The Advanced Research WRF (ARW) Dynamics Solver

1. What is a dynamics solver?
2. Variables and coordinates
3. Equations
4. Time integration scheme
5. Grid staggering
6. Advection (transport) and conservation
7. Time step parameters
8. Filters
9. Map projections and global configuration
10. Boundary condition options

WRF ARW Tech Note

A Description of the Advanced Research WRF Version 3 (June 2008, 2012 update)

http://www.mmm.ucar.edu/wrf/users/pub-doc.html

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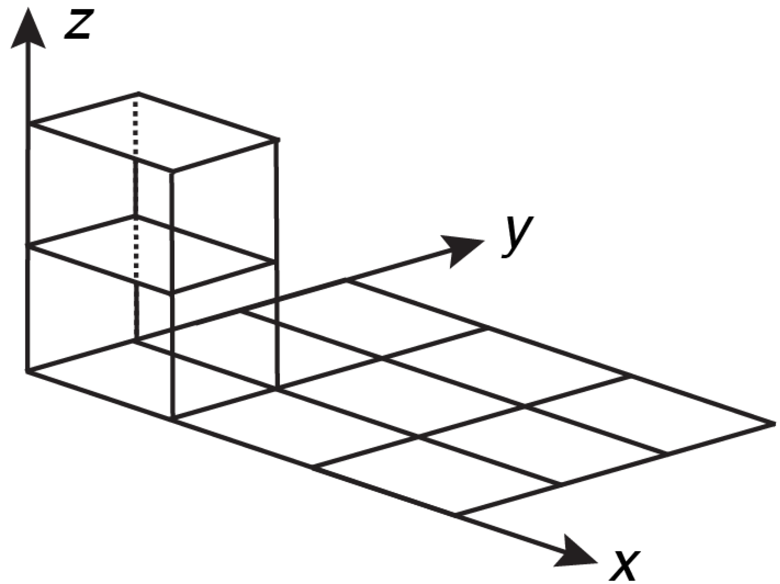
A *dynamical solver* (or a *dynamical core*, or *dycore*) performs a time (t) and space (x,y,z) integration of the equations of motion.

Given the 3D atmospheric state at time t, $S(x,y,z,t)$, we integrate the equations forward in time from $t \longrightarrow T$, i.e. we run the model and produce a forecast.

The equations cannot be solved analytically, so we *discretize* the equations on a *grid* and compute *approximate* solutions.



The accuracy of the solutions depend on the numerical method and the mesh spacing (grid).

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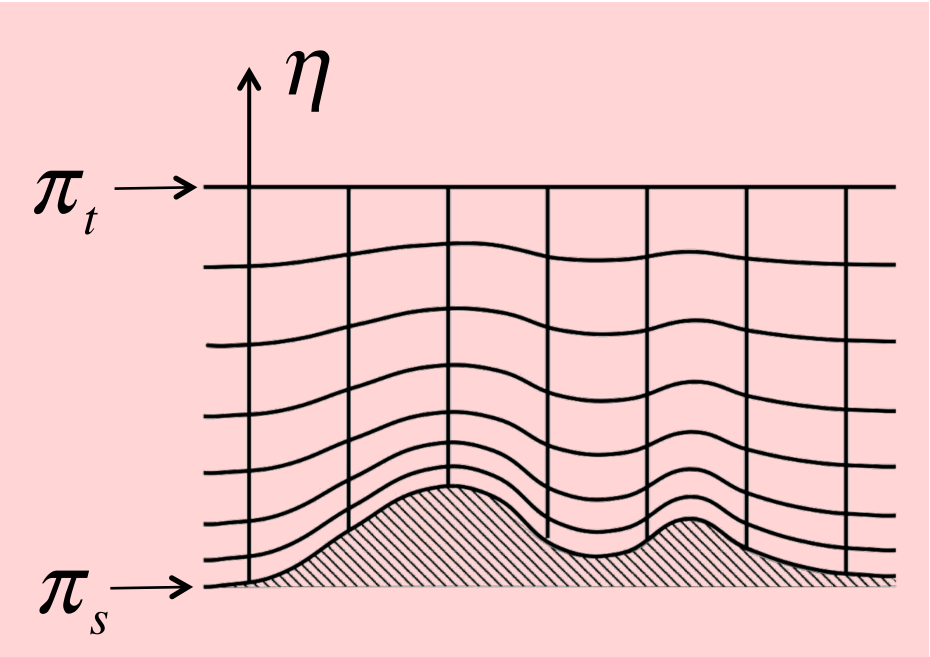
Dynamics: 2. Variables and coordinates

Vertical coordinates: (1) Traditional terrain-following mass coordinate

Hydrostatic pressure π

Column mass (per unit area) $\mu = \pi_s - \pi_t$

Vertical coordinate $\eta = \frac{(\pi - \pi_t)}{\mu}$



Layer mass (per unit area) $\mu \Delta \eta = \Delta \pi = g \rho \Delta z$

Pressure $\pi(\eta) = \eta \mu + \pi_t$

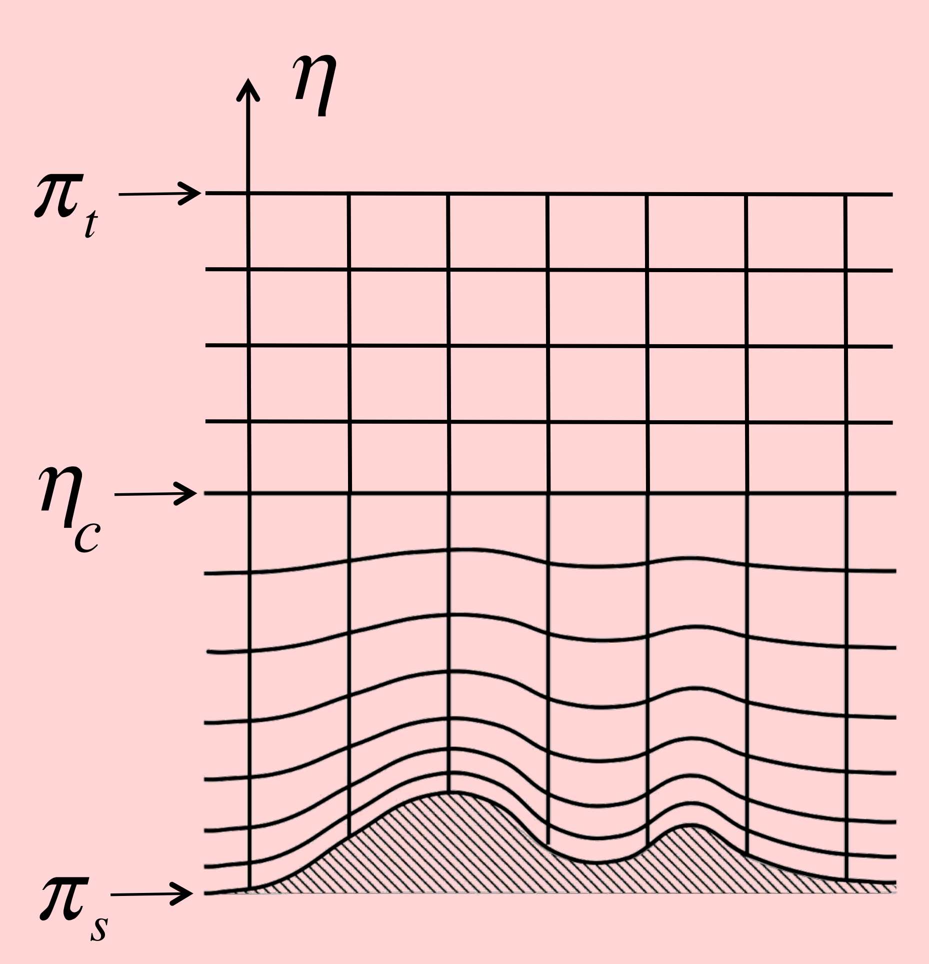
Vertical coordinates: (2) Hybrid terrain-following mass coordinate

Isobaric (constant pressure)

Pressure $\pi = \pi(\eta)$

Terrain following

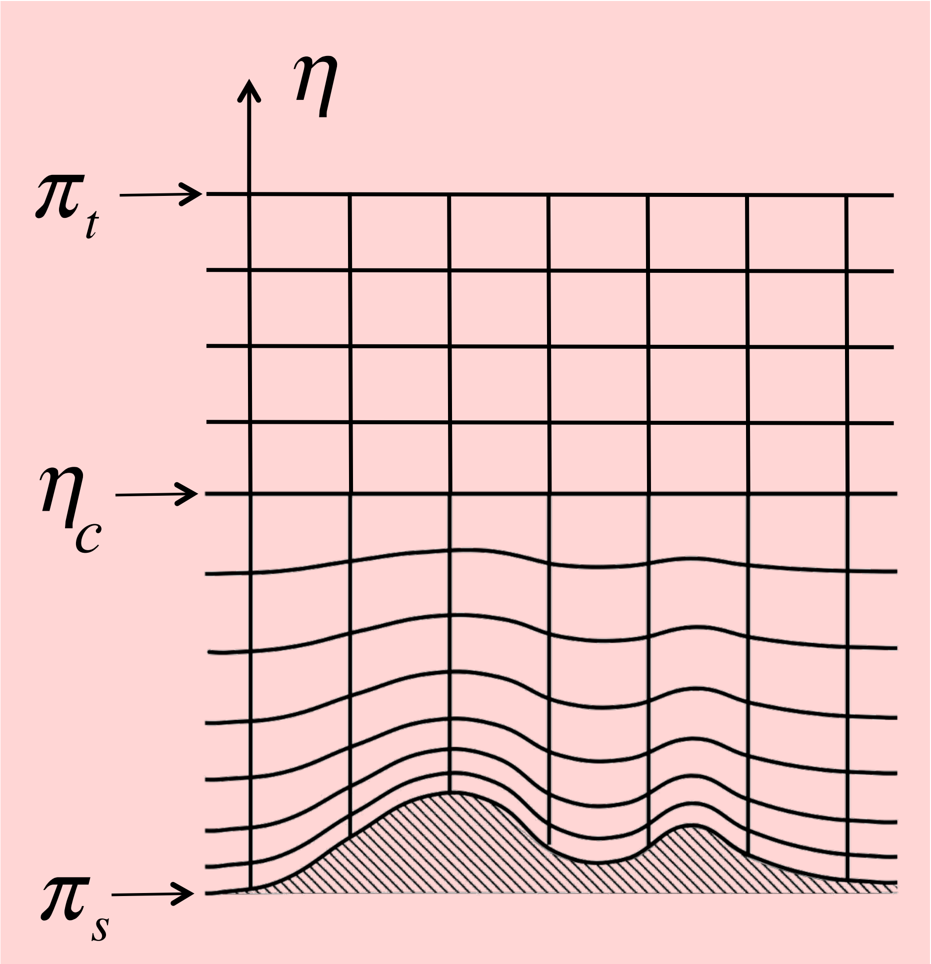
Pressure $\pi(\eta) = \eta\mu + \pi_t$



Variables:

Column mass (per unit area)

$$\mu = \pi_s - \pi_t$$



Conserved state (prognostic) variables:

$$\mu, \quad U = \mu u, \quad V = \mu v,$$

$$W = \mu w, \quad \Theta = \mu \theta$$

Non-conserved state variable:

$$\phi = gz$$

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Dynamics: 2. Variables and coordinates

Vertical momentum eqn.

$$\frac{\partial W}{\partial t} + g\left(\mu_d - \frac{\alpha}{\alpha_d}\frac{\partial p}{\partial \eta}\right) = -\frac{\partial U w}{\partial x} - \frac{\partial \Omega w}{\partial \eta}$$

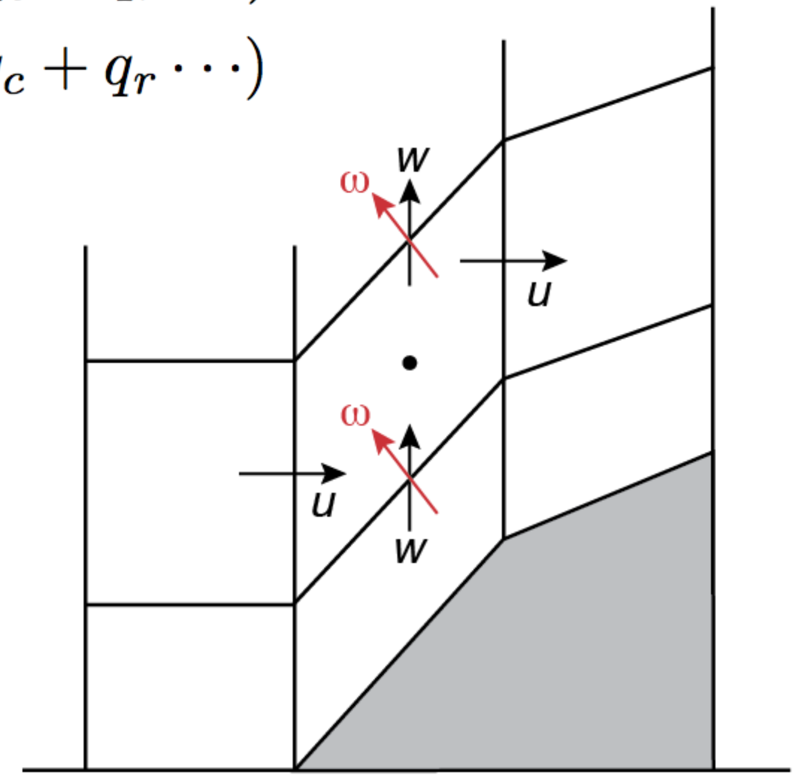
Subscript *d* denotes *dry*, and

$$\alpha_d = \frac{1}{\rho_d} \qquad \begin{aligned} \alpha &= \alpha_d \left(1 + q_v + q_c + q_r \cdots\right)^{-1} \\ \rho &= \rho_d \left(1 + q_v + q_c + q_r \cdots\right) \end{aligned}$$

covariant (*u, ω*) and
contravariant *w* velocities

$$u = \frac{dx}{dt}, \quad w = \frac{dz}{dt}, \quad \omega = \frac{d\eta}{dt}$$

$$U = \mu u, \quad W \mu w, \quad \Omega = \mu\omega$$



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Dynamics: 3. Equations

transport | pressure gradient

$$\frac{\partial U}{\partial t} = -\frac{\partial U u}{\partial x} - \frac{\partial V u}{\partial y} - \frac{\partial \Omega u}{\partial \eta} - \alpha\mu_d \frac{\partial p}{\partial x} - \frac{\alpha}{\alpha_d}\frac{\partial p}{\partial \eta}\frac{\partial \phi}{\partial x} + R_u + Q_u$$

$$\frac{\partial V}{\partial t} = -\frac{\partial U v}{\partial x} - \frac{\partial V v}{\partial y} - \frac{\partial \Omega v}{\partial \eta} - \alpha\mu_d \frac{\partial p}{\partial y} - \frac{\alpha}{\alpha_d}\frac{\partial p}{\partial \eta}\frac{\partial \phi}{\partial y} + R_v + Q_v$$

$$\frac{\partial W}{\partial t} = -\frac{\partial U w}{\partial x} - \frac{\partial V w}{\partial y} - \frac{\partial \Omega w}{\partial \eta} - g\left(\mu_d - \frac{\alpha}{\alpha_d}\frac{\partial p}{\partial \eta}\right) + R_w + Q_w$$

$$\frac{\partial \mu_d}{\partial t} = -\frac{\partial U}{\partial x} - \frac{\partial V}{\partial y} - \frac{\partial \Omega}{\partial \eta}$$

$$\frac{\partial \Theta}{\partial t} = -\frac{\partial U \theta}{\partial x} - \frac{\partial V \theta}{\partial y} - \frac{\partial \Omega \theta}{\partial \eta} + R_\theta + Q_\theta$$

$$\frac{\partial \mu_d q_j}{\partial t} = -\frac{\partial U q_j}{\partial x} - \frac{\partial V q_j}{\partial y} - \frac{\partial \Omega q_j}{\partial \eta} + R_{q_j} + Q_{q_j}$$

$$\frac{\partial \phi}{\partial t} = -u\frac{\partial \phi}{\partial x} - v\frac{\partial \phi}{\partial y} - \omega\frac{\partial \phi}{\partial \eta} + gw$$

↑ ← numerical filters, physics, projection terms

← geopotential eqn term

Diagnostic relations: $\frac{\partial \phi}{\partial \eta} = -\alpha_d \mu_d, p = \left(\frac{R_d \Theta_m}{p_o \mu_d \alpha_d}\right)^\gamma, \ \Theta_m = \Theta\left(1 + \frac{R_v}{R_d} q_v\right)$

3rd Order Runge-Kutta time integration

$$\frac{\partial U}{\partial t} = RHS_u$$
$$\frac{\partial V}{\partial t} = RHS_v$$
$$\frac{\partial W}{\partial t} = RHS_w$$

•
•
•

advance $\phi^t \rightarrow \phi^{t+\Delta t}$

$$\phi^* = \phi^t + \frac{\Delta t}{3} RHS(\phi^t)$$
$$\phi^{**} = \phi^t + \frac{\Delta t}{2} RHS(\phi^*)$$
$$\phi^{t+\Delta t} = \phi^t + \Delta t\, RHS(\phi^{**})$$

Amplification factor $\phi_t = i k \phi; \quad \phi^{n+1} = A \phi^n; \quad |A| = 1 - \frac{(k\Delta t)^4}{24}$

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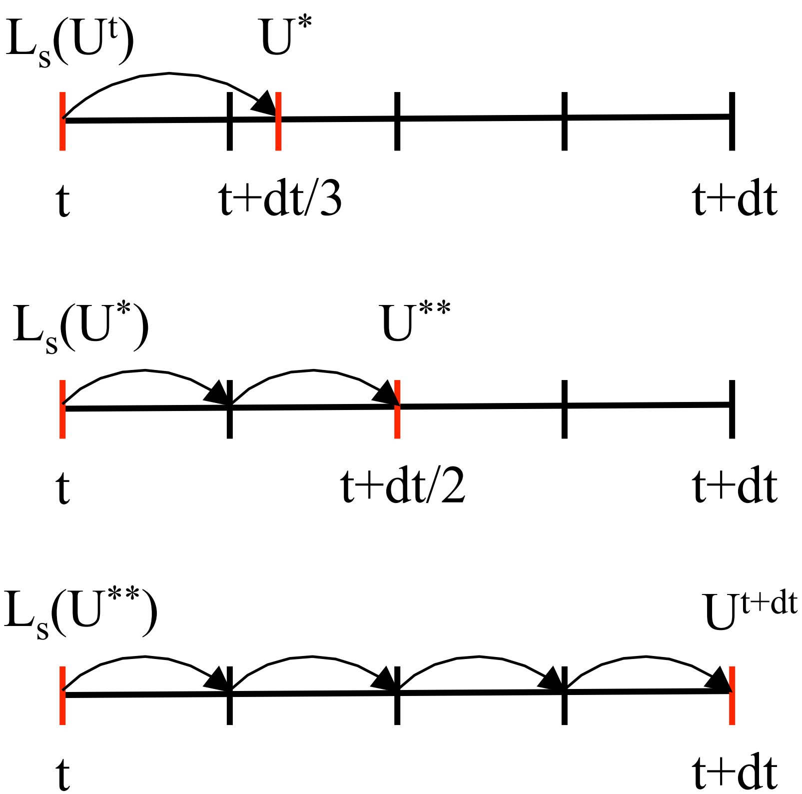
$$U_t = L_{fast}(U) + L_{slow}(U)$$

3rd order Runge-Kutta, 3 steps



fast: acoustic and gravity wave terms.
slow: everything else.

- RK3 is 3rd order accurate for linear eqns, 2nd order accurate for nonlinear eqns.
- Stable for centered and upwind advection schemes.
- Stable for Courant number $Udt/dx < 1.73$
- Three $L_{slow}(U)$ evaluations per timestep.

Dynamics: 4. Time integration scheme – acoustic step

$$U^{t+\Delta t},\quad V^{t+\Delta t} \qquad \frac{\partial U}{\partial t} + \left(\mu_d \alpha \frac{\partial p}{\partial x} + \frac{\alpha}{\alpha_d}\frac{\partial p}{\partial \eta}\frac{\partial \phi}{\partial x}\right)^{\tau} = R_U^t$$

$$\mu_d^{\tau+\Delta\tau} \quad \Omega^{\tau+\Delta\tau} \qquad \frac{\partial \mu_d}{\partial t} + \frac{\partial U}{\partial x}^{\tau+\Delta\tau} + \frac{\partial \Omega}{\partial \eta}^{\tau+\Delta\tau} = 0$$

$$\Theta^{\tau+\Delta\tau} \qquad \frac{\partial \Theta}{\partial t} + \left(\frac{\partial U\theta^t}{\partial x} + \frac{\partial \Omega\theta^t}{\partial \eta}\right)^{\tau+\Delta\tau} = R_\Theta^t$$

$$\begin{array}{ll} W^{\tau+\Delta\tau} & \left\{ \begin{array}{l} \dfrac{\partial W}{\partial t} + g\overline{\left(\mu_d - \dfrac{\alpha}{\alpha_d}\dfrac{\partial p}{\partial \eta}\right)}^{\tau} = R_W^t \\ \mu_d^t \dfrac{\partial \phi}{\partial t} + U^{\tau+\Delta\tau}\dfrac{\partial \phi}{\partial x}^t + \Omega^{\tau+\Delta\tau}\dfrac{\partial \phi}{\partial \eta}^t - g\overline{W}^{\tau} = R_\phi^t \end{array} \right. \\ \phi^{\tau+\Delta\tau} & \end{array}$$

- Forward-backward differencing on U, Θ, and μ equations
- Vertically implicit differencing on W and ϕ equations

Begin time step

Runge-Kutta loop (steps 1, 2, and 3)

(i) advection, p-grad, buoyancy using (ϕ^t, ϕ^*, ϕ^{**})

(ii) physics if step 1, save for steps 2 and 3

(iii) mixing, other non-RK dynamics, save…

(iv) assemble dynamics tendencies

Acoustic step loop

(i) advance U,V, then μ, Θ, then w, ϕ

(ii) time-average U,V, Ω

End acoustic loop

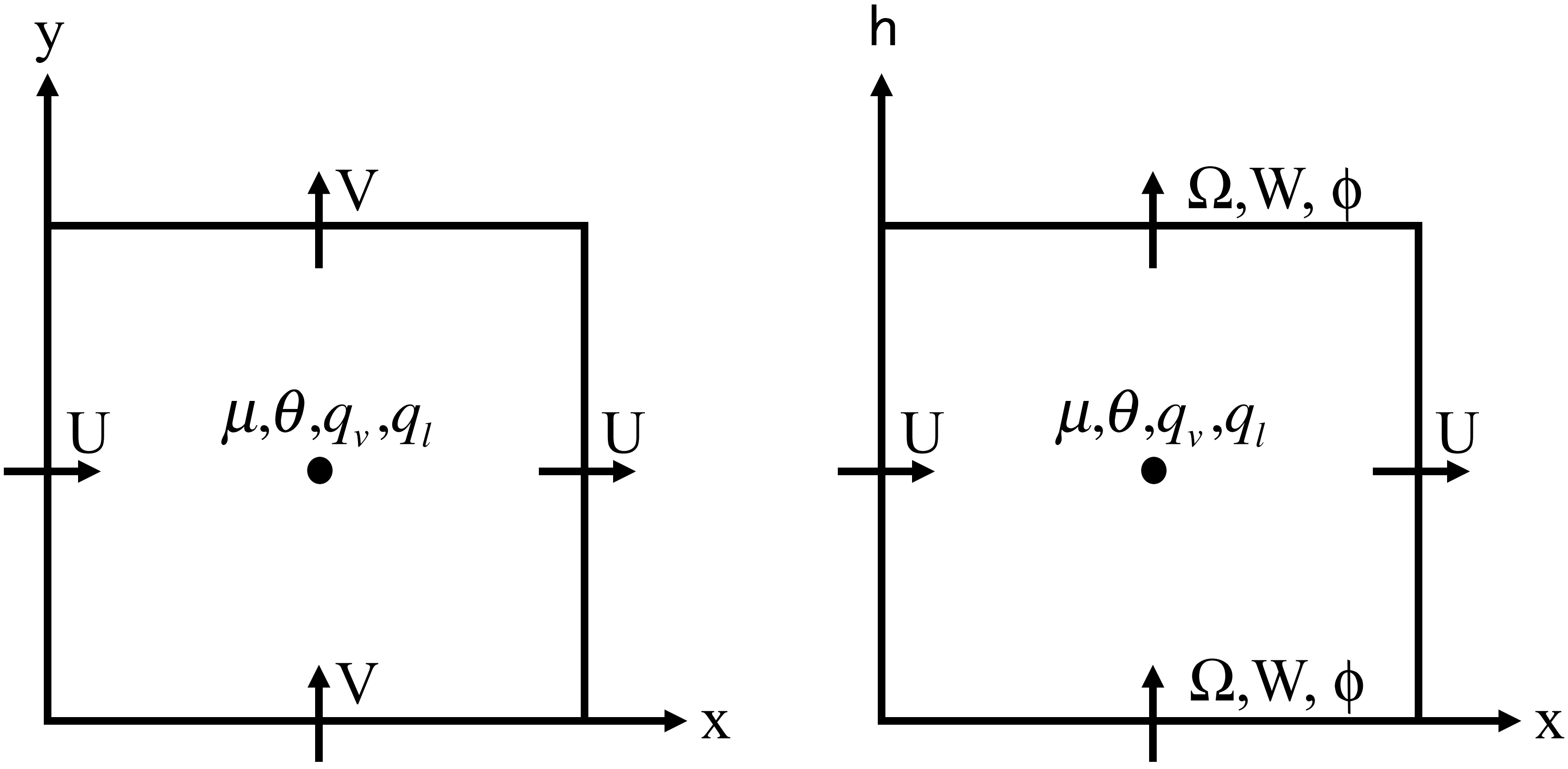
Advance scalars using time-averaged U,V, Ω

End Runge-Kutta loop

Adjustment physics (currently microphysics)

End time step

C-grid staggering



horizontal

vertical

2nd, 3rd, 4th, 5th and 6th order centered and upwind-biased schemes are available in the ARW model.

Example: 5th order scheme

$$\frac{\partial(U\psi)}{\partial x} = \frac{1}{\Delta x}\left(F_{i+\frac{1}{2}}(U\psi) - F_{i-\frac{1}{2}}(U\psi) \right)$$

where

$$F_{i-\frac{1}{2}}(U\psi) = U_{i-\frac{1}{2}}\left\{ \frac{37}{60}(\psi_i + \psi_{i-1}) - \frac{2}{15}(\psi_{i+1} + \psi_{i-2}) + \frac{1}{60}(\psi_{i+2} + \psi_{i-3}) \right\}$$

$$- sign(1,U)\,\frac{1}{60}\left\{ (\psi_{i+2} - \psi_{i-3}) - 5(\psi_{i+1} - \psi_{i-2}) + 10(\psi_i - \psi_{i-1}) \right\}$$

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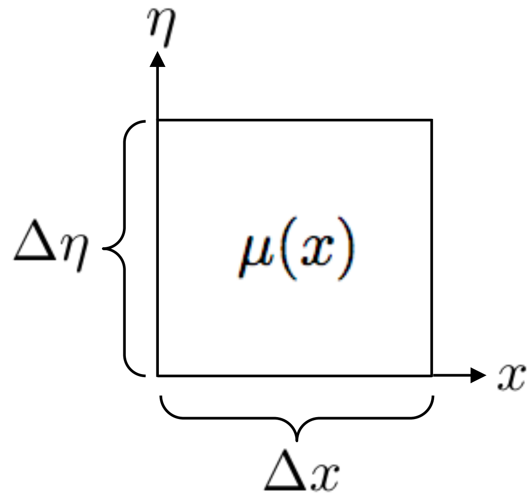
For constant U, the 5th order flux divergence tendency becomes

$$\Delta t \frac{\delta(U\psi)}{\Delta x}\bigg|_{5th} = \Delta t \frac{\delta(U\psi)}{\Delta x}\bigg|_{6th} - \underbrace{\left|\frac{U\Delta t}{\Delta x}\right| \frac{1}{60}\left(-\psi_{i-3} + 6\psi_{i-2} - 15\psi_{i-1} + 20\psi_i - 15\psi_{i+1} + 6\psi_{i+2} - \psi_{i+3}\right)}_{\frac{Cr}{60}\frac{\partial^6 \psi}{\partial x^6} + H.O.T}$$

The odd-ordered flux divergence schemes are equivalent to the next higher ordered (even) flux-divergence scheme plus a dissipation term of the higher even order with a coefficient proportional to the Courant number.

Dynamics: 6. Advection (transport) and conservation – dry-air mass

control volume
(2D example)



Mass in a control volume is proportional to

$$(\Delta x \Delta \eta)(\mu)^t$$

since $\mu(x)\Delta\eta = \Delta\pi = -g\rho\Delta z$

Dynamics: 6. Advection (transport) and conservation – dry-air mass

Mass in a control volume $(\Delta x \Delta \eta)(\mu)^t$
2D example

Mass conservation equation

$$\Delta t^{-1}(\Delta x \Delta \eta) \cdot \left[(\mu)^{t+\Delta t} - (\mu)^t)\right] = \left[(\mu u \Delta \eta)_{x-\Delta x/2,\eta} - (\mu u \Delta \eta)_{x+\Delta x/2,\eta})\right] + \left[(\mu \omega \Delta x)_{x,\eta-\Delta \eta/2} - (\mu \omega \Delta x)_{x,\eta+\Delta \eta/2})\right]$$

Change in mass over a time step

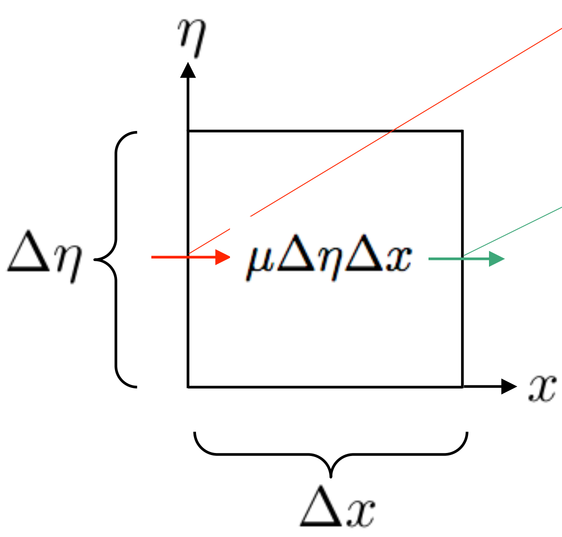
mass fluxes through control volume faces

Dynamics: 6. Advection (transport) and conservation – dry-air mass

Mass in a control volume $(\Delta x \Delta \eta)(\mu)^t$

Mass conservation equation

$$\Delta t^{-1}(\Delta x \Delta \eta) \cdot \left[(\mu)^{t+\Delta t} - (\mu)^t)\right] = \left[(\mu u \Delta \eta)_{x-\Delta x/2,\eta} - (\mu u \Delta \eta)_{x+\Delta x/2,\eta}\right] + \left[(\mu \omega \Delta x)_{x,\eta-\Delta \eta/2} - (\mu \omega \Delta x)_{x,\eta+\Delta \eta/2})\right]$$



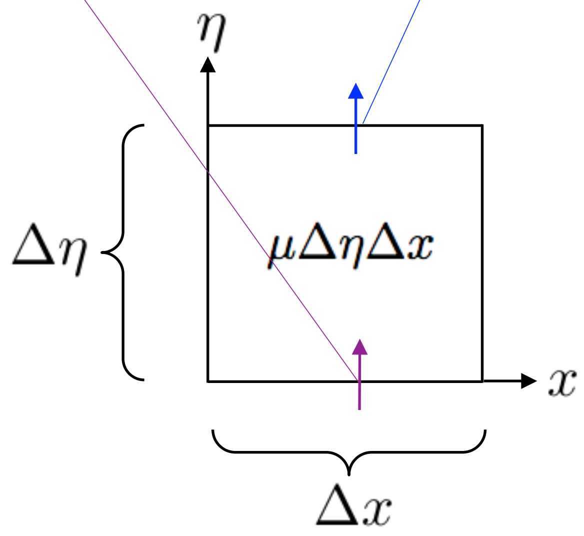
Horizontal fluxes through the vertical control-volume faces

Mass in a control volume $(\Delta x \Delta \eta)(\mu)^t$

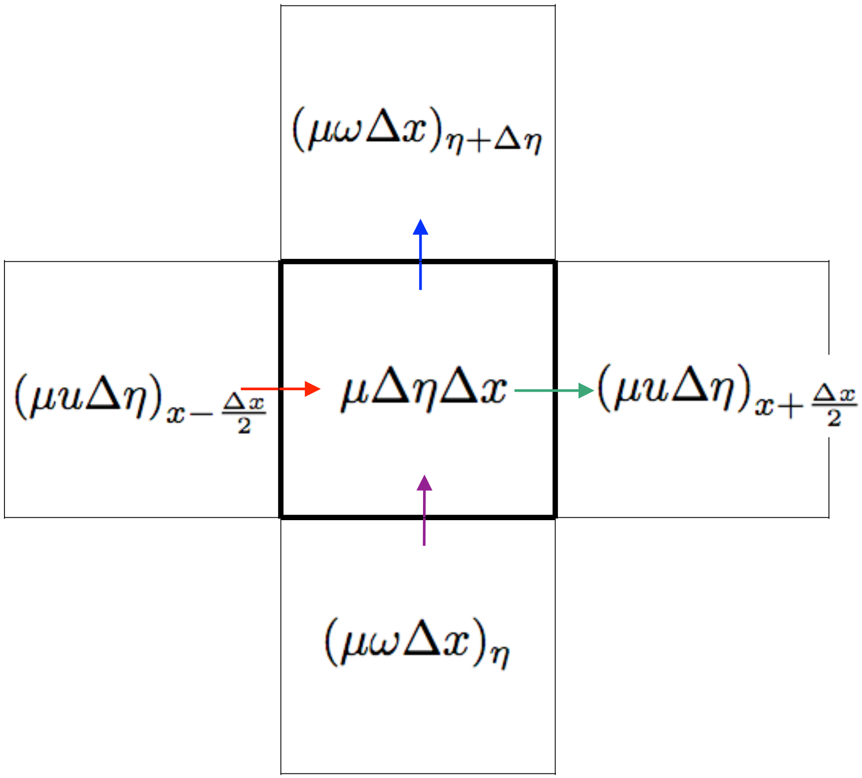
Mass conservation equation

$$\Delta t^{-1}(\Delta x \Delta \eta)\cdot\left[(\mu)^{t+\Delta t} - (\mu)^t)\right] = \left[(\mu u \Delta \eta)_{x-\Delta x/2,\eta} - (\mu u \Delta \eta)_{x+\Delta x/2,\eta})\right] + \left[(\mu \omega \Delta x)_{x,\eta-\Delta \eta/2} - (\mu \omega \Delta x)_{x,\eta+\Delta \eta/2})\right]$$

Vertical fluxes through the horizontal control-volume faces



The same mass fluxes are used for neighboring grid cells - hence mass is conserved locally and globally.



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Dynamics: 6. Advection (transport) and conservation – scalars

Mass in a control volume $(\Delta x \Delta \eta)(\mu)^t$

Scalar mass $(\Delta x \Delta \eta)(\mu\phi)^t$

Mass conservation equation:

$$\Delta t^{-1}(\Delta x \Delta \eta) \cdot \left[(\mu)^{t+\Delta t} - (\mu)^t)\right] = \left[(\mu u \Delta \eta)_{x-\Delta x/2,\eta} - (\mu u \Delta \eta)_{x+\Delta x/2,\eta})\right] + \left[(\mu \omega \Delta x)_{x,\eta-\Delta\eta/2} - (\mu \omega \Delta x)_{x,\eta+\Delta\eta/2})\right]$$

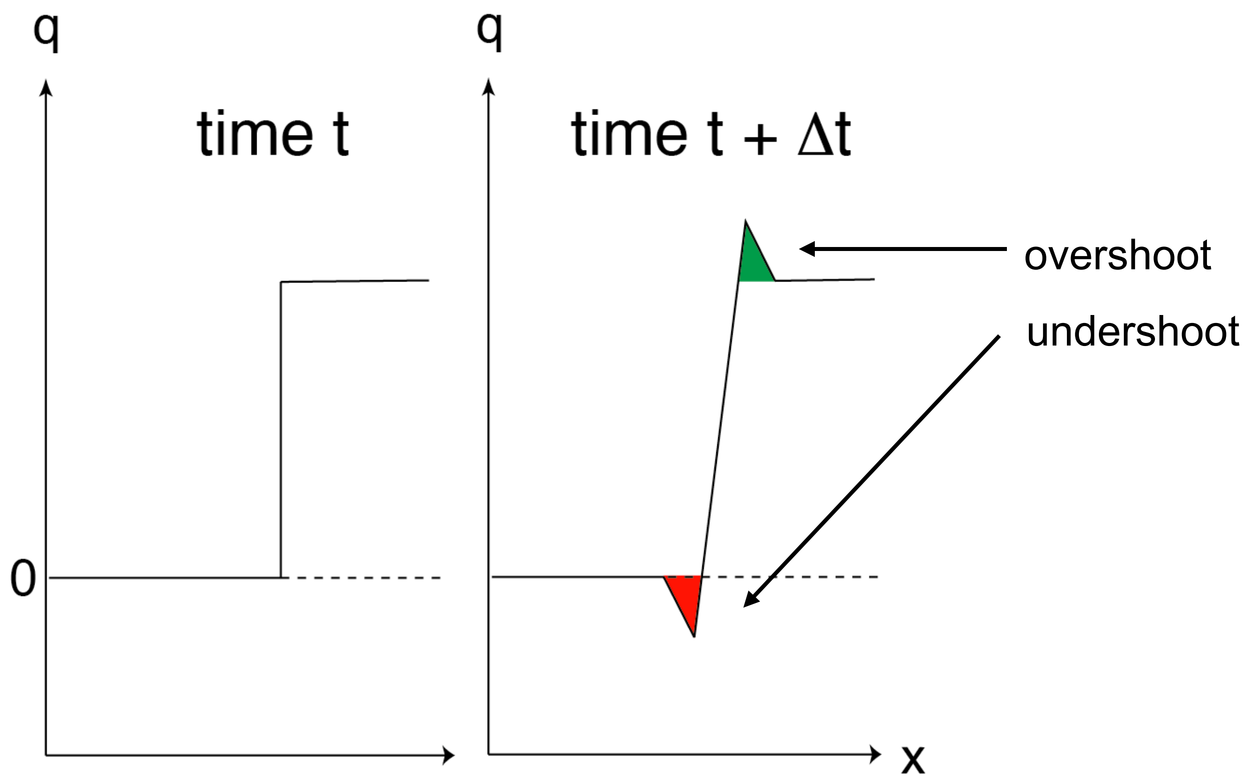
change in mass over a time step | mass fluxes through control volume faces

Scalar mass conservation equation:

$$\Delta t^{-1}(\Delta x \Delta \eta) \cdot \left[(\mu\phi)^{t+\Delta t} - (\mu\phi)^t)\right] = \left[(\mu u \phi \Delta \eta)_{x-\Delta x/2,\eta} - (\mu u \phi \Delta \eta)_{x+\Delta x/2,\eta})\right] + \left[(\mu \omega \phi \Delta x)_{x,\eta-\Delta\eta/2} - (\mu \omega \phi \Delta x)_{x,\eta+\Delta\eta/2})\right]$$

change in tracer mass over a time step | tracer mass fluxes through control volume faces

1D advection



ARW transport is conservative,
but not positive definite nor monotonic.
Removal of negative q ■
results in spurious source of q ■ .

Scalar update, last RK3 step

$$(\mu\phi)^{t+\Delta t} = (\mu\phi)^t - \Delta t \sum_{i=1}^{n} \delta_{x_i}[f_i] \qquad (1)$$

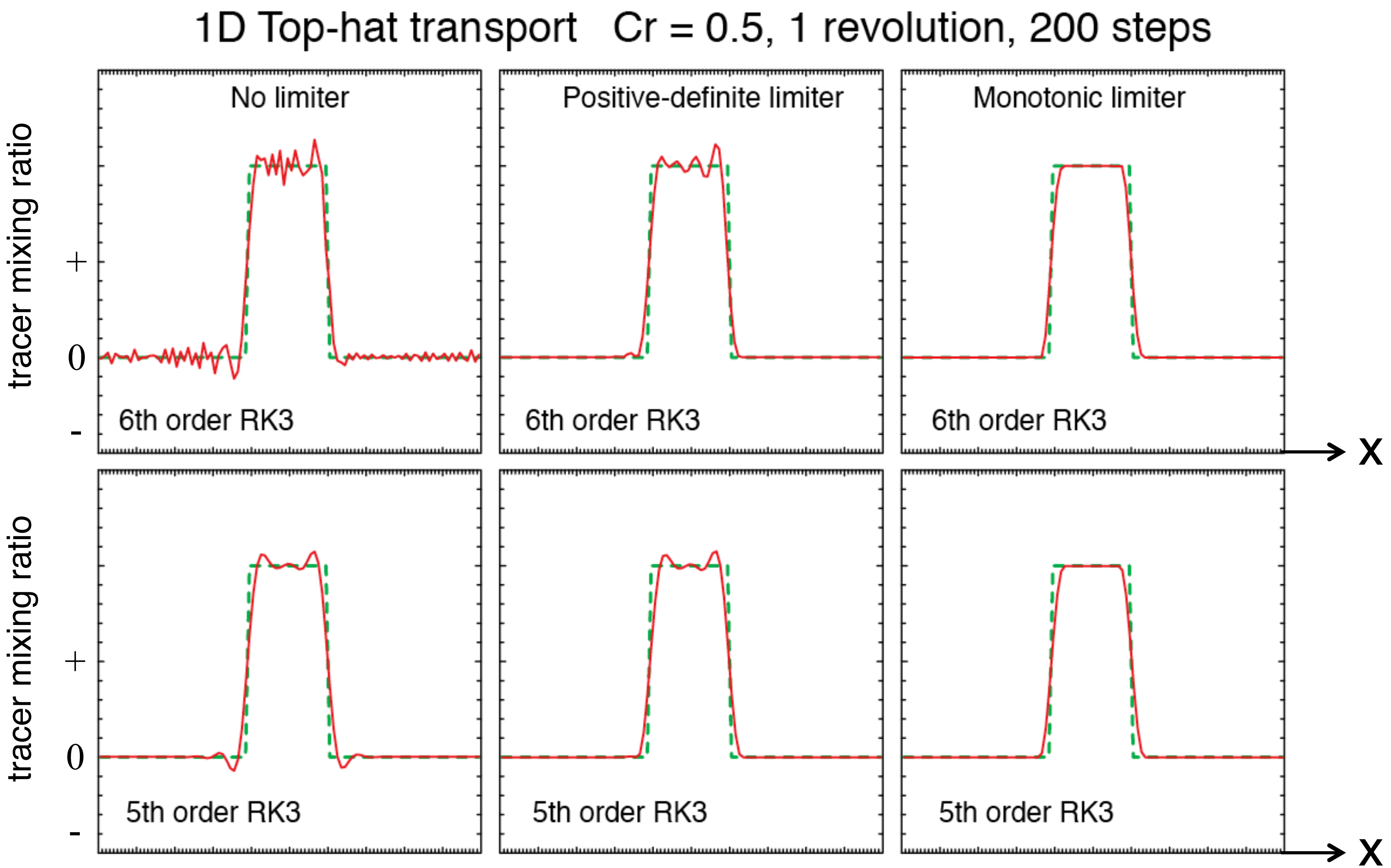
(1) Decompose flux: $f_i = f_i^{upwind} + f_i^c$

(2) Renormalize high-order correction fluxes f_i^c such that solution is positive definite or monotonic: $f_i^c = R(f_i^c)$

(3) Update scalar eqn. (1) using $f_i = f_i^{upwind} + R(f_i^c)$

Skamarock, MWR 2006, 2241-2250

1D Example: Top-Hat Advection



Dynamics: 6. Advection (transport) and conservation

Where are the transport-scheme parameters?

The namelist.input file:

&dynamics

h_mom_adv_order
v_mom_adv_order
h_sca_adv_order
v_sca_adv_order

<u>scheme order</u> (2, 3, 4, or 5)
defaults:
horizontal (h_*) = 5
vertical (v_*) = 3

momentum_adv_opt →

= 1 standard scheme
= 3 5th order WENO
default: 1

moist_adv_opt
scalar_adv_opt
chem_adv_opt
tracer_adv_opt
tke_adv_opt

<u>options:</u>
= 1, 2, 3 : no limiter,
positive definite (PD),
montonic
= 4 : 5th order WENO
= 5 : 5th order PD WENO

3rd order Runge-Kutta time step Δt_{RK}

Courant number limited, 1D: $C_r = \frac{U\Delta t}{\Delta x} < 1.43$

Generally stable using a timestep approximately twice as large as used in a leapfrog model.

Where?

The namelist.input file:

&domains

time_step (integer seconds)

time_step_fract_num

time_step_fract_den

3rd order Runge-Kutta time step Δt_{RK} (&domains *time_step*)

Acoustic time step

2D horizontal Courant number limited: $C_r = \frac{C_s \Delta \tau}{\Delta h} < \frac{1}{\sqrt{2}}$

$\Delta \tau_{sound} = \Delta t_{RK} / (\text{number of acoustic steps})$

Where?

The namelist.input file:

&dynamics

time_step_sound (integer)

Dynamics: 7. Time step parameters

3rd order Runge-Kutta time step Δt_{RK} (&domains *time_step*)

Acoustic time step [&dynamics *time_step_sound* (integer)]

Guidelines for time step

Δt_{RK} in seconds should be about 6*Δx (grid size in kilometers). Larger Δt can be used in smaller-scale dry situations, but *time_step_sound* (default = 4) should increase proportionately if larger Δt is used.

If ARW blows up (aborts) quickly, try:

Decreasing Δt_{RK} (that also decreases Δt_{sound}),

Or increasing time_step_sound (that decreases Δt_{sound} but does not change Δt_{RK})

Dynamics: 8. Filters – divergence damping

Purpose: filter acoustic modes (3-D divergence, $D = \nabla \cdot \rho \mathbf{V}$ *)*

$$\left\{ \frac{\partial \rho \mathbf{V}}{\partial t} + \nabla p + \ldots = \gamma_d' \nabla D \right\}$$

$$\nabla \cdot \left\{ \quad \right\} \quad \rightarrow \quad \frac{\partial D}{\partial t} + \nabla^2 p + \ldots = \gamma_d' \nabla^2 D$$

From the pressure equation: $p_t \simeq c^2 D$

$$\frac{\partial \rho \mathbf{V}}{\partial t} + \nabla [p_\tau + \gamma_d (p^\tau - p^{\tau - \Delta \tau})] + \ldots = 0$$

$\gamma_d = 0.1$ recommended (default) (&dynamics *smdiv*)

(Illustrated in height coordinates for simplicity)

Dynamics: 8. Filters – time off-centering the vertical acoustic modes

Purpose: damp vertically-propagating acoustic modes

$$\frac{\partial W}{\partial t} + g\overline{\left(\mu_d - \frac{\alpha}{\alpha_d}\frac{\partial p}{\partial \eta}\right)}^{\tau} = \ldots$$

$$\frac{\partial \phi}{\partial t} - \frac{g}{\mu_d^t}\overline{W}^{\tau} = \ldots$$

$$\overline{(\quad)}^{\tau} = \frac{1+\beta}{2}\overline{(\quad)}^{\tau+\Delta\tau} + \frac{1-\beta}{2}\overline{(\quad)}^{\tau}$$

Slightly forward centering the vertical pressure gradient damps 3-D divergence as demonstrated for the divergence damper

$\beta = 0.1$ recommended (default) [&dynamics *epssm*]

Dynamics: 8. Filters – external mode filter

Purpose: filter the external mode

Vertically integrated horizontal divergence, $D_h = \int_1^0 (\nabla_\eta \cdot \mu \mathbf{V}_h) d\eta$

$$\left\{ \frac{\partial \mu \mathbf{V}_h}{\partial t} + \ldots = -\gamma_e \nabla_\eta D_h \right\}$$

$$\int_1^0 \nabla_\eta \cdot \left\{ \quad \right\} d\eta \quad \rightarrow \quad \frac{\partial D_h}{\partial t} + \ldots = \gamma_e \nabla^2 D_h$$

Continuity equation: $\frac{\partial \mu}{\partial t} = -\nabla_\eta \cdot \mu \mathbf{V}_h - \frac{\partial \mu \dot{\eta}}{\partial \eta} = D_h$

$$\boxed{\frac{\partial \mu \mathbf{V}_h}{\partial \tau} + \ldots = -\gamma_e \frac{\Delta x^2}{\Delta \tau^2} \nabla_\eta (\mu^\tau - \mu^{\tau - \Delta\tau})}$$

γ_e = 0.01 recommended (default) [&dynamics *emdiv*]

(Primarily for real-data applications)

Purpose: damp anomalously-large vertical velocities

(usually associated with anomalous physics tendencies)

Additional term:

$$\partial_t W = \ldots \underline{- \mu_d \, \mathrm{sign}(W) \gamma_w (Cr - Cr_\beta)}$$

$$Cr = \left| \frac{\Omega dt}{\mu d\eta} \right|$$

Cr_β= 1.0 typical value (default)

[share/module_model_constants.F *w_beta*]

γ_w = 0.3 m/s^2 recommended (default)

[share/module_model_constants.F *w_alpha*]

[&dynamics *w_damping* 0 (off; default) 1 (on)]

2nd-Order Horizontal Mixing, Horizontal-Deformation-Based K_h

Purpose: mixing on horizontal coordinate surfaces (real-data applications) [&dynamics *diff_opt=1, km_opt=4*]

$$K_h = C_s^2\, l^2 \left[0.25(D_{11} - D_{22})^2 + \overline{D_{12}^2}^{xy} \right]^{\frac{1}{2}}$$

where $l = (\Delta x \Delta y)^{1/2}$

$$D_{11} = 2\, m^2 \left[\partial_x (m^{-1} u) - z_x \partial_z (m^{-1} u) \right]$$
$$D_{22} = 2\, m^2 \left[\partial_y (m^{-1} v) - z_y \partial_z (m^{-1} v) \right]$$
$$D_{12} = m^2 \left[\partial_y (m^{-1} u) - z_y \partial_z (m^{-1} u) + \partial_x (m^{-1} v) - z_x \partial_z (m^{-1} v) \right]$$

$C_s = 0.25$ (Smagorinsky coefficient, default value) [&dynamics *c_s*]

Implicit Rayleigh w Damping Layer for Split-Explicit Nonhydrostatic NWP Models (gravity-wave absorbing layer)

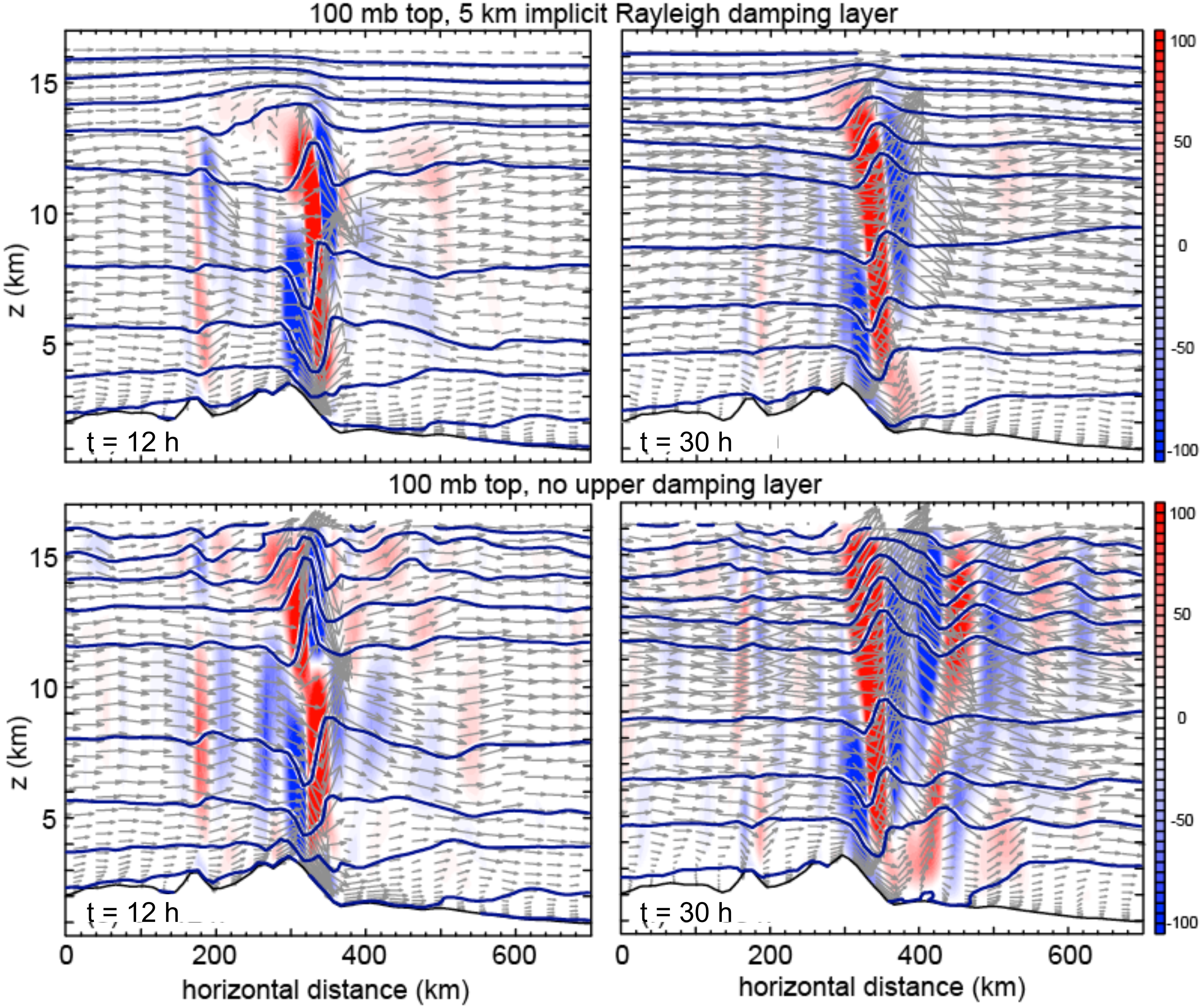
$$W^{\tau+\Delta\tau} = W^{*\tau+\Delta\tau} - \Delta\tau R_w(\eta) W^{\tau+\Delta\tau}$$

$$R_w(\eta) = \begin{cases} \gamma_r \sin^2\left[\frac{\pi}{2}\left(1 - \frac{z_{top}-z}{z_d}\right)\right] & \text{for } z \geq (z_{top} - z_d); \\ 0 & \text{otherwise,} \end{cases}$$

$R_w(\eta)$- damping rate (t^{-1})
z_d - depth of the damping layer
γ_r - damping coefficient

[&dynamics *damp_opt = 3* (default = 0)]
[&dynamics *damp_coef = 0.2* (recommended, = 0. default)]
[&dynamics *zdamp = 5000.* (z_d *(meters)*; default); height below model top where damping begins]

Model Initialized 04 Dec 2007 00 UTC



w (cm/s)

w (cm/s)

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ARW Model: projection options

1. Cartesian geometry:
   idealized cases
2. Lambert Conformal:
   mid-latitude applications
3. Polar Stereographic:
   high-latitude applications
4. Mercator:
   low-latitude applications
5. Latitude-Longitude global, regional

Projections 1-4 are isotropic ($m_x = m_y$)
Latitude-longitude projection is anistropic ($m_x \neq m_y$)

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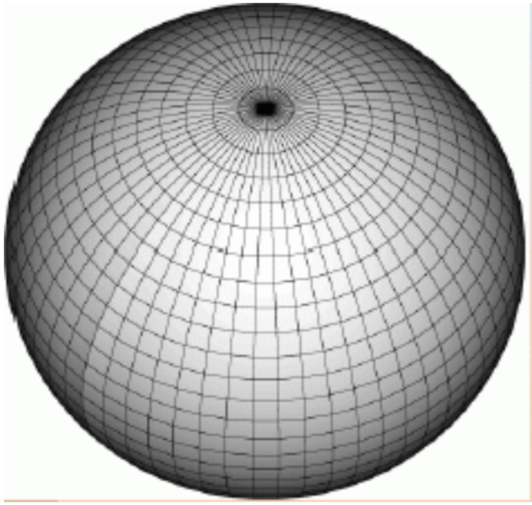
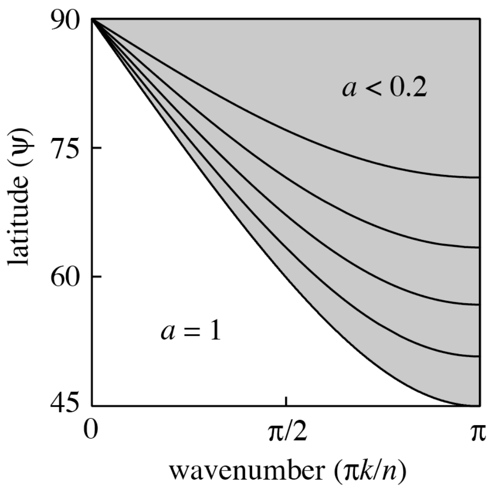
Global ARW – Polar filters



Converging gridlines severely limit timestep.
The polar filter removes this limitation.

Filter procedure - Along a grid latitude circle:
1. Fourier transform variable.
2. Filter Fourier coefficients.
3. Transform back to physical space.

Filter Coefficient $a(k)$, $\psi_o = 45^o$



$$\hat{\phi}(k)_{filtered} = a(k)\,\hat{\phi}(k), \quad \text{for all } k$$

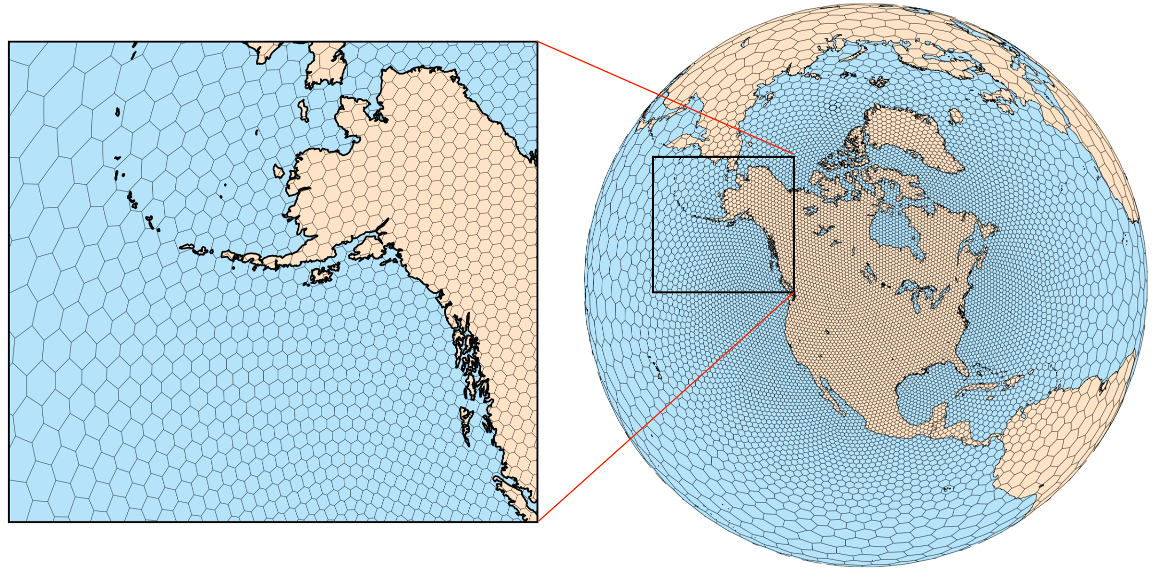
$$a(k) = \min\left[1., \max\left(0., \left(\frac{\cos\psi}{\cos\psi_o}\right)^2 \frac{1}{\sin^2(\pi k/n)}\right)\right]$$

$k =$ dimensionless wavenumber
$\hat{\phi}(k) =$ Fourier coefficients from forward transform
$a(k) =$ filter coefficients
$\psi =$ latitude $\psi_o =$ polar filter latitude, filter when $|\psi| > \psi_o$

Bill Skamarock (skamaroc@ucar.edu)

An alternative to global ARW…

MPAS
Model for Prediction Across Scales



- Global, nonhydrostatic, C-grid Voronoi mesh
- Numerics similar to WRF; WRF-NRCM physics
- No pole problems
- Variable-resolution mesh – no nested BC problems

Available at: http://mpas-dev.github.io/

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ARW Model: Boundary Condition Options

Lateral boundary conditions

1. Specified (Coarse grid, real-data applications).
2. Open lateral boundaries (gravity-wave radiative).
3. Symmetric lateral boundary condition (free-slip wall).
4. Periodic lateral boundary conditions.
5. Nested boundary conditions (specified).

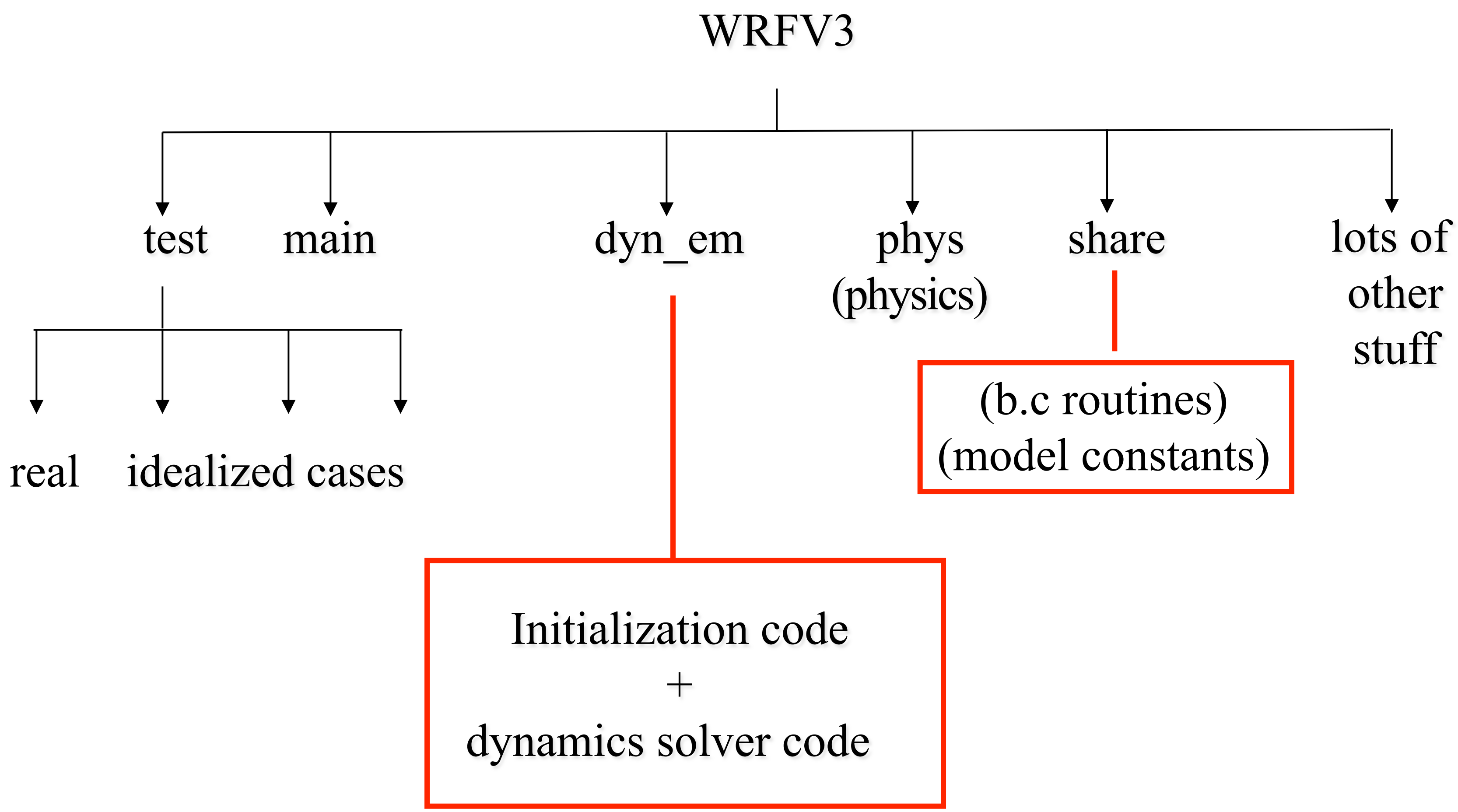
Top boundary conditions

1. Constant pressure.

Bottom boundary conditions

1. Free slip.
2. Various B.L. implementations of surface drag, fluxes.

Dynamics: Where are things?



WRF ARW Tech Note
A Description of the Advanced Research WRF Version 3 (June 2008, 2012 update)
http://www.mmm.ucar.edu/wrf/users/pub-doc.html