

## Dynamics: Introduction

# The Advanced Research WRF (ARW) Dynamics Solver

1. What is a dynamics solver?
2. Variables and coordinates
3. Equations
4. Time integration scheme
5. Grid staggering
6. Advection (transport) and conservation
7. Time step parameters
8. Filters
9. Map projections and global configuration
10. Boundary condition options

### WRF ARW Tech Note

A Description of the Advanced Research WRF Version 3 (June 2008, 2012 update)

<http://www.mmm.ucar.edu/wrf/users/pub-doc.html>

We are in the process of updating the technical note to Version 4

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## Dynamics: 2. Variables and coordinates

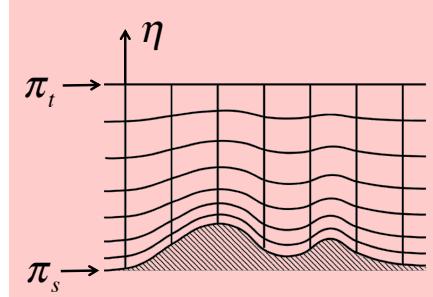
### Vertical coordinates: (1) Traditional terrain-following mass coordinate

$$\text{Dry hydrostatic pressure } \pi_d$$

$$\text{Column mass (per unit area)} \mu_d = \pi_s - \pi_t$$

$$\text{Vertical coordinate } \eta = \frac{(\pi_d - \pi_t)}{\mu_d}$$

$$\text{Layer mass (per unit area)} \mu_d \Delta \eta = \Delta \pi_d = -g \rho_d \Delta z, \quad \text{Pressure } \pi_d(\eta) = \eta \mu_d + \pi_t$$



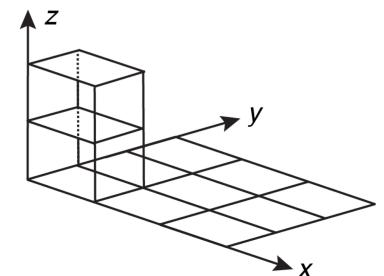
## Dynamics: 1. What is a *dynamics solver*?

A *dynamical solver* (or a *dynamical core*, or *dycore*) performs a time ( $t$ ) and space ( $x,y,z$ ) integration of the equations of motion.

Given the 3D atmospheric state at time  $t$ ,  $S(x,y,z,t)$ , we integrate the equations forward in time from  $t \rightarrow T$ , i.e. we run the model and produce a forecast.

The equations cannot be solved analytically, so we *discretize* the equations on a *grid* and compute *approximate* solutions.

The accuracy of the solutions depend on the numerical method and the mesh spacing (grid).



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## Dynamics: 2. Variables and coordinates

### Vertical coordinates: (2) Hybrid terrain-following mass coordinate

Isobaric coordinate (constant pressure):

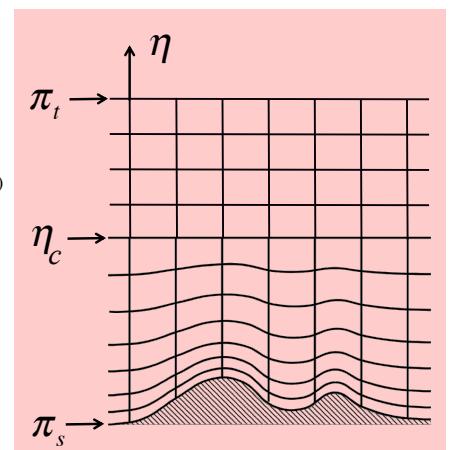
$$\eta = \frac{\pi_d}{\pi_0 - \pi_t}$$

Hybrid terrain-following coordinate:

$$\pi_d(\eta) = B(\eta) \mu_d + \pi_t \quad (\text{Terrain-following}) \\ + [\eta - B(\eta)] (\pi_0 - \pi_t) \quad (\text{Isobaric})$$

level at which  $B \rightarrow 0$ ,  
i.e. transition between  
isobaric and  
terrain-following coordinate.

Default WRFV4 configuration:  
Hybrid coordinate is enabled,  $\eta_c = 0.2$



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## Dynamics: 2. Variables and coordinates

Variables:

Grid volume mass (per unit area):

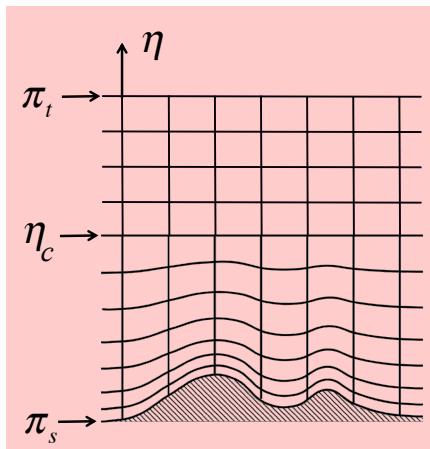
$$\mu_d = \frac{\partial \pi_d}{\partial \eta} = B_\eta (\pi_s - \pi_t) + (1 - B_\eta)(\pi_0 - \pi_t)$$

Conserved state (prognostic) variables:

$$\begin{aligned} \mu_d, \quad U &= \mu_d u, \quad V = \mu_d v, \\ W &= \mu_d w, \quad \Theta = \mu_d \theta \end{aligned}$$

Non-conserved state variable:

$$\phi = gz$$



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## Dynamics: 3. Equations

$$\begin{aligned} \frac{\partial U}{\partial t} &= \\ \frac{\partial V}{\partial t} &= \\ \frac{\partial W}{\partial t} &= \\ \frac{\partial \mu_d}{\partial t} &= \\ \frac{\partial \Theta}{\partial t} &= \\ \frac{\partial \mu_d q_j}{\partial t} &= \\ \frac{\partial \phi}{\partial t} &= \end{aligned}$$

## Dynamics: 2. Variables and coordinates

$$\text{Vertical momentum eqn.} \quad \frac{\partial W}{\partial t} + g \left( \mu_d - \frac{\alpha}{\alpha_d} \frac{\partial p}{\partial \eta} \right) = - \frac{\partial U w}{\partial x} - \frac{\partial \Omega w}{\partial \eta}$$

Subscript  $d$  denotes dry, and

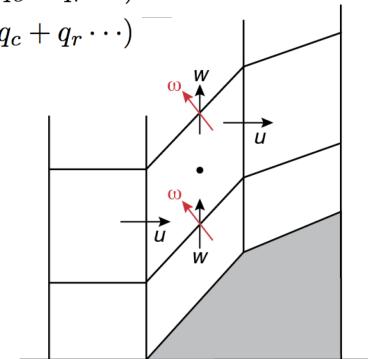
$$\alpha_d = \frac{1}{\rho_d} \quad \alpha = \alpha_d (1 + q_v + q_c + q_r \dots)^{-1}$$

$$\rho = \rho_d (1 + q_v + q_c + q_r \dots)$$

covariant ( $u, \omega$ ) and  
contravariant  $w$  velocities

$$u = \frac{dx}{dt}, \quad w = \frac{dz}{dt}, \quad \omega = \frac{d\eta}{dt}$$

$$U = \mu u, \quad W \mu w, \quad \Omega = \mu \omega$$



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## Dynamics: 3. Equations

transport

$$\begin{aligned} \frac{\partial U}{\partial t} &= - \frac{\partial U u}{\partial x} - \frac{\partial V u}{\partial y} - \frac{\partial \Omega u}{\partial \eta} \\ \frac{\partial V}{\partial t} &= - \frac{\partial U v}{\partial x} - \frac{\partial V v}{\partial y} - \frac{\partial \Omega v}{\partial \eta} \\ \frac{\partial W}{\partial t} &= - \frac{\partial U w}{\partial x} - \frac{\partial V w}{\partial y} - \frac{\partial \Omega w}{\partial \eta} \\ \frac{\partial \mu_d}{\partial t} &= - \frac{\partial U}{\partial x} - \frac{\partial V}{\partial y} - \frac{\partial \Omega}{\partial \eta} \\ \frac{\partial \Theta}{\partial t} &= - \frac{\partial U \theta}{\partial x} - \frac{\partial V \theta}{\partial y} - \frac{\partial \Omega \theta}{\partial \eta} \\ \frac{\partial \mu_d q_j}{\partial t} &= - \frac{\partial U q_j}{\partial x} - \frac{\partial V q_j}{\partial y} - \frac{\partial \Omega q_j}{\partial \eta} \\ \frac{\partial \phi}{\partial t} &= - u \frac{\partial \phi}{\partial x} - v \frac{\partial \phi}{\partial y} - \omega \frac{\partial \phi}{\partial \eta} \end{aligned}$$

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## Dynamics: 3. Equations

### transport      pressure gradient

$$\begin{aligned}
 \frac{\partial U}{\partial t} &= -\frac{\partial U u}{\partial x} - \frac{\partial V u}{\partial y} - \frac{\partial \Omega u}{\partial \eta} - \alpha \mu_d \frac{\partial p}{\partial x} - \frac{\alpha}{\alpha_d} \frac{\partial p}{\partial \eta} \frac{\partial \phi}{\partial x} \\
 \frac{\partial V}{\partial t} &= -\frac{\partial U v}{\partial x} - \frac{\partial V v}{\partial y} - \frac{\partial \Omega v}{\partial \eta} - \alpha \mu_d \frac{\partial p}{\partial y} - \frac{\alpha}{\alpha_d} \frac{\partial p}{\partial \eta} \frac{\partial \phi}{\partial y} \\
 \frac{\partial W}{\partial t} &= -\frac{\partial U w}{\partial x} - \frac{\partial V w}{\partial y} - \frac{\partial \Omega w}{\partial \eta} - g \left( \mu_d - \frac{\alpha}{\alpha_d} \frac{\partial p}{\partial \eta} \right) \\
 \frac{\partial \mu_d}{\partial t} &= -\frac{\partial U}{\partial x} - \frac{\partial V}{\partial y} - \frac{\partial \Omega}{\partial \eta} \\
 \frac{\partial \Theta}{\partial t} &= -\frac{\partial U \theta}{\partial x} - \frac{\partial V \theta}{\partial y} - \frac{\partial \Omega \theta}{\partial \eta} \\
 \frac{\partial \mu_d q_j}{\partial t} &= -\frac{\partial U q_j}{\partial x} - \frac{\partial V q_j}{\partial y} - \frac{\partial \Omega q_j}{\partial \eta} \\
 \frac{\partial \phi}{\partial t} &= -u \frac{\partial \phi}{\partial x} - v \frac{\partial \phi}{\partial y} - \omega \frac{\partial \phi}{\partial \eta}
 \end{aligned}$$

## Dynamics: 3. Equations

### transport      pressure gradient

$$\begin{aligned}
 \frac{\partial U}{\partial t} &= -\frac{\partial U u}{\partial x} - \frac{\partial V u}{\partial y} - \frac{\partial \Omega u}{\partial \eta} - \alpha \mu_d \frac{\partial p}{\partial x} - \frac{\alpha}{\alpha_d} \frac{\partial p}{\partial \eta} \frac{\partial \phi}{\partial x} + R_u + Q_u \\
 \frac{\partial V}{\partial t} &= -\frac{\partial U v}{\partial x} - \frac{\partial V v}{\partial y} - \frac{\partial \Omega v}{\partial \eta} - \alpha \mu_d \frac{\partial p}{\partial y} - \frac{\alpha}{\alpha_d} \frac{\partial p}{\partial \eta} \frac{\partial \phi}{\partial y} + R_v + Q_v \\
 \frac{\partial W}{\partial t} &= -\frac{\partial U w}{\partial x} - \frac{\partial V w}{\partial y} - \frac{\partial \Omega w}{\partial \eta} - g \left( \mu_d - \frac{\alpha}{\alpha_d} \frac{\partial p}{\partial \eta} \right) + R_w + Q_w \\
 \frac{\partial \mu_d}{\partial t} &= -\frac{\partial U}{\partial x} - \frac{\partial V}{\partial y} - \frac{\partial \Omega}{\partial \eta} \\
 \frac{\partial \Theta}{\partial t} &= -\frac{\partial U \theta}{\partial x} - \frac{\partial V \theta}{\partial y} - \frac{\partial \Omega \theta}{\partial \eta} + R_\theta + Q_\theta \\
 \frac{\partial \mu_d q_j}{\partial t} &= -\frac{\partial U q_j}{\partial x} - \frac{\partial V q_j}{\partial y} - \frac{\partial \Omega q_j}{\partial \eta} + R_{q_j} + Q_{q_j} \\
 \frac{\partial \phi}{\partial t} &= -u \frac{\partial \phi}{\partial x} - v \frac{\partial \phi}{\partial y} - \omega \frac{\partial \phi}{\partial \eta} + gw
 \end{aligned}$$

↑  
numerical filters,  
physics,  
projection terms  
← geopotential eqn term

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## Dynamics: 3. Equations

### transport      pressure gradient

$$\begin{aligned}
 \frac{\partial U}{\partial t} &= -\frac{\partial U u}{\partial x} - \frac{\partial V u}{\partial y} - \frac{\partial \Omega u}{\partial \eta} - \alpha \mu_d \frac{\partial p}{\partial x} - \frac{\alpha}{\alpha_d} \frac{\partial p}{\partial \eta} \frac{\partial \phi}{\partial x} + R_u + Q_u \\
 \frac{\partial V}{\partial t} &= -\frac{\partial U v}{\partial x} - \frac{\partial V v}{\partial y} - \frac{\partial \Omega v}{\partial \eta} - \alpha \mu_d \frac{\partial p}{\partial y} - \frac{\alpha}{\alpha_d} \frac{\partial p}{\partial \eta} \frac{\partial \phi}{\partial y} + R_v + Q_v \\
 \frac{\partial W}{\partial t} &= -\frac{\partial U w}{\partial x} - \frac{\partial V w}{\partial y} - \frac{\partial \Omega w}{\partial \eta} - g \left( \mu_d - \frac{\alpha}{\alpha_d} \frac{\partial p}{\partial \eta} \right) + R_w + Q_w \\
 \frac{\partial \mu_d}{\partial t} &= -\frac{\partial U}{\partial x} - \frac{\partial V}{\partial y} - \frac{\partial \Omega}{\partial \eta} \\
 \frac{\partial \Theta}{\partial t} &= -\frac{\partial U \theta}{\partial x} - \frac{\partial V \theta}{\partial y} - \frac{\partial \Omega \theta}{\partial \eta} + R_\theta + Q_\theta \\
 \frac{\partial \mu_d q_j}{\partial t} &= -\frac{\partial U q_j}{\partial x} - \frac{\partial V q_j}{\partial y} - \frac{\partial \Omega q_j}{\partial \eta} + R_{q_j} + Q_{q_j} \\
 \frac{\partial \phi}{\partial t} &= -u \frac{\partial \phi}{\partial x} - v \frac{\partial \phi}{\partial y} - \omega \frac{\partial \phi}{\partial \eta} + gw
 \end{aligned}$$

↑  
numerical filters,  
physics,  
projection terms  
← geopotential eqn term

$$\text{Diagnostic relations: } \frac{\partial \phi}{\partial \eta} = -\alpha_d \mu_d, p = \left( \frac{R_d \Theta_m}{p_o \mu_d \alpha_d} \right)^{\gamma}, \Theta_m = \Theta \left( 1 + \frac{R_v}{R_d} q_v \right)$$

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## Dynamics: 4. Time integration scheme

### 3<sup>rd</sup> Order Runge-Kutta time integration

$$\begin{aligned}
 \frac{\partial U}{\partial t} &= RHS_u && \text{advance } \phi^t \rightarrow \phi^{t+\Delta t} \\
 \frac{\partial V}{\partial t} &= RHS_v && \phi^* = \phi^t + \frac{\Delta t}{3} RHS(\phi^t) \\
 \frac{\partial W}{\partial t} &= RHS_w && \phi^{**} = \phi^t + \frac{\Delta t}{2} RHS(\phi^*) \\
 &\vdots && \phi^{t+\Delta t} = \phi^t + \Delta t RHS(\phi^{**})
 \end{aligned}$$

$$\text{Amplification factor } \phi_t = ik\phi; \quad \phi^{n+1} = A\phi^n; \quad |A| = 1 - \frac{(k\Delta t)^4}{24}$$

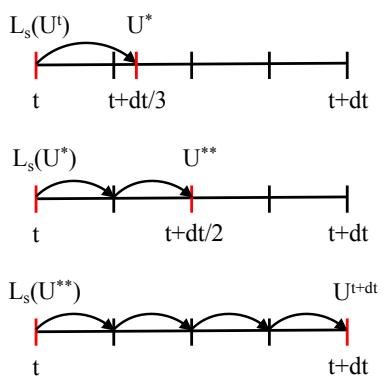
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## Dynamics: 4. Time integration scheme – time splitting

$$U_t = L_{\text{fast}}(U) + L_{\text{slow}}(U)$$

3rd order Runge-Kutta, 3 steps



*fast*: acoustic and gravity wave terms.

*slow*: everything else.

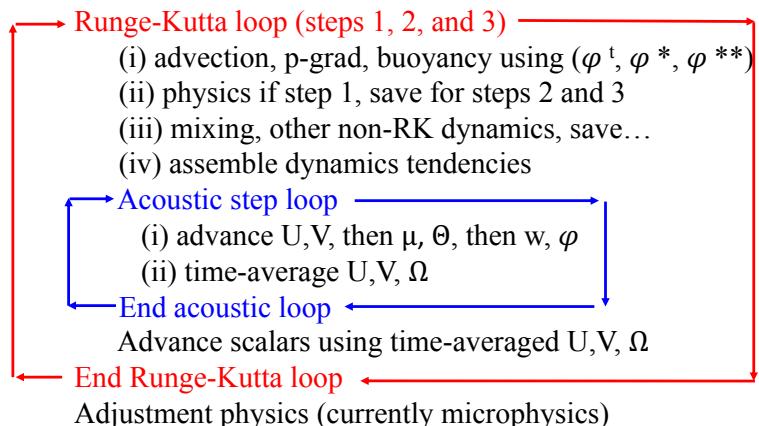
- RK3 is 3rd order accurate for linear eqns, 2nd order accurate for nonlinear eqns.
- Stable for centered and upwind advection schemes.
- Stable for Courant number  $Udt/dx < 1.73$
- Three  $L_{\text{slow}}(U)$  evaluations per timestep.

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## Dynamics: 4. Time integration scheme - implementation

Begin time step



End time step

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## Dynamics: 4. Time integration scheme – acoustic step

$$U^{t+\Delta t}, \quad V^{t+\Delta t}$$

$$\mu_d^{\tau+\Delta\tau} \quad \Omega^{\tau+\Delta\tau}$$

$$\Theta^{\tau+\Delta\tau}$$

$$W^{\tau+\Delta\tau}$$

$$\phi^{\tau+\Delta\tau}$$

$$\frac{\partial U}{\partial t} + \left( \mu_d \alpha \frac{\partial p}{\partial x} + \frac{\alpha}{\alpha_d} \frac{\partial p}{\partial \eta} \frac{\partial \phi}{\partial x} \right)^{\tau} = R_U^t$$

$$\frac{\partial \mu_d}{\partial t} + \frac{\partial U^{\tau+\Delta\tau}}{\partial x} + \frac{\partial \Omega^{\tau+\Delta\tau}}{\partial \eta} = 0$$

$$\frac{\partial \Theta}{\partial t} + \left( \frac{\partial U \theta^t}{\partial x} + \frac{\partial \Omega \theta^t}{\partial \eta} \right)^{\tau+\Delta\tau} = R_\Theta^t$$

$$\frac{\partial W}{\partial t} + g \left( \mu_d - \frac{\alpha}{\alpha_d} \frac{\partial p}{\partial \eta} \right)^{\tau} = R_W^t$$

$$\mu_d^t \frac{\partial \phi}{\partial t} + U^{\tau+\Delta\tau} \frac{\partial \phi^t}{\partial x} + \Omega^{\tau+\Delta\tau} \frac{\partial \phi^t}{\partial \eta} - g \bar{W}^\tau = R_\phi^t$$

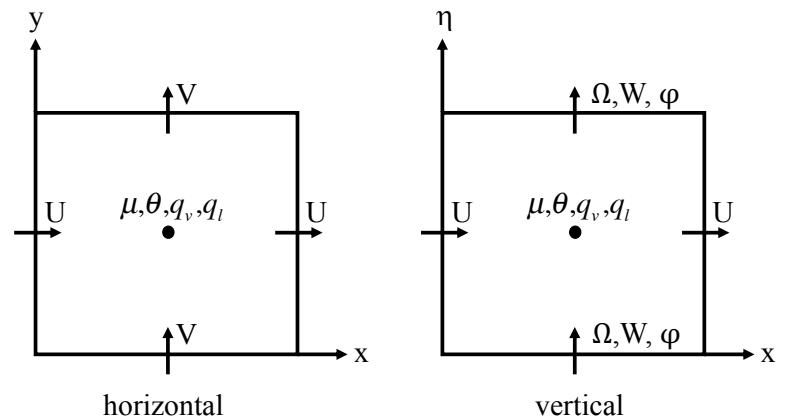
- Forward-backward differencing on  $U$ ,  $\Theta$ , and  $\mu$  equations
- Vertically implicit differencing on  $W$  and  $\phi$  equations

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## Dynamics: 5. Grid staggering – horizontal and vertical

C-grid staggering



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## Dynamics: 6. Advection (transport) and conservation – dry-air mass

transport      pressure gradient

$$\begin{aligned}
 \frac{\partial U}{\partial t} &= -\frac{\partial U u}{\partial x} - \frac{\partial V u}{\partial y} - \frac{\partial \Omega u}{\partial \eta} - \alpha \mu_d \frac{\partial p}{\partial x} - \frac{\alpha}{\alpha_d} \frac{\partial p}{\partial \eta} \frac{\partial \phi}{\partial x} + R_u + Q_u \\
 \frac{\partial V}{\partial t} &= -\frac{\partial U v}{\partial x} - \frac{\partial V v}{\partial y} - \frac{\partial \Omega v}{\partial \eta} - \alpha \mu_d \frac{\partial p}{\partial y} - \frac{\alpha}{\alpha_d} \frac{\partial p}{\partial \eta} \frac{\partial \phi}{\partial y} + R_v + Q_v \\
 \frac{\partial W}{\partial t} &= -\frac{\partial U w}{\partial x} - \frac{\partial V w}{\partial y} - \frac{\partial \Omega w}{\partial \eta} - g \left( \mu_d - \frac{\alpha}{\alpha_d} \frac{\partial p}{\partial \eta} \right) + R_w + Q_w \\
 \frac{\partial \mu_d}{\partial t} &= -\frac{\partial U}{\partial x} - \frac{\partial V}{\partial y} - \frac{\partial \Omega}{\partial \eta} \quad \text{Next:} \\
 \frac{\partial \Theta}{\partial t} &= -\frac{\partial U \theta}{\partial x} - \frac{\partial V \theta}{\partial y} - \frac{\partial \Omega \theta}{\partial \eta} + R_\theta + Q_\theta \\
 \frac{\partial \mu_d q_j}{\partial t} &= -\frac{\partial U q_j}{\partial x} - \frac{\partial V q_j}{\partial y} - \frac{\partial \Omega q_j}{\partial \eta} + R_{q_j} + Q_{q_j} \\
 \frac{\partial \phi}{\partial t} &= -u \frac{\partial \phi}{\partial x} - v \frac{\partial \phi}{\partial y} - \omega \frac{\partial \phi}{\partial \eta} + gw
 \end{aligned}$$

Diagnostic relations:  $\frac{\partial \phi}{\partial \eta} = -\alpha_d \mu_d, p = \left( \frac{R_d \Theta_m}{p_o \mu_d \alpha_d} \right)^r, \Theta_m = \Theta \left( 1 + \frac{R_v}{R_d} q_v \right)$

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## Dynamics: 6. Advection (transport) and conservation – dry-air mass

Mass in a control volume  $(\Delta x \Delta \eta)(\mu)^t$   
2D example

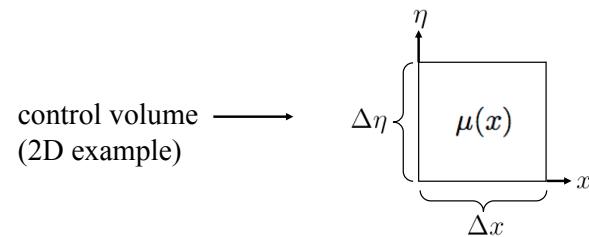
Mass conservation equation

$$\Delta t^{-1}(\Delta x \Delta \eta) \cdot [(\mu)^{t+\Delta t} - (\mu)^t] = [(\mu u \Delta \eta)_{x-\Delta x/2, \eta} - (\mu u \Delta \eta)_{x+\Delta x/2, \eta}] + [(\mu \omega \Delta x)_{x, \eta-\Delta \eta/2} - (\mu \omega \Delta x)_{x, \eta+\Delta \eta/2}]$$

Change in mass over a time step

mass fluxes through  
control volume faces

## Dynamics: 6. Advection (transport) and conservation – dry-air mass



Mass in a control volume is proportional to

$$(\Delta x \Delta \eta)(\mu)^t$$

since  $\mu(x) \Delta \eta = \Delta \pi = -g \rho \Delta z$

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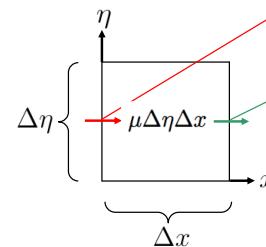
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## Dynamics: 6. Advection (transport) and conservation – dry-air mass

Mass in a control volume  $(\Delta x \Delta \eta)(\mu)^t$

Mass conservation equation

$$\Delta t^{-1}(\Delta x \Delta \eta) \cdot [(\mu)^{t+\Delta t} - (\mu)^t] = [(\mu u \Delta \eta)_{x-\Delta x/2, \eta} - (\mu u \Delta \eta)_{x+\Delta x/2, \eta}] + [(\mu \omega \Delta x)_{x, \eta-\Delta \eta/2} - (\mu \omega \Delta x)_{x, \eta+\Delta \eta/2}]$$



Horizontal fluxes through the  
vertical control-volume faces

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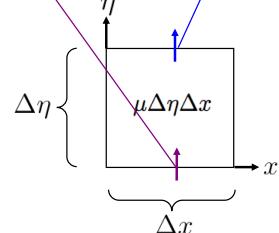
## Dynamics: 6. Advection (transport) and conservation – dry-air mass

Mass in a control volume  $(\Delta x \Delta \eta)(\mu)^t$

Mass conservation equation

$$\Delta t^{-1}(\Delta x \Delta \eta) \cdot [(\mu)^{t+\Delta t} - (\mu)^t] = [(\mu u \Delta \eta)_{x-\Delta x/2, \eta} - (\mu u \Delta \eta)_{x+\Delta x/2, \eta}] + [(\mu \omega \Delta x)_{x, \eta-\Delta \eta/2} - (\mu \omega \Delta x)_{x, \eta+\Delta \eta/2}]$$

Vertical fluxes through the horizontal control-volume faces

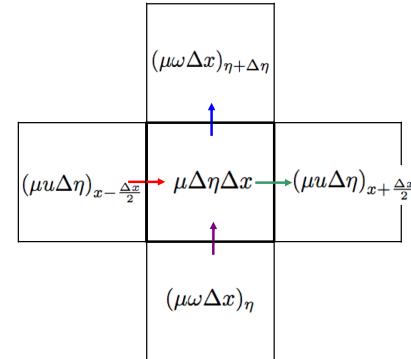


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## Dynamics: 6. Advection (transport) and conservation – dry-air mass

The same mass fluxes are used for neighboring grid cells - hence mass is conserved locally and globally.



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## Dynamics: 6. Advection (transport) and conservation

### transport

$$\begin{aligned} \frac{\partial U}{\partial t} &= -\frac{\partial U u}{\partial x} - \frac{\partial V u}{\partial y} - \frac{\partial W u}{\partial \eta} - \alpha \mu_d \frac{\partial p}{\partial x} - \frac{\alpha}{\alpha_d} \frac{\partial p}{\partial \eta} \frac{\partial \phi}{\partial x} + R_u + Q_u \\ \frac{\partial V}{\partial t} &= -\frac{\partial U v}{\partial x} - \frac{\partial V v}{\partial y} - \frac{\partial W v}{\partial \eta} - \alpha \mu_d \frac{\partial p}{\partial y} - \frac{\alpha}{\alpha_d} \frac{\partial p}{\partial \eta} \frac{\partial \phi}{\partial y} + R_v + Q_v \\ \frac{\partial W}{\partial t} &= -\frac{\partial U w}{\partial x} - \frac{\partial V w}{\partial y} - \frac{\partial \Omega w}{\partial \eta} - g \left( \mu_d - \frac{\alpha}{\alpha_d} \frac{\partial p}{\partial \eta} \right) + R_w + Q_w \\ \frac{\partial \mu_d}{\partial t} &= -\frac{\partial U}{\partial x} - \frac{\partial V}{\partial y} - \frac{\partial \Omega}{\partial \eta} \\ \frac{\partial \Theta}{\partial t} &= -\frac{\partial U \theta}{\partial x} - \frac{\partial V \theta}{\partial y} - \frac{\partial \Omega \theta}{\partial \eta} + R_\theta + Q_\theta \\ \frac{\partial \mu_d q_j}{\partial t} &= -\frac{\partial U q_j}{\partial x} - \frac{\partial V q_j}{\partial y} - \frac{\partial \Omega q_j}{\partial \eta} + R_{q_j} + Q_{q_j} \\ \frac{\partial \phi}{\partial t} &= -u \frac{\partial \phi}{\partial x} - v \frac{\partial \phi}{\partial y} - \omega \frac{\partial \phi}{\partial \eta} + gw \end{aligned}$$

Entropy and scalar mass conservation in WRF

$$\text{Diagnostic relations: } \frac{\partial \phi}{\partial \eta} = -\alpha_d \mu_d, p = \left( \frac{R_d \Theta_m}{p_o \mu_d \alpha_d} \right)^{\gamma}, \Theta_m = \Theta \left( 1 + \frac{R_v}{R_d} q_v \right)$$

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## Dynamics: 6. Advection (transport) and conservation – scalars

Mass in a control volume  $(\Delta x \Delta \eta)(\mu)^t$

Scalar mass  $(\Delta x \Delta \eta)(\mu \phi)^t$

Mass conservation equation:

$$\Delta t^{-1}(\Delta x \Delta \eta) \cdot [(\mu \phi)^{t+\Delta t} - (\mu \phi)^t] = [(\mu u \Delta \eta)_{x-\Delta x/2, \eta} - (\mu u \Delta \eta)_{x+\Delta x/2, \eta}] + [(\mu \omega \Delta x)_{x, \eta-\Delta \eta/2} - (\mu \omega \Delta x)_{x, \eta+\Delta \eta/2}]$$

↑ change in mass over a time step

mass fluxes through control volume faces

Scalar mass conservation equation:

$$\Delta t^{-1}(\Delta x \Delta \eta) \cdot [(\mu \phi)^{t+\Delta t} - (\mu \phi)^t] = [(\mu u \phi \Delta \eta)_{x-\Delta x/2, \eta} - (\mu u \phi \Delta \eta)_{x+\Delta x/2, \eta}] + [(\mu \omega \phi \Delta x)_{x, \eta-\Delta \eta/2} - (\mu \omega \phi \Delta x)_{x, \eta+\Delta \eta/2}]$$

↑ change in tracer mass over a time step

tracer mass fluxes through control volume faces

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## Dynamics: 6. Advection (transport) and conservation

transport	pressure gradient
$\frac{\partial U}{\partial t} = -\frac{\partial U u}{\partial x} - \frac{\partial V u}{\partial y} - \frac{\partial \Omega u}{\partial \eta}$	$-\alpha \mu_d \frac{\partial p}{\partial x} - \frac{\alpha}{\alpha_d} \frac{\partial p}{\partial \eta} \frac{\partial \phi}{\partial x} + R_u + Q_u$
$\frac{\partial V}{\partial t} = -\frac{\partial U v}{\partial x} - \frac{\partial V v}{\partial y} - \frac{\partial \Omega v}{\partial \eta}$	$-\alpha \mu_d \frac{\partial p}{\partial y} - \frac{\alpha}{\alpha_d} \frac{\partial p}{\partial \eta} \frac{\partial \phi}{\partial y} + R_v + Q_v$
$\frac{\partial W}{\partial t} = -\frac{\partial U w}{\partial x} - \frac{\partial V w}{\partial y} - \frac{\partial \Omega w}{\partial \eta}$	$-g \left( \mu_d - \frac{\alpha}{\alpha_d} \frac{\partial p}{\partial \eta} \right) + R_w + Q_w$
$\frac{\partial \mu_d}{\partial t} = -\frac{\partial U}{\partial x} - \frac{\partial V}{\partial y} - \frac{\partial \Omega}{\partial \eta}$	
$\frac{\partial \Theta}{\partial t} = -\frac{\partial U \theta}{\partial x} - \frac{\partial V \theta}{\partial y} - \frac{\partial \Omega \theta}{\partial \eta}$	$+ R_\theta + Q_\theta$
$\frac{\partial \mu_d q_j}{\partial t} = -\frac{\partial U q_j}{\partial x} - \frac{\partial V q_j}{\partial y} - \frac{\partial \Omega q_j}{\partial \eta}$	$+ R_{q_j} + Q_{q_j}$
$\frac{\partial \phi}{\partial t} = -u \frac{\partial \phi}{\partial x} - v \frac{\partial \phi}{\partial y} - \omega \frac{\partial \phi}{\partial \eta}$	$+ gw$

Diagnostic relations:  $\frac{\partial \phi}{\partial \eta} = -\alpha_d \mu_d, p = \left( \frac{R_d \Theta_m}{p_o \mu_d \alpha_d} \right)^r, \Theta_m = \Theta \left( 1 + \frac{R_v}{R_d} q_v \right)$

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Transport  
schemes: flux  
divergence  
(transport) options  
in WRF

## Dynamics: 6. Advection (transport) and conservation

2<sup>nd</sup>, 3<sup>rd</sup>, 4<sup>th</sup>, 5<sup>th</sup> and 6<sup>th</sup> order centered and upwind-biased schemes are available in the ARW model.

Example: 5<sup>th</sup> order scheme

$$\frac{\partial(U\psi)}{\partial x} = \frac{1}{\Delta x} \left( F_{i+\frac{1}{2}}(U\psi) - F_{i-\frac{1}{2}}(U\psi) \right)$$

where

$$F_{i-\frac{1}{2}}(U\psi) = U_{i-\frac{1}{2}} \left\{ \frac{37}{60}(\psi_i + \psi_{i-1}) - \frac{2}{15}(\psi_{i+1} + \psi_{i-2}) + \frac{1}{60}(\psi_{i+2} + \psi_{i-3}) \right\} \\ - sign(1,U) \frac{1}{60} \left\{ (\psi_{i+2} - \psi_{i-3}) - 5(\psi_{i+1} - \psi_{i-2}) + 10(\psi_i - \psi_{i-1}) \right\}$$

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## Dynamics: 6. Advection (transport) and conservation

For constant U, the 5<sup>th</sup> order flux divergence tendency becomes

$$\Delta t \frac{\delta(U\psi)}{\Delta x} \Big|_{5th} = \Delta t \frac{\delta(U\psi)}{\Delta x} \Big|_{6th} \\ - \underbrace{\left[ \frac{U \Delta t}{\Delta x} \right] \frac{1}{60} (-\psi_{i-3} + 6\psi_{i-2} - 15\psi_{i-1} + 20\psi_i - 15\psi_{i+1} + 6\psi_{i+2} - \psi_{i+3})}_{Cr \frac{\partial^6 \psi}{\partial x^6} + H.O.T}$$

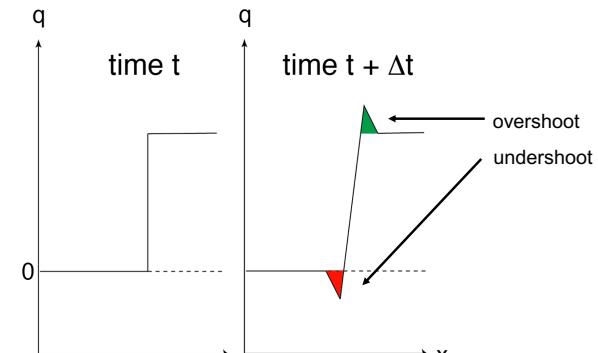
The odd-ordered flux divergence schemes are equivalent to the next higher ordered (even) flux-divergence scheme plus a dissipation term of the higher even order with a coefficient proportional to the Courant number.

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## Dynamics: 6. Advection (transport) and conservation – shape preserving

1D advection



ARW transport is conservative,  
but not positive definite nor monotonic.  
Removal of negative q █  
results in spurious source of q █.

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## Dynamics: 6. Advection (transport) and conservation – shape preserving

Scalar update, last RK3 step

$$(\mu\phi)^{t+\Delta t} = (\mu\phi)^t - \Delta t \sum_{i=1}^n \delta_{x_i}[f_i] \quad (1)$$

(1) Decompose flux:  $f_i = f_i^{upwind} + f_i^c$

(2) Renormalize high-order correction fluxes  $f_i^c$  such that solution is positive definite or monotonic:  $f_i^c = R(f_i^c)$

(3) Update scalar eqn. (1) using  $f_i = f_i^{upwind} + R(f_i^c)$

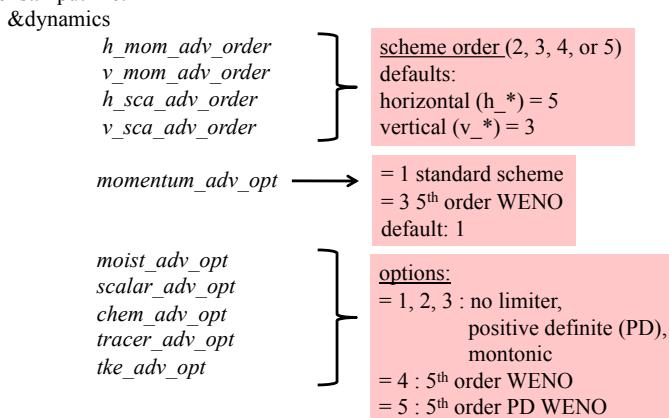
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## Dynamics: 6. Advection (transport) and conservation

Where are the transport-scheme parameters?

The namelist.input file:



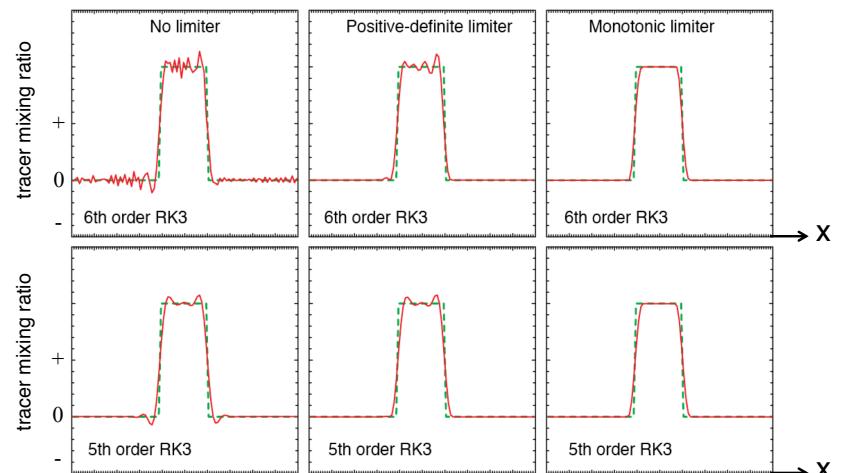
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## Dynamics: 6. Advection (transport) and conservation – examples

### 1D Example: Top-Hat Advection

1D Top-hat transport Cr = 0.5, 1 revolution, 200 steps



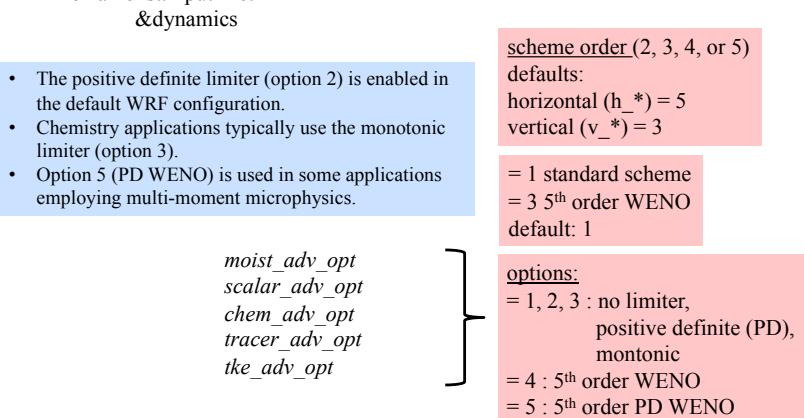
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## Dynamics: 6. Advection (transport) and conservation

Where are the transport-scheme parameters?

The namelist.input file:



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## Dynamics: 7. Time step parameters

3<sup>rd</sup> order Runge-Kutta time step  $\Delta t_{RK}$

Courant number limited, 1D:  $C_r = \frac{U\Delta t}{\Delta x} < 1.43$  (5<sup>th</sup> order adv.)

Generally stable using a timestep approximately twice as large as used in a leapfrog model.

Where?

The namelist.input file:

&domains

*time\_step* (integer seconds)

*time\_step\_fract\_num*

*time\_step\_fract\_den*

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## Dynamics: 7. Time step parameters

3<sup>rd</sup> order Runge-Kutta time step  $\Delta t_{RK}$  (&domains *time\_step*)

Acoustic time step [&dynamics *time\_step\_sound* (integer)]

Guidelines for time step

$\Delta t_{RK}$  in seconds should be about  $6 * \Delta x$  (grid size in kilometers). Larger  $\Delta t$  can be used in smaller-scale dry situations, but *time\_step\_sound* (default = 4) should increase proportionately if larger  $\Delta t$  is used.

If ARW blows up (aborts) quickly, try:

Decreasing  $\Delta t_{RK}$  (that also decreases  $\Delta t_{sound}$ ),

Or increasing *time\_step\_sound* (that decreases  $\Delta t_{sound}$  but does not change  $\Delta t_{RK}$ )

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## Dynamics: 7. Time step parameters

3<sup>rd</sup> order Runge-Kutta time step  $\Delta t_{RK}$  (&domains *time\_step*)

Acoustic time step

2D horizontal Courant number limited:  $C_r = \frac{C_s \Delta \tau}{\Delta h} < \frac{1}{\sqrt{2}}$   
 $\Delta \tau_{sound} = \Delta t_{RK} / (\text{number of acoustic steps})$

Where?

The namelist.input file:

&dynamics

*time\_step\_sound* (integer)

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## Dynamics: 8. Filters – divergence damping

Purpose: filter acoustic modes (3-D divergence,  $D = \nabla \cdot \rho \mathbf{V}$ )

$$\left\{ \frac{\partial \rho \mathbf{V}}{\partial t} + \nabla p + \dots = \gamma'_d \nabla D \right\}$$

$$\nabla \cdot \left\{ \quad \right\} \rightarrow \frac{\partial D}{\partial t} + \nabla^2 p + \dots = \gamma'_d \nabla^2 D$$

From the pressure equation:  $p_t \simeq c^2 D$

$$\frac{\partial \rho \mathbf{V}}{\partial t} + \nabla [p_\tau + \gamma_d (p^\tau - p^{\tau - \Delta \tau})] + \dots = 0$$

$\gamma_d = 0.1$  recommended (default) (&dynamics *smdiv*)

(Illustrated in height coordinates for simplicity)

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## Dynamics: 8. Filters – time off-centering the vertical acoustic modes

Purpose: damp vertically-propagating acoustic modes

$$\frac{\partial W}{\partial t} + g \left( \mu_d - \frac{\alpha}{\alpha_d} \frac{\partial p}{\partial \eta} \right)^\tau = \dots$$

$$\frac{\partial \phi}{\partial t} - \frac{g}{\mu_d^t} \bar{W}^\tau = \dots$$

$$(\overline{\phantom{x}})^\tau = \frac{1+\beta}{2} (\overline{\phantom{x}})^{\tau+\Delta\tau} + \frac{1-\beta}{2} (\overline{\phantom{x}})^\tau$$

Slightly forward centering the vertical pressure gradient damps 3-D divergence as demonstrated for the divergence damper

$\beta = 0.1$  recommended (default) [&dynamics epssm]

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## Dynamics: 8. Filters – vertical velocity damping

Purpose: damp anomalously-large vertical velocities  
(usually associated with anomalous physics tendencies)

Additional term:

$$\partial_t W = \dots - \mu_d \text{ sign}(W) \gamma_w (Cr - Cr_\beta)$$

$$Cr = \left| \frac{\Omega dt}{\mu d\eta} \right|$$

$Cr_\beta = 1.0$  typical value (default)  
[share/module\_model\_constants.F w\_beta]  
 $\gamma_w = 0.3 \text{ m/s}^2$  recommended (default)  
[share/module\_model\_constants.F w\_alpha]  
[&dynamics w\_damping 0 (off; default) 1 (on)]

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## Dynamics: 8. Filters – external mode filter

Purpose: filter the external mode

$$\text{Vertically integrated horizontal divergence, } D_h = \int_1^0 (\nabla_\eta \cdot \mu \mathbf{V}_h) d\eta$$

$$\left\{ \frac{\partial \mu \mathbf{V}_h}{\partial t} + \dots = -\gamma_e \nabla_\eta D_h \right\}$$

$$\int_1^0 \nabla_\eta \cdot \left\{ \dots \right\} d\eta \rightarrow \frac{\partial D_h}{\partial t} + \dots = \gamma_e \nabla^2 D_h$$

$$\text{Continuity equation: } \frac{\partial \mu}{\partial t} = -\nabla_\eta \cdot \mu \mathbf{V}_h - \frac{\partial \mu \dot{\eta}}{\partial \eta} = D_h$$

$$\frac{\partial \mu \mathbf{V}_h}{\partial \tau} + \dots = -\gamma_e \frac{\Delta x^2}{\Delta \tau^2} \nabla_\eta (\mu^\tau - \mu^{\tau-\Delta\tau})$$

$\gamma_e = 0.01$  recommended (default) [&dynamics emdiv]

(Primarily for real-data applications)

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## Dynamics: 8. Filters – 2D Smagorinsky

2nd-Order Horizontal Mixing,  
Horizontal-Deformation-Based  $K_h$

Purpose: mixing on horizontal coordinate surfaces  
(real-data applications) [&dynamics diff\_opt=1, km\_opt=4]

$$K_h = C_s^2 l^2 \left[ 0.25(D_{11} - D_{22})^2 + \overline{D_{12}^{2xy}} \right]^{\frac{1}{2}}$$

$$\text{where } l = (\Delta x \Delta y)^{1/2}$$

$$D_{11} = 2m^2 [\partial_x(m^{-1}u) - z_x \partial_z(m^{-1}u)]$$

$$D_{22} = 2m^2 [\partial_y(m^{-1}v) - z_y \partial_z(m^{-1}v)]$$

$$D_{12} = m^2 [\partial_y(m^{-1}u) - z_y \partial_z(m^{-1}u) + \partial_x(m^{-1}v) - z_x \partial_z(m^{-1}v)]$$

$C_s = 0.25$  (Smagorinsky coefficient, default value)  
[&dynamics c\_s]

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## Dynamics: 8. Filters – gravity-wave absorbing layer

Implicit Rayleigh w Damping Layer for Split-Explicit Nonhydrostatic NWP Models (gravity-wave absorbing layer)

$$W^{\tau+\Delta\tau} = W^{*\tau+\Delta\tau} - \Delta\tau R_w(\eta) W^{\tau+\Delta\tau}$$

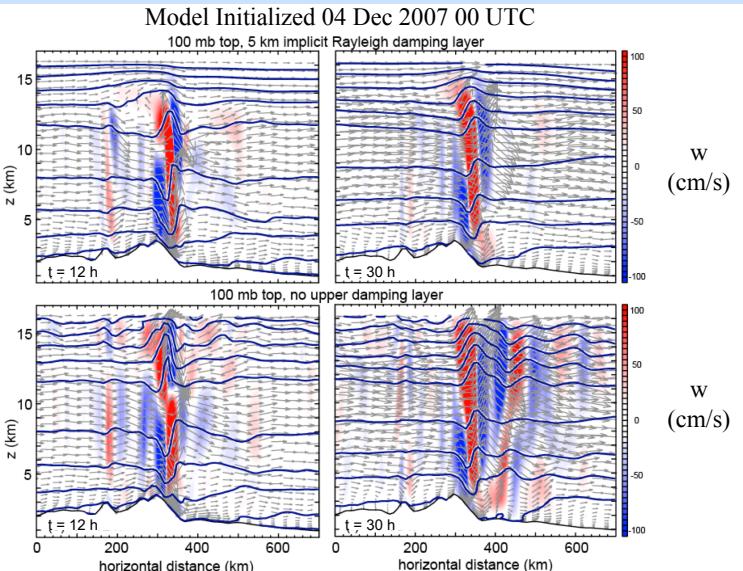
$$R_w(\eta) = \begin{cases} \gamma_r \sin^2 \left[ \frac{\pi}{2} \left( 1 - \frac{z_{top}-z}{z_d} \right) \right] & \text{for } z \geq (z_{top} - z_d); \\ 0 & \text{otherwise,} \end{cases} \quad \begin{array}{l} R_w(\eta) - \text{damping rate (t}^{-1}) \\ z_d - \text{depth of the damping layer} \\ \gamma_r - \text{damping coefficient} \end{array}$$

[&dynamics *damp\_opt* = 3 (default = 0)]  
 [&dynamics *damp\_coef* = 0.2 (recommended, = 0. default)]  
 [&dynamics *zdampl* = 5000. (*z\_d* (meters); default); height below model top where damping begins]

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## Dynamics: 8. Filters – gravity-wave absorbing layer example



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## Dynamics: 9. Map projections and global configuration

### ARW Model: projection options

1. Cartesian geometry:  
idealized cases
2. Lambert Conformal:  
mid-latitude applications
3. Polar Stereographic:  
high-latitude applications
4. Mercator:  
low-latitude applications
5. Latitude-Longitude global, regional

Projections 1-4 are isotropic ( $m_x = m_y$ )

Latitude-longitude projection is anisotropic ( $m_x \neq m_y$ )

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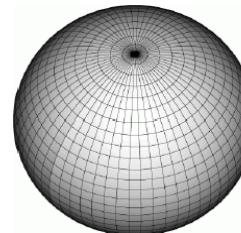
## Dynamics: 9. Map projections and global configuration

### Global ARW – Polar filters

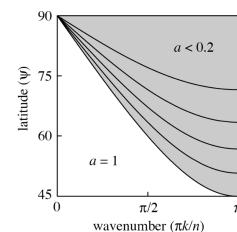
Converging gridlines severely limit timestep.  
The polar filter removes this limitation.

Filter procedure - Along a grid latitude circle:

1. Fourier transform variable.
2. Filter Fourier coefficients.
3. Transform back to physical space.



Filter Coefficient  $a(k), \psi_o = 45^\circ$



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$$\hat{\phi}(k)_{\text{filtered}} = a(k) \hat{\phi}(k), \quad \text{for all } k$$

$$a(k) = \min \left[ 1., \max \left( 0., \left( \frac{\cos \psi}{\cos \psi_o} \right)^2 \frac{1}{\sin^2(\pi k/n)} \right) \right]$$

$k$  = dimensionless wavenumber

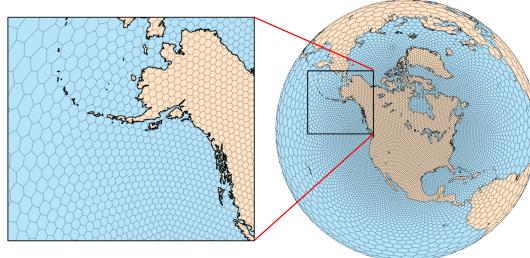
$\hat{\phi}(k)$  = Fourier coefficients from forward transform

$a(k)$  = filter coefficients

$\psi$  = latitude  $\psi_o$  = polar filter latitude, filter when  $|\psi| > \psi_o$

## Dynamics: 9. Map projections and global configuration

An alternative to  
global ARW...



- Global, nonhydrostatic, C-grid Voronoi mesh
- Numerics similar to WRF; WRF-NRCM physics
- No pole problems
- Variable-resolution mesh – no nested BC problems

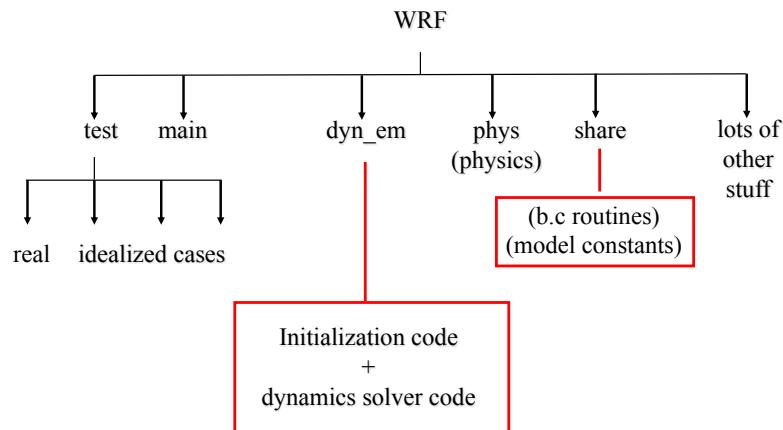
Available at: <http://mpas-dev.github.io/>



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## Dynamics: Where are things?



### WRF ARW Tech Note

A Description of the Advanced Research WRF Version 3 (June 2008, 2012 update)

<http://www.mmm.ucar.edu/wrf/users/pub-doc.html>

We are in the process of updating the technical note to Version 4

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## Dynamics: 10. Boundary condition options

ARW Model: Boundary Condition Options

### Lateral boundary conditions

1. Specified (Coarse grid, real-data applications).
2. Open lateral boundaries (gravity-wave radiative).
3. Symmetric lateral boundary condition (free-slip wall).
4. Periodic lateral boundary conditions.
5. Nested boundary conditions (specified).

### Top boundary conditions

1. Constant pressure.

### Bottom boundary conditions

1. Free slip.
2. Various B.L. implementations of surface drag, fluxes.

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