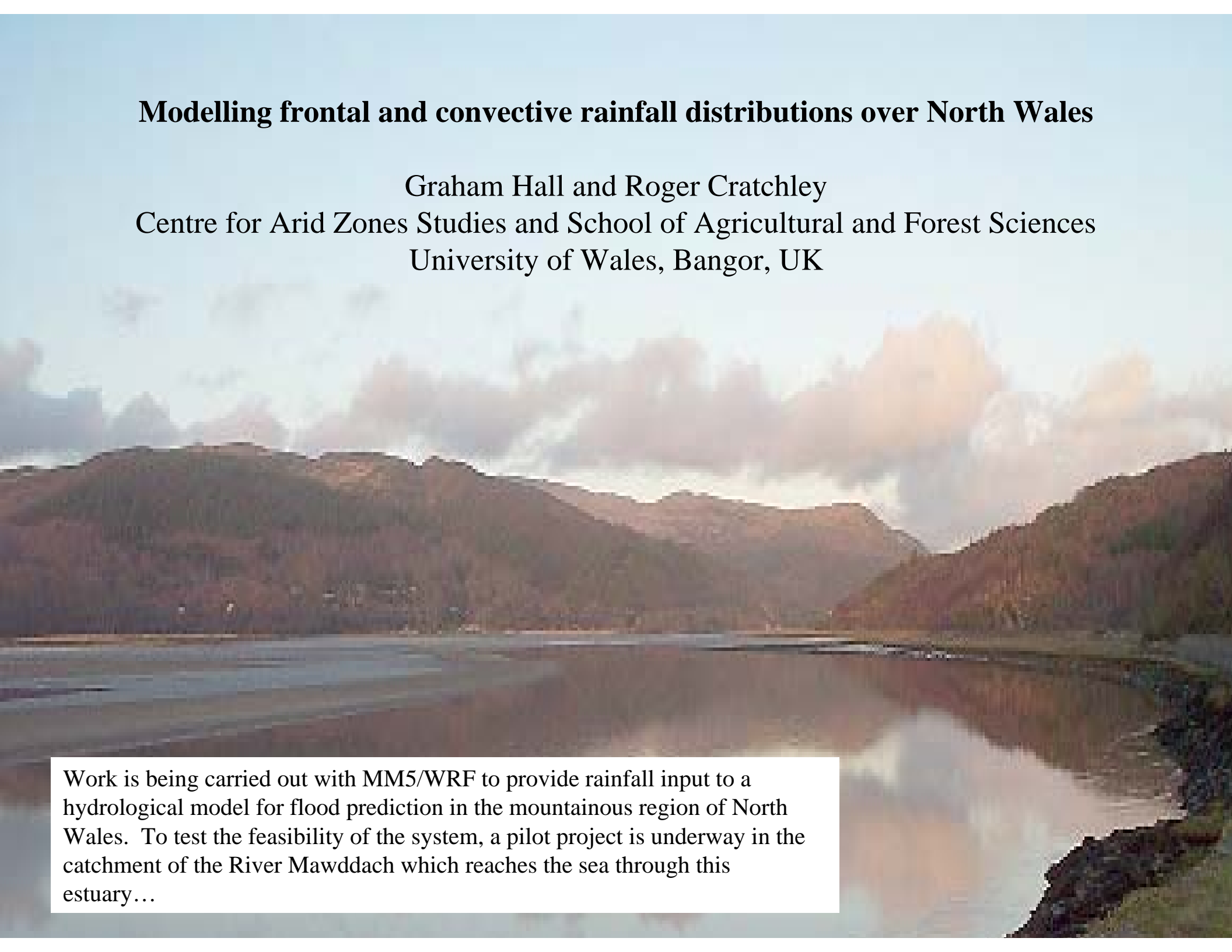


# **Modelling frontal and convective rainfall distributions over North Wales**

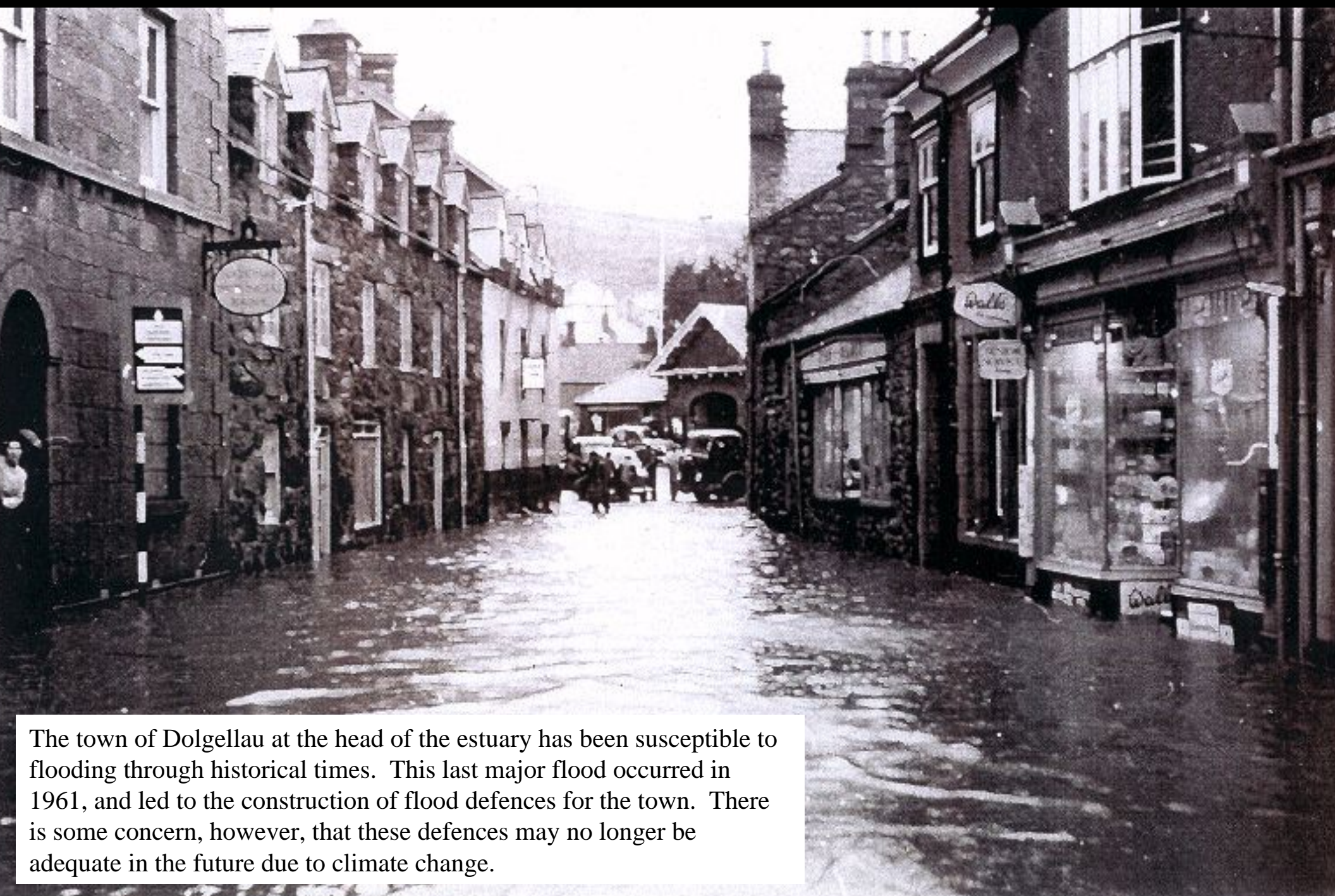
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Work is being carried out with MM5/WRF to provide rainfall input to a hydrological model for flood prediction in the mountainous region of North Wales. To test the feasibility of the system, a pilot project is underway in the catchment of the River Mawddach which reaches the sea through this estuary...



The town of Dolgellau at the head of the estuary has been susceptible to flooding through historical times. This last major flood occurred in 1961, and led to the construction of flood defences for the town. There is some concern, however, that these defences may no longer be adequate in the future due to climate change.



At the commencement of the project there were six rain gauges in and around the Mawddach catchment operated by government organisations. An additional 24 data-logging rain gauges have been added, particularly in the upland areas, so that accurate hourly rainfall maps of the catchment can now be produced for individual storm events.

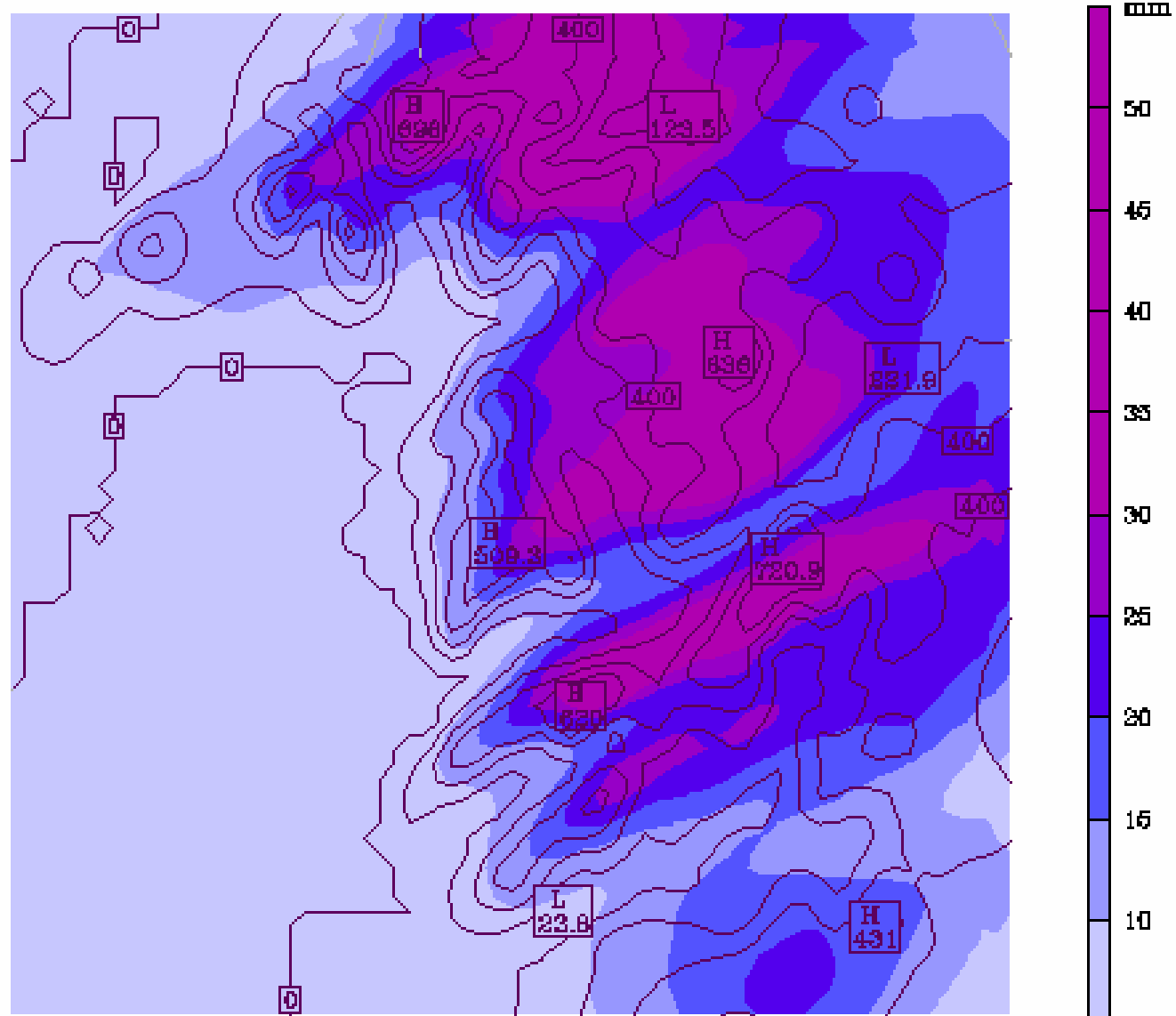


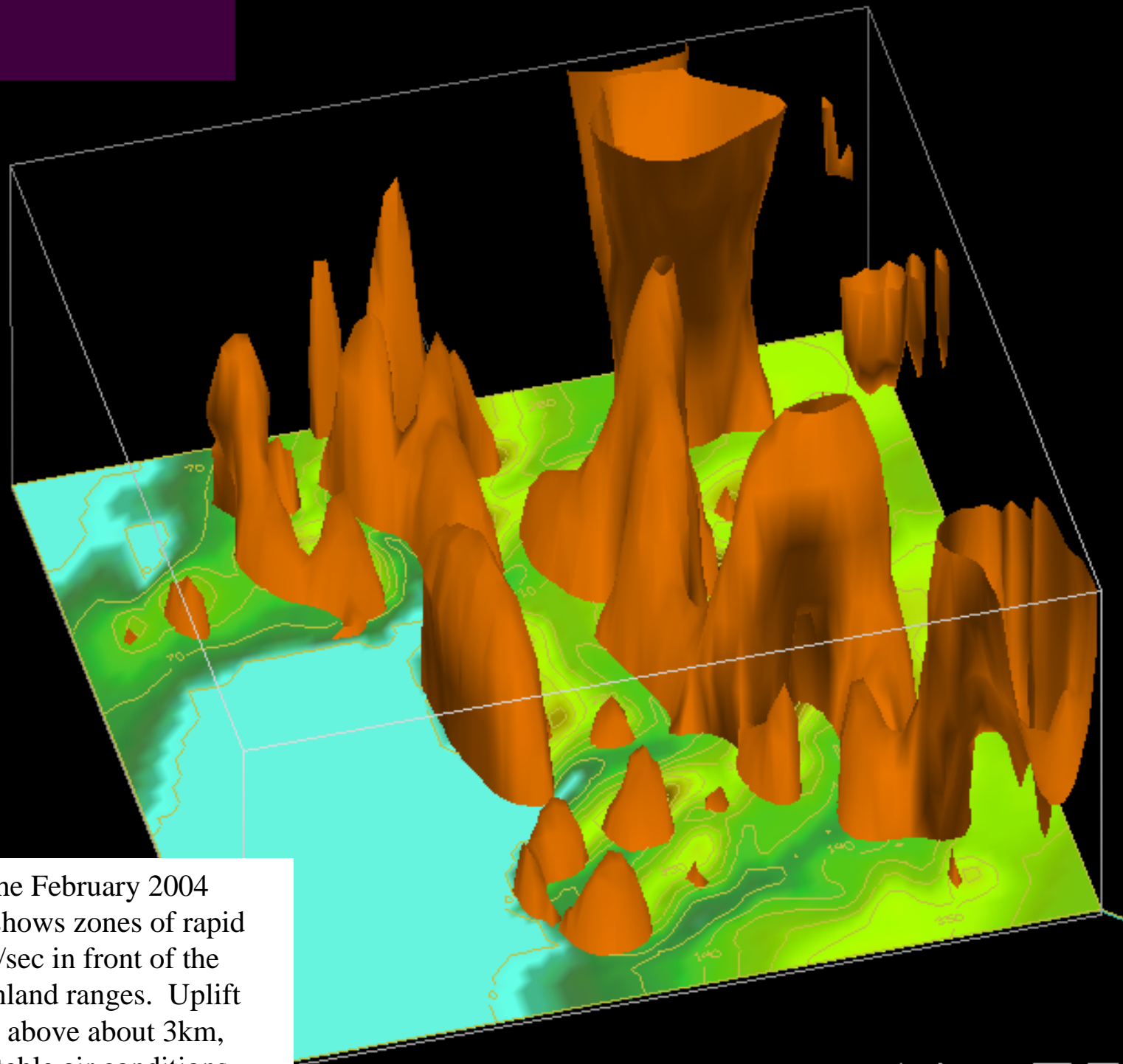


It has become clear from field data that large differences in rainfall occur across the catchment. This range of mountains, the Rhinogs, along the coast receives relatively little rainfall, whereas a second range of mountains some 20km further inland, the Arenigs, is the focus of maximum rainfall and the main source area for floods.

Dataset: febplot RIP: febplot Init: 0000 UTC Tue 03 Feb 04  
 Fcst: 12.00 Valid: 1200 UTC Tue 03 Feb 04 (1200 LST Tue 03 Feb 04)  
 Total precip. in past 6 h  
 Terrain height AMSL

Modelling has been carried out with MM5 for five winter storm events of low CAPE associated with frontal systems from the Atlantic. In all cases the agreement with rain gauge data was within 10%, and often better than 5% error. The pattern of low rainfall over coastal mountains and rainfall increasing inland is correctly modelled, as in this example from a flood event of 3 February 2004

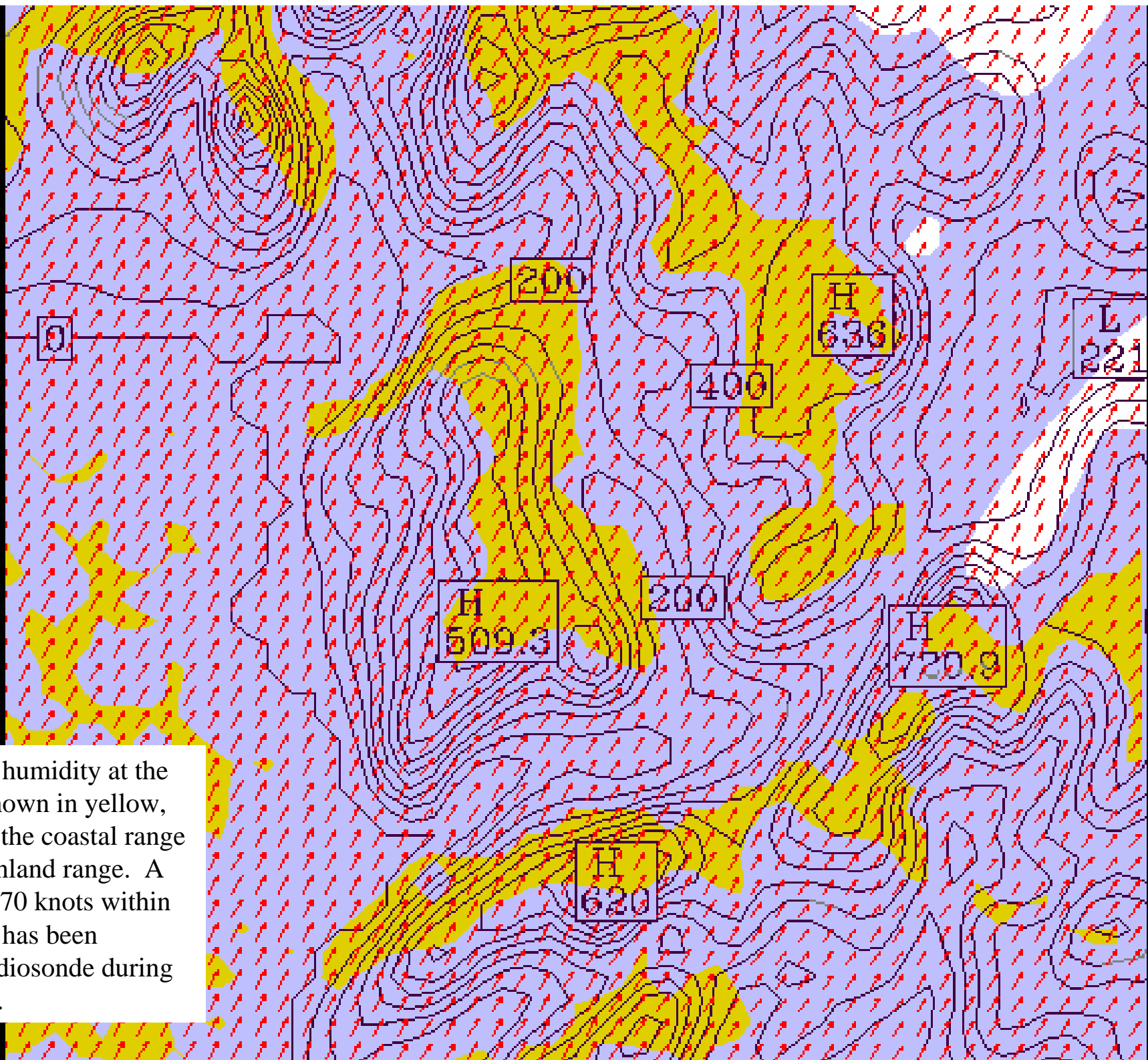


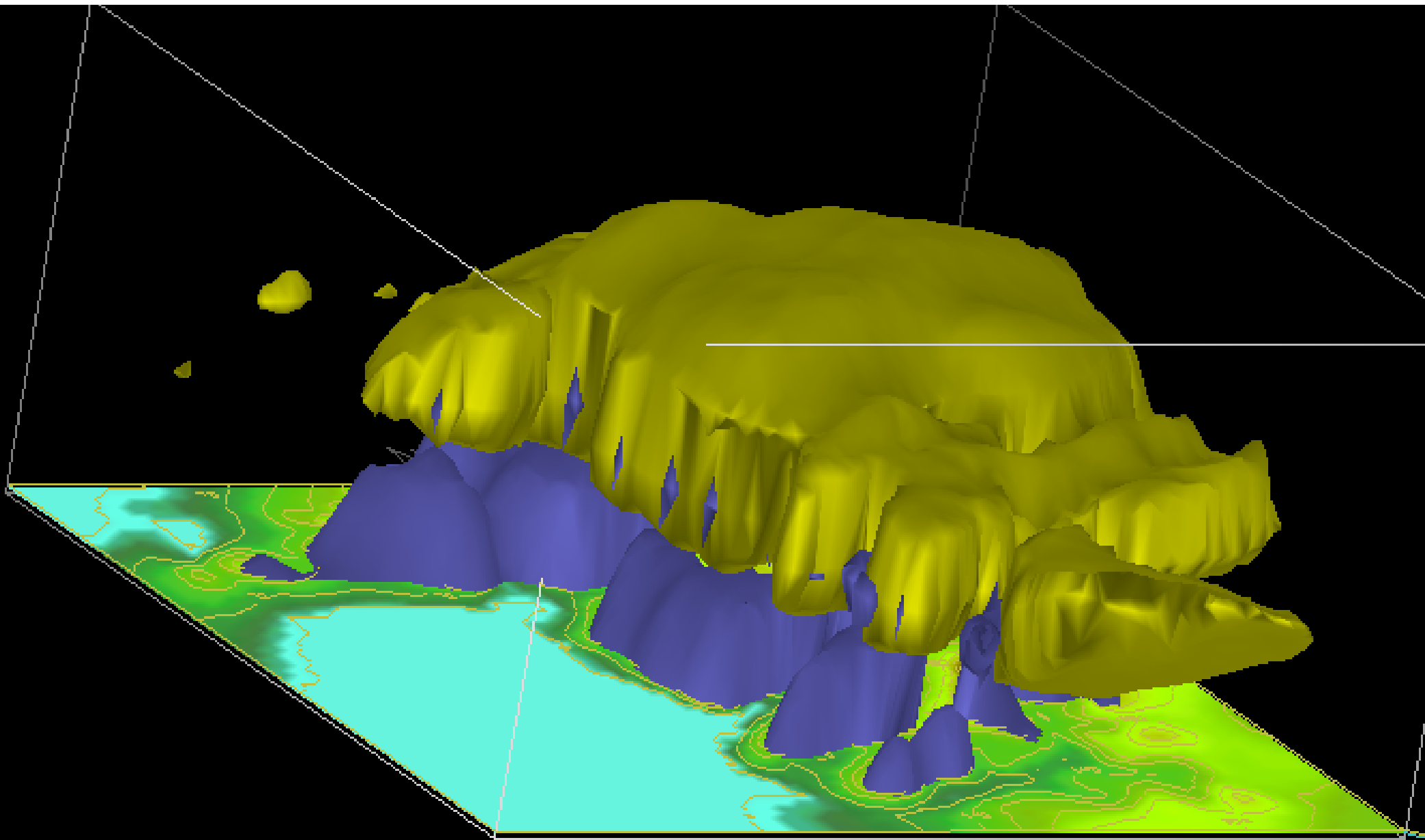


Analysis of the February 2004 storm event shows zones of rapid uplift  $>40\text{cm/sec}$  in front of the coastal and inland ranges. Uplift zones die out above about 3km, indicating stable air conditions.

Vis5D

Zones of 100% humidity at the 900mb level, shown in yellow, occur inland of the coastal range and above the inland range. A low level jet of 70 knots within the SW airflow has been identified by radiosonde during the storm event.





This diagram shows isosurfaces for cloud mixing ratio  $>0.4$  (blue) and precipitation mixing ratio  $>0.4$  (yellow). A seeder – feeder mechanism has been suggested to explain the rainfall enhancement inland. Droplets falling from higher level seeder clouds enter the saturated low level jet of stratiform cloud. The droplets grow within this feeder zone and are advected inland to fall as rain over the inland mountain range.





Flooding around the estuary caused by the February 2004 storm. It has been possible to use the rainfall distribution generated by MM5 as an input to a MODFLOW groundwater model, and in turn to use river discharge from MODFLOW as an input to RIVER-2D software for modelling the depth and extent of flooding.



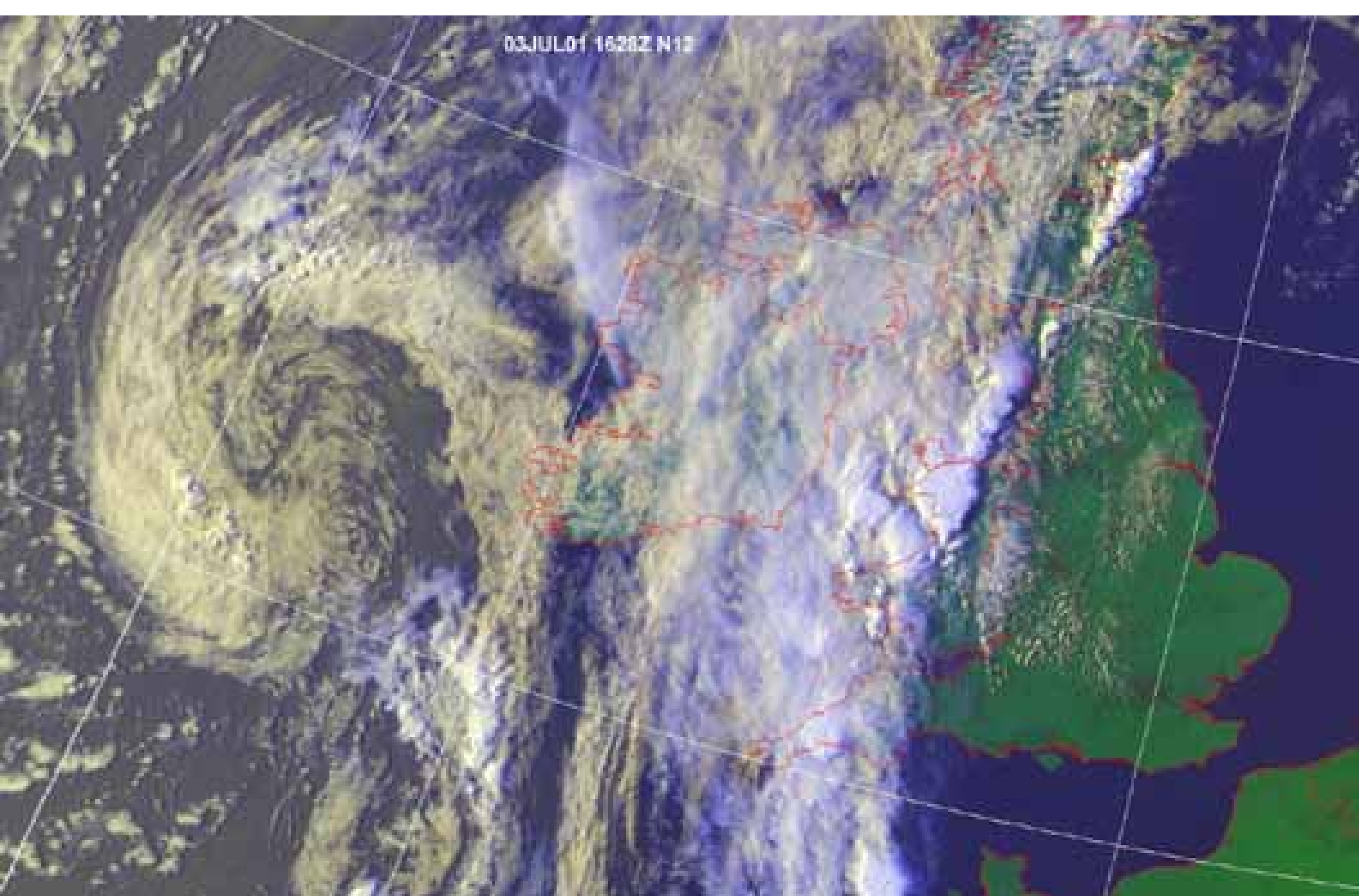
The next phase of the project has been to examine summer storms over the region. The largest storm event ever recorded was on 3 July 2001, when an intense line of thunderstorms developed ahead of a cold front. Photo: John Mason



At the height of the storm, roads became impassable and golf-ball sized hail was experienced. Lightning strikes occurred at a rate of one per second. 35mm of rainfall was recorded in one hour.  
Photo: Bob Chilton



Bridges were severely damaged or washed away by flood waters...



03JUL01 1628Z N13

Satellite image of the July 2001 storm, with a depression over the Atlantic and cold front along the coast of Wales. Squall line storms occurred ahead of the cold front at the mixing zone with warm moist air over Britain.



Dataset: julyplot RIP: julyplot

Init: 0000 UTC Tue 03 Jul 01

Fcst: 19.00

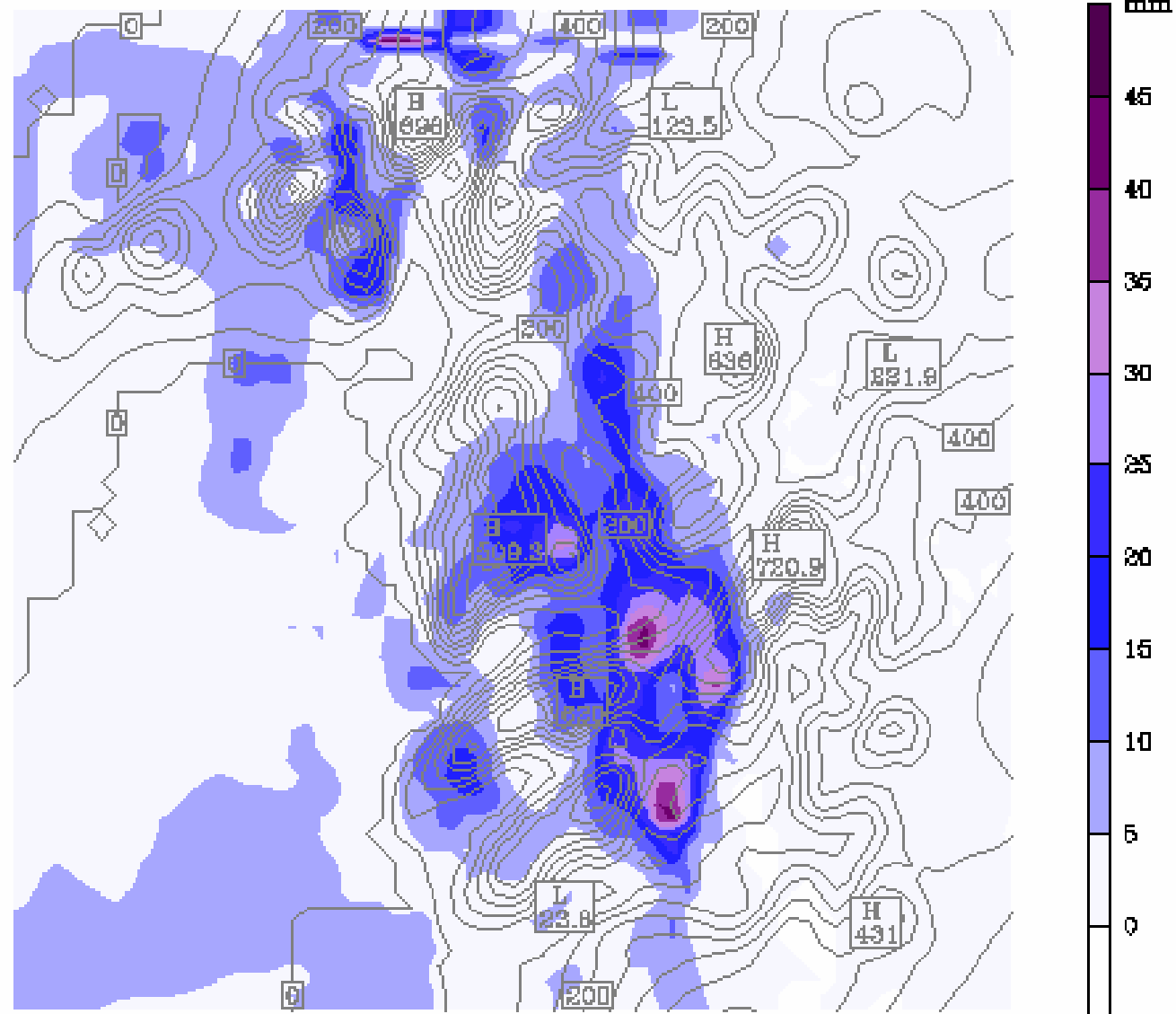
Valid: 1900 UTC Tue 03 Jul 01 (1900 LST Tue 03 Jul 01)

Total precip. in past 1 h

Terrain height AMSL

Modelling of the July 2001 storm was carried out with MM5:

- An initial model with no convective parameterisation on the inner 1km domain failed to generate the storm rainfall at all.
- A model using the Anthes-Kuo scheme generated a good representation of the storm as shown here... Location, timing and rainfall intensity all have good correspondence with field data. The squall line illustrates a bow echo structure with rear inflow notch and central supercell.
- Models using the Grell and Kain-Fritsch schemes again failed to generate any storm rainfall in the required area.



Dataset: julyplot RIP: julyplot

Init: 0000 UTC Tue 03 Jul 01

Fest: 19.00

Valid: 1900 UTC Tue 03 Jul 01 (1900 LST Tue 03 Jul 01)

Vertical velocity

XY= 3.0, 40.0 to 65.0, 40.0

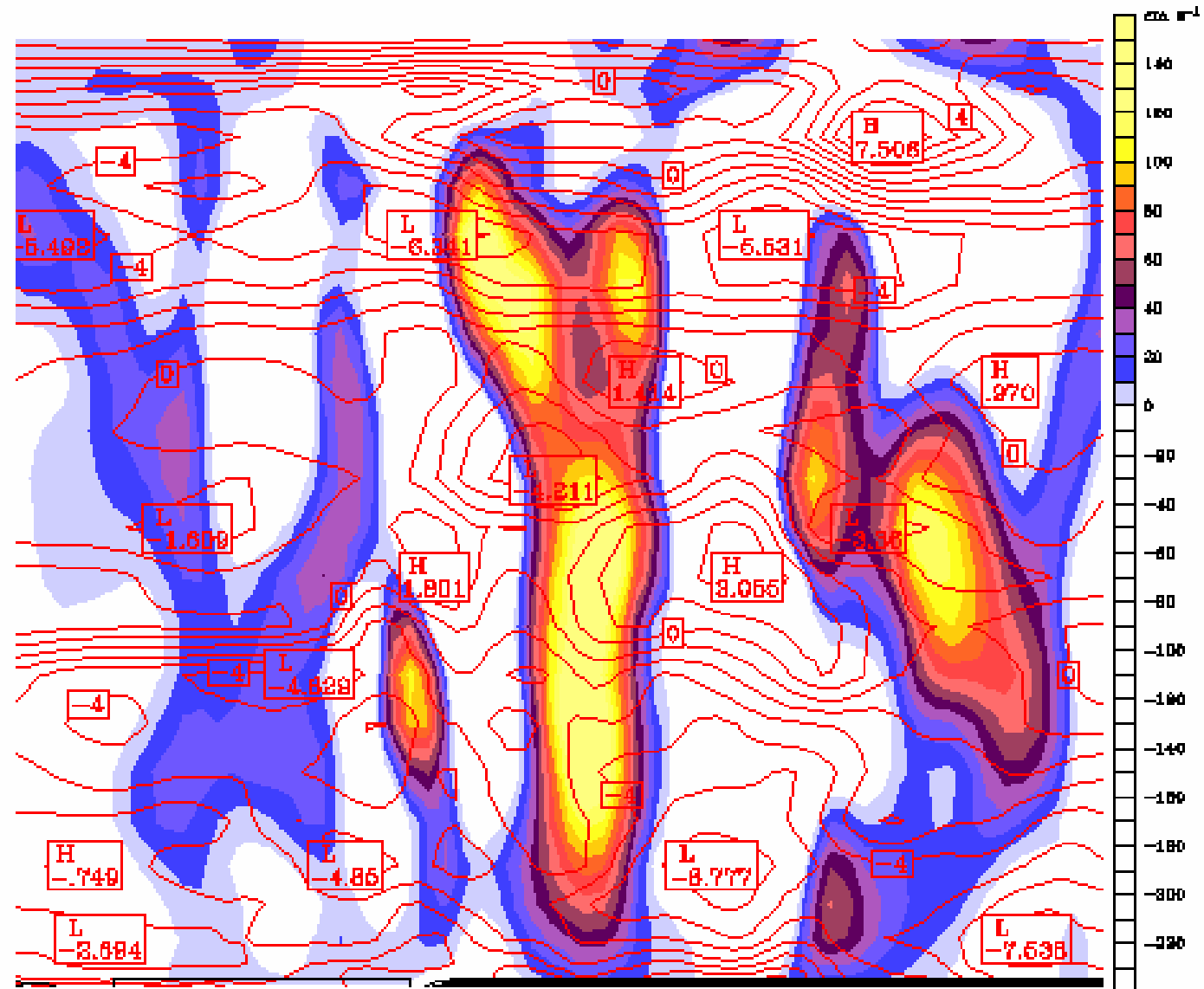
Horizontal wind (x-comp.)

XY= 3.0, 40.0 to 65.0, 40.0

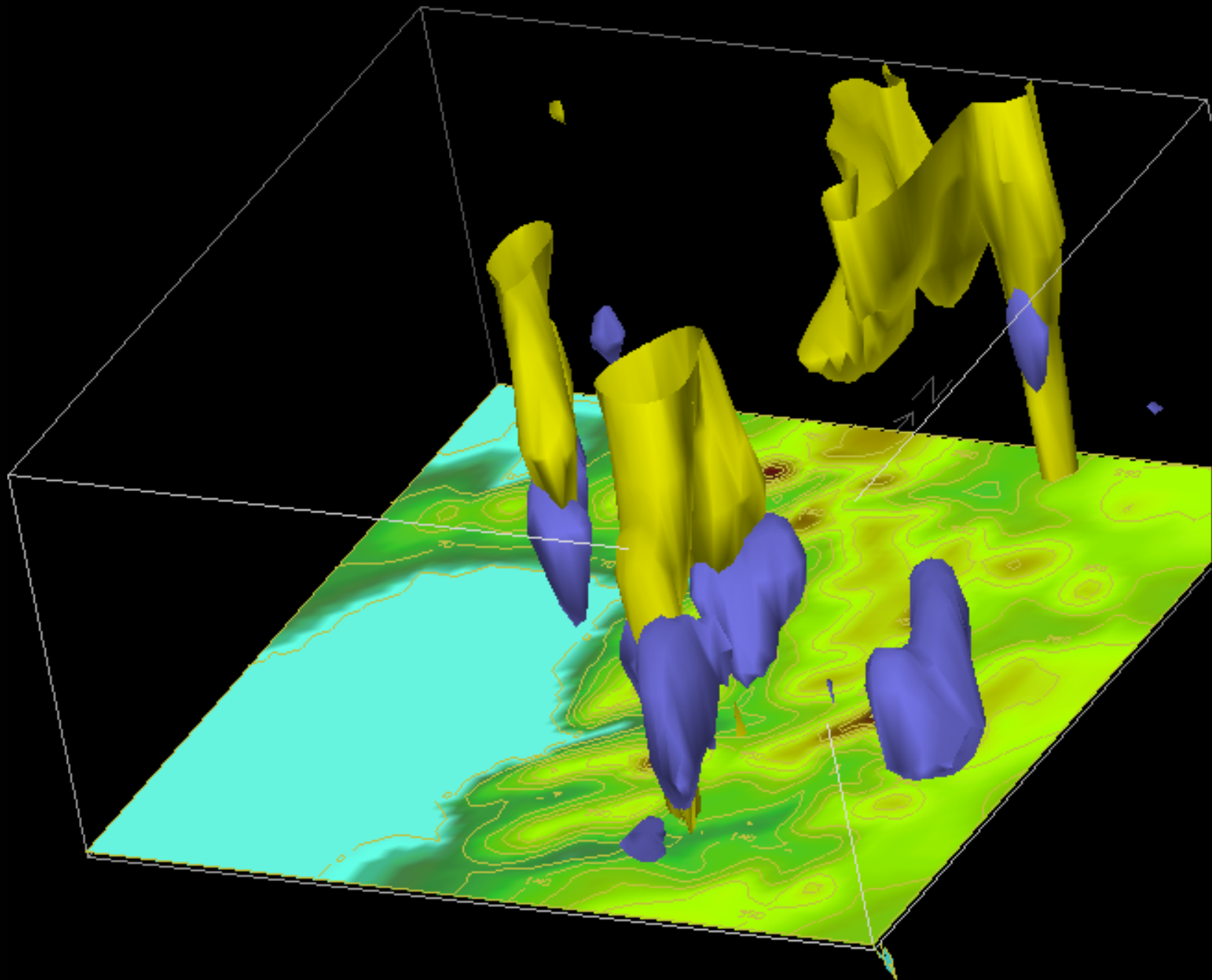
Terrain height AMSL

XY= 3.0, 40.0 to 65.0, 40.0

W-E Cross section of the Anthes-Kuo model showing the convective supercell. Colour fill indicates vertical velocities, and red contours indicate horizontal velocities. There is evidence of mid altitude coupling of regional wind shear to generate cyclonic motion...

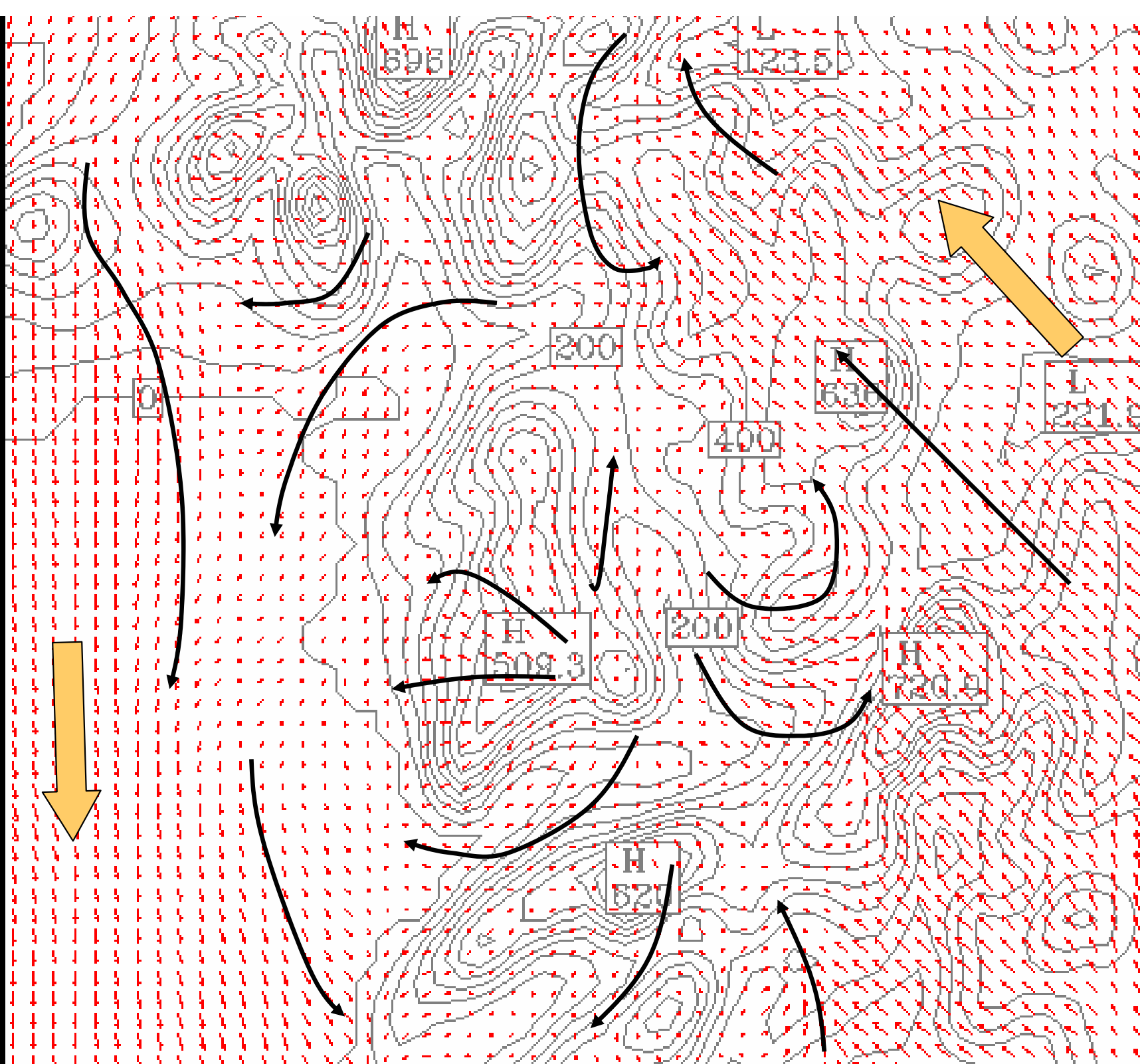


CONTOURS: UNITS= $m\ s^{-1}$  LOW= 0.0000 HIGH= 440.00 INTERVAL= 80.000  
 CONTOURS: UNITS= $m\ s^{-1}$  LOW= -7.0000 HIGH= 7.0000 INTERVAL= 1.0000  
 Model Info: V2.6.3 Anthes-Kuo Blackadar Simple Ice 1 km, 28 levels, 2 sec



Isosurface (yellow) for precipitation mixing ratio  $> 0.4$  indicates localisation of rainfall within the squall line

Air flows at the 900mb level indicate cyclonic outflows of descending air around the convection centre



Dataset: julyplot2 RIP: julyplot2

Init: 0000 UTC Tue 03 Jul 01

Fcst: 21.00

Valid: 2100 UTC Tue 03 Jul 01 (2100 LST Tue 03 Jul 01)

Vertical velocity

XY= 3.0, 40.0 to 65.0, 40.0

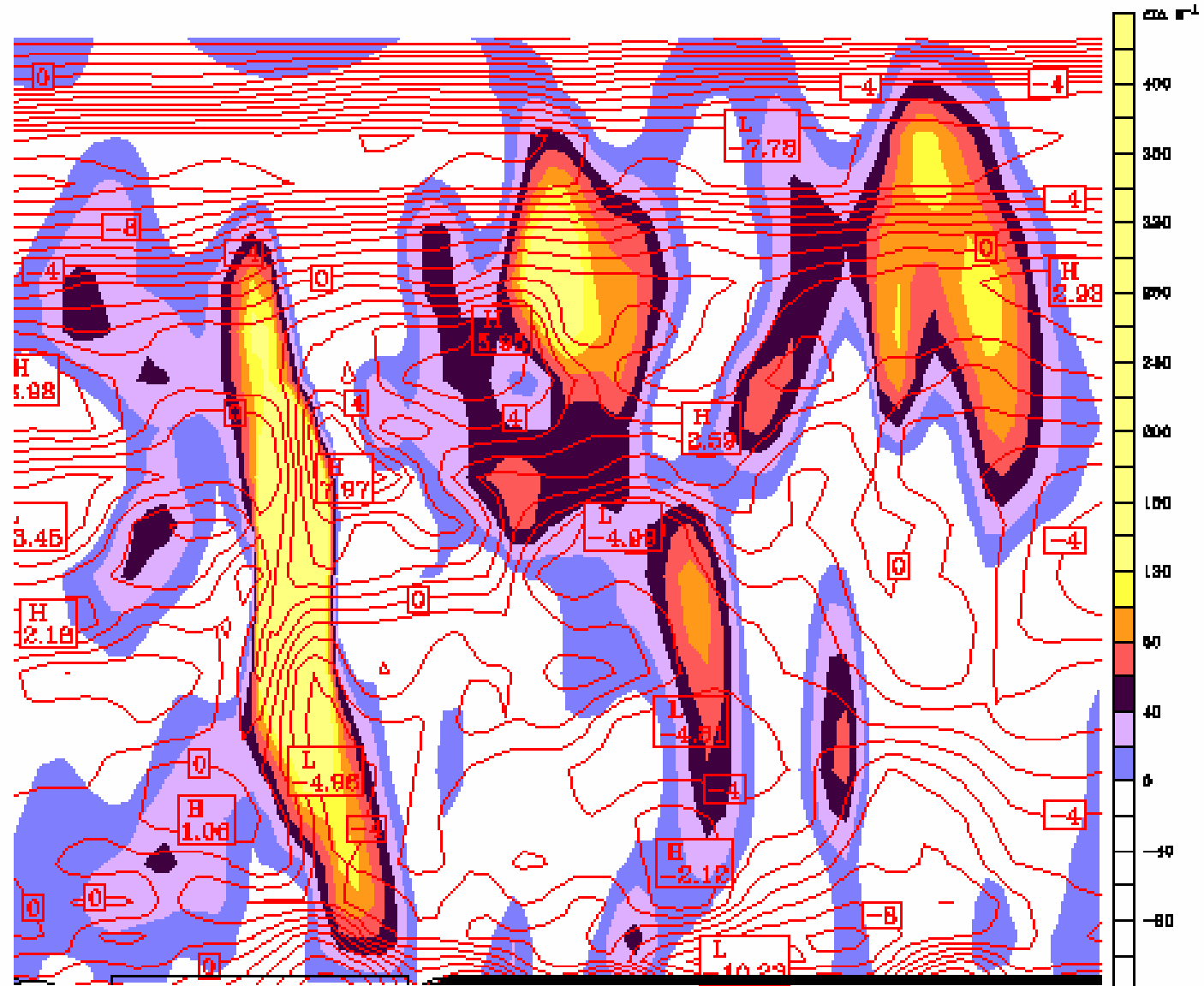
Horizontal wind (x-comp.)

XY= 3.0, 40.0 to 65.0, 40.0

Terrain height AMSL

XY= 3.0, 40.0 to 65.0, 40.0

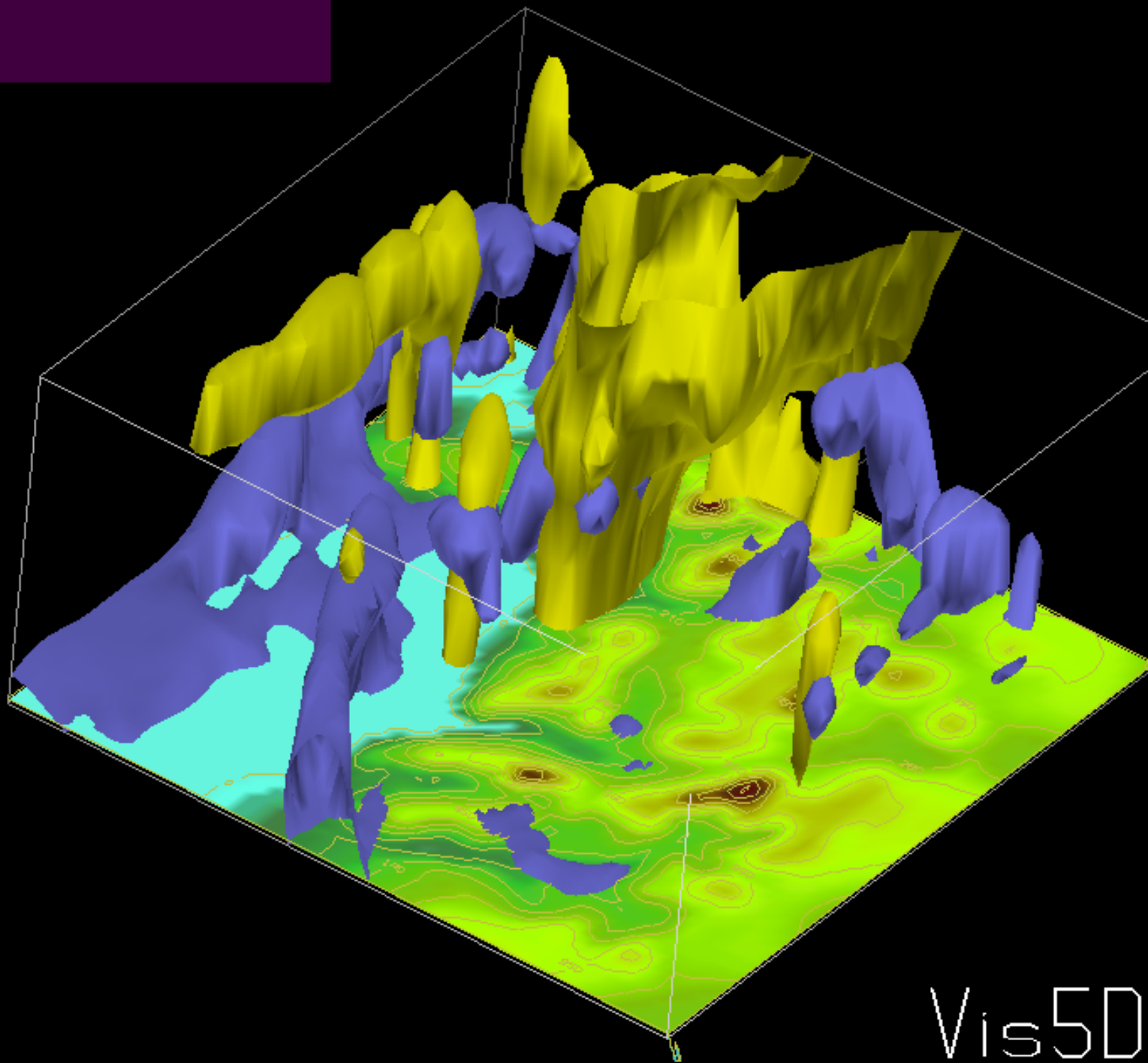
Results using the  
Grell cumulus  
scheme generate a  
number of small  
convective cells  
over the region...



CONTOURS: UNITS= $m\ s^{-1}$  LOW= 0.00000 HIGH= 440.00 INTERVAL= 80.000  
 CONTOURS: UNITS= $m\ s^{-1}$  LOW= -11.000 HIGH= 7.0000 INTERVAL= 1.0000  
 Model info: V3.6.3 Grell Blackadar Simple ice 1 km, 23 levels, 2 sec



Cloud formation  
under the Grell  
scheme is  
evident, but  
rainfall is  
restricted to the  
cold front along  
the coast...



Dataset: julyplot2 RIF: julyplot2

Init: 0000 UTC Tue 03 Jul 01

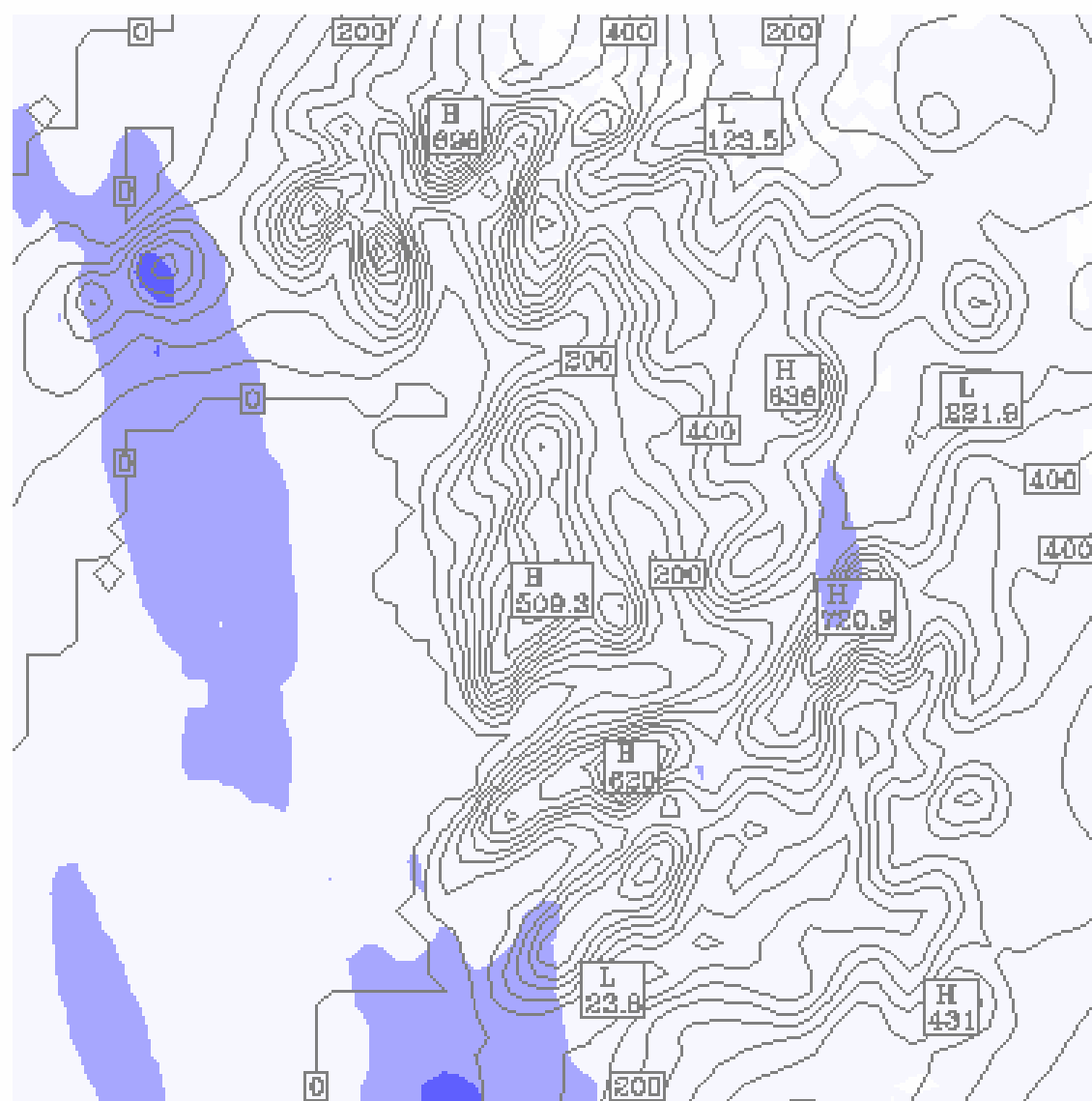
Fest: 21.00

Valid: 2100 UTC Tue 03 Jul 01 (2100 LST Tue 03 Jul 01)

Total precip. in past 1 h

Terrain height AMSL

No rainfall is  
generated along the  
inland axis of the  
squall line...





High CAPE storm events are becoming increasingly violent in Britain – this example is the disastrous flooding of Boscastle in Cornwall in August 2004. It is clearly important to have a convective scheme available in WRF which can reliably predict such events.

Preliminary results of the Mawddach project suggest that a simple bulk convective scheme such as the Anthes-Kuo scheme may be the most effective in generating rainfall patterns on a high resolution 1km grid. It is necessary to model the convergence of air masses, condensation and release of latent heat to realistically drive convective cells.