Fundamentals in Atmospheric Modeling

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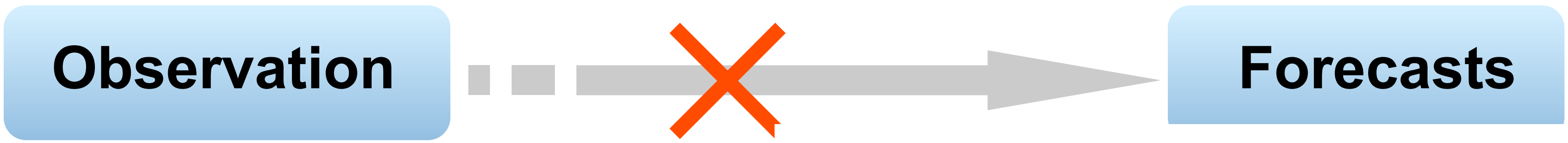
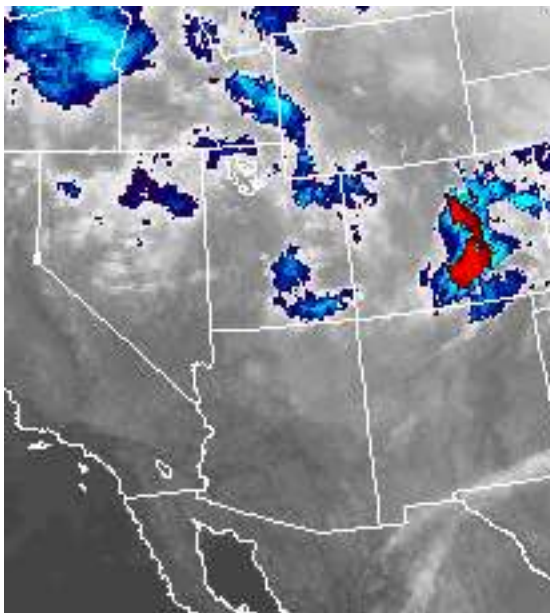
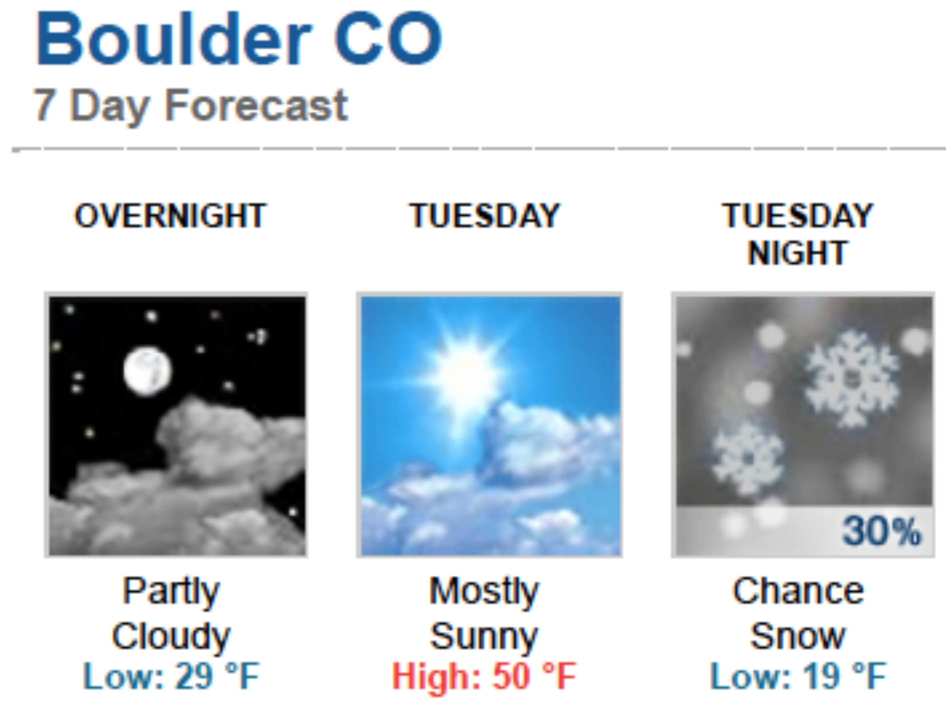
List of presentations

- Concept of modeling
- Structure of models
- Predictability

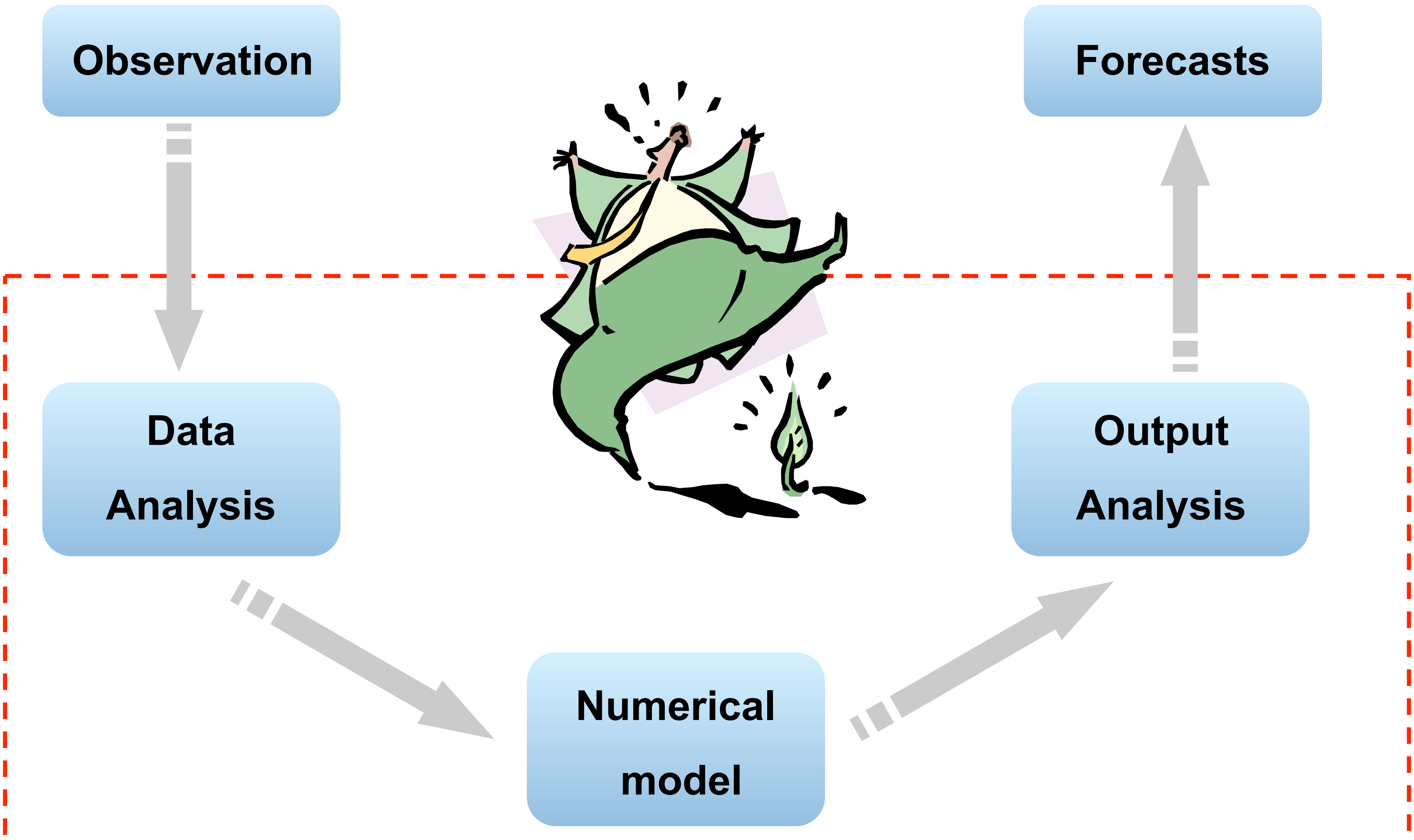
How were the today's forecasts made ?







Numerical model is a crucial component



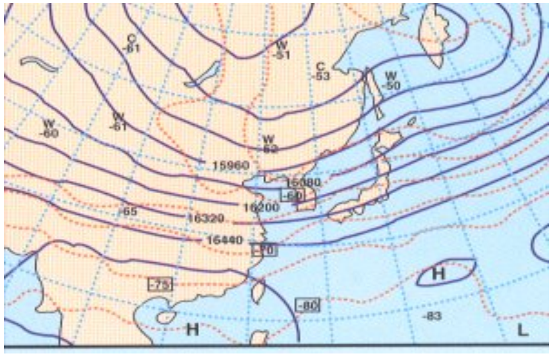
Then, how ?

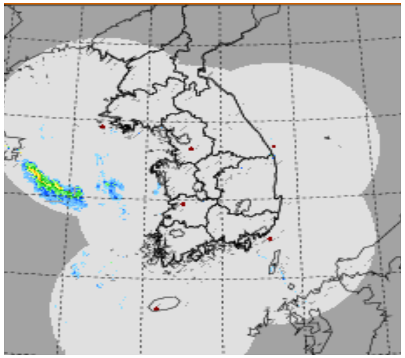
Step1:

Observation

Step2:

Data analysis





Theory of NWP

Thermodynamics

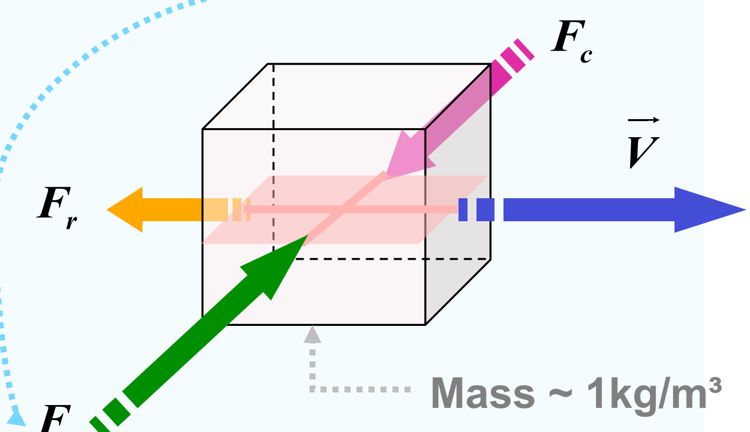
Heat = Energy + Work

$$\Delta H = c_p \Delta T - \alpha \Delta p$$
$$= c_v \Delta T + p \Delta \alpha$$

Dynamics

Force = Mass × Acceleration

- Mass ≒ 1 kg/m³
- Force: **PGF, CO, Friction**…



Theory of NWP : Atmosphere is conserved

- **Momentum** $$F = ma$$
- **Mass** $$\frac{1}{M}\frac{dM}{dt} = 0$$
- **Moisture** $$\frac{dq}{dt} = E - C$$
- **Ideal gas** $$p\alpha = RT$$
- **Energy** $$Q = C_v \frac{dT}{dt} + p\frac{d\alpha}{dt}$$

CONSERVATION

The governing equations

V. Bjerknes** (1904) pointed out for the first time that there is a complete set of* ***7 equations with 7 unknowns *that governs the evolution of the atmosphere:*

$$\frac{d\mathbf{v}}{dt} = -\alpha\nabla p - \nabla\phi + \mathbf{F} - 2\Omega\times\mathbf{v} \qquad \text{(1-3)}$$

$$\frac{\partial\rho}{\partial t} = -\nabla.(\rho\mathbf{v}) \qquad \text{(4)}$$

$$p = \rho RT \qquad \text{(5)}$$

$$\frac{ds}{dt} = C_p \frac{1}{\theta}\frac{d\theta}{dt} = \frac{Q}{T} \qquad \text{(6)}$$

$$\frac{dq}{dt} = E - C \qquad \text{(7)}$$

7 equations, 7 unknown (u,v,w,T, p, den and q)

solvable

History of numerical weather forecasts

1904 : Norwegian **V.** Bjerknes (1862-1951) :
Setup the governing equations

1922 : British **L. F.** Richardson (1881-1953) :
Integrate model ➔ failed

1939 : Swedish **C.-G. Rossby** :

1948, 1949, J. G. Charney (1917-1981)

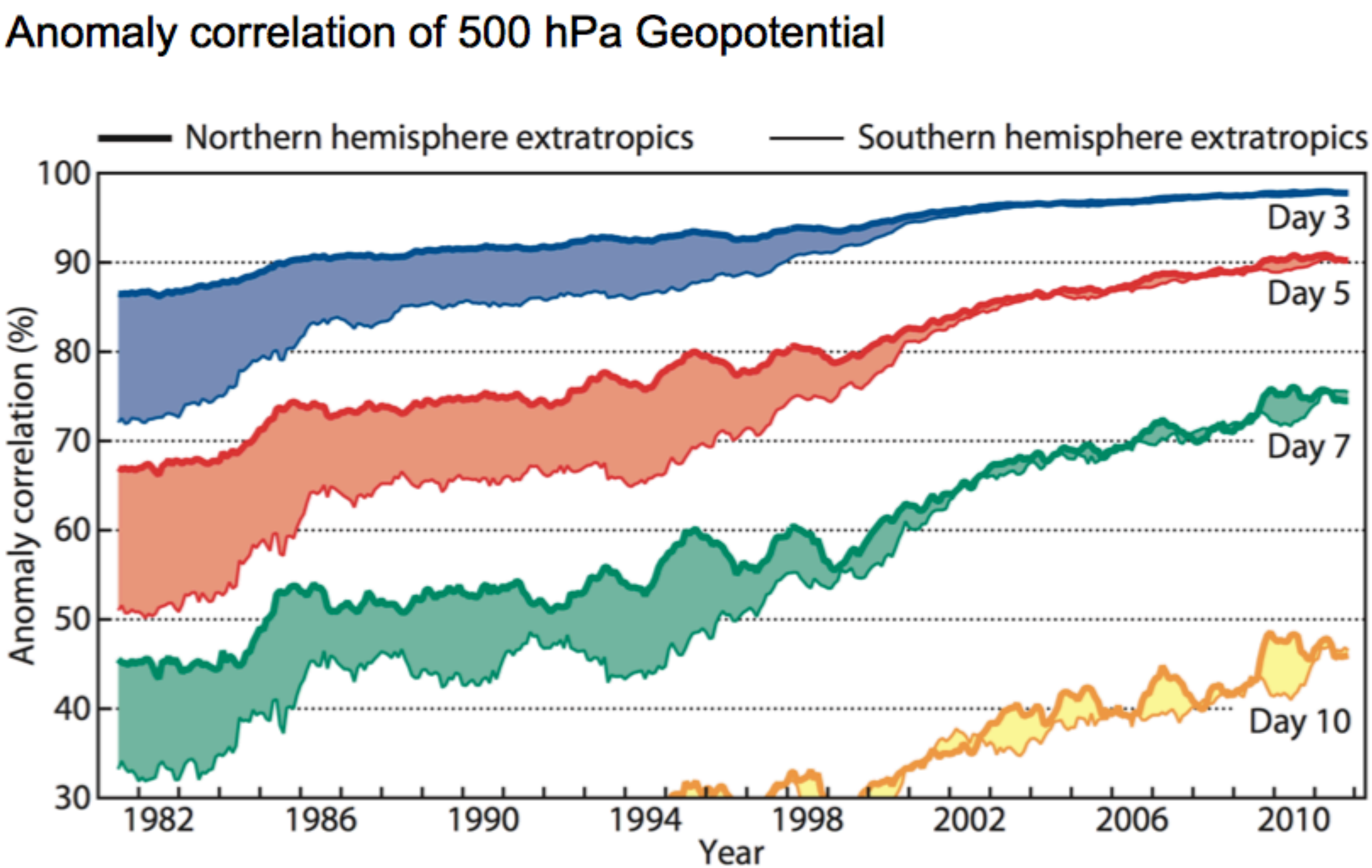
1950 : **Princeton Group**
(Charney, Fjortoft,
von Newman)
ENIAC
(Electrical Numerical
Integrator and Computer)
→ **first success**

Computer Age (1946~)

- von Neumann and Charney
  - Applied ENIAC to weather prediction

- Carl-Gustaf Rossby
  - The Swedish Institute of Meteorology
  - First routine real-time numerical weather forecasting. (1954)
    (US in 1958, Japan in 1959)

History of NWP skill : ECMWF



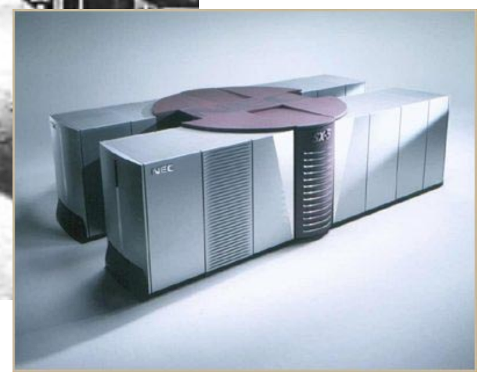
1day / 10 yrs

Factors for the improvement (Kalnay 2002)

- Supercomputers
- Physical processes
- Initial conditions

Super-computer for weather models

ENIAC,1946



NEC SX-5

Cray T3E

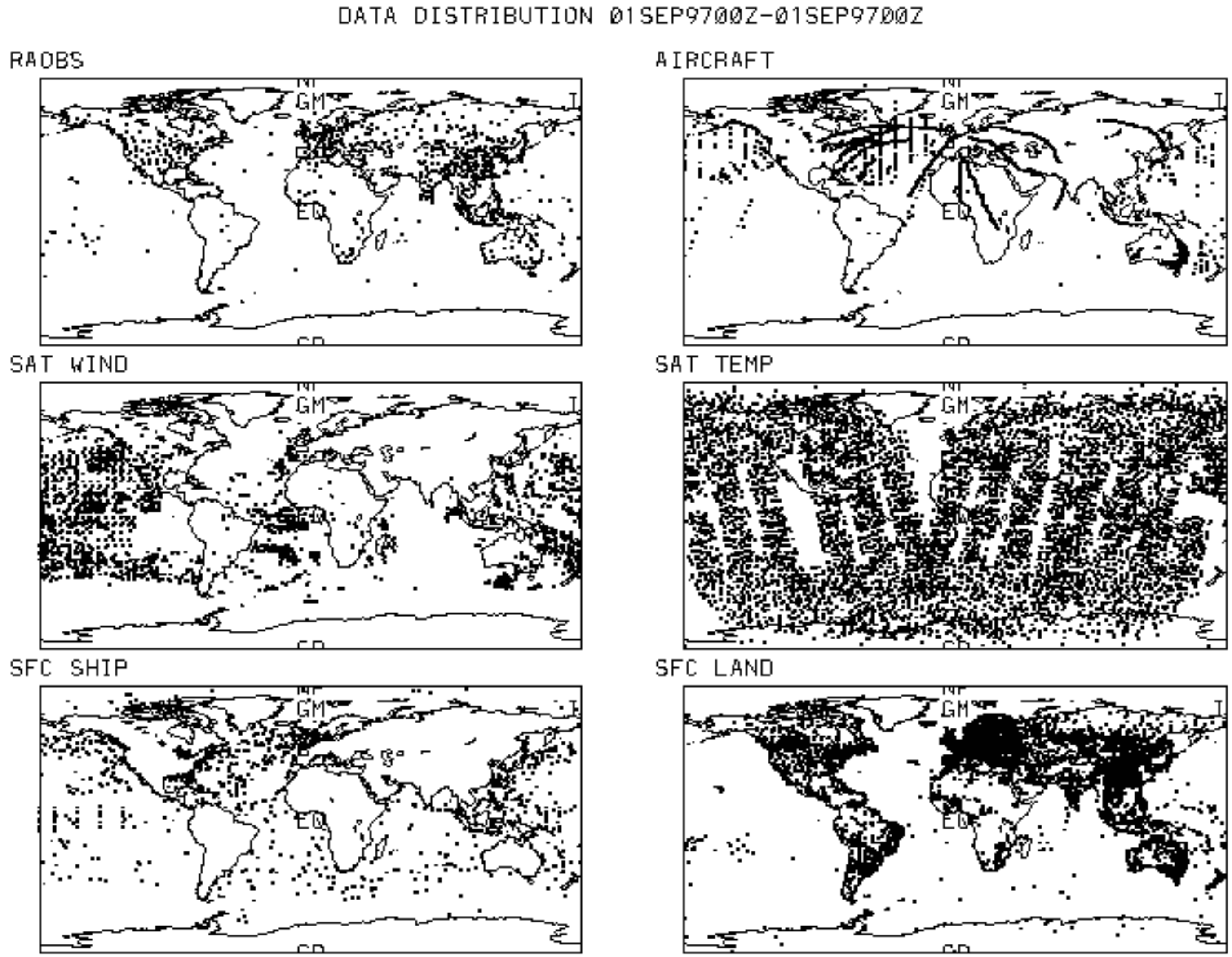
Cray SV1

Fujitsu VPP700E

Cray T90

Initial condition
(data assimilation)

Various observations



heterogeneous in space and time….

Data Assimilation

- Model 1°X 1° resolution, 20 levels

  u, v, T, q, Ps, Tg

$$360\times180\times20 = 1.3\times10^6 \times 4\ \mathrm{var}\,iabl = 5\times10^6$$

- observation : $10^4 \sim 10^5$ non-uniform distribution

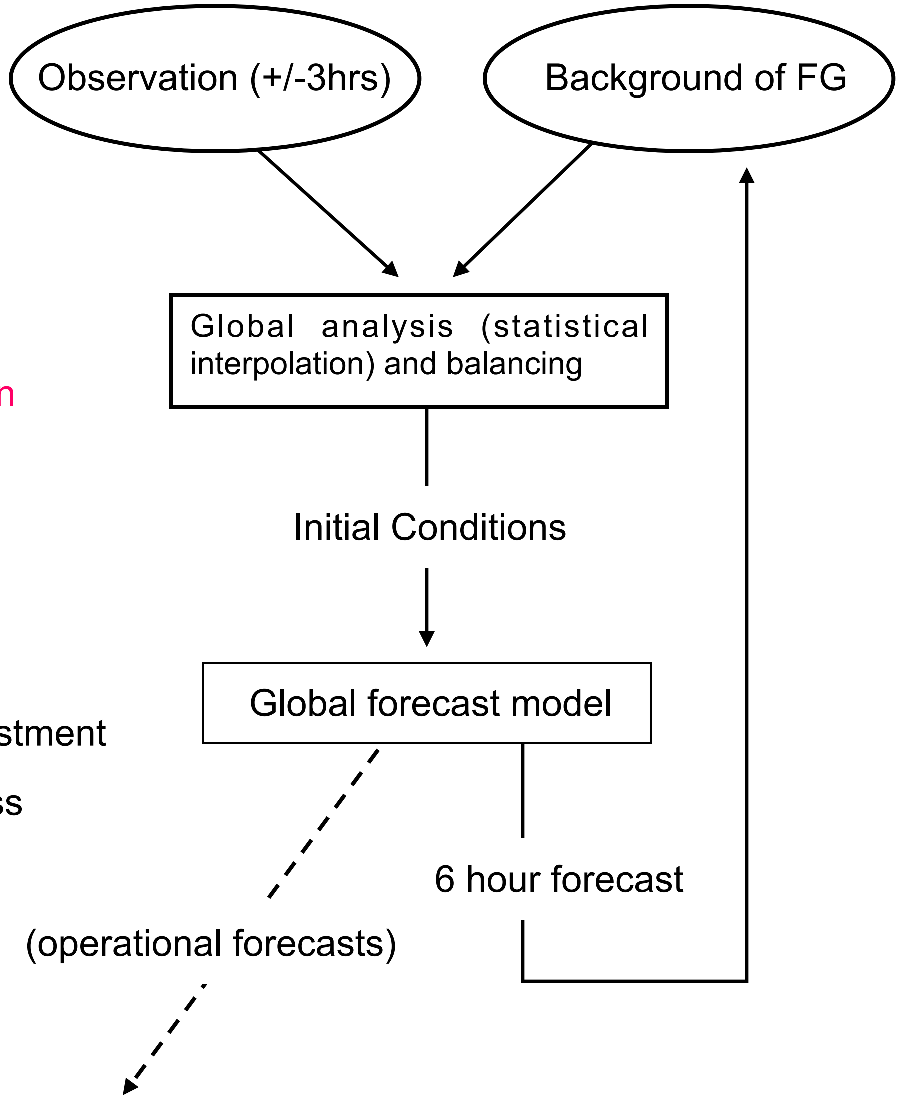
  ± 3 *hour window*

- Data assimilation cycle 1) data checking

  2) objective analysis

  3) Initialization: dynamical adjustment

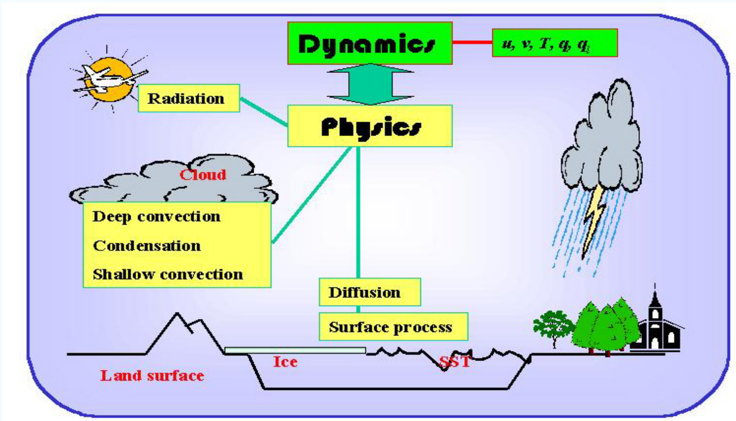
  4) short-range fcst for first guess

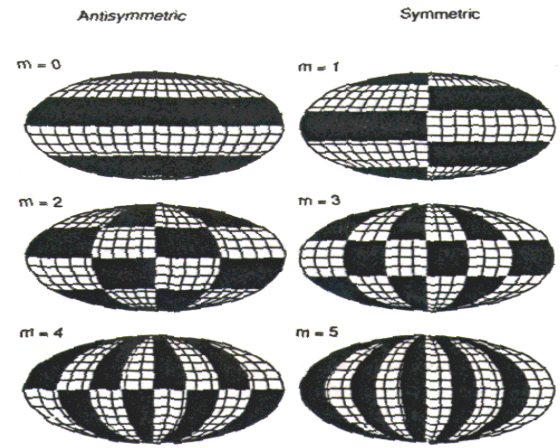


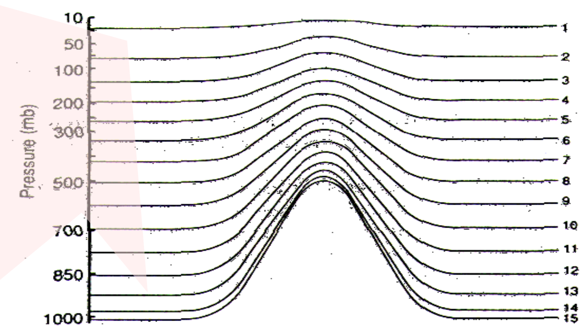
Model

- Dynamics : Identity (Speed)
- Physics : Components (Predictability)

Step3: Integration







15-level sigma-coordinate model

$$\frac{\partial u}{\partial t} = -u\frac{\partial u}{\partial x} - v\frac{\partial u}{\partial y} - \omega\frac{\partial u}{\partial p} - \frac{\partial \Phi}{\partial x} + fv + F_x$$

$$\frac{\partial v}{\partial t} = -u\frac{\partial v}{\partial x} - v\frac{\partial v}{\partial y} - \omega\frac{\partial v}{\partial p} - \frac{\partial \Phi}{\partial y} - fu + F_y$$

$$\frac{\partial \Phi}{\partial t} = -\frac{RT}{p}$$

$$\frac{\partial T}{\partial t} = -u\frac{\partial T}{\partial x} - v\frac{\partial T}{\partial y} + \omega\left(\frac{\kappa T}{p} - \frac{\partial T}{\partial p}\right) + \frac{\dot{H}}{c_p}$$

$$\frac{\partial \omega}{\partial p} = -\left(\frac{\partial u}{\partial x} + \frac{\partial v}{\partial y}\right)$$

Dynamics : Numerical method (spatial)

Finite difference method (FDM) :

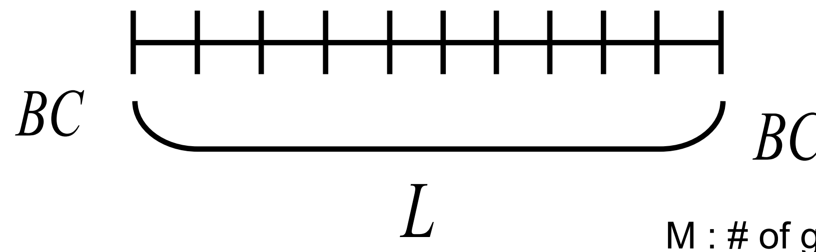
Spectral method (SPM) :

Finite element method (FEM) :

Ex) $\frac{\partial \phi}{\partial t} = -c\frac{\partial \phi}{\partial x}$; advection eq.

1) FDM (Finite difference)

$$\frac{\Delta \phi}{\Delta t} = \frac{\phi_2 - \phi_1}{t_2 - t_1}$$



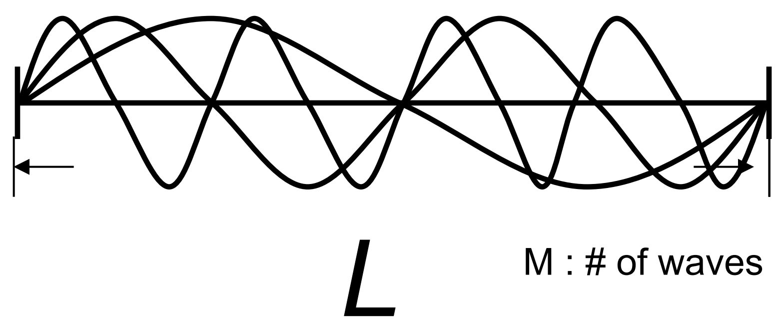
M : # of grids

2) Spectral Method

- Determine basis function to get $H(\phi(x))$
- Expand ϕ in terms of a time series

$e_m(x)$ (basis funct), $m = m_1 \ldots m_n$. → infinite

$$\Rightarrow \phi(x,t) = \sum_{m=m_1}^{m_2} \phi_m(t) e_m(x)$$



M : # of waves

* Resolution Increases $\begin{cases} \Delta x \to decreases \\ m \to increases \end{cases}$

Dynamics : Numerical method (temporal)

a) $\dfrac{u^{n+1}-u^{n-1}}{2\Delta t}=F\left(u^{n}\right)$: leap-frog **good for hyperbolic**

unstable for parabolic

b) $\dfrac{u^{n+1}-u^{n}}{\Delta t}=F\left(u^{n}\right)$: Euler-forward **good for diffusion**

unstable for hyperbolic

c) $\dfrac{u^{n+1}-u^{n}}{\Delta t}=F\left(\dfrac{u^{n}+u^{n+1}}{2}\right)$: **Crank-Nicholson**

d) $\dfrac{u^{n+1}-u^{n}}{\Delta t}=F\left(u^{n+1}\right)$: **Fully implicit, backward**

e) $\dfrac{u^{*}-u^{n}}{\Delta t}=F\left(u^{n}\right)$: $\dfrac{u^{n+1}-u^{n}}{\Delta t}=F\left(u^{*}\right)$: **Euler-backward (Matzuno)**

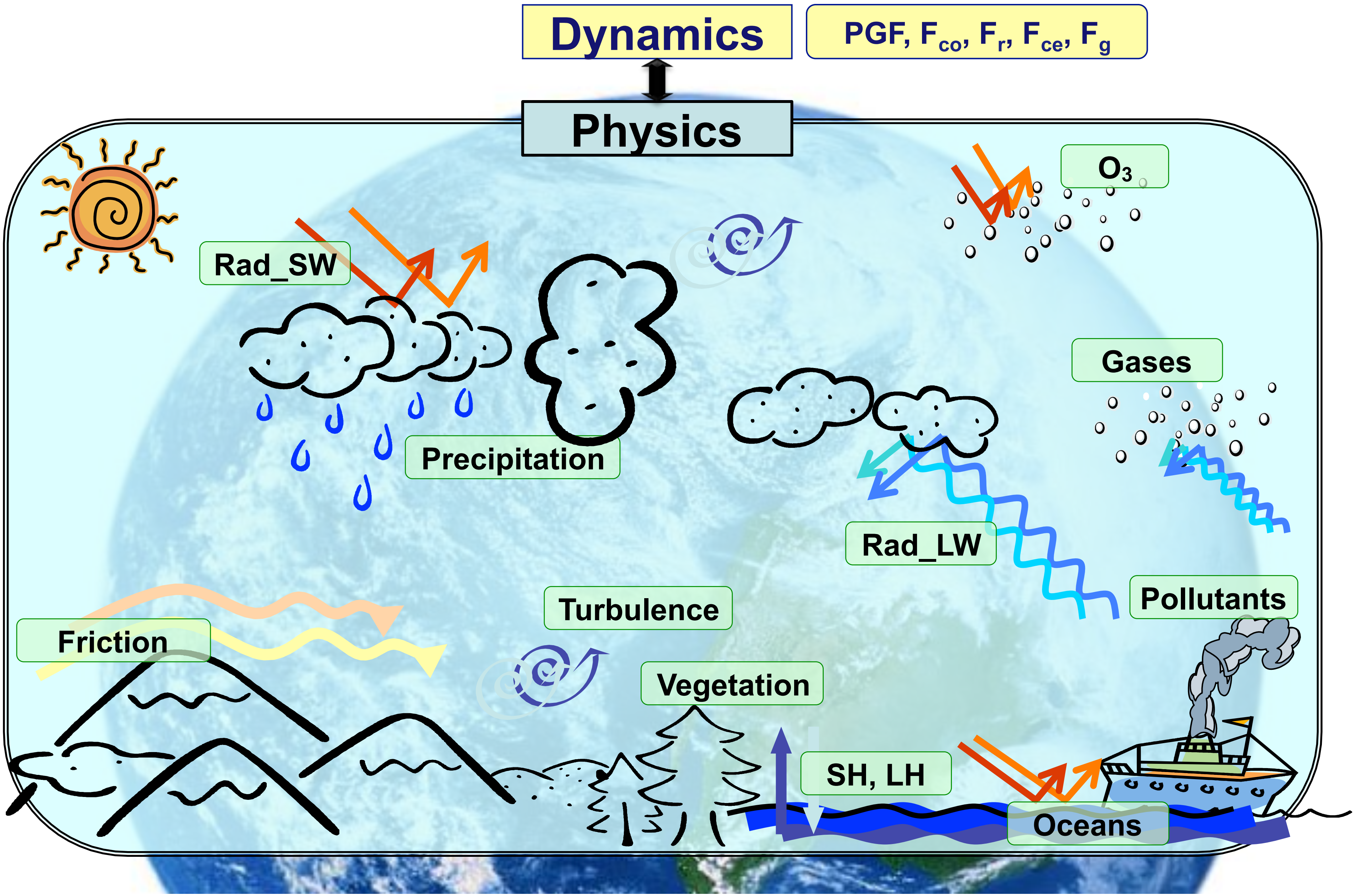
f) $\dfrac{u^{n+\frac{1}{2}*}-u^{n}}{\Delta t/2}=F\left(u^{n}\right)$: $\dfrac{u^{n+\frac{1}{2}**}-u^{n}}{\Delta t/2}=F\left(u^{n+\frac{1}{2}*}\right)$

$$\frac{u^{n+1}-u^{n}}{\Delta t}=\frac{1}{6}\left[F\left(u^{n}\right)+2F\left(u^{n+\frac{1}{2}*}\right)+2F\left(u^{n+\frac{1}{2}**}\right)+F\left(u^{n+1*}\right)\right]$$: **RK(Runge-Kuta)-4th order**

g) $\dfrac{u^{n+1}-u^{n-1}}{2\Delta t}=F_{1}\left(u^{n}\right)+F_{2}\left(\dfrac{u^{n+1}-u^{n-1}}{2}\right)$: **Semi-Implicit**

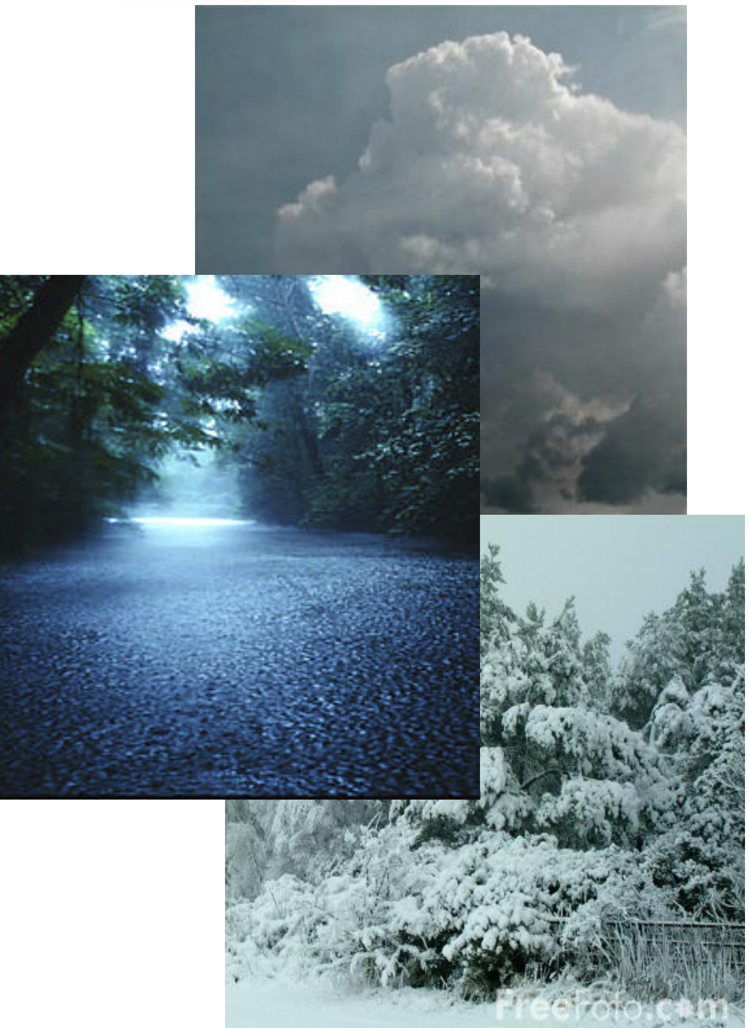
h) $\dfrac{u^{*}-u^{n}}{\Delta t}=F_{1}\left(u^{n}\right);\quad \dfrac{u^{n+1}-u^{*}}{\Delta t}=F_{2}\left(u^{*}\right)$: **Fractional steps**

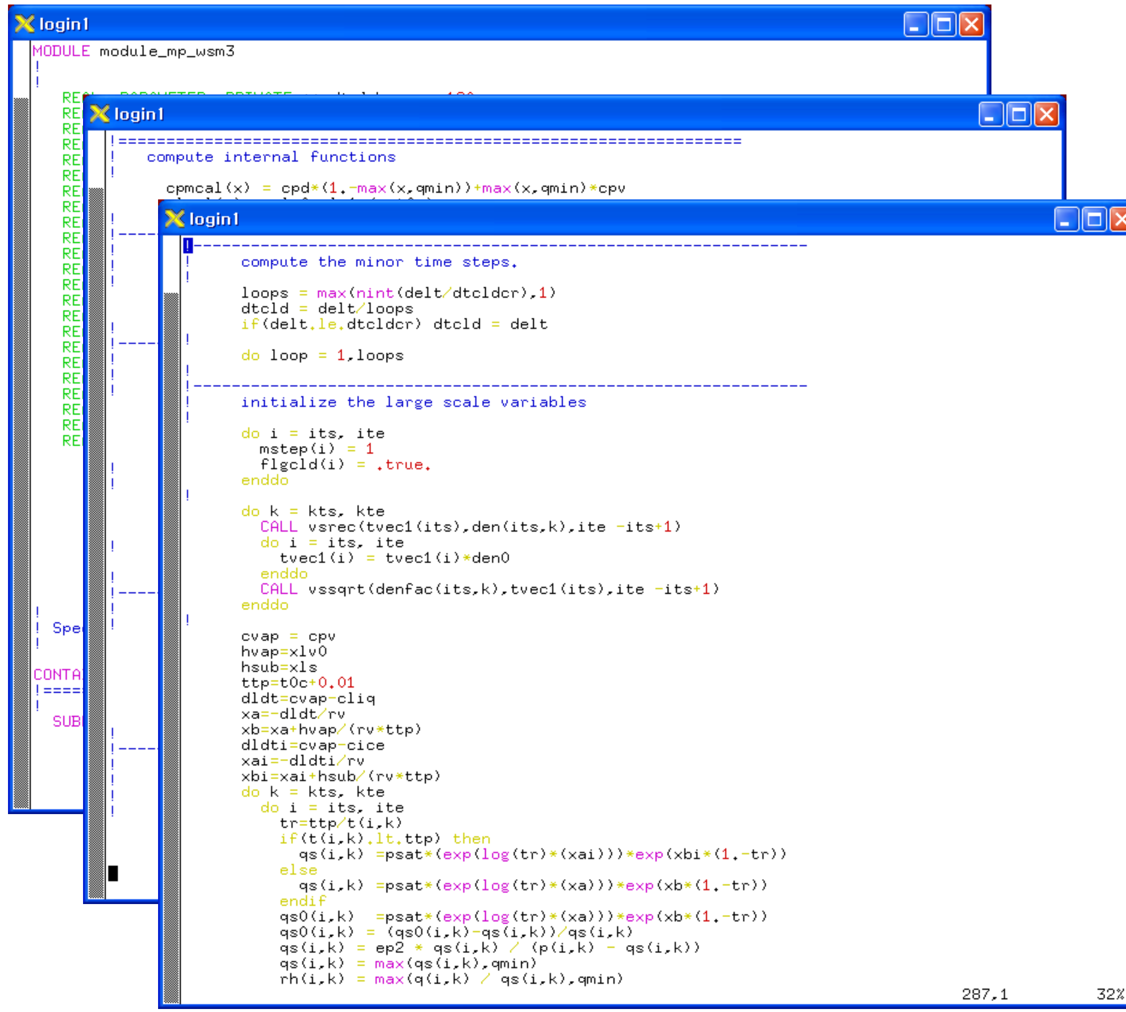
Physics modules : Branches of atmospheric sciences



Physics module (example): Cloud and precipitation

Real atmosphere → Theory / Formulation → Model (computer program)





Physics module (example): Cloud and precipitation

login1

```
!
      do k = kts, kte
        do i = its, ite
          supsat = max(q(i,k),qmin)-qs(i,k)
          satdt = supsat/dtcld
          if(t(i,k).ge.t0c) then
!
!===============================================================
!
! warm rain processes
!
! - follows the processes in RH83 and LFO except for autoconcersion
!
!===============================================================
!---------------------------------------------------------------
! paut1: auto conversion rate from cloud to rain [HDC 16]
!        (C->R)
!---------------------------------------------------------------
            if(qci(i,k).gt.qc0) then
              paut(i,k) = qck1*exp(log(qci(i,k))*((7./3.)))
              paut(i,k) = min(paut(i,k),qci(i,k)/dtcld)
            endif
!---------------------------------------------------------------
! pracw: accretion of cloud water by rain [D89 B15]
!        (C->R)
!---------------------------------------------------------------
            if(qrs(i,k).gt.qcrmin.and.qci(i,k).gt.qmin) then
                pacr(i,k) = min(pacrr*rslope3(i,k)*rslopeb(i,k)           &
                     *qci(i,k)*denfac(i,k),qci(i,k)/dtcld)
            endif
!---------------------------------------------------------------
! pres1: evaporation/condensation rate of rain [HDC 14]
!        (V->R or R->V)
!---------------------------------------------------------------
            if(qrs(i,k).gt.0.) then
                coeres = rslope2(i,k)*sqrt(rslope(i,k)*rslopeb(i,k))
                pres(i,k) = (rh(i,k)-1.)*(precr1*rslope2(i,k)             &
                         +precr2*work2(i,k)*coeres)/work1(i,k)
              if(pres(i,k).lt.0.) then
                pres(i,k) = max(pres(i,k),-qrs(i,k)/dtcld)
                pres(i,k) = max(pres(i,k),satdt/2)
              else
                pres(i,k) = min(pres(i,k),satdt/2
              endif
            endif
          else
!
!===============================================
```

T>0°C

$$Paut1 = \min\left(\frac{0.104 g E_C \rho^{\frac{4}{3}}}{\mu (N_c \rho_w)^{\frac{1}{3}}} q_c^{\frac{7}{3}},\ \frac{q_c}{dt} \right)$$

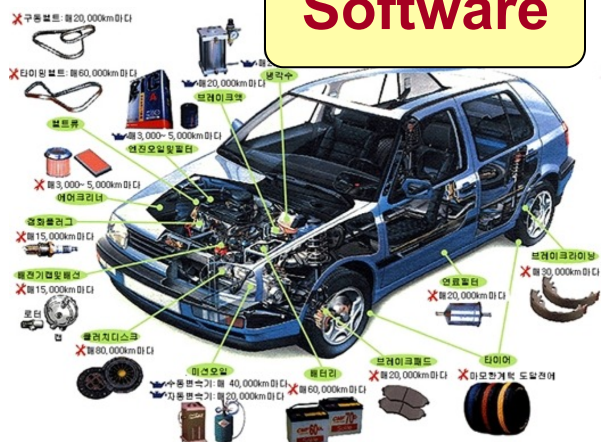
$$P_{racw} = \frac{\pi a_r E_{CR} N_{0r} q_c}{4} \left(\frac{\rho_0}{\rho} \right)^{\frac{1}{2}} \frac{\Gamma(3+b_r)}{\lambda_r^{3+b_r}}$$

$$Pres1 = \frac{2\pi N_{0r}(S_w - 1)}{(A_w + B_w)} \left[\frac{0.78}{\lambda_r^2} + \frac{a_r^{\frac{1}{2}} 0.31 \Gamma(b_r/2 + 5/2)}{\lambda_r^{b_r/2+5/2}} \left(\frac{\mu}{D} \right)^{\frac{1}{3}} \left(\frac{1}{\mu} \right)^{\frac{1}{2}} \left(\frac{\rho_0}{\rho} \right)^{\frac{1}{4}} \right]$$

Car and model

Dynamics

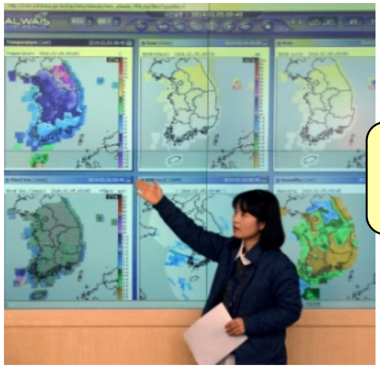
Software



Physics

Data assimilation

Driver



Forecaster

Classification of models

- **Dynamic frame**

Hydrostatic	Non-hydrostatic
Large-scale	Small-scale (heavy rainfall, complex mountain)

- **Scale**

Global	Regional
10 km – 100 km	1 km-10 km

- **Purpose**

Initial data-> FORECAST	Forcing → RESPONSE
NWP : upto 2 weeks	**GCM** (General circulation model)

Predictability

Chaos theory (Lorenz)

Charney (1951) : Uncertainties in initial condition and model

↓

Lorenz (1962,1963) : Unstable nature of atmosphere

↓

Purpose : NWP is better than statistical forecast

Tool : 4 K memory computer

Model : 12 variables (heating and dissipation forcing)

Results : differences -> non-periodicity

↓

Initial condition (3 decimal point) : different after 2 month

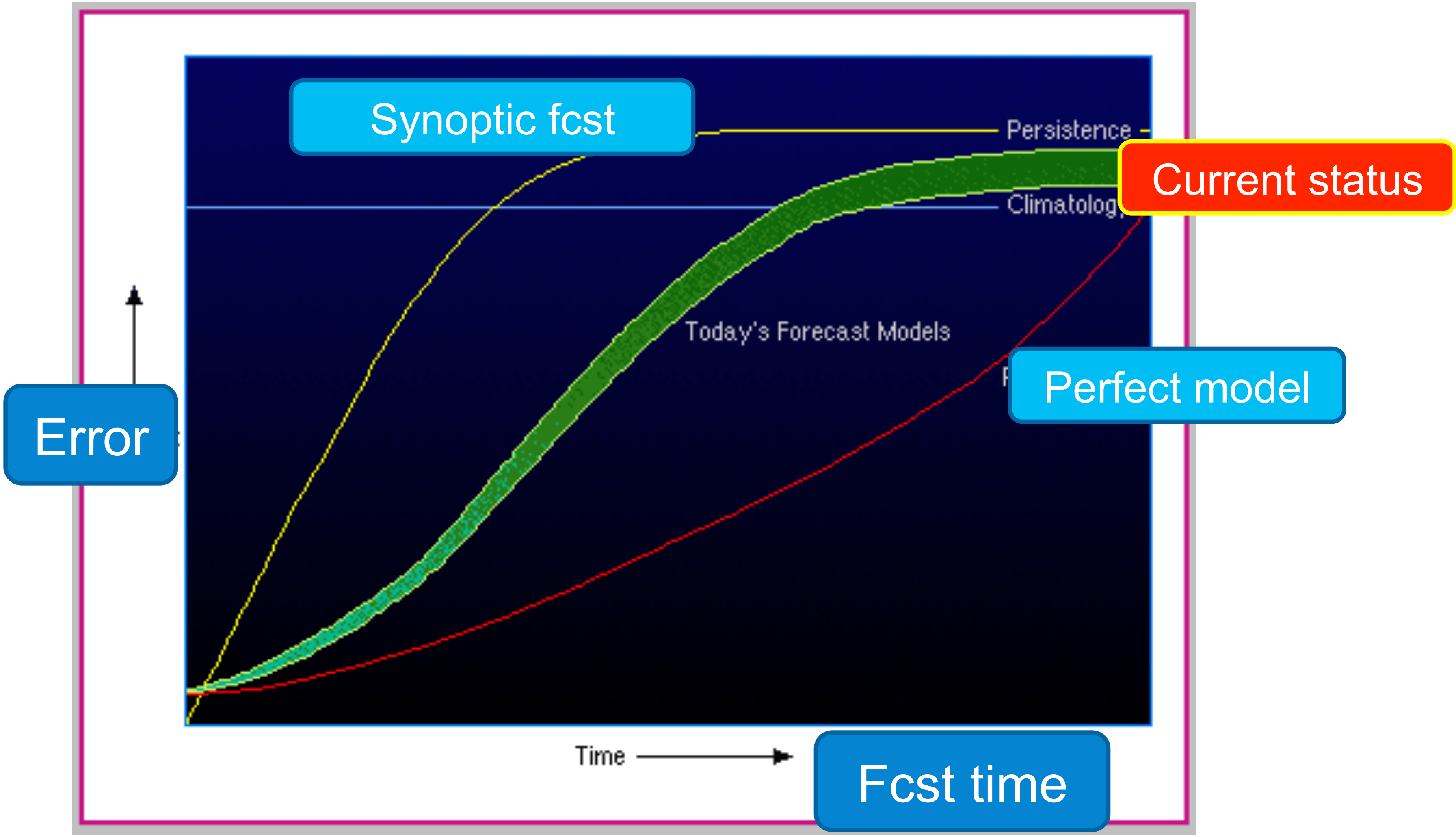
↓

Round-off error -> cause of non-periodcity

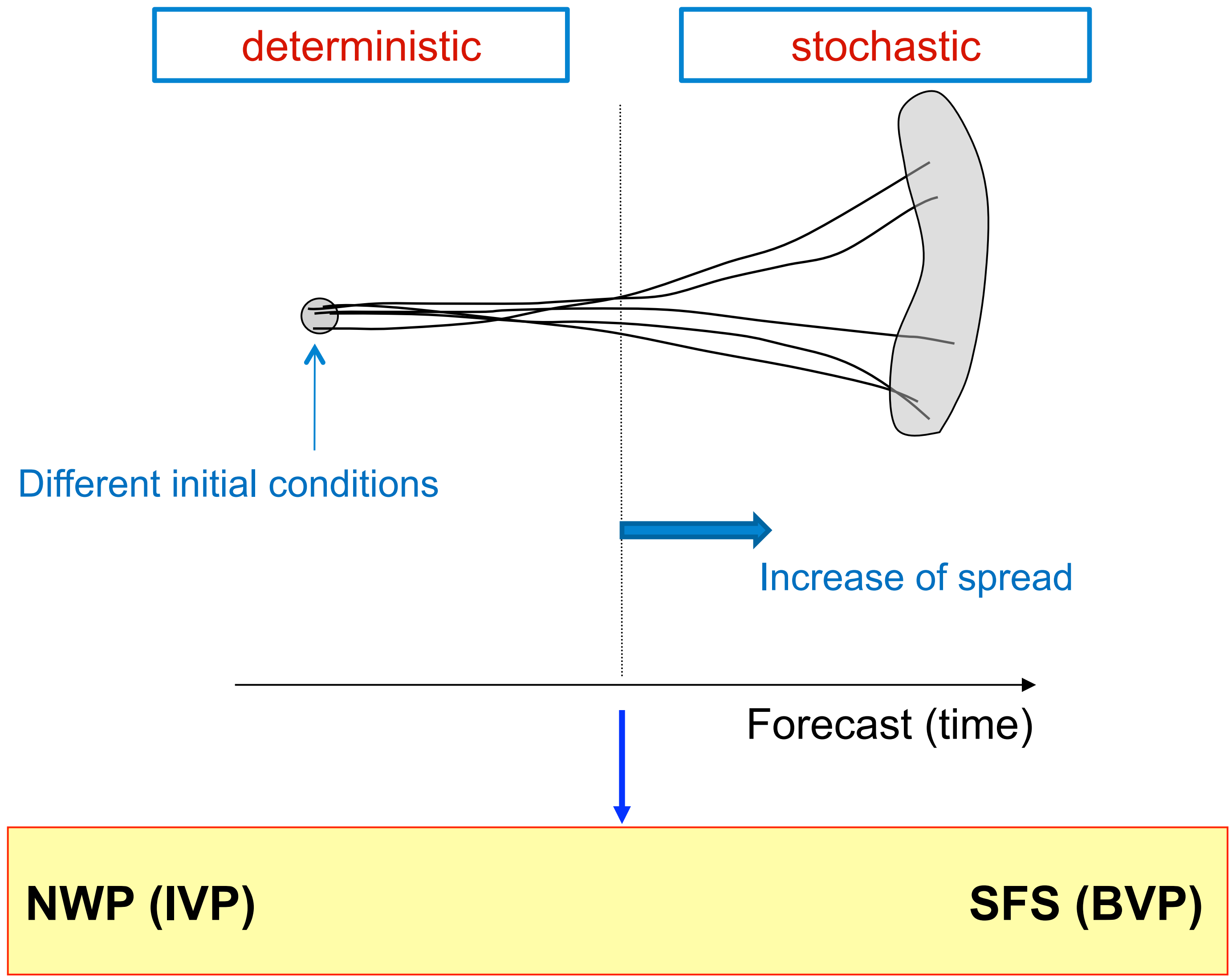
↓

Chaos theory– two weeks

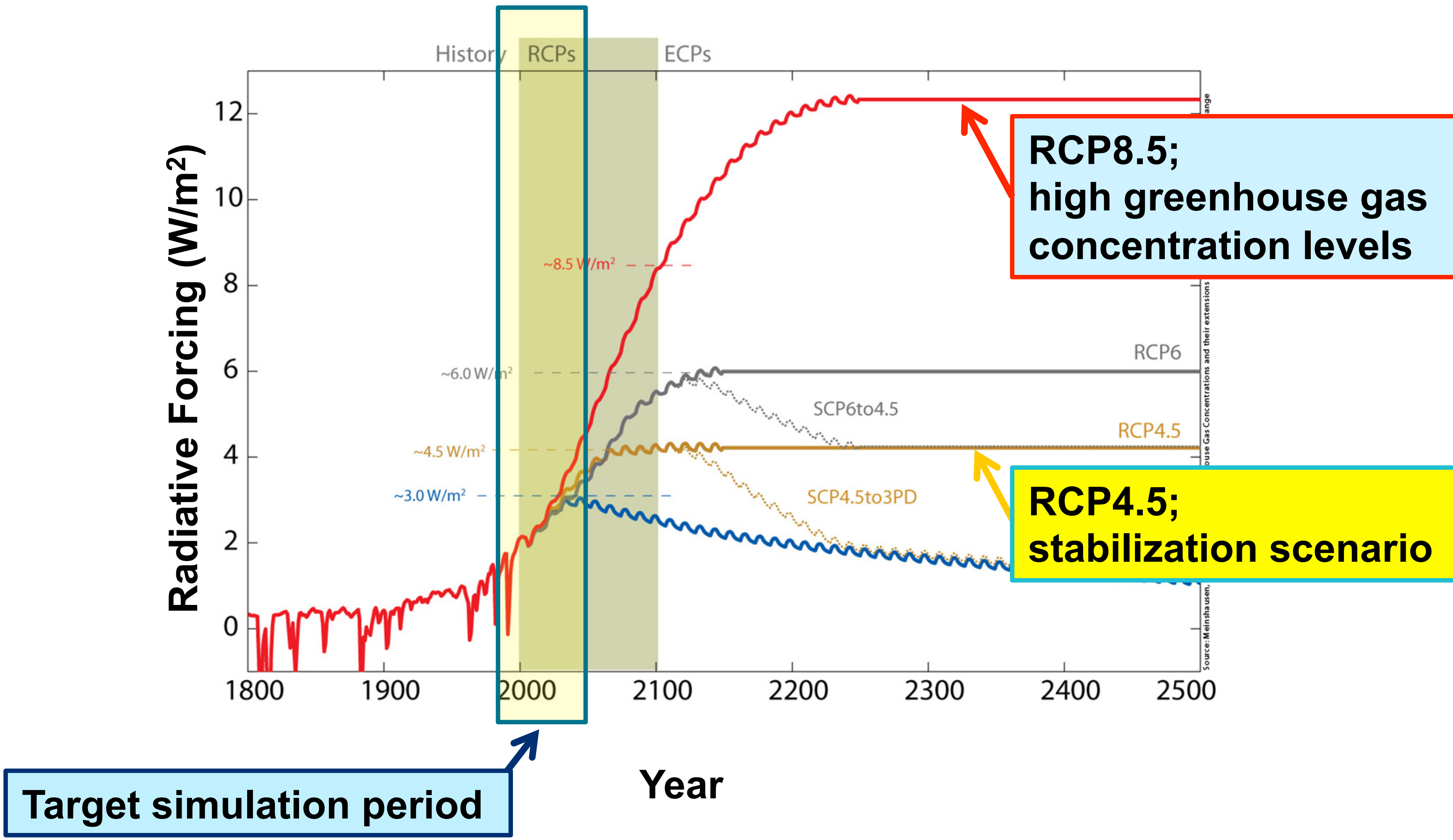
Predictability : Atmosphere is unstable



Ensemble forecasts : Seasonal and beyond

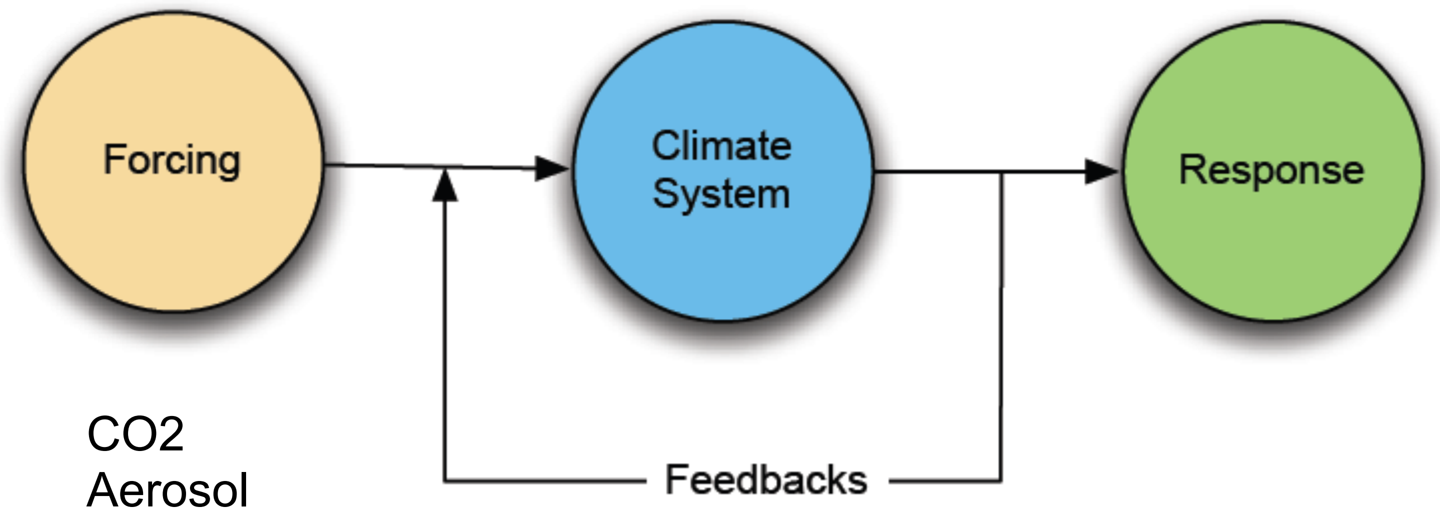


Climate prediction : RCP scenarios



(Meinshausen et al. 2011)

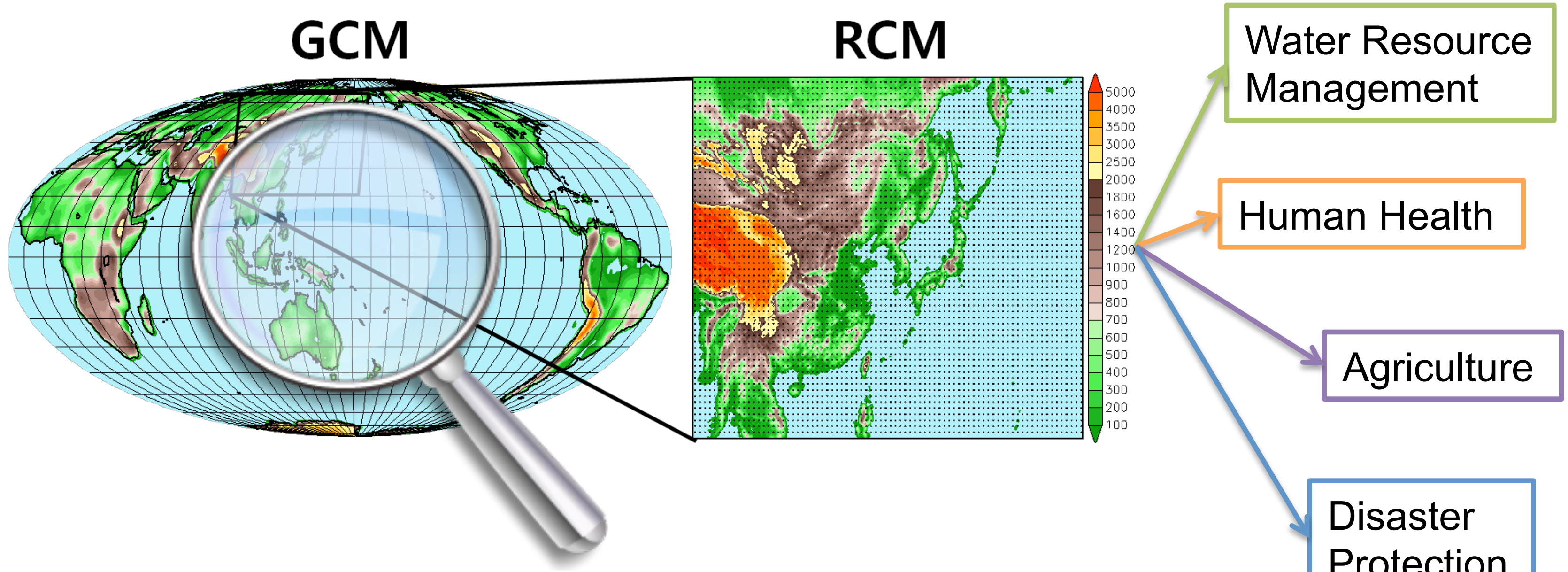
Climate prediction : Climate system sensitivity



CO2
Aerosol
Volcanic

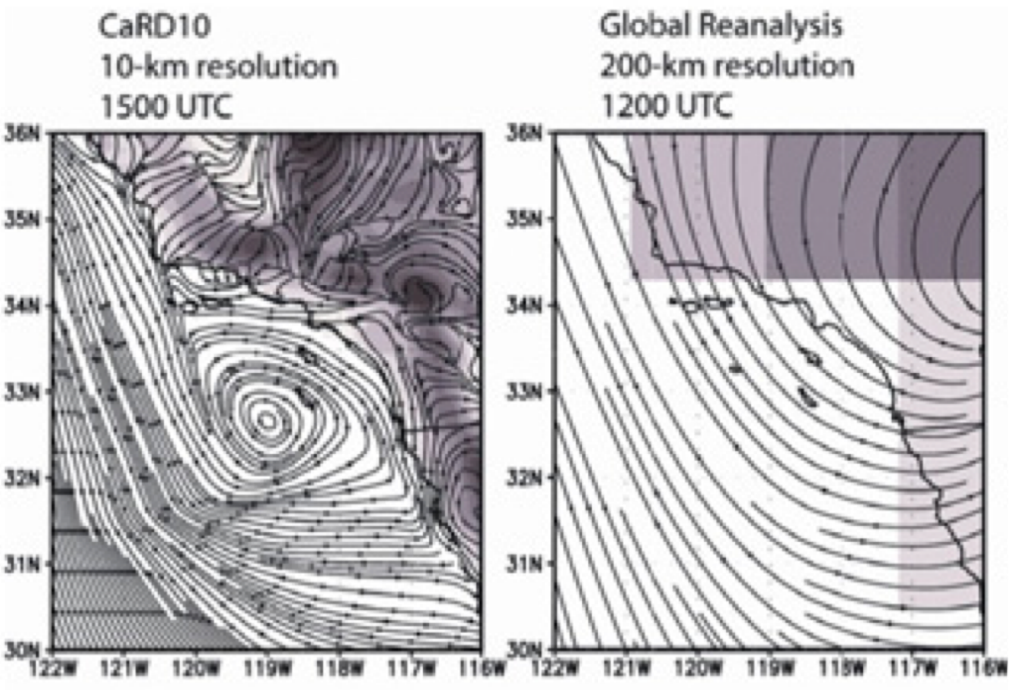
Water vapor feedback
Ice-albedo feedback
Vegetation feedbacks
Cloud (radiative) feedback
(Great debate!, Mostly still uncertain)
…

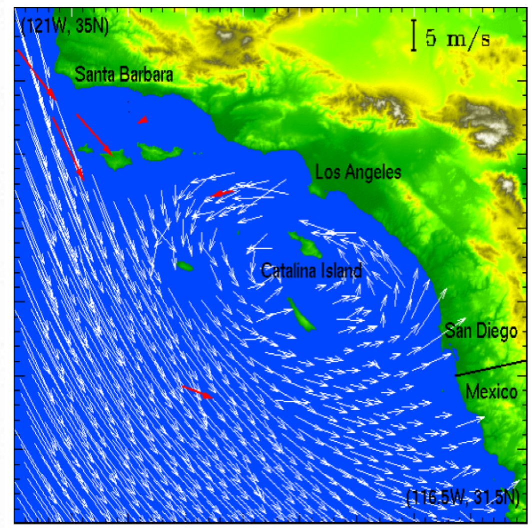
Global versus Regional



Regional model is a magnifying glass

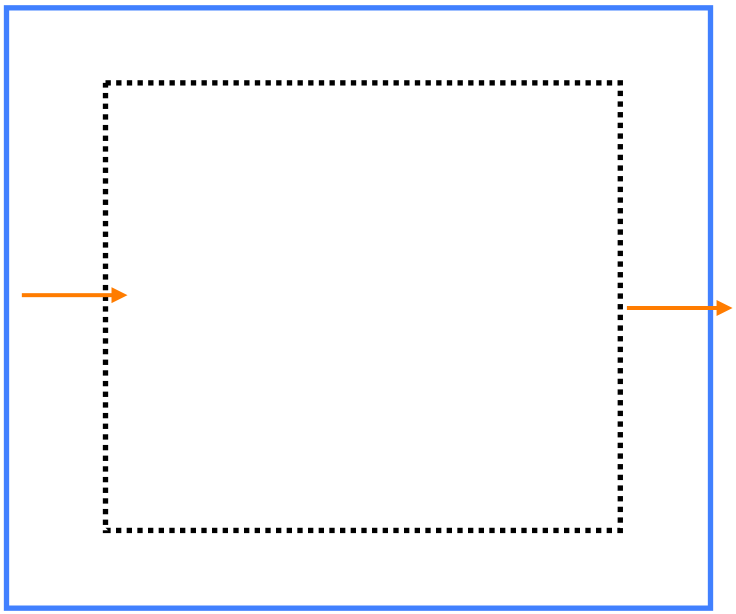
Benefit ?   ---- Very clear !



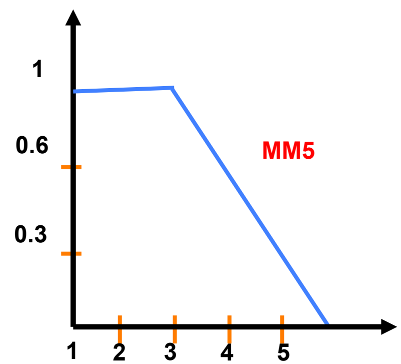


Another inherent issue in regional modeling
: lateral boundary treatment is empirical

Buffer zone



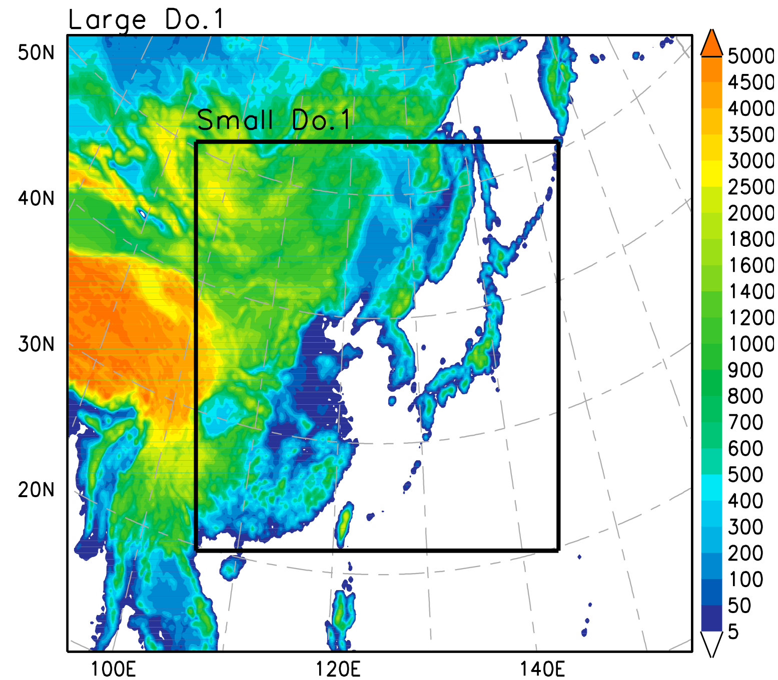
F(n) : weighting of global

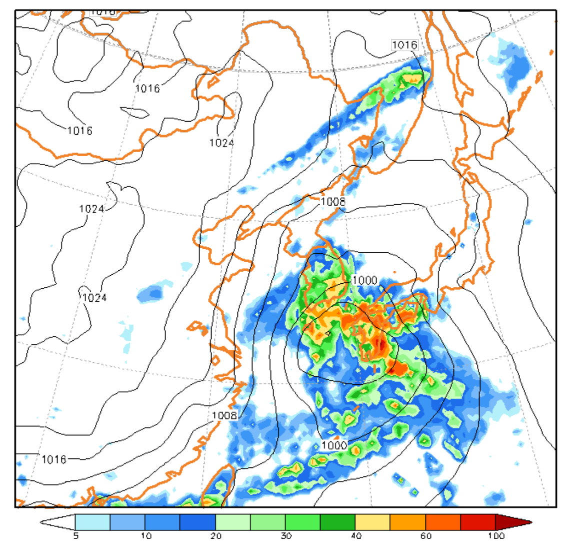


$$\left.\frac{\partial A}{\partial t}\right|_n = F(n)F_1(A_{CM}-A_{FM}) - F(n)F_2\nabla^2(A_{CM}-A_{FM})$$

So, empirical

Domain size sensitivity : A mid-latitude cyclone

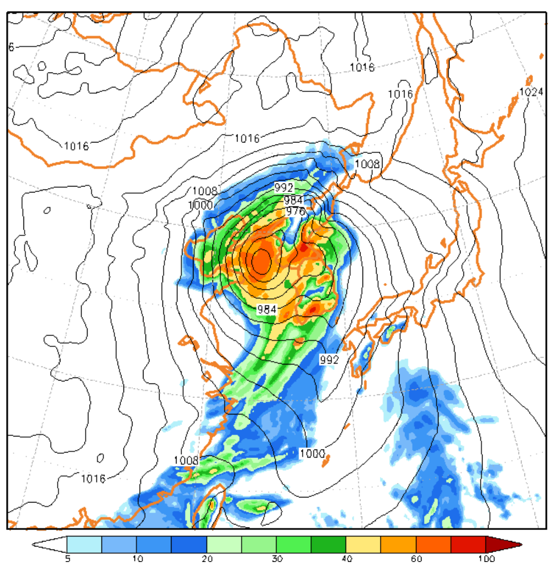


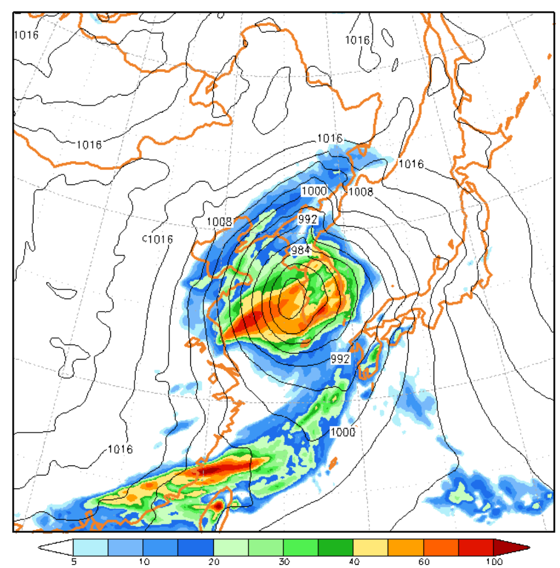


Observed

Large

Away from OBS

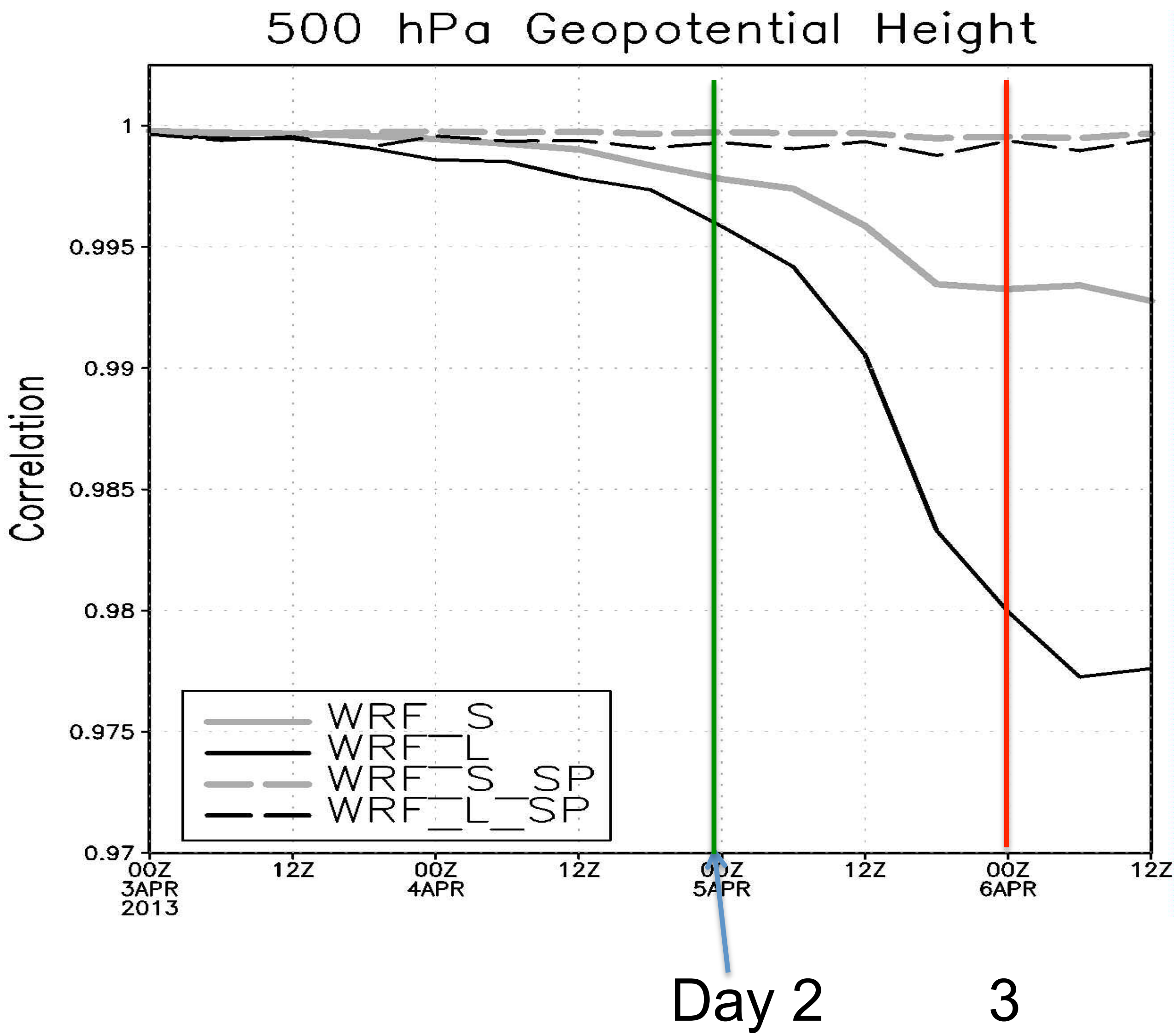




Small

Close to OBS

Domain size sensitivity : Pattern correlation with global



Fundamental limit of the regional model : low resolution global and mathematically ill-posed setup

Small domain keeps the large-scale from the global but loses its freedom

Spectral nudging keeps the large-scale, but may lose the regional details

Thanks for your attention !

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Hong, S.-Y., and M. Kanamitsu, 2014: Dynamical downscaling: Fundamental issues from an NWP point of view and recommendations. *Asia-Pac. J. Atmos. Sci*., **50,** 83-104, doi: 10.1007/s13143-014-0029-2.

Dudhia, J., 2014: A history of mesoscale model Development. *Asia-Pac. J. Atmos. Sci*., **50,** 121-131.