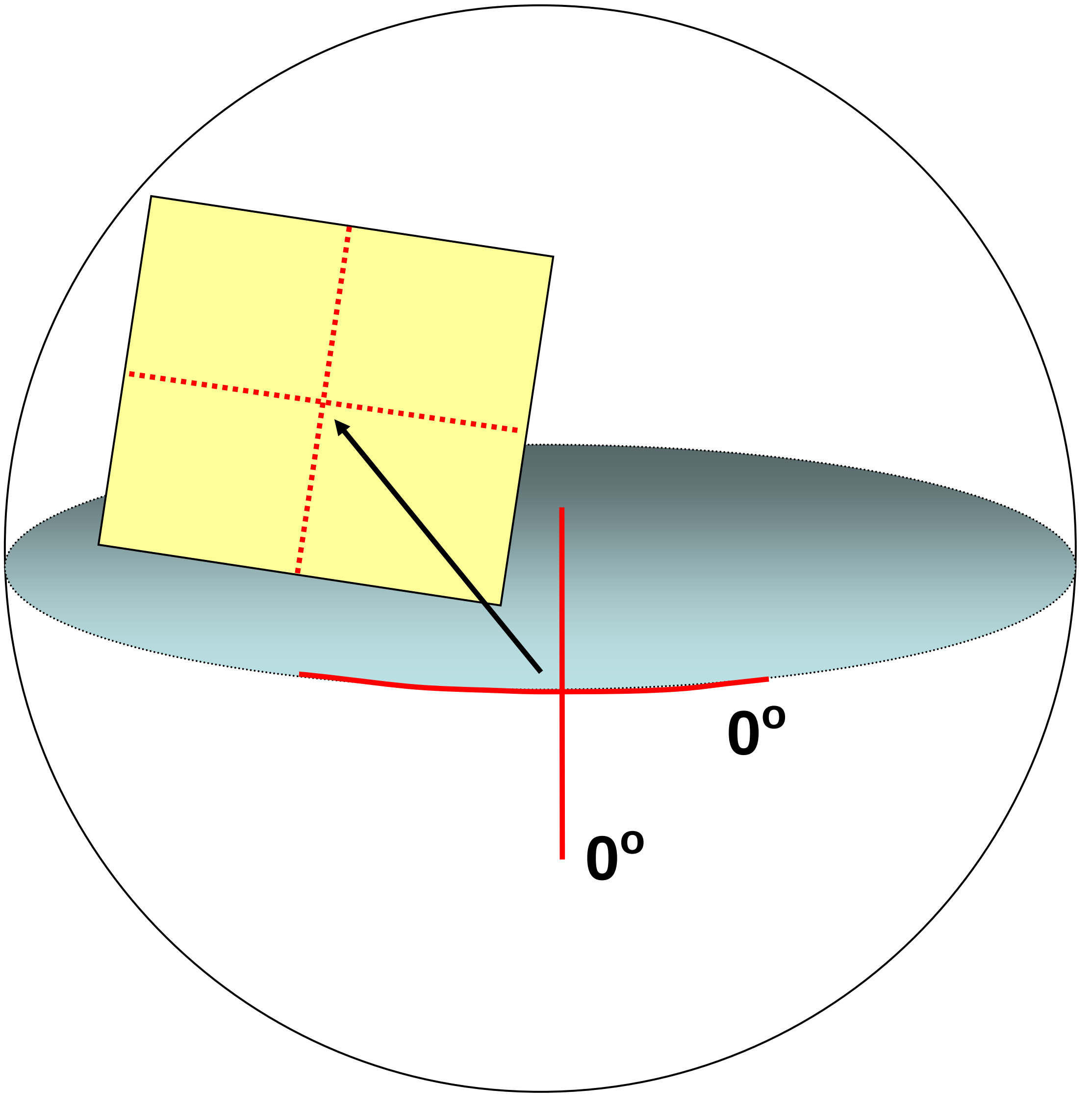


For a domain spanning 10N to 70N:

$$\Delta x \propto \cos(lat)$$

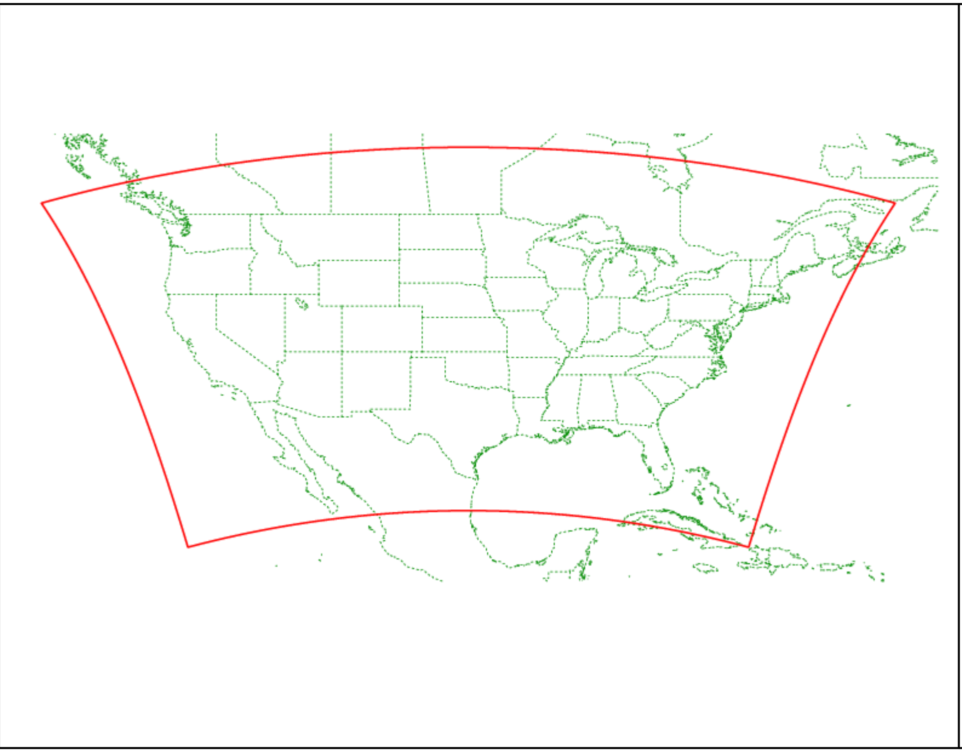
Regular lat-lon grid

$$\cos(70^0)/\cos(10^0) = 0.347$$

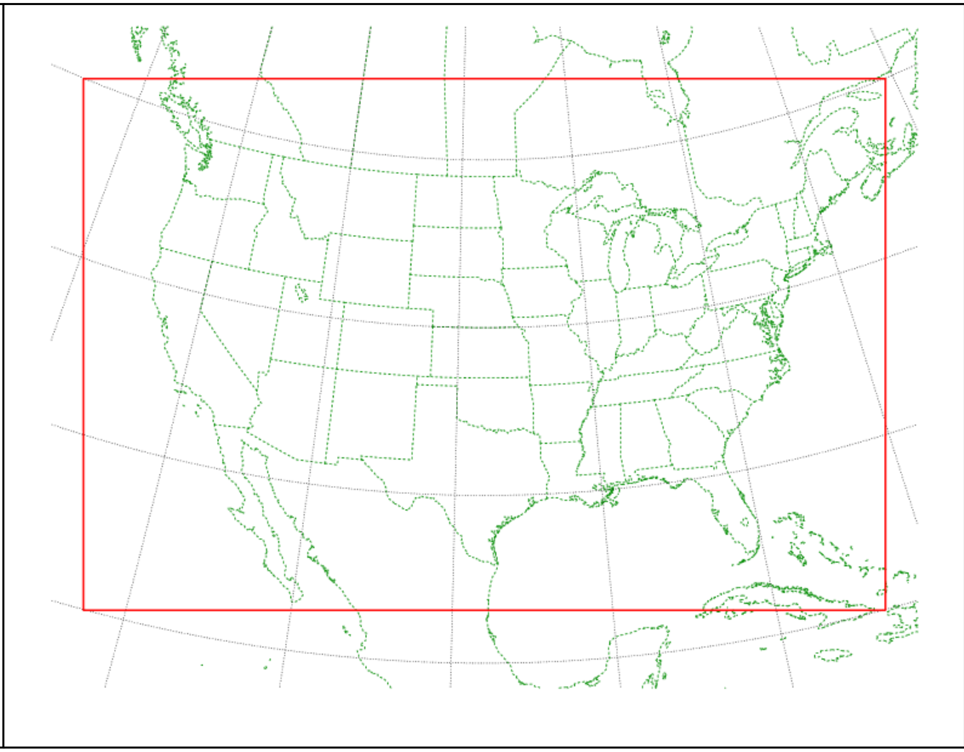
Rotated lat-lon grid

$$\cos(30^0)/\cos(0^0) = 0.866$$

Sample rotated lat-lon domain



Plotted on a regular latitude longitude map background



Plotted on a rotated latitude longitude map background (same rotation as model grid).

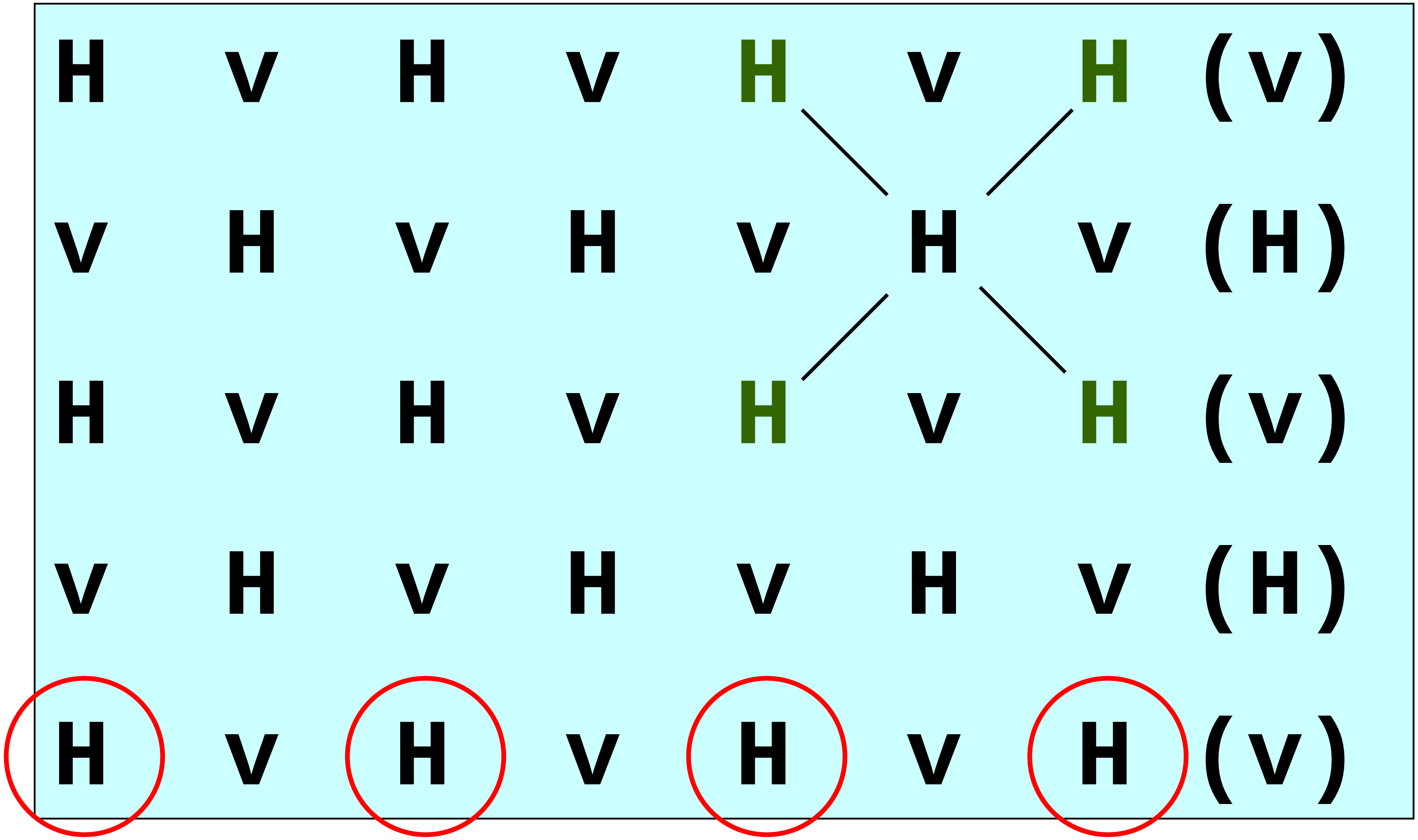
The E-grid Stagger

H	v	H	v	H	v	H	(v)
v	H	v	H	v	H	v	(H)
H	v	H	v	H	v	H	(v)
v	H	v	H	v	H	v	(H)
H	v	H	v	H	v	H	(v)

H=mass point, v=wind point

red=(1,1) ; blue=(1,2)

The E-grid Stagger



XDIM=4 (# of mass points on odd numbered row)

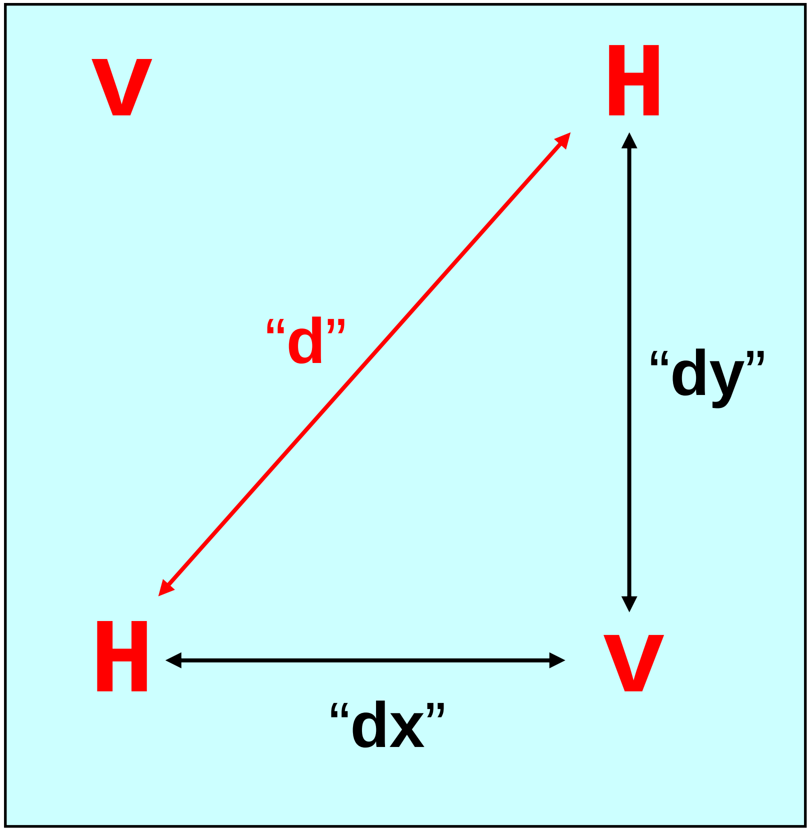
YDIM=5 (number of rows)

The E-grid Stagger - properties

- Due to the indexing convention, the X-dimension is half as large as would be expected from a C-grid domain (typically XDIM < YDIM for the E-grid).
- “Think diagonally” –the shortest distance between adjacent like points is along the diagonals of the grid.

- E-grid energy and enstrophy conserving momentum advection scheme (Janjic, 1984, MWR) controls the spurious nonlinear energy cascade (accumulation of small scale computational noise due to nonlinearity) more effectively than schemes on the C grid – an argument in favor of the E grid.

The E-grid Stagger

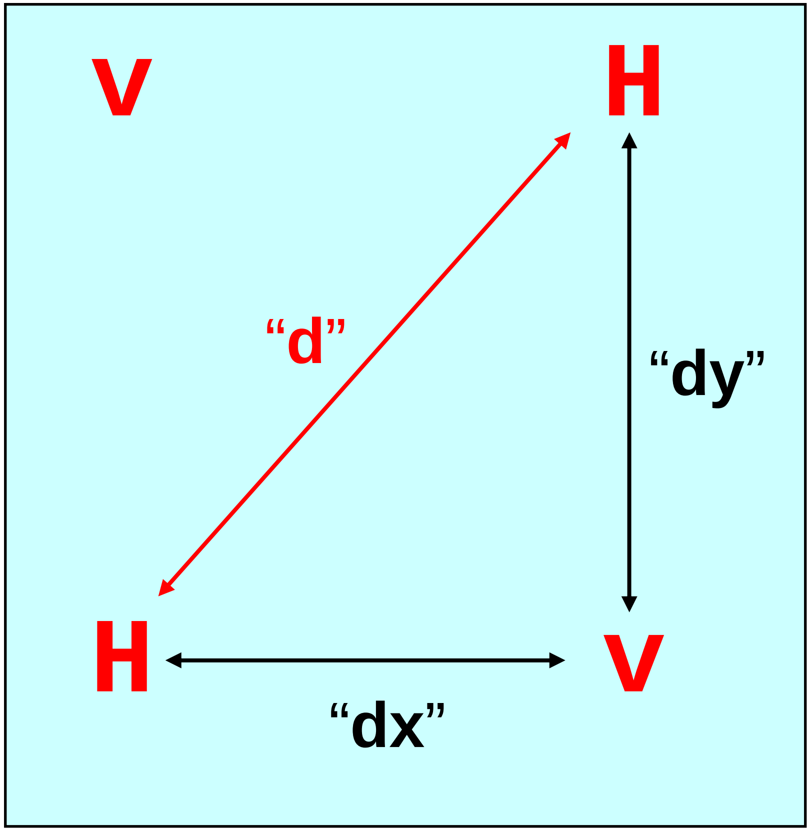


- Conventional grid spacing is the diagonal distance “**d**”.

- Grid spacings in the WPS and WRF namelists are the “**dx**” and “**dy**” values, *specified in fractions of a degree for the WRF-NMM*.

- “WRF domain wizard” takes input grid spacing “**d**” in km and computes the angular distances “**dx**” and “**dy**” for the namelist.

The E-grid Stagger - examples



"d" (km)	"dx" (deg)	"dy" (deg)	~dt (s)
4	.026	.0256	7.5
12	.078	.0768	26.66
32	.208	.2048	72

Note that dx > dy traditionally. Not a requirement, but helps offset the slight meridional convergence in the x-direction, and keeps the grid cells "more square" on average. More important for domains covering a large latitudinal expanse.

Spatial Discretization

- Basic discretization principle is conservation of important properties of the continuous system.
  - "Mimetic" approach

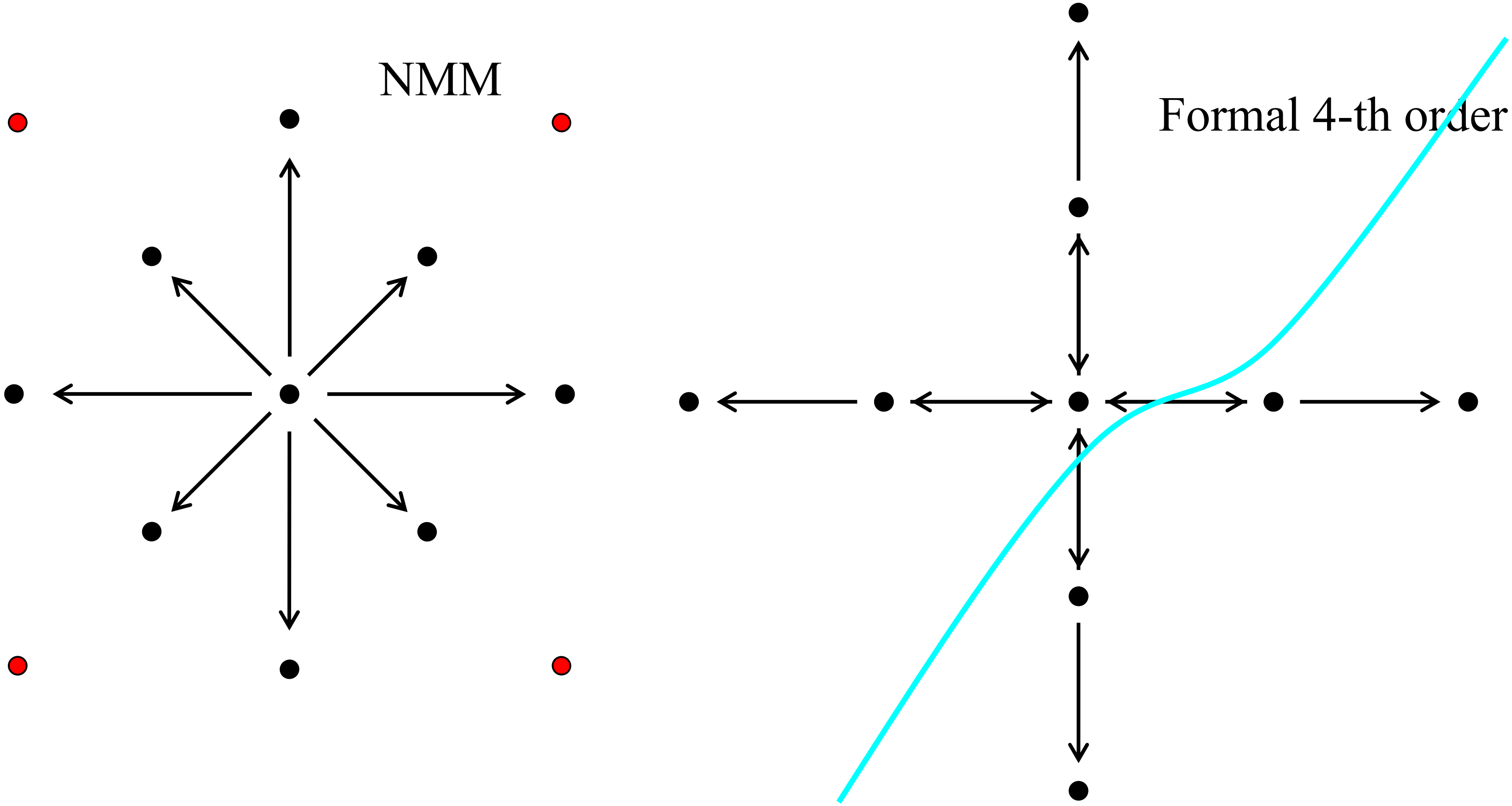
    http://www.math.unm.edu/~stanly/mimetic/mimetic.html

  - Something of a novelty in applied mathematics, …
  - … but well established in atmospheric modeling (Arakawa, 1966, 1972, 1977 …; Sadourny, 1975 …; Janjic, 1977, 1984 …; Tripoli, 1992, …)

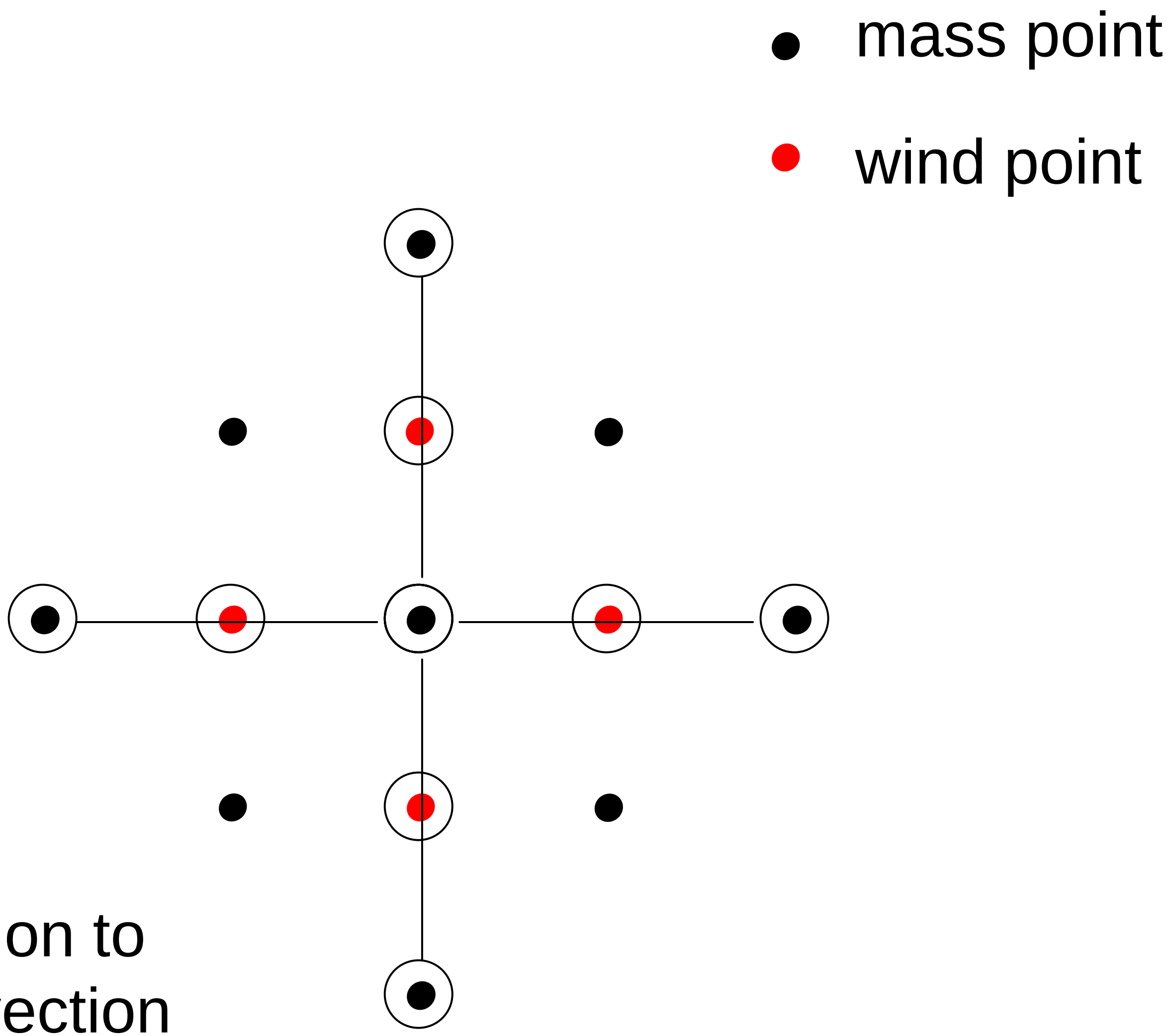
Spatial Discretization

General Philosophy

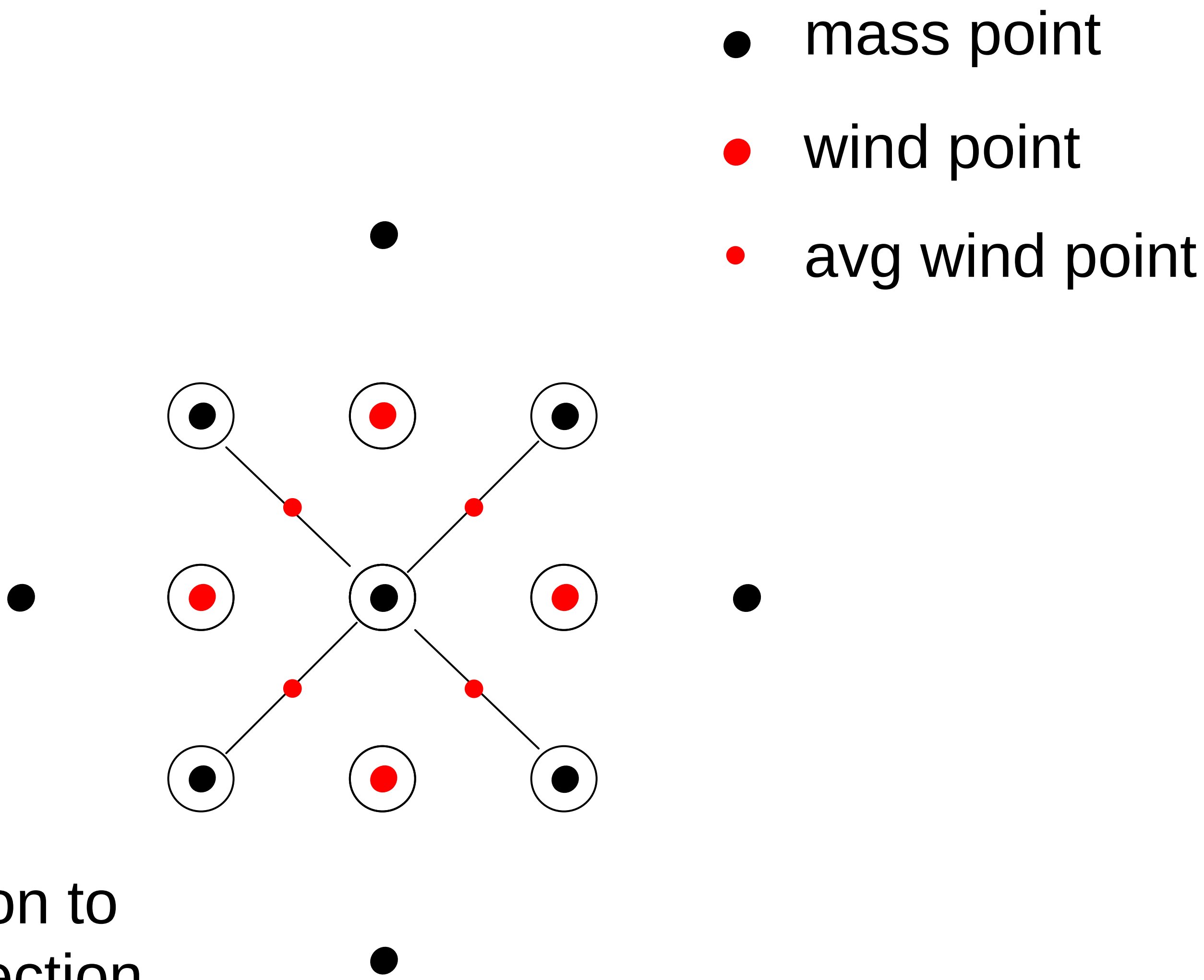
- Conserve energy and enstrophy in order to control nonlinear energy cascade; eliminate the need for numerical filtering to the extent possible.
- Conserve a number of first order and quadratic quantities (mass, momentum, energy, …).
- Use consistent order of accuracy for advection and divergence operators and the omega-alpha term; consistent transformations between KE and PE in the hydrostatic limit.
- Preserve properties of differential operators.



Advection and divergence operators – each point talks to all eight neighboring points (isotropic)

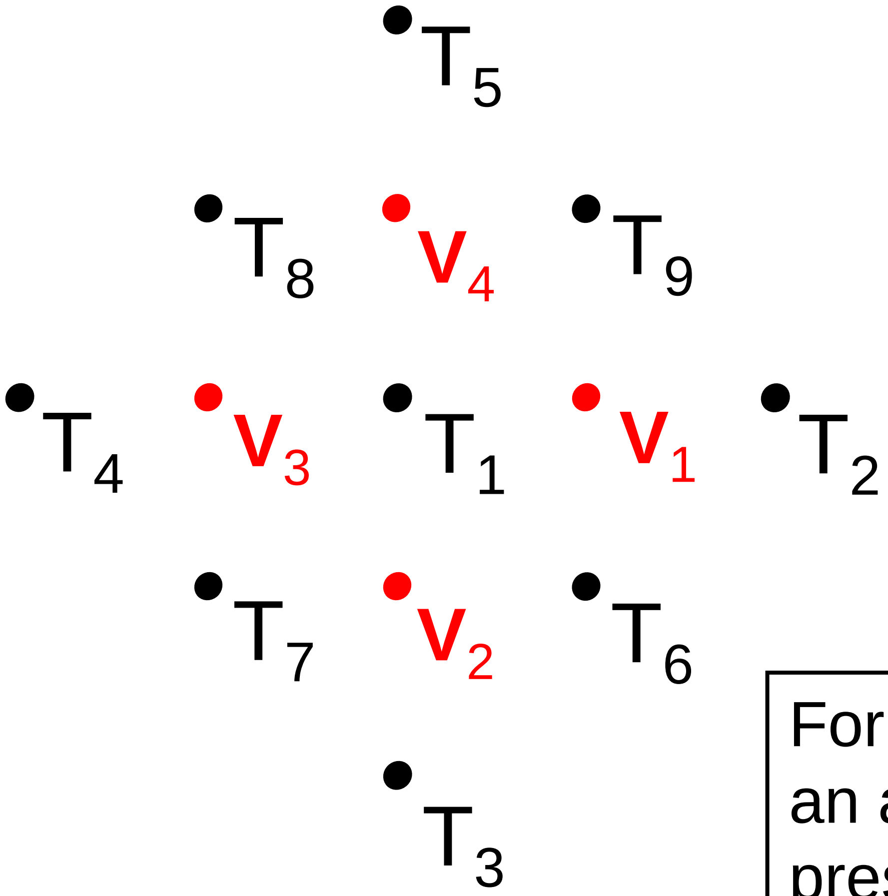


1/3 of contribution to divergence/advection comes from these N/S and E/W fluxes.



2/3 of contribution to divergence/advection comes from these diagonal fluxes.

Example of horizontal temperature advection in detail



For each T_n, there is an associated layer pressure depth (dp_n). There also is a dx_n specific to each point

Example of horizontal temperature advection in detail

Temperature fluxes in E/W, N/S, and diagonal directions:

$$TEW = u_3 dy\ (dp_1+dp_4)(T_1-T_4) + u_1 dy\ (dp_1+dp_2)(T_2-T_1)$$

$$TNS = v_2 dx_2\ (dp_1+dp_3)(T_1-T_3) + v_4 dx_4\ (dp_1+dp_5)(T_5-T_1)$$

$$TNE = [(u_1 dy + v_1 dx_1 + u_4 dy + v_4 dx_4)\ (dp_1+dp_9)\ (T_9-T_1)$$
$$+ \quad (u_3 dy + v_3 dx_3 + u_2 dy + v_2 dx_2)\ (dp_1+dp_7)\ (T_1-T_7)]$$

$$TSE = [(u_1 dy - v_1 dx_1 + u_2 dy - v_2 dx_2)\ (dp_1+dp_6)\ (T_6-T_1)$$
$$+ \quad (u_3 dy - v_3 dx_3 + u_4 dy - v_4 dx_4)\ (dp_1+dp_8)\ (T_1-T_8)]$$

Advective tendency, ADT, combines the fluxes:

$$ADT = (TEW + TNS + TNE + TSE) * (-dt/24) * (1/dx_1 * dy * dp_1)$$

NMM Vertical Coordinate

Pressure-sigma hybrid (Arakawa and Lamb, 1977)

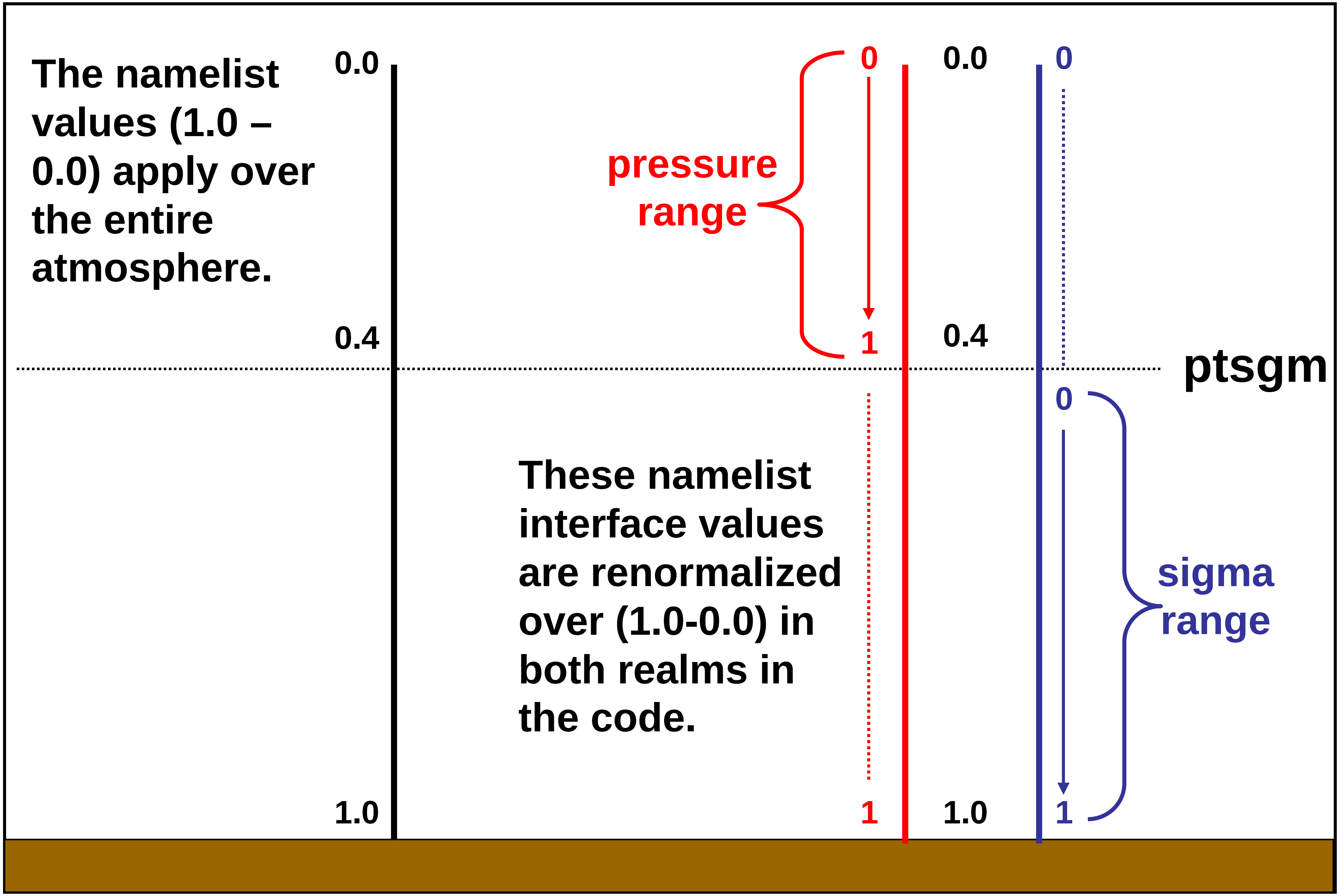
Has the desirable properties of a terrain-following, pressure coordinate:

- Exact mass (etc.) conservation
- Nondivergent flow on pressure surfaces
- No problems with weak static stability
- No discontinuities or internal boundary conditions

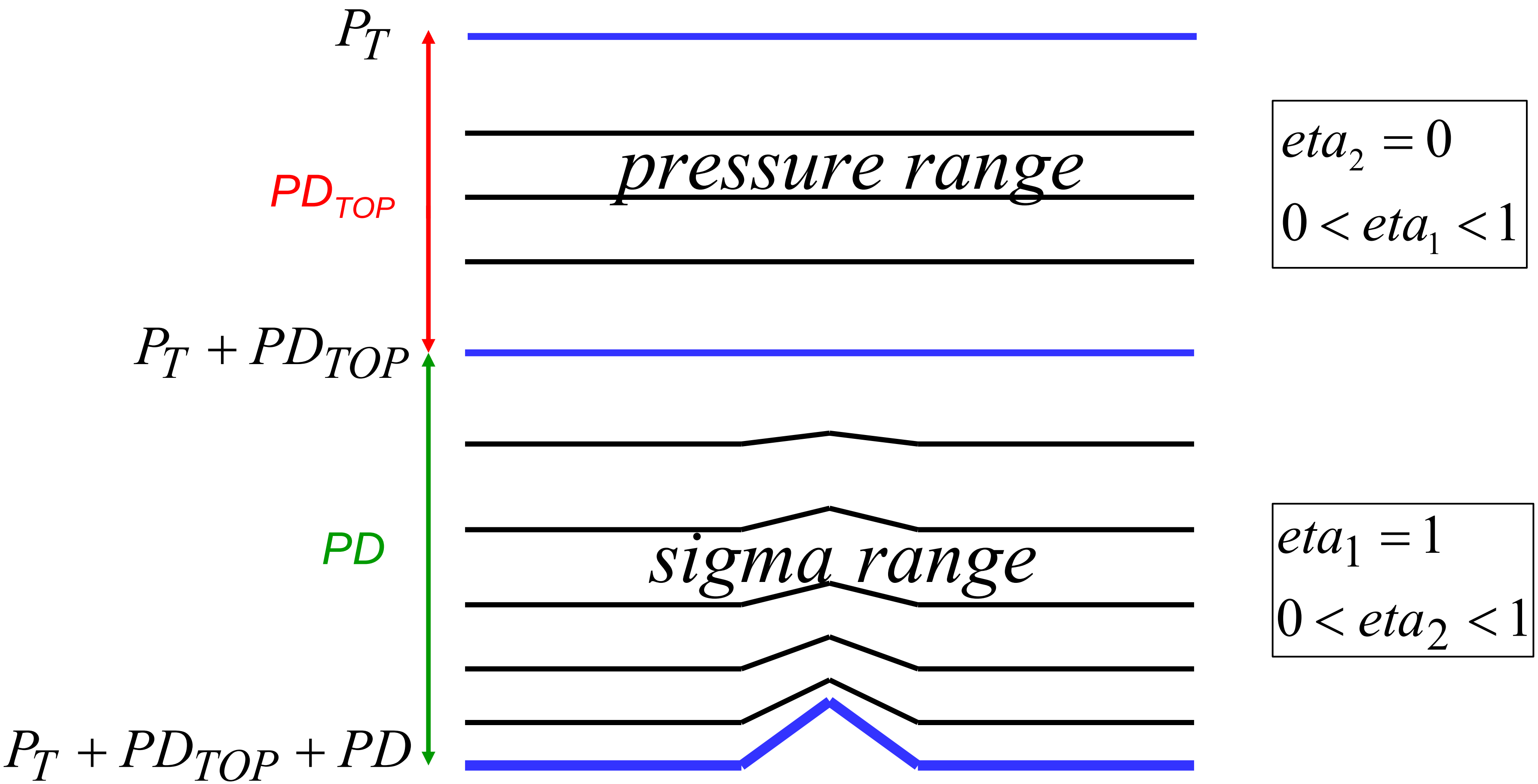
And an additional benefit from the hybrid:

- Flat coordinate surfaces at high altitudes where sigma problems worst (e.g., Simmons and Burridge, 1981)

Pressure-Sigma Hybrid Vertical Coordinate



Pressure-Sigma Hybrid Vertical Coordinate



$$\pi = eta_1 * PD_{TOP} + eta_2 * PD + P_T$$

Equations in Hybrid Coordinate

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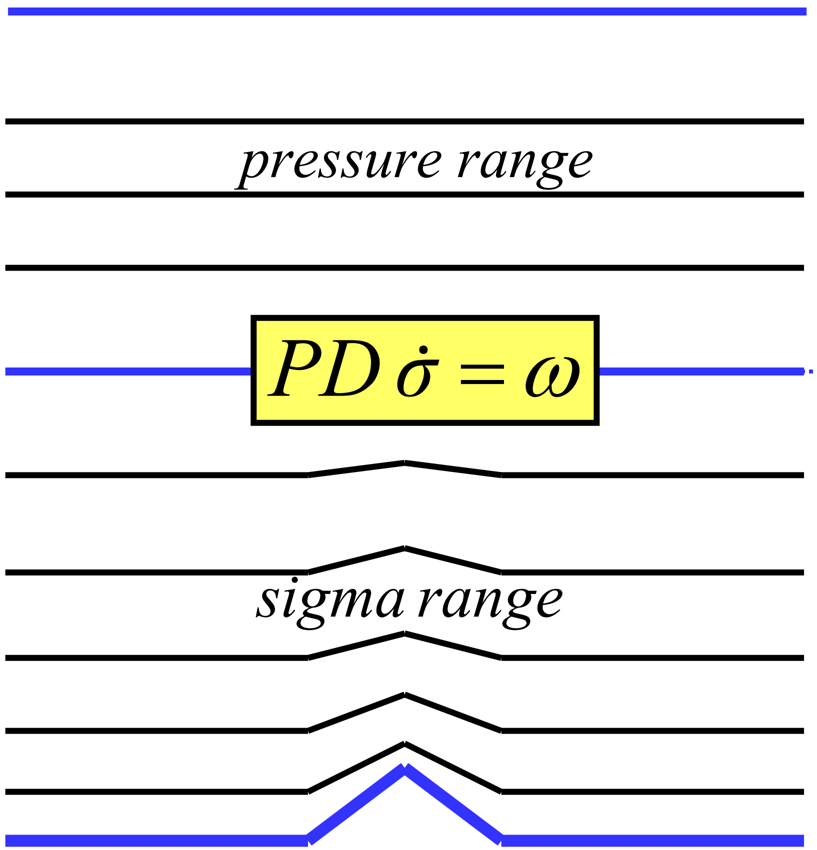
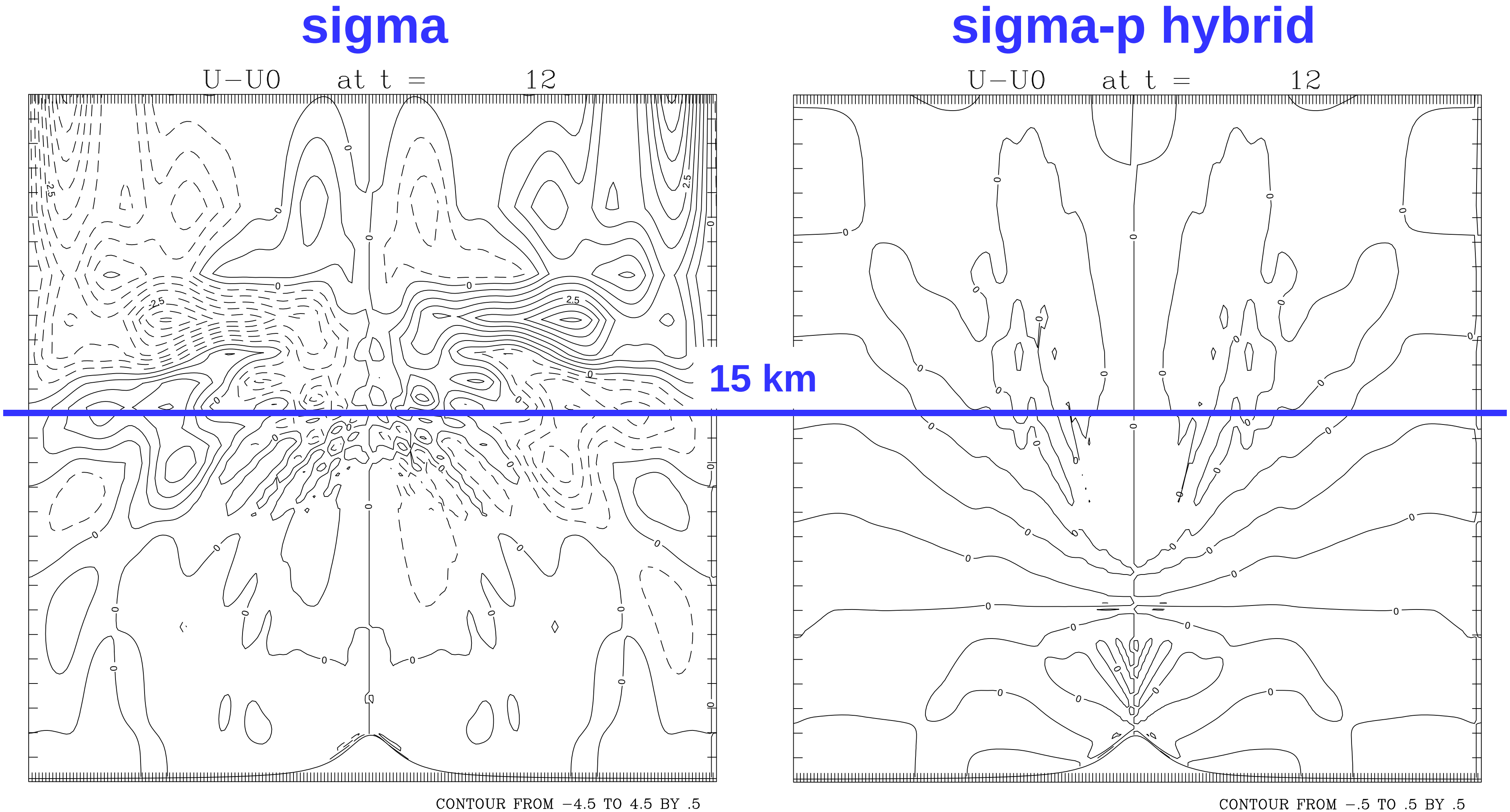


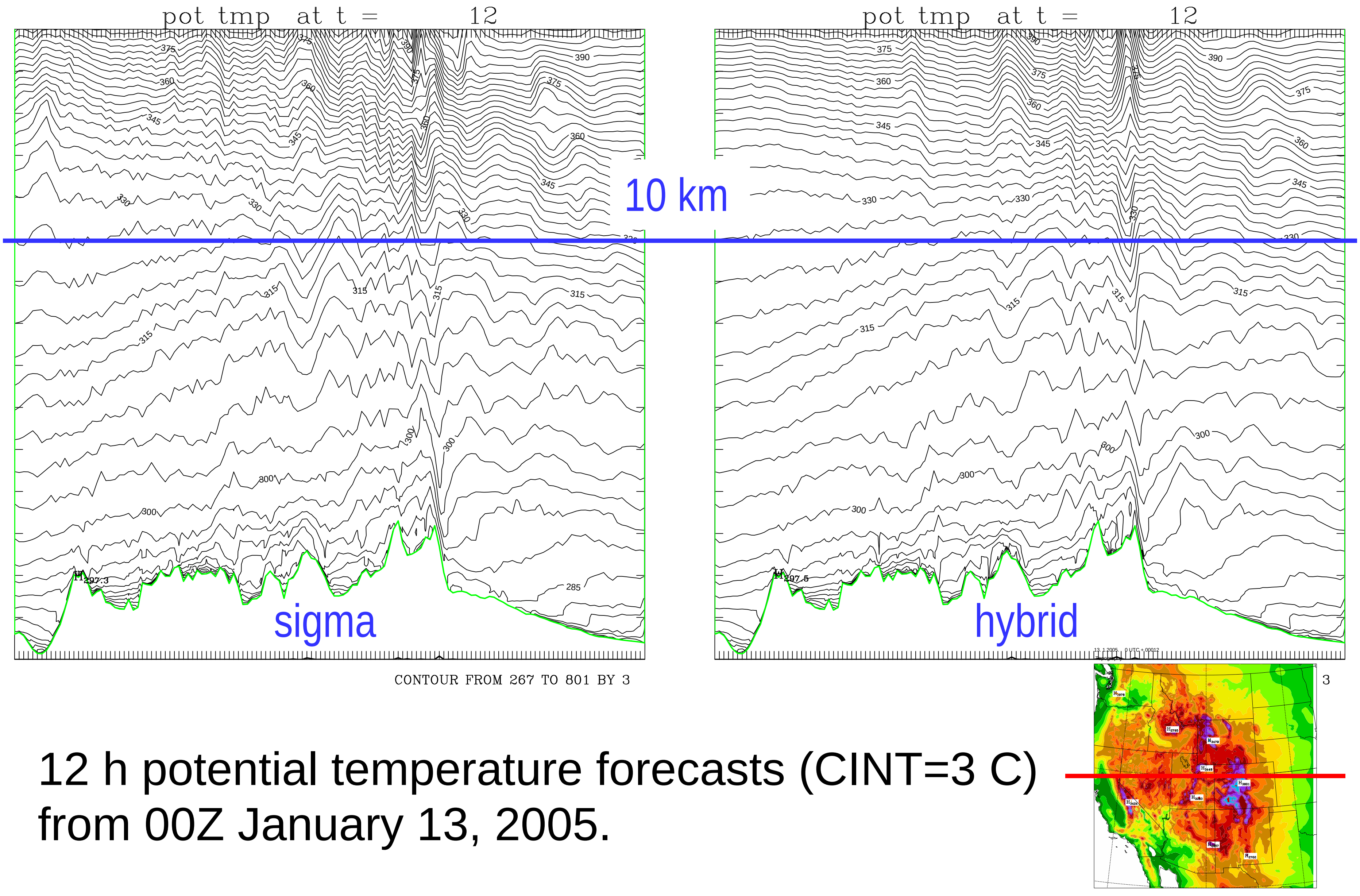
$$\nabla_p \bullet (\mathbf{v}) + \frac{\partial \omega}{\partial p} = 0$$

$$\frac{\partial PD}{\partial t} + \nabla_\sigma \bullet (PD\ \mathbf{v}) + \frac{\partial (PD\,\dot{\sigma})}{\partial \sigma} = 0$$



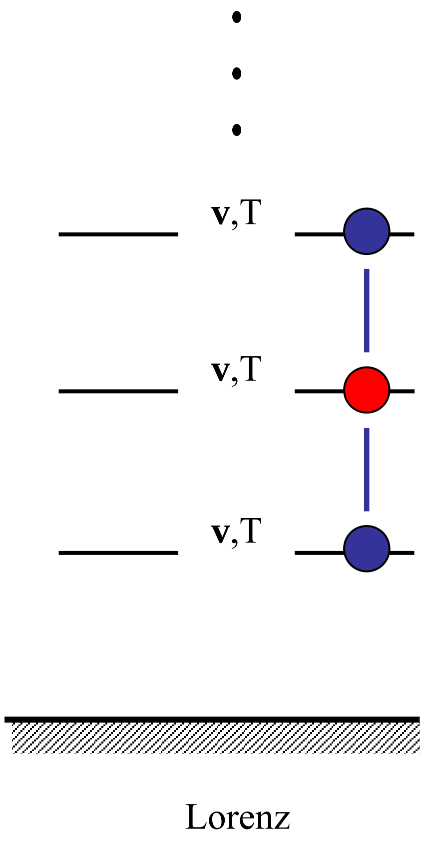
Wind developing due to the spurious pressure gradient force in an idealized integration. The hybrid coordinate boundary between the pressure and sigma domains is at ~400 hPa.

Example of nonphysical small scale energy source



12 h potential temperature forecasts (CINT=3 C) from 00Z January 13, 2005.

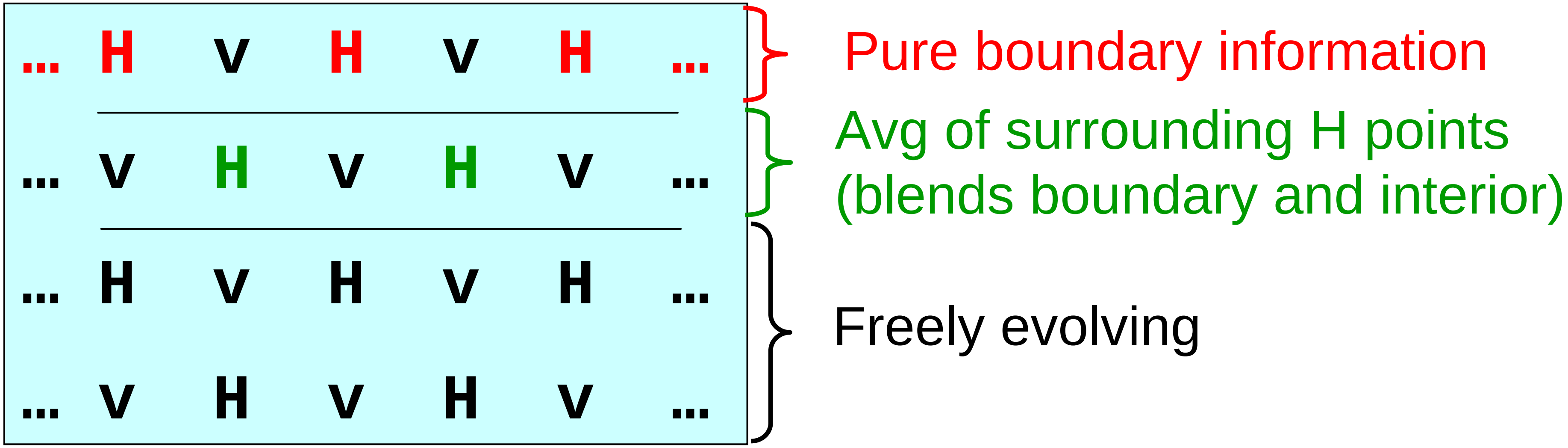
Vertical discretization



Vertical advection combines the advective fluxes computed above and below the layer of interest.

Lateral Boundary Conditions

- Lateral boundary information prescribed only on outermost row:



- Upstream advection in three rows next to the boundary
  - No computational outflow boundary condition for advection

- Enhanced divergence damping close to the boundaries.

Dissipative Processes – lateral diffusion

A 2nd order, nonlinear Smagorinsky-type horizontal diffusion is utilized:

- Diffusion strength a function of the local TKE, deformation of the 3D flow, and a code-specified constant (COAC*).

- Lateral diffusion is zeroed for sloping model surfaces (> ~ 54 m / 12 km grid point).

* COAC has a default value of 1.6 and is specified in ./dyn_nmm/module_initialize_real.F. Larger values generate more diffusive smoothing.

Dissipative Processes - divergence damping *

- Internal mode damping (on each vertical layer)

$$\mathrm{v}_j = \mathrm{v}_j + \frac{(\nabla \cdot dp_{j+1}\vec{\mathbf{v}}_{j+1} - \nabla \cdot dp_{j-1}\vec{\mathbf{v}}_{j-1})}{(dp_{j+1} + dp_{j-1})} \cdot DDMPV$$

- External mode damping (vertically integrated)

$$\mathrm{v}_j = \mathrm{v}_j + \frac{(\int \nabla \cdot dp_{j+1}\vec{\mathbf{v}}_{j+1} - \int \nabla \cdot dp_{j-1}\vec{\mathbf{v}}_{j-1})}{(\int dp_{j+1} + \int dp_{j-1})} \cdot DDMPV$$

$$DDMPV \approx \sqrt{2} \cdot dt \cdot CODAMP$$

CODAMP is a code-specified parameter = 6.4 by default.

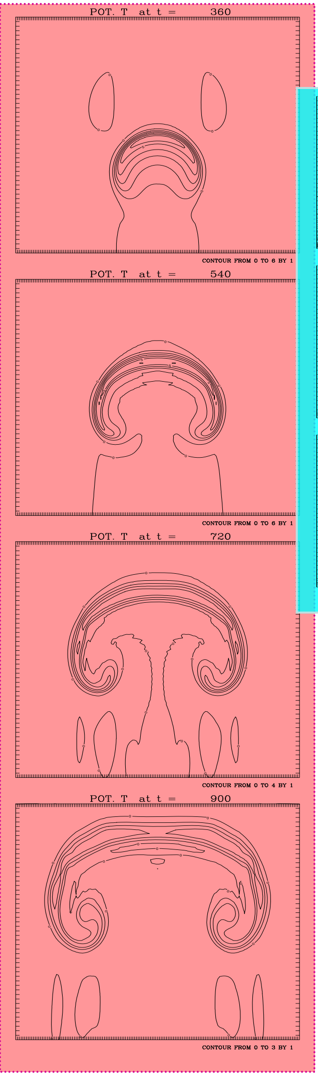
Gravity Wave Drag & Mountain Blocking

- Accounts for sub-grid scale mountain effects: mountain waves (GWD) and stability-dependent blocking of low-level flow around topography (MB).
- More important for coarser grid spacing (> ~10 km) and longer (multi-day) integrations.
- gwd_opt=2 in physics namelist to invoke for the WRF-NMM.
- Benefits overall synoptic patterns and near-surface wind and temperature forecasts.
- Based on the GFS model package for GWD (Alpert et al., 1988, 1996; Kim & Arakawa,1995) and MB (Lott & Miller, 1997).

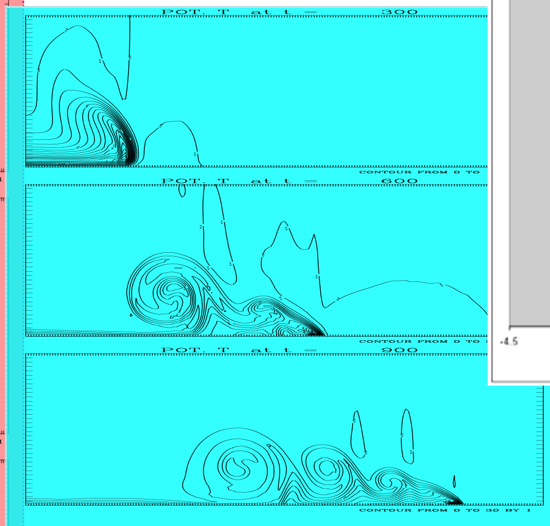
Courtesy of Brad Ferrier

Dynamics formulation tested on various scales

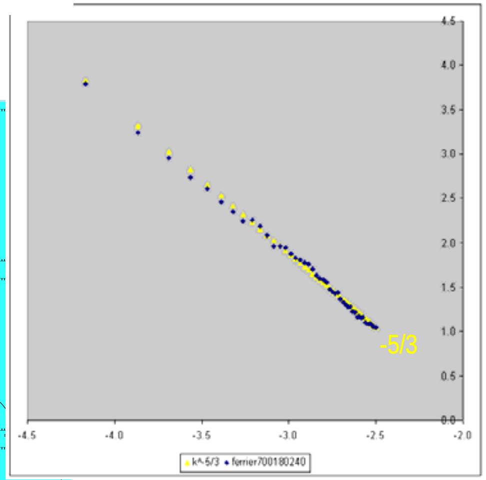
Warm bubble



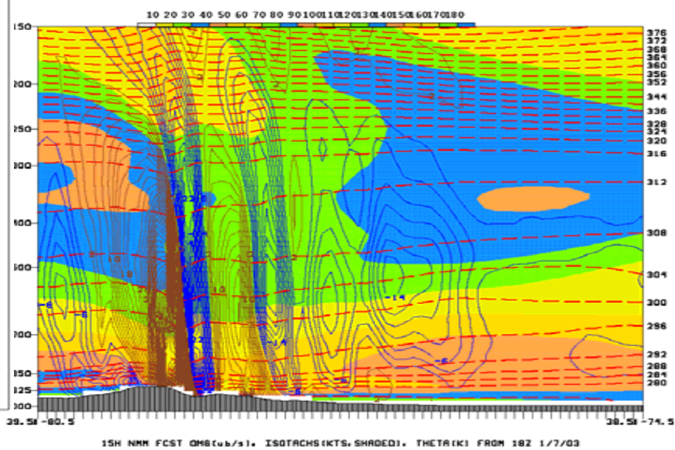
Cold bubble



Decaying 3D turbulence

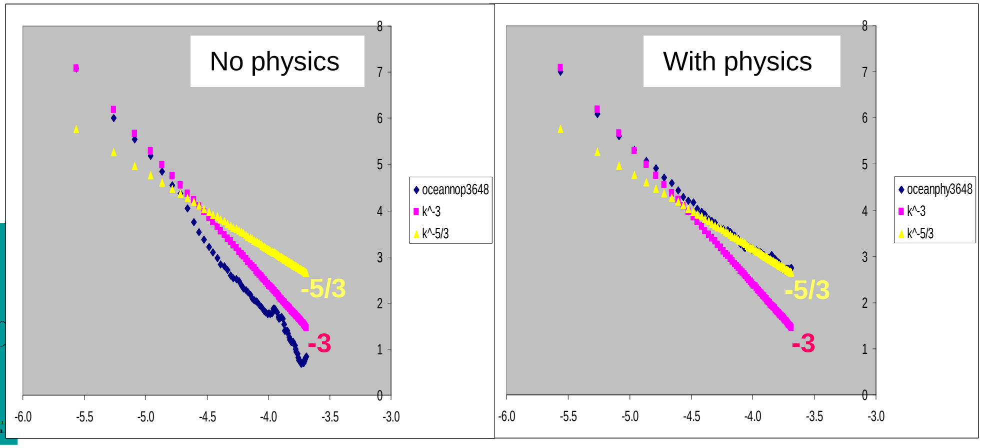


Mountain waves



Convection

Atmospheric spectra



Summary

- Robust, reliable, fast
- Represents an extension of NWP methods developed and refined over a decades-long period into the nonhydrostatic realm.
- Utilized at NCEP in the NAM, HWRF, Hires Window* and Short Range Ensemble Forecast (SREF*) operational systems.

* = WRF-ARW used in these systems as well