

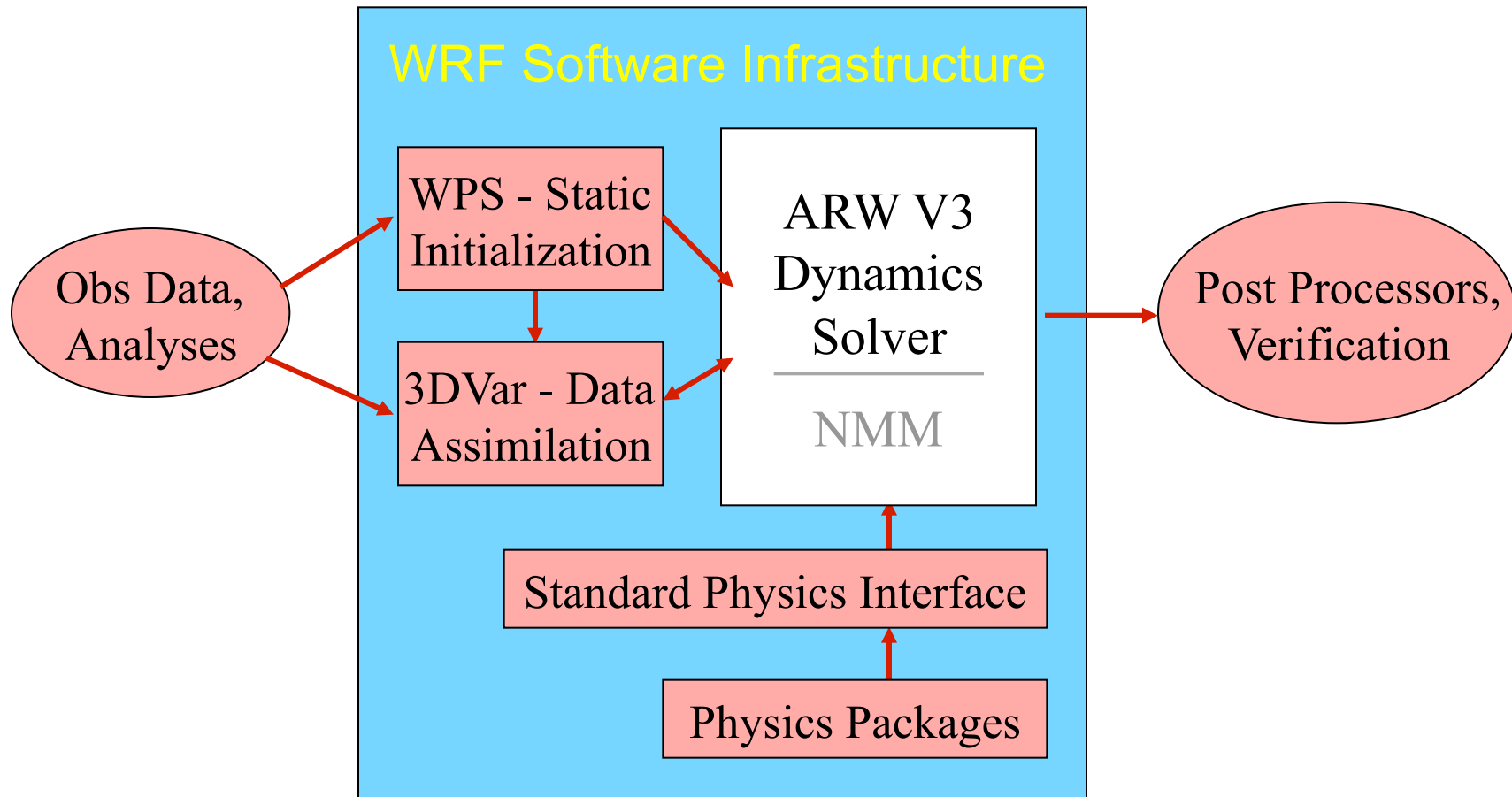
# The Advanced Research WRF (ARW) Dynamics Solver

Bill Skamarock

[skamaroc@ucar.edu](mailto:skamaroc@ucar.edu)

Jimmy Dudhia

[dudhia@ucar.edu](mailto:dudhia@ucar.edu)



### WRF ARW Tech Note

A Description of the Advanced Research WRF Version 3

<http://www.mmm.ucar.edu/wrf/users/pub-doc.html>

# ARW Dynamical Solver

- Terrain representation
- Vertical coordinate
- Equations / variables
- Time integration scheme
- Grid staggering
- Advection scheme
- Time step parameters
- Filters
- Boundary conditions
- Nesting
- Map projections

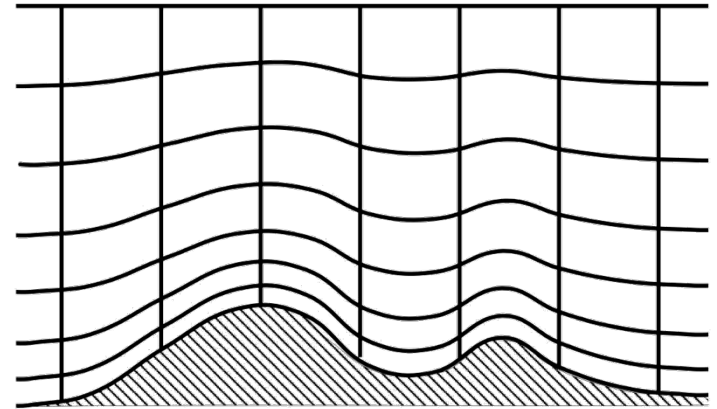
# Vertical Coordinate and Prognostic Variables

Hydrostatic pressure  $\pi$

Column mass  $\mu = \pi_s - \pi_t$   
(per unit area)

Vertical coordinate  $\eta = \frac{(\pi - \pi_t)}{\mu}$

Layer mass  $\mu \Delta \eta = \Delta \pi = g \rho \Delta z$   
(per unit area)



Conserved state (prognostic) variables:

$$\mu, \quad U = \mu u, \quad V = \mu v, \quad W = \mu w, \quad \Theta = \mu \theta$$

Non-conserved state variable:  $\phi = gz$

## 2D Flux-Form Moist Equations in ARW

Moist Equations:

$$\frac{\partial U}{\partial t} + \alpha \mu_d \frac{\partial p}{\partial x} + \frac{\alpha}{\alpha_d} \frac{\partial p}{\partial \eta} \frac{\partial \phi}{\partial x} = - \frac{\partial Uu}{\partial x} - \frac{\partial \Omega u}{\partial \eta}$$

$$\frac{\partial W}{\partial t} + g \left( \mu_d - \frac{\alpha}{\alpha_d} \frac{\partial p}{\partial \eta} \right) = - \frac{\partial Uw}{\partial x} - \frac{\partial \Omega w}{\partial \eta}$$

$$\frac{\partial \mu_d}{\partial t} + \frac{\partial U}{\partial x} + \frac{\partial \Omega}{\partial \eta} = 0$$

$$\frac{\partial \Theta}{\partial t} + \frac{\partial U\theta}{\partial x} + \frac{\partial \Omega\theta}{\partial \eta} = \mu Q$$

$$\frac{d\phi}{dt} = gw$$

$$\frac{\partial (\mu_d q_{v,l})}{\partial t} + \frac{\partial (U q_{v,l})}{\partial x} + \frac{\partial (\Omega q_{v,l})}{\partial \eta} = \mu Q_{v,l}$$

Diagnostic relations:

$$\frac{\partial \phi}{\partial \eta} = -\alpha_d \mu_d, \quad p = \left( \frac{R\Theta}{p_o \mu_d \alpha_v} \right)^\gamma$$

# Time Integration in ARW

## 3<sup>rd</sup> Order Runge-Kutta time integration

advance  $\phi^t \rightarrow \phi^{t+\Delta t}$

$$\phi^* = \phi^t + \frac{\Delta t}{3} R(\phi^t)$$

$$\phi^{**} = \phi^t + \frac{\Delta t}{2} R(\phi^*)$$

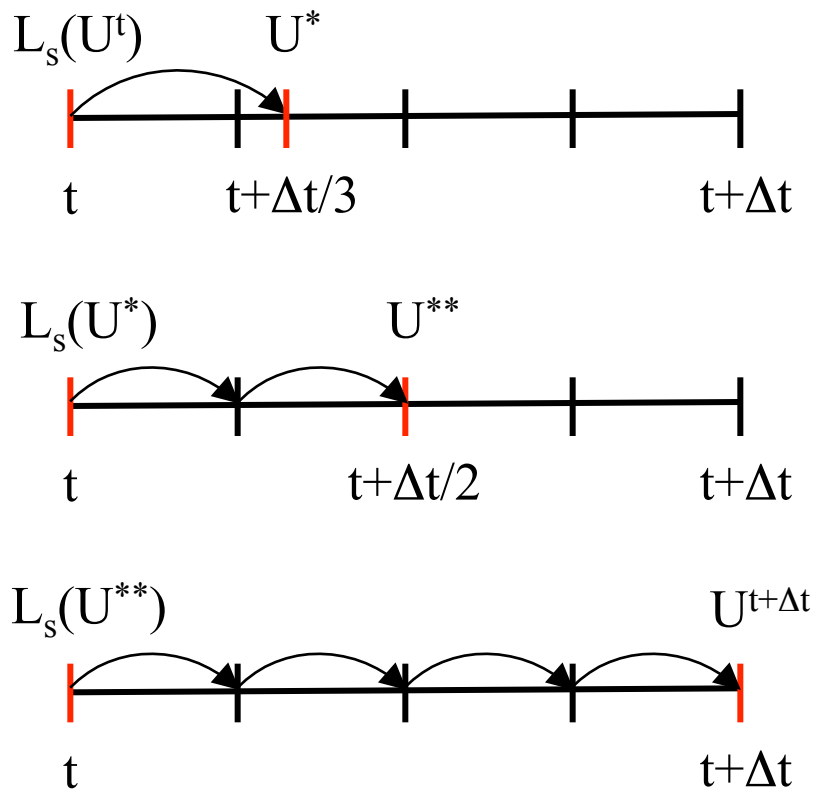
$$\phi^{t+\Delta t} = \phi^t + \Delta t R(\phi^{**})$$

Amplification factor  $\phi_t = i k \phi$ ;  $\phi^{n+1} = A \phi^n$ ;  $|A| = 1 - \frac{(k\Delta t)^4}{24}$

# Time-Split Runge-Kutta Integration Scheme

$$U_t = L_{\text{fast}}(U) + L_{\text{slow}}(U)$$

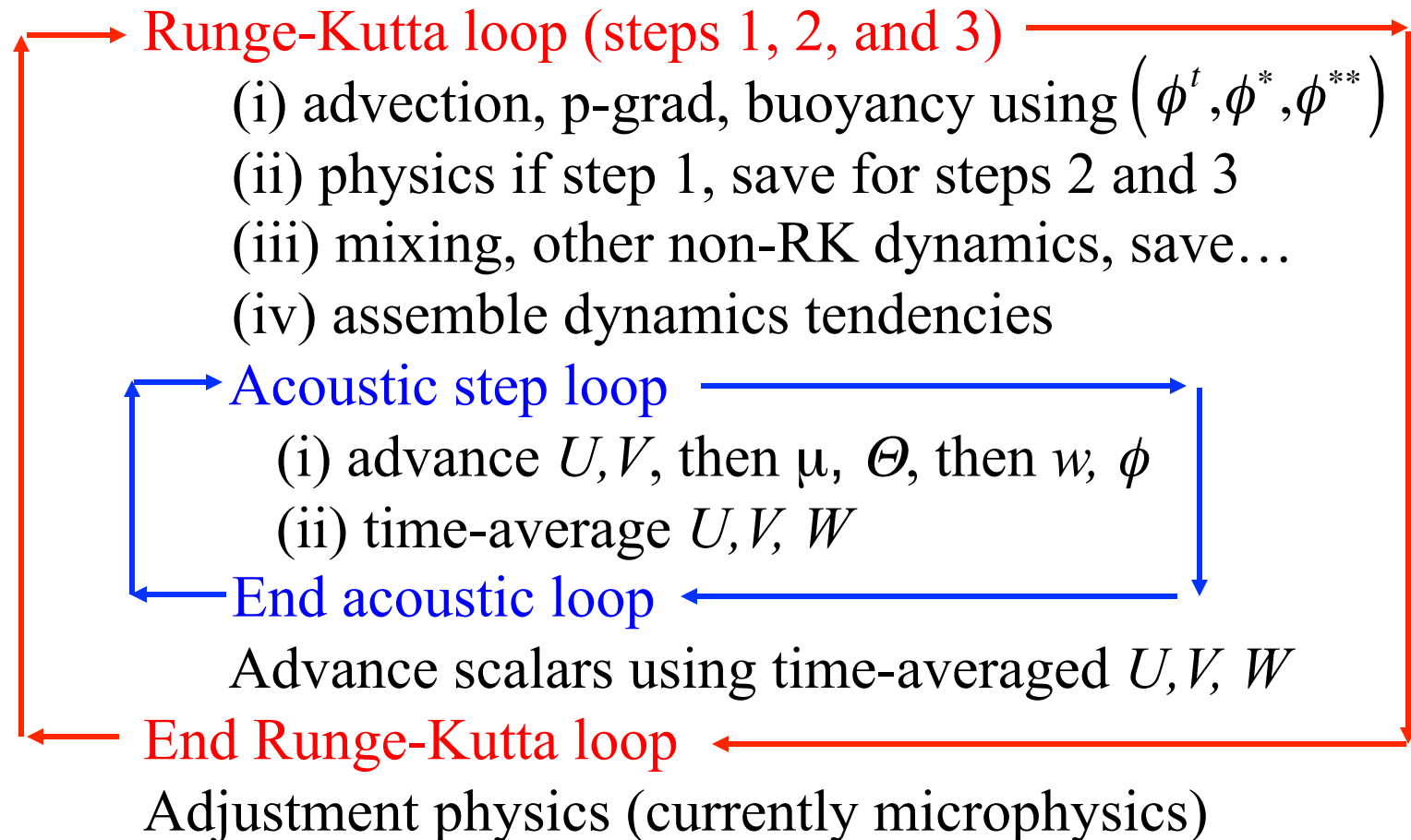
3rd order Runge-Kutta, 3 steps



- RK3 is 3rd order accurate for linear eqns, 2nd order accurate for nonlinear eqns.
- Stable for centered and upwind advection schemes.
- Stable for Courant number  $U\Delta t/\Delta x < 1.73$
- Three  $L_{\text{slow}}(U)$  evaluations per timestep.

# WRF ARW Model Integration Procedure

Begin time step

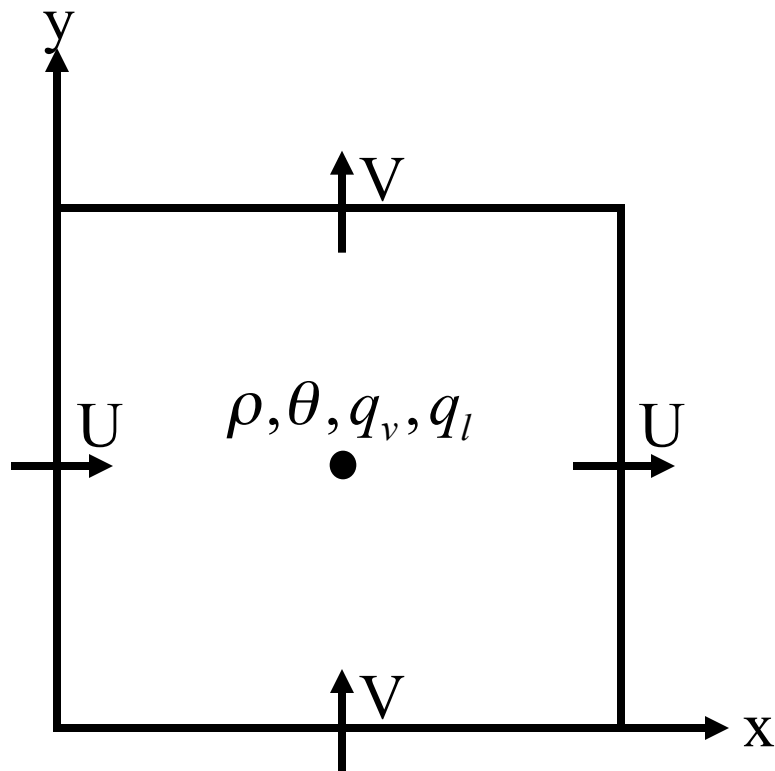


End time step

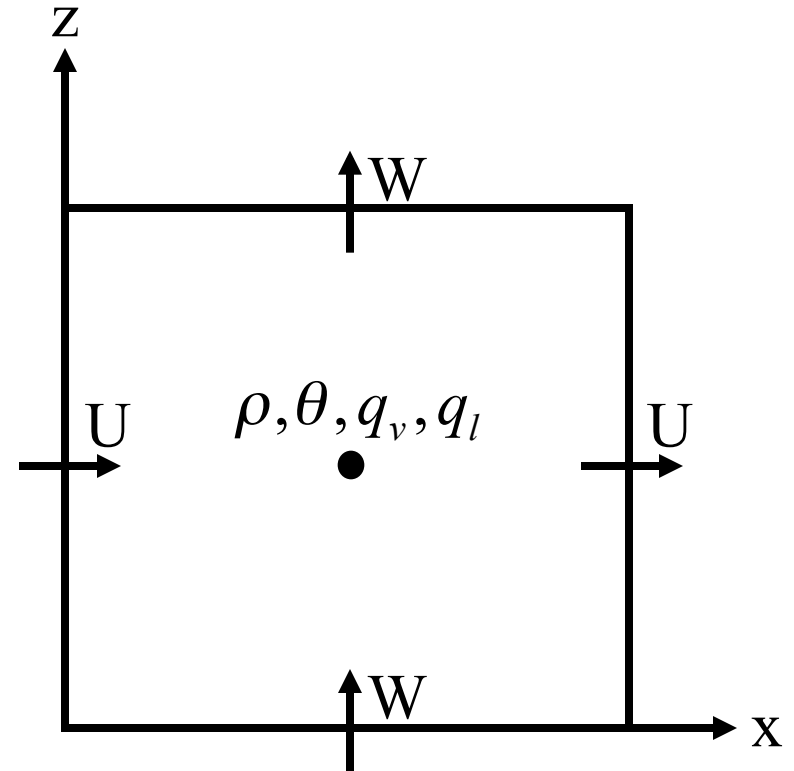


# ARW model, grid staggering

## C-grid staggering



horizontal



vertical

## Advection in the ARW Model

2<sup>nd</sup>, 3<sup>rd</sup>, 4<sup>th</sup>, 5<sup>th</sup> and 6<sup>th</sup> order centered and upwind-biased schemes are available in the ARW model.

Example: 5<sup>th</sup> order scheme

$$\frac{\partial(U\phi)}{\partial x} = \frac{1}{\Delta x} \left( F_{i+\frac{1}{2}}(U\phi) - F_{i-\frac{1}{2}}(U\phi) \right)$$

where

$$F_{i-\frac{1}{2}}(U\phi) = U_{i-\frac{1}{2}} \left\{ \frac{37}{60}(\phi_i + \phi_{i-1}) - \frac{2}{15}(\phi_{i+1} + \phi_{i-2}) + \frac{1}{60}(\phi_{i+2} + \phi_{i-3}) \right\} \\ - \text{sign}(1, U) \frac{1}{60} \left\{ (\phi_{i+2} - \phi_{i-3}) - 5(\phi_{i+1} - \phi_{i-2}) + 10(\phi_i - \phi_{i-1}) \right\}$$

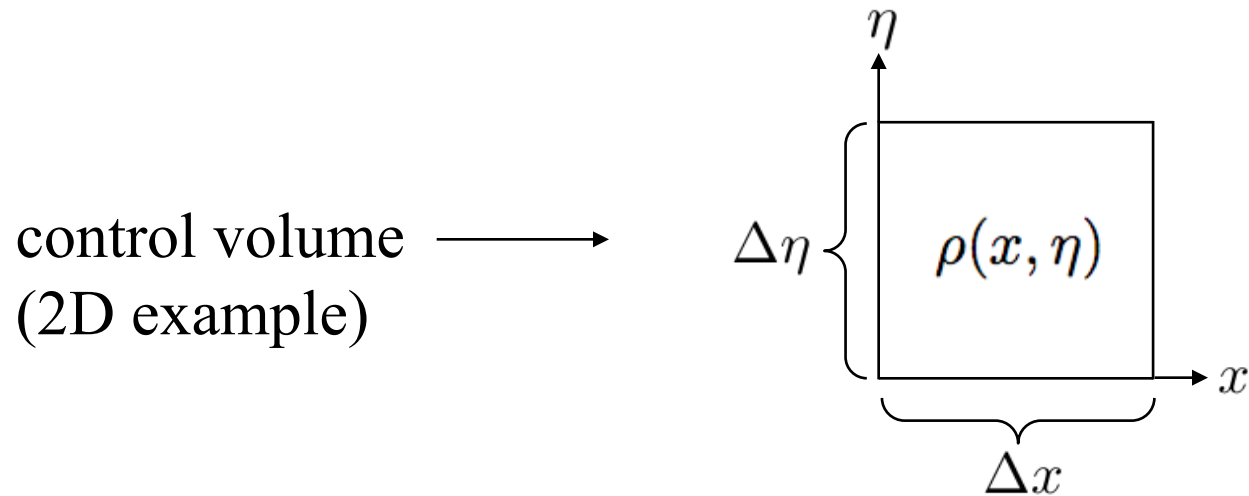
## Advection in the ARW Model

For constant  $U$ , the 5<sup>th</sup> order flux divergence tendency becomes

$$\begin{aligned} \Delta t \frac{\delta(U\phi)}{\Delta x} \Big|_{5th} &= \Delta t \frac{\delta(U\phi)}{\Delta x} \Big|_{6th} \\ &- \underbrace{\left| \frac{U\Delta t}{\Delta x} \right| \frac{1}{60} (-\phi_{i-3} + 6\phi_{i-2} - 15\phi_{i-1} + 20\phi_i - 15\phi_{i+1} + 6\phi_{i+2} - \phi_{i+3})}_{\frac{Cr}{60} \frac{\partial^6 \phi}{\partial x^6} + H.O.T} \end{aligned}$$

The odd-ordered flux divergence schemes are equivalent to the next higher ordered (even) flux-divergence scheme plus a dissipation term of the higher even order with a coefficient proportional to the Courant number.

# Mass Conservation in the ARW Model



Mass in a control volume is proportional to

$$(\Delta x \Delta \eta)(\mu)^t$$

since  $\mu(x)\Delta\eta = \Delta\pi = -g\rho\Delta z$

# Mass Conservation in the ARW Model

Mass in a control volume  $(\Delta x \Delta \eta)(\mu)^t$   
2D example

Mass conservation equation

$$\Delta t^{-1}(\Delta x \Delta \eta) \cdot [(\mu)^{t+\Delta t} - (\mu)^t] = [(\mu u \Delta \eta)_{x-\Delta x/2, \eta} - (\mu u \Delta \eta)_{x+\Delta x/2, \eta}] + [(\mu \omega \Delta x)_{x, \eta-\Delta \eta/2} - (\mu \omega \Delta x)_{x, \eta+\Delta \eta/2}]$$

Change in mass over a time step

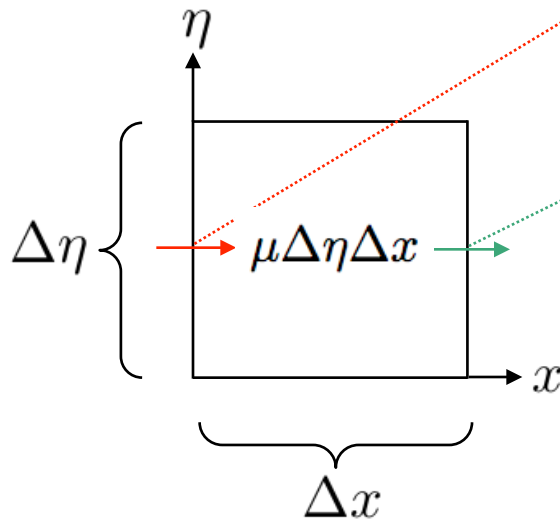
mass fluxes through  
control volume faces

# Mass Conservation in the ARW Model

Mass in a control volume  $(\Delta x \Delta \eta)(\mu)^t$

Mass conservation equation

$$\Delta t^{-1}(\Delta x \Delta \eta) \cdot [(\mu)^{t+\Delta t} - (\mu)^t] = [(\mu u \Delta \eta)_{x-\Delta x/2, \eta} - (\mu u \Delta \eta)_{x+\Delta x/2, \eta}] + [(\mu \omega \Delta x)_{x, \eta-\Delta \eta/2} - (\mu \omega \Delta x)_{x, \eta+\Delta \eta/2}]$$



Horizontal fluxes through the vertical control-volume faces

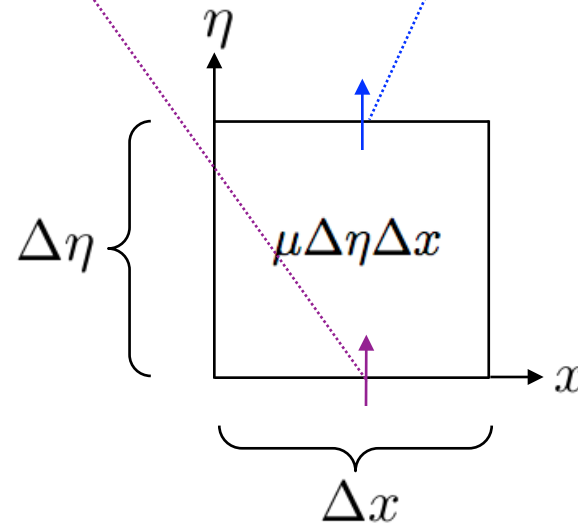
# Mass Conservation in the ARW Model

Mass in a control volume  $(\Delta x \Delta \eta)(\mu)^t$

Mass conservation equation

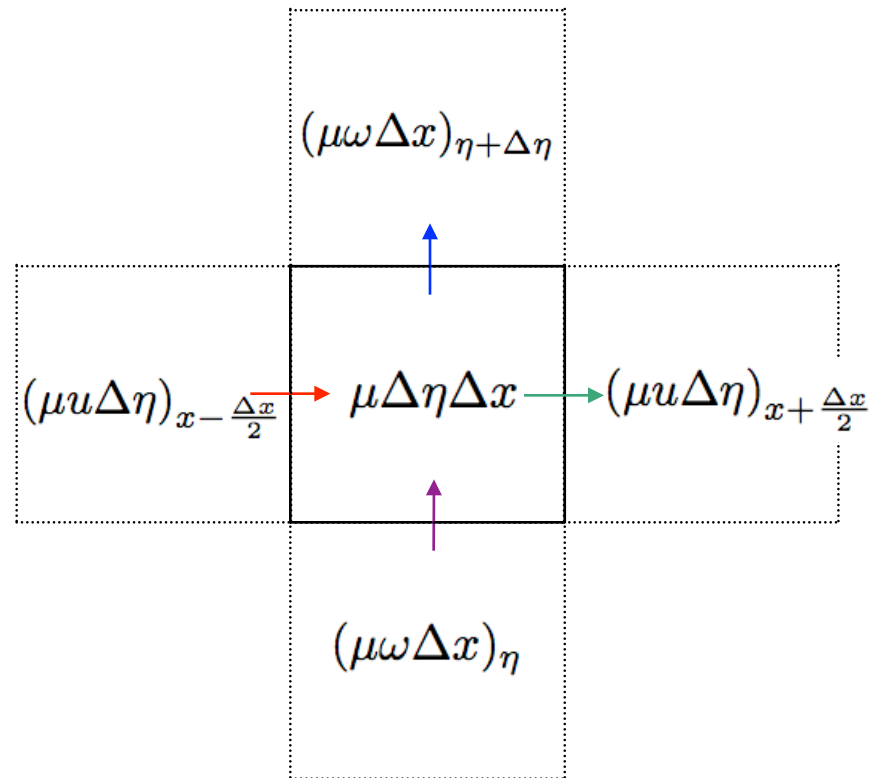
$$\Delta t^{-1}(\Delta x \Delta \eta) \cdot [(\mu)^{t+\Delta t} - (\mu)^t] = [(\mu u \Delta \eta)_{x-\Delta x/2,\eta} - (\mu u \Delta \eta)_{x+\Delta x/2,\eta}] + [(\mu \omega \Delta x)_{x,\eta-\Delta \eta/2} - (\mu \omega \Delta x)_{x,\eta+\Delta \eta/2}]$$

Vertical fluxes through the  
horizontal control-volume faces



# Mass Conservation in the ARW Model

The same mass fluxes are used for neighboring grid cells - hence mass is conserved locally and globally.





# Scalar Mass Conservation in the ARW Model

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Mass in a control volume	$(\Delta x \Delta \eta)(\mu)^t$
Scalar mass	$(\Delta x \Delta \eta)(\mu \phi)^t$

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Mass conservation equation:

$\Delta t^{-1}(\Delta x \Delta \eta) \cdot [(\mu)^{t+\Delta t} - (\mu)^t]$	$= [(\mu u \Delta \eta)_{x-\Delta x/2, \eta} - (\mu u \Delta \eta)_{x+\Delta x/2, \eta}] +$ $[(\mu \omega \Delta x)_{x, \eta-\Delta \eta/2} - (\mu \omega \Delta x)_{x, \eta+\Delta \eta/2}]$
↑ change in mass over a time step	mass fluxes through control volume faces

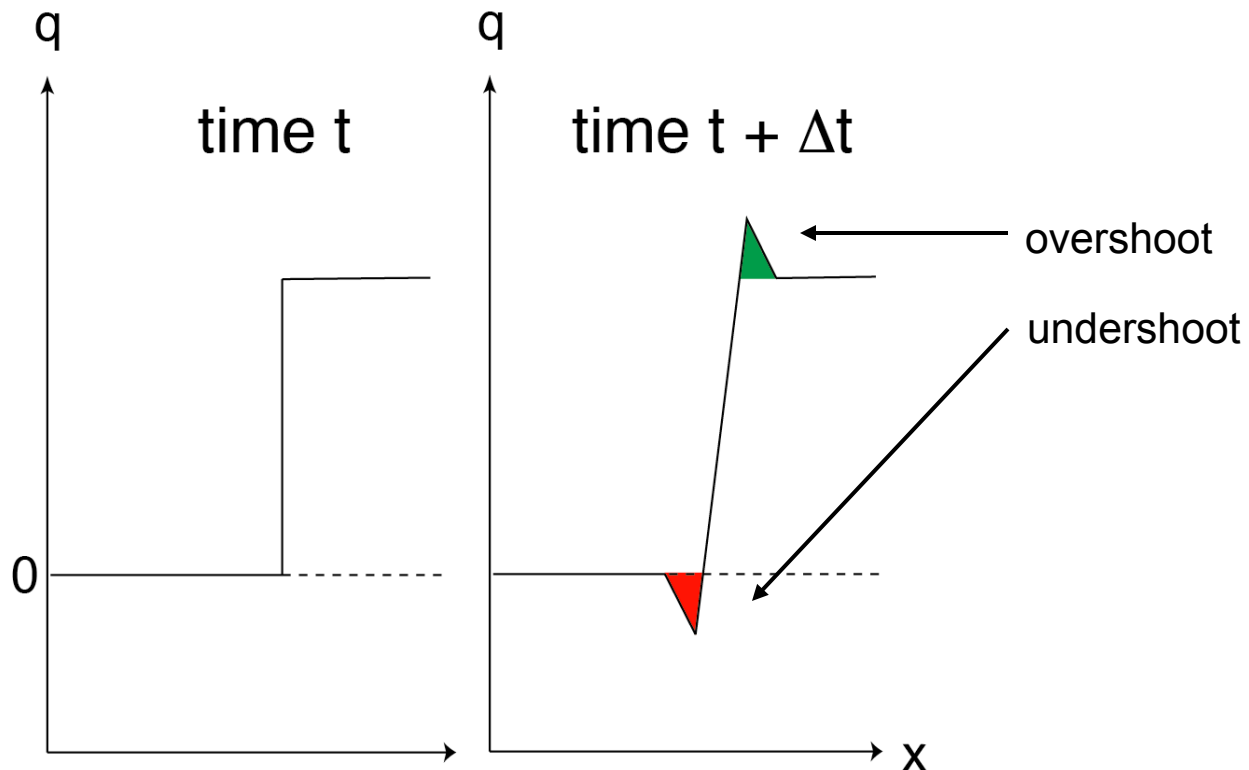
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Scalar mass conservation equation:

$\Delta t^{-1}(\Delta x \Delta \eta) \cdot [(\mu \phi)^{t+\Delta t} - (\mu \phi)^t]$	$= [(\mu u \phi \Delta \eta)_{x-\Delta x/2, \eta} - (\mu u \phi \Delta \eta)_{x+\Delta x/2, \eta}] +$ $[(\mu \omega \phi \Delta x)_{x, \eta-\Delta \eta/2} - (\mu \omega \phi \Delta x)_{x, \eta+\Delta \eta/2}]$
↑ change in tracer mass over a time step	tracer mass fluxes through control volume faces

# Moisture Transport in ARW

## 1D advection



ARW scheme is conservative,  
but not positive definite nor monotonic.  
Removal of negative  $q$  ■  
results in spurious source of  $q$  ■ .

# Positive-Definite/Monotonic Flux Renormalization

Scalar update, last RK3 step

$$(\mu\phi)^{t+\Delta t} = (\mu\phi)^t - \Delta t \sum_{i=1}^n \delta_{x_i} [f_i] \quad (1)$$

(1) Decompose flux:  $f_i = f_i^{upwind} + f_i^c$

(2) Renormalize high-order correction fluxes  $f_i^c$  such that solution is positive definite or monotonic:  $f_i^c = R(f_i^c)$

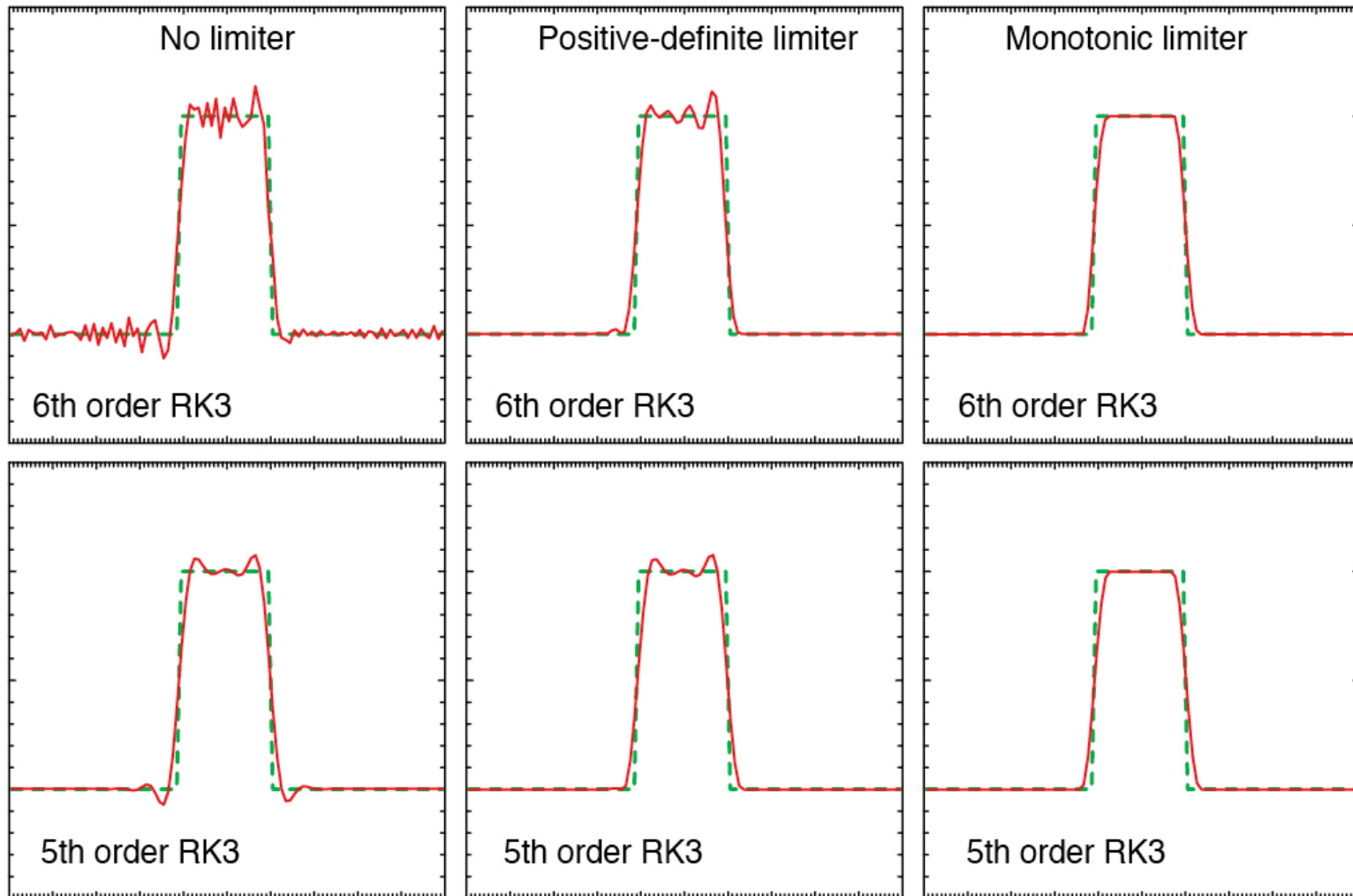
(3) Update scalar eqn. (1) using  $f_i = f_i^{upwind} + R(f_i^c)$

Skamarock, MWR 2006, 2241-2250

# PD/Monotonic Limiters in ARW - 1D Example

## Top-Hat Advection

1D Top-hat transport  $Cr = 0.5$ , 1 revolution, 200 steps



# ARW Model: Dynamics Parameters

## 3<sup>rd</sup> order Runge-Kutta time step

Courant number limited, 1D:  $C_r = \frac{U\Delta t}{\Delta x} < 1.73$

Generally stable using a timestep approximately twice as large as used in a leapfrog model.

## Acoustic time step

2D horizontal Courant number limited:  $C_r = \frac{C_s\Delta\tau}{\Delta x} < \frac{1}{\sqrt{2}}$

$\Delta\tau_{sound} = \Delta t_{RK} / (\text{number of acoustic steps})$

## Guidelines for time step

$\Delta t$  in seconds should be about  $6 * \Delta x$  (grid size in kilometers). Larger  $\Delta t$  can be used in smaller-scale dry situations, but *time\_step\_sound* (default = 4) should increase proportionately if larger  $\Delta t$  is used.

# Maximum Courant Number for Advection

$$C_a = U \Delta t / \Delta x$$

<i>Time Integration Scheme</i>	<i>Advection Scheme</i>				
	<i>2<sup>nd</sup></i>	<i>3<sup>rd</sup></i>	<i>4<sup>th</sup></i>	<i>5<sup>th</sup></i>	<i>6<sup>th</sup></i>
Leapfrog ( $\alpha=0.1$ )	0.91	U	0.66	U	0.57
RK2	U	0.90	U	0.39	U
RK3	1.73	1.63	1.26	1.43	1.09

U = unstable

(Wicker & Skamarock, 2002)

# ARW Filters: Divergence Damping

Purpose: filter acoustic modes

$$p^{*\tau} = p^\tau + \gamma_d (p^\tau - p^{\tau - \Delta\tau}) \quad \text{since} \quad p_t \sim c^2 \nabla \cdot \rho \mathbf{V}$$

$$\begin{aligned} \delta_\tau U'' + \mu^{t*} \alpha^{t*} \partial_x p''^\tau + (\mu^{t*} \partial_x \bar{p}) \alpha''^\tau \\ + (\alpha/\alpha_d) [\mu^{t*} \partial_x \phi''^\tau + (\partial_x \phi^{t*}) (\partial_\eta p'' - \mu'')^\tau] = R_U^{t*} \end{aligned}$$

$$\begin{aligned} \delta_\tau V'' + \mu^{t*} \alpha^{t*} \partial_y p''^\tau + (\mu^{t*} \partial_y \bar{p}) \alpha''^\tau + \\ + (\alpha/\alpha_d) [\mu^{t*} \partial_y \phi''^\tau + (\partial_y \phi^{t*}) (\partial_\eta p'' - \mu'')^\tau] = R_V^{t*} \end{aligned}$$

$\gamma_d = 0.1$  recommended (default)

# ARW Filters: External Mode Filter

Purpose: filter the external mode  
(primarily for real-data applications)

Additional terms:

$$\delta_{\tau} U'' = \dots - \gamma_e (\Delta x^2 / \Delta \tau) \delta_x (\delta_{\tau - \Delta \tau} \mu_d'')$$

$$\delta_{\tau} V'' = \dots - \gamma_e (\Delta y^2 / \Delta \tau) \delta_y (\delta_{\tau - \Delta \tau} \mu_d'')$$

$$\delta_{\tau} \mu_d = m^2 \int_1^0 [\partial_x U'' + \partial_y V'']^{\tau + \Delta \tau} d\eta$$

$\gamma_e = 0.01$  recommended (default)



# ARW Filters: Vertically Implicit Off-Centered Acoustic Step

Purpose: damp vertically-propagating acoustic modes

$$\delta_\tau W'' - m^{-1} g \overline{\left[ (\alpha/\alpha_d)^{t^*} \partial_\eta (C \partial_\eta \phi'') + \partial_\eta \left( \frac{c_s^2}{\alpha^{t^*}} \frac{\Theta''}{\Theta^{t^*}} \right) - \mu_d'' \right]^\tau} = R_W^{t^*}$$
$$\delta_\tau \phi'' + \frac{1}{\mu_d^{t^*}} [m \Omega^{\tau+\Delta\tau} \phi_\eta - \overline{g W''^\tau}] = R_\phi^{t^*}.$$

$$\bar{a}^\tau = \frac{1+\beta}{2} a^{\tau+\Delta\tau} + \frac{1-\beta}{2} a^\tau$$

$\beta = 0.1$  recommended (default)

# ARW Filters: Vertical Velocity Damping

Purpose: damp anomalously-large vertical velocities  
(usually associated with anomalous physics tendencies)

Additional term:

$$\partial_t W = \dots - \mu_d \operatorname{sign}(W) \gamma_w (Cr - Cr_\beta)$$

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$$Cr = \left| \frac{\Omega dt}{\mu d \eta} \right|$$

$Cr_\beta = 1.0$  typical value (default)

$\gamma_w = 0.3 \text{ m/s}^2$  recommended (default)

# ARW Filters: 2nd-Order Horizontal Mixing, Horizontal-Deformation-Based $K_h$

Purpose: mixing on horizontal coordinate surfaces  
(real-data applications)

$$K_h = C_s^2 l^2 \left[ 0.25(D_{11} - D_{22})^2 + \overline{D_{12}^2}^{xy} \right]^{\frac{1}{2}}$$

where  $l = (\Delta x \Delta y)^{1/2}$

$$D_{11} = 2 m^2 [\partial_x(m^{-1}u) - z_x \partial_z(m^{-1}u)]$$

$$D_{22} = 2 m^2 [\partial_y(m^{-1}v) - z_y \partial_z(m^{-1}v)]$$

$$D_{12} = m^2 [\partial_y(m^{-1}u) - z_y \partial_z(m^{-1}u) \\ + \partial_x(m^{-1}v) - z_x \partial_z(m^{-1}v)]$$

$C_s = 0.25$  (Smagorinsky coefficient, default value)

# ARW Filters: Upper Level Gravity-Wave Absorbers

(1) Implicit Rayleigh W - damping (nonhydrostatic equations only!)

$$\tilde{W}''^{\tau+\Delta\tau} - W''^{\tau} - g\Delta\tau \frac{\alpha}{\alpha_d} \frac{\partial}{\partial\eta} \left( C \frac{\partial \overline{\phi}''^{\tau}}{\partial\eta} \right) = \Delta\tau R_W^* \quad (1)$$

$$W''^{\tau+\Delta\tau} - \tilde{W}''^{\tau+\Delta\tau} = -\tau(z)\Delta\tau W''^{\tau+\Delta\tau} \quad (2)$$

$$\phi''^{\tau+\Delta\tau} - \phi''^{\tau} - g\Delta\tau \frac{1}{\mu} \overline{W}''^{\tau} = \Delta\tau R_{\phi}^* \quad (3)$$

**Vertically implicit solution procedure:**

Eliminate  $\phi''^{\tau+\Delta\tau}$  from (1) using (3), solve for  $\tilde{W}''^{\tau+\Delta\tau}$ .

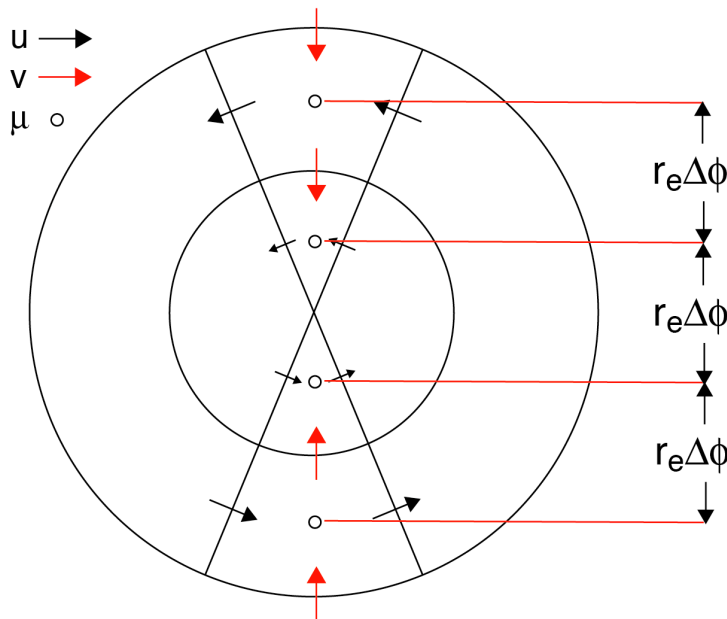
Apply implicit Rayleigh damping - solve for  $W''^{\tau+\Delta\tau}$  using (2).

Recover the geopotential using  $W''^{\tau+\Delta\tau}$  in (3).

$$\tau(z) = \begin{cases} \gamma_r \sin^2 \left[ \frac{\pi}{2} \left( 1 - \frac{z_{top}-z}{z_d} \right) \right] & \text{for } z \geq (z_{top} - z_d); \\ 0 & \text{otherwise,} \end{cases} \quad \begin{array}{l} \tau(z) - \text{damping rate (t}^{-1}\text{)} \\ z_d - \text{depth of the damping layer} \\ \gamma_r - \text{dimensionless damping coefficient} \end{array}$$

# Global ARW - Latitude-Longitude Grid

- Map factors -  $m_x$  and  $m_y$ 
  - Computational grid poles need not be geographic poles.
  - Limited area and nesting capable.
- Polar boundary conditions
- Polar filtering

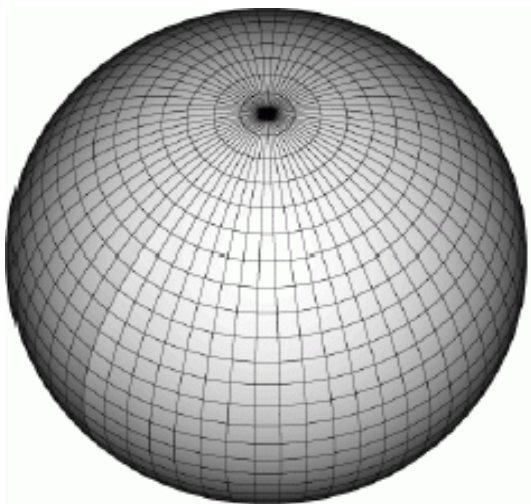


Zero meridional flux at the poles (cell-face area is zero).

$v$  (poles) only needed for meridional derivative of  $v$  near the poles (we interpolate).

All other meridional derivatives are well-defined near/at poles.

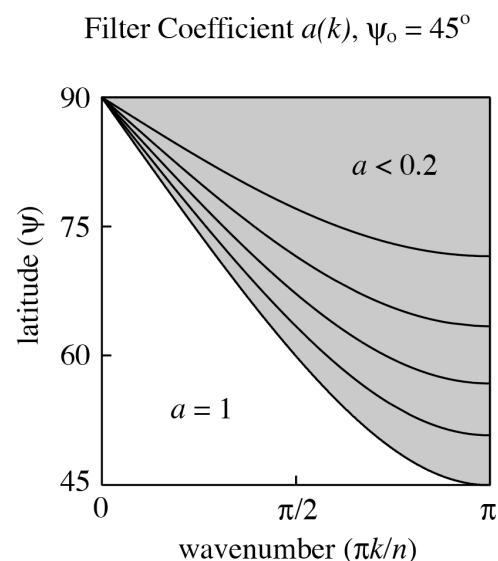
# ARW Filters: Polar Filter



Converging gridlines severely limit timestep.  
The polar filter removes this limitation.

Filter procedure - Along a grid latitude circle:

1. Fourier transform variable.
2. Filter Fourier coefficients.
3. Transform back to physical space.



$$\hat{\phi}(k)_{filtered} = a(k) \hat{\phi}(k), \quad \text{for all } k$$

$$a(k) = \min \left[ 1., \max \left( 0., \left( \frac{\cos \psi}{\cos \psi_o} \right)^2 \frac{1}{\sin^2(\pi k/n)} \right) \right]$$

$k$  = dimensionless wavenumber

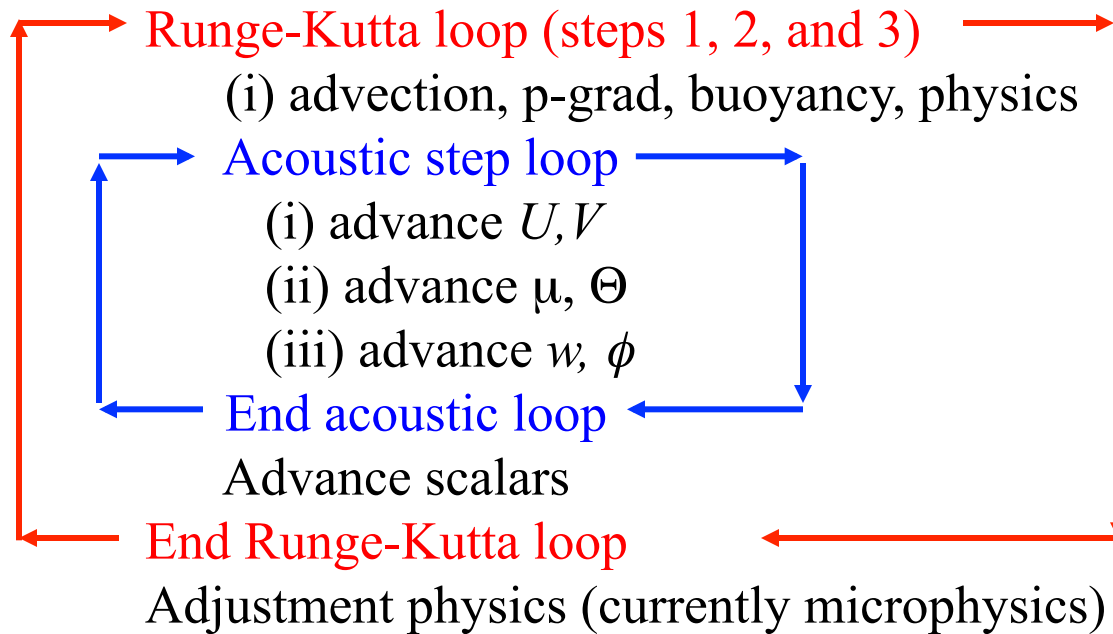
$\hat{\phi}(k)$  = Fourier coefficients from forward transform

$a(k)$  = filter coefficients

$\psi$  = latitude  $\psi_o$  = polar filter latitude, filter when  $|\psi| > \psi_o$

# WRF ARW Model Integration Procedure

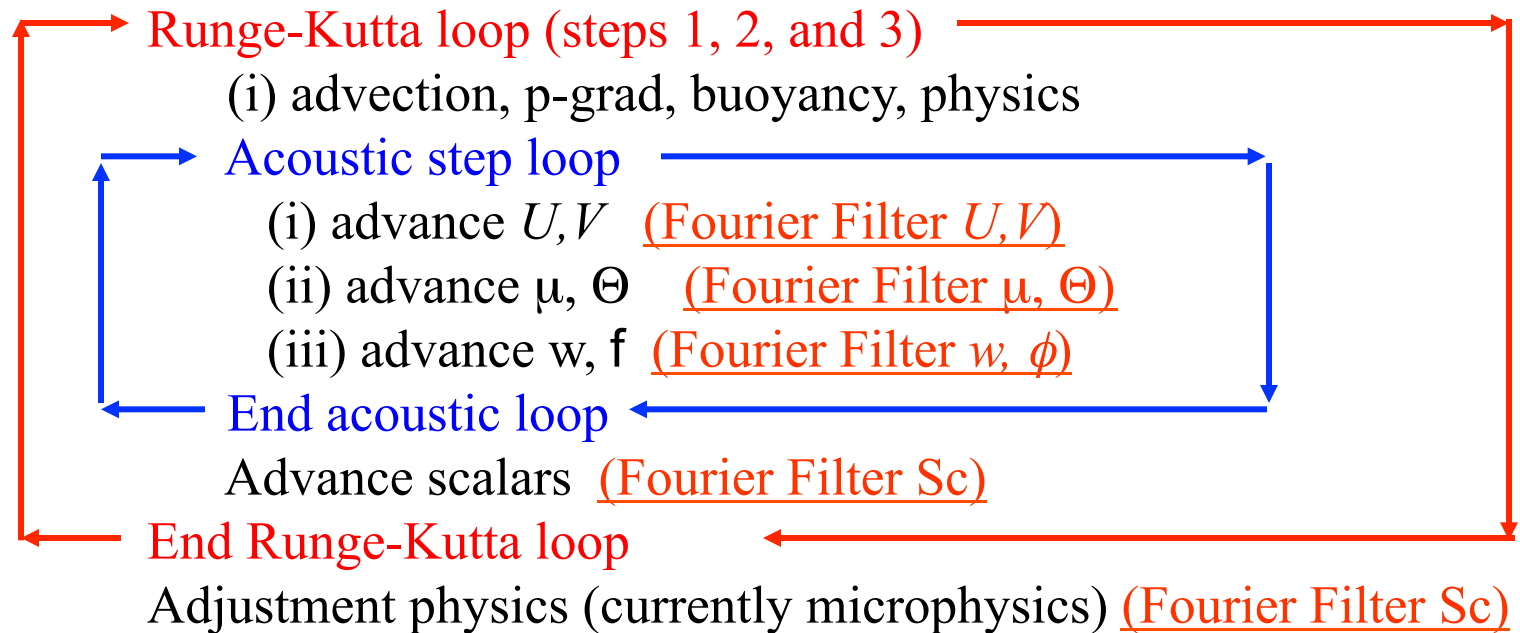
Begin time step



End time step

# WRF ARW Model Integration Procedure

Begin time step



End time step

Timestep limited by minimum  $\Delta x$  outside of polar-filter region.



# ARW Model: Coordinate Options

1. Cartesian geometry:  
idealized cases
2. Lambert Conformal:  
mid-latitude applications
3. Polar Stereographic:  
high-latitude applications
4. Mercator:  
low-latitude applications
5. Latitude-Longitude  
global  
regional

Projections 1-4 are isotropic ( $m_x = m_y$ )

Latitude-Longitude projection is anisotropic ( $m_x \neq m_y$ )

# ARW Model: Boundary Condition Options

## Lateral boundary conditions

1. Specified (Coarse grid, real-data applications).
2. Open lateral boundaries (gravity-wave radiative).
3. Symmetric lateral boundary condition (free-slip wall).
4. Periodic lateral boundary conditions.
5. Nested boundary conditions (specified).

## Top boundary conditions

1. Constant pressure.

## Bottom boundary conditions

1. Free slip.
2. Various B.L. implementations of surface drag, fluxes.

# ARW Model: Nesting

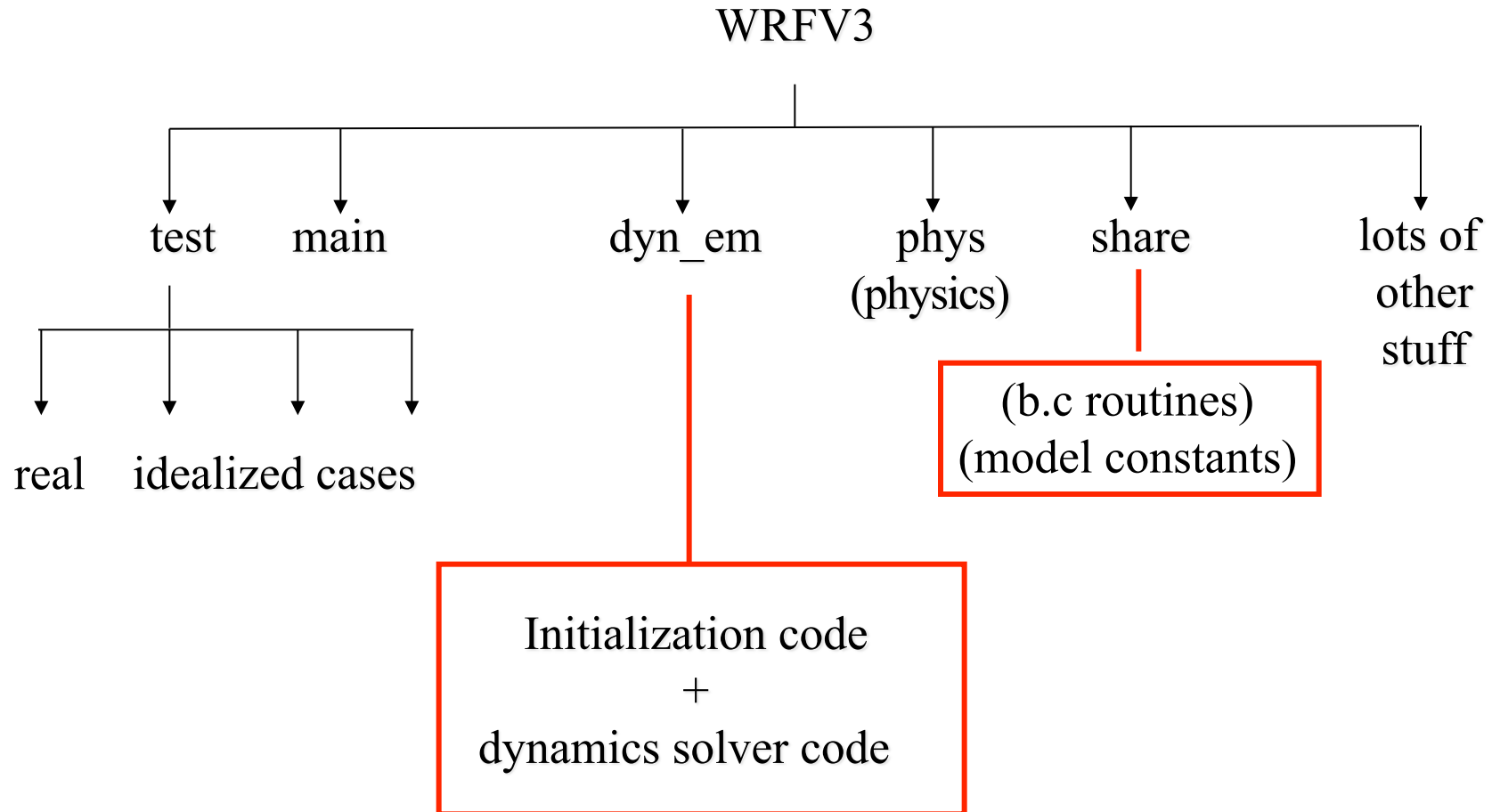
## 2-way nesting

1. Multiple domains run concurrently
2. Multiple levels, multiple nests per level
3. Any integer ratio grid size and time step
4. Parent domain provides nest boundaries
5. Nest feeds back interior values to parent

## 1-way nesting

1. Parent domain is run first
2. *ndown* uses coarse output to generate nest boundary conditions
3. Nest initial conditions from fine-grid input file
4. Nest is run after *ndown*

# WRF ARW code



## WRF ARW Tech Note

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