



Tuning of 3DVAR and its application to high-resolution radar data assimilation system

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Introduction

- ❖ Heavy rainfall is one of the major severe weathers over East Asia producing devastating flash flood, and consequently causing fatalities and property damage. Heavy rainfall is usually resulted from individual mesoscale storms or mesoscale convective systems (MCSs) embedded in synoptic-scale disturbances (Lee et al., 1998).
- ❖ We need high-resolution observations and radar data assimilation techniques to understand the evolution and development mechanisms of mesoscale convective storms responsible for heavy rainfall and better predict heavy rainfall events.
- ❖ Assimilation of radar data is a key scientific issue in numerical weather prediction of convective systems for very short-range forecasting (Wilson et al., 1998). In recent years considerable progress has been made in the retrieval of boundary-layer winds from Doppler radar observations, the assimilation of radar observations into convective-scale numerical models for heavy rainfall prediction, and the assimilation of radar rainfall estimates.

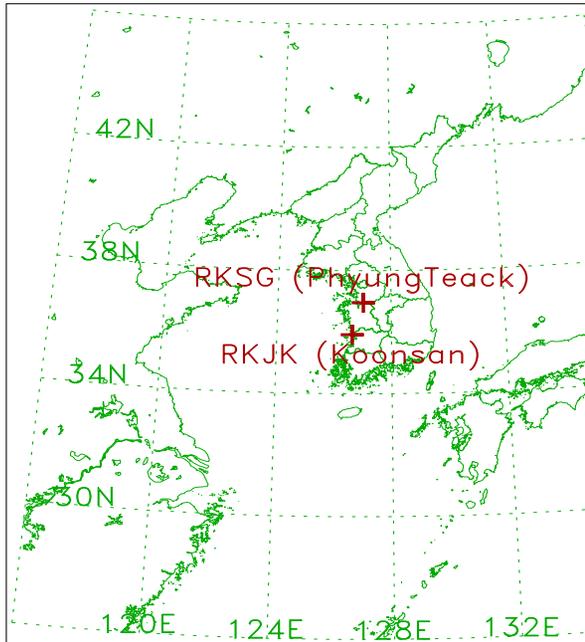
Introduction

- ❖ The objective of this study is to investigate **very short-range forecasting of the WRF model through the 3DVAR data assimilation of Dual-Doppler radar data** (radial velocity and reflectivity) for a heavy rainfall case accompanying mesoscale convective systems (MCSs) over the Korean Peninsula.
- ❖ The 3DVAR system is modified by tuning the scale lengths and observation error statistics. We compare the results of the increment analysis update (IAU) and the rapid update cycle (RUC) on the modified 3DVAR system.

Radar Data

- **Radar position**

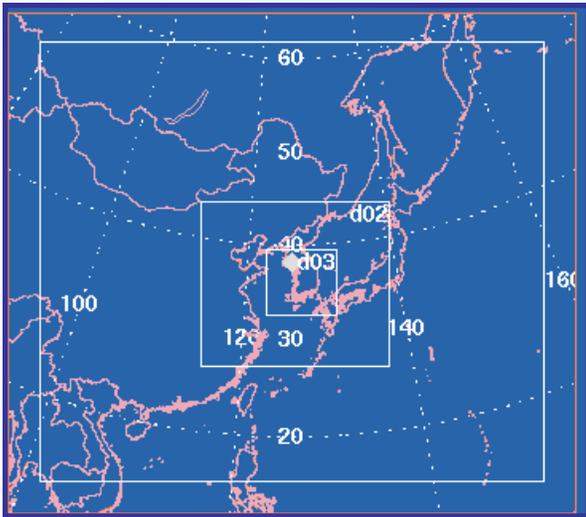
Map from IDL



- **Radar Information**

WSR-88D(NEXRAD)	
Radar Beam	S Band (wave length : 10cm)
Detection range	Reflectivity : 450 km Radial velocity : 240 km
Format	Raw data Level II

Model Configuration



Physics processes	domain 1 (30 km)	domain 2 (10 km)	domain 3 (3.3 km)
Horizontal Dimensions	191 X 171	160 X 178	241 X 229
Time interval (Δt)	90 sec	30 sec	10 sec
Cumulus Parameterization	Kain-Fritsch scheme	Kain-Fritsch scheme	none
Explicit moisture	Lin et al. scheme	Lin et al. scheme	Lin et al. scheme
PBL	YSU scheme	YSU scheme	YSU scheme
Radiation	RRTM/ Dudhia scheme	RRTM/ Dudhia scheme	RRTM/ Dudhia scheme
Surface-Land	Noah LSM	Noah LSM	Noah LSM
Initial and Boundary data	NCEP / FNL analysis	NCEP / FNL analysis	NCEP / FNL analysis

Data Assimilation Experiment

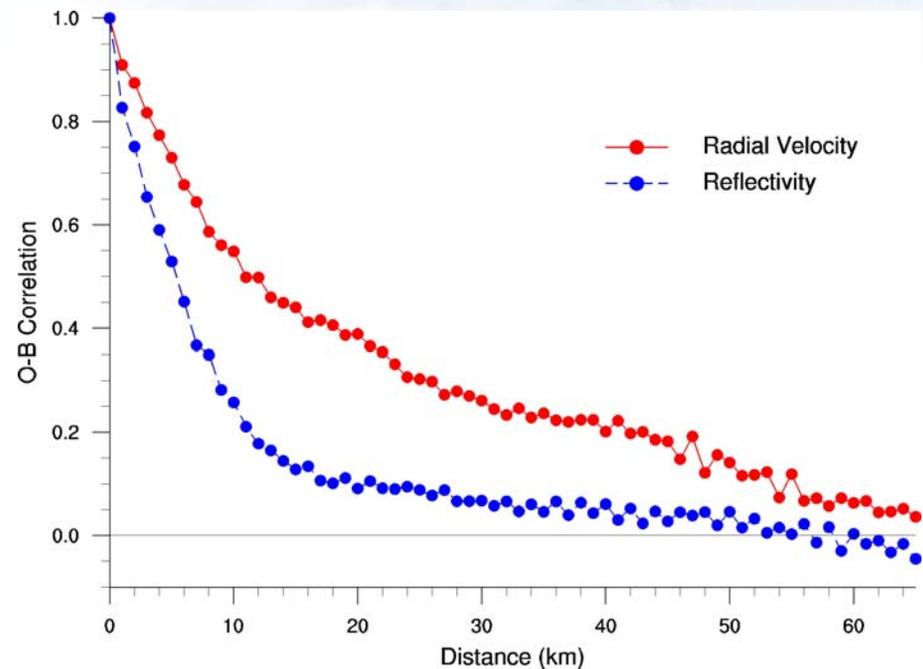
Tuning of 3DVAR (Version 2.1) and its application to high-resolution radar data assimilation system.

- ❖ The current scale-lengths of background error (about 110 km for wind and 40 km for mixing ratio) are not proper for high resolution asynoptic observations such as radar and AWS.
- ❖ The average minimization ratio of cost function, about 25 %, suggests the further minimization possible by adjusting the observation errors.
- ❖ Development of a radar data assimilation system for Increment Analysis Update (IAU) and Rapid Update Cycle (RUC) with 3DVAR.

Tuning of Background Scale-lengths

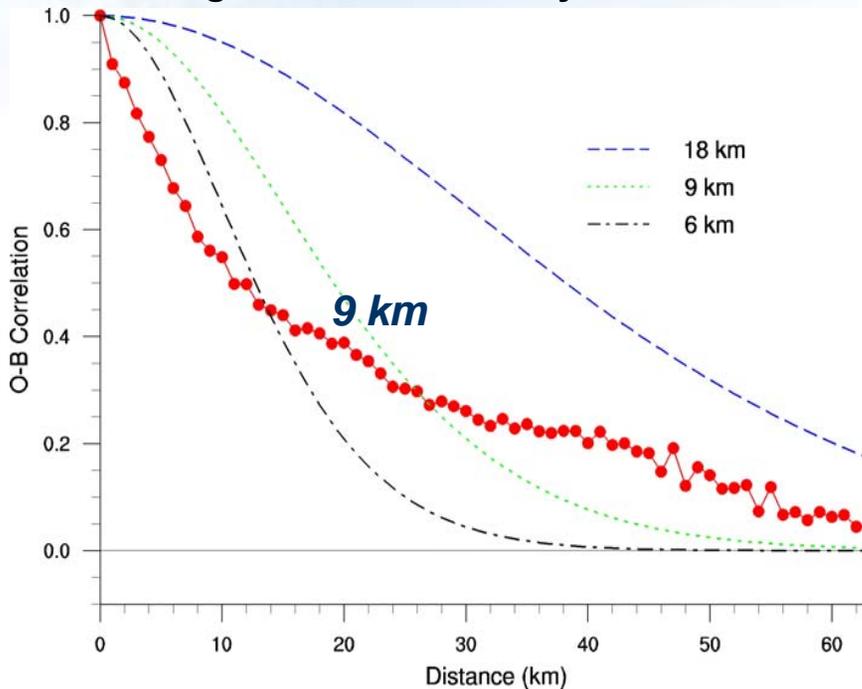
The O-B (observation – background value) of radar data is calculated from two MCS cases and one frontal case. The O-B correlation decreases according to short distance. It means that the radar observation detects meso- as well as micro- scale phenomena. The locality of the radar reflectivity is higher than radical velocity

O-B Statistics

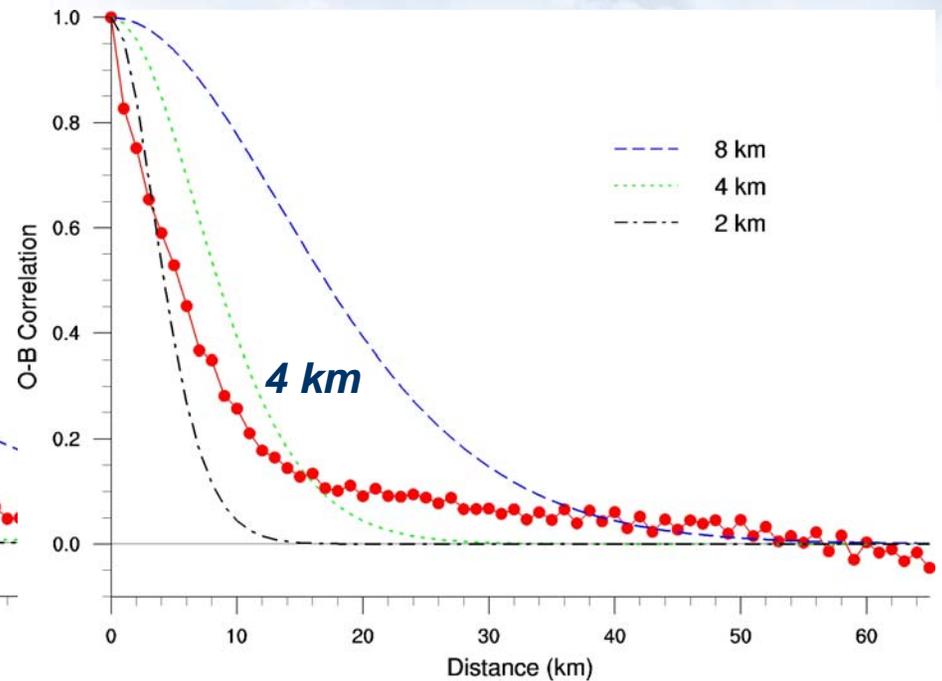


Scale lengths

Scale length for radial velocity



Scale length for reflectivity



- The locality of the radar observation can be reflected by tuning the scale length of the background error using a recursive filter.
- 9 km and 4 km of scale lengths are proper for radial velocity and reflectivity, respectively.

3DVAR Error Tuning

- **The expectation value of the minimized cost function is given by a half of the effective number of observations (Desroziers and Ivanov, 2001)**
- **Further minimization of cost function is possible by scaling each cost function term to satisfy the expectation values of the minimized cost function term, respectively**
- **The adjustment of scaling parameters are iteratively done/**

3DVAR Error tuning

- ❖ To adjust the real cost function to the expectation value, scaling parameters (or error factors) will be applied to each cost function term and to each observation type, and determined iteratively.

$$J = \frac{1}{s_b^2} J^b + \frac{1}{s_o^2} J^o$$

$$s_o^2 \Big|^{i+1} = \left(\frac{J^o \Big|^{i+1}}{E(J^o)} \right) \quad s_b^2 \Big|^{i+1} = \left(\frac{J^b \Big|^{i+1}}{E(J^b)} \right)$$

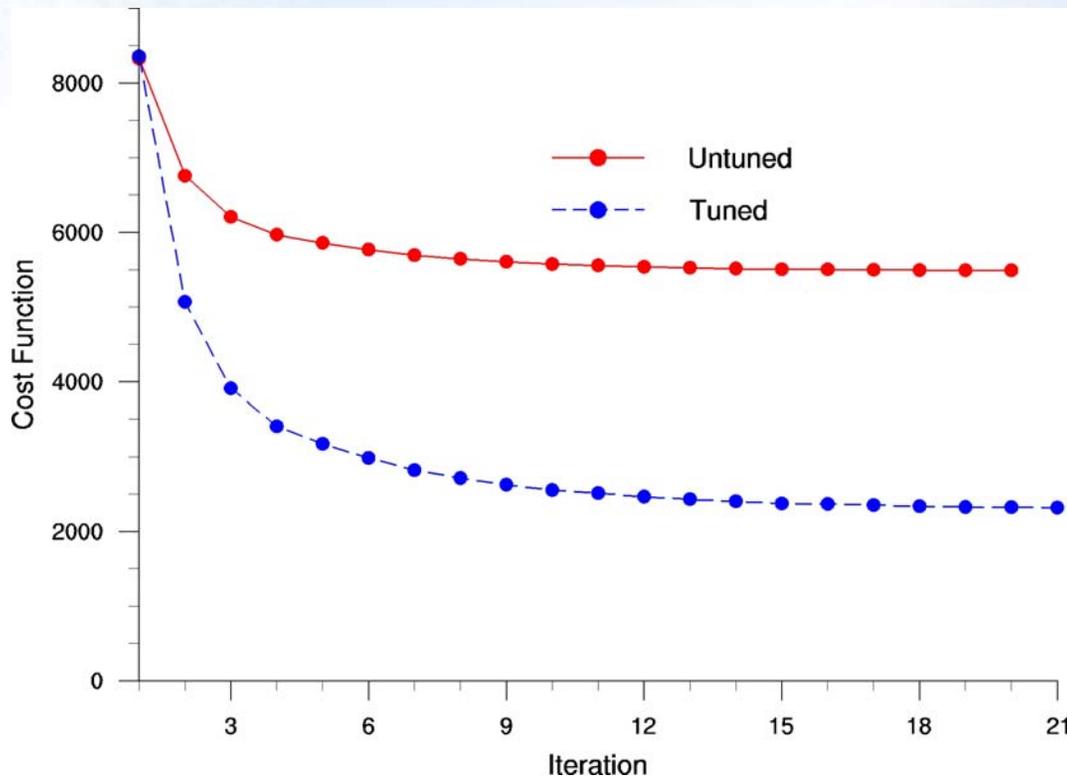
where i means iteration number. The estimation of $E(J^o)$ and $E(J^b)$ (Desroziers and Ivanov, 2002).

3DVAR Error Factors

Observation	P	S(0)	S(1)	S(2)	S(3)
Radial Velocity	16820	1.00	0.72	0.69	0.68
Reflectivity	26977	1.00	1.95	1.95	1.95
J ^b		1.00	1.04	1.11	1.13

The numbers of effective radar radial velocity and reflectivity are 16,820 and 26,977, respectively, and the tuning parameters of the observation error converge rapidly.

Impact of tuning on minimization



J/p for ideal: 0.5

J/p for untuned: 1.38

J/p for tuned: 0.53

The decrease in the cost function of the untuned 3DVAR and the tuned 3DVAR is 25 % and 64 %, respectively.

In particular, the ratio of the minimized cost function (J) and the number of used observation (p) is 1.38 and 0.53 for the untuned 3DVAR and the tuned 3DVAR, respectively.

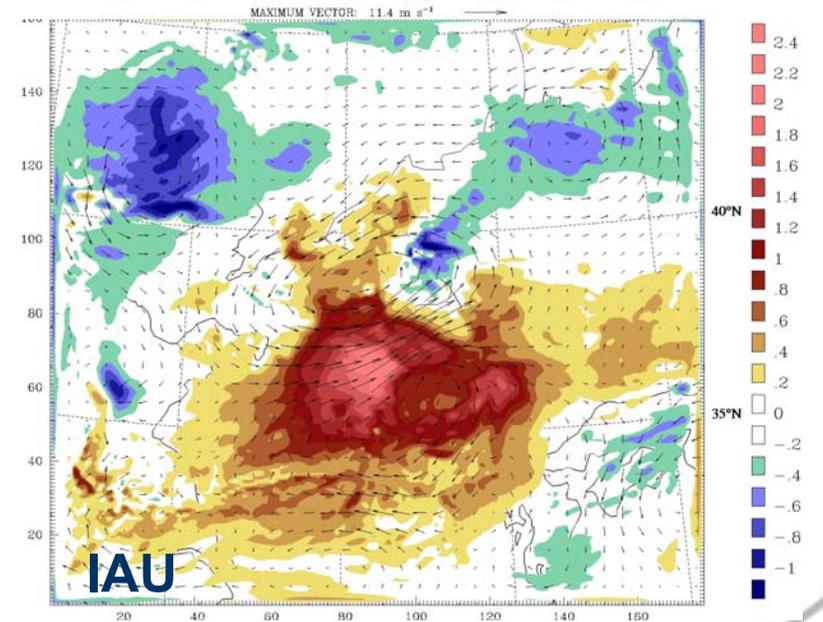
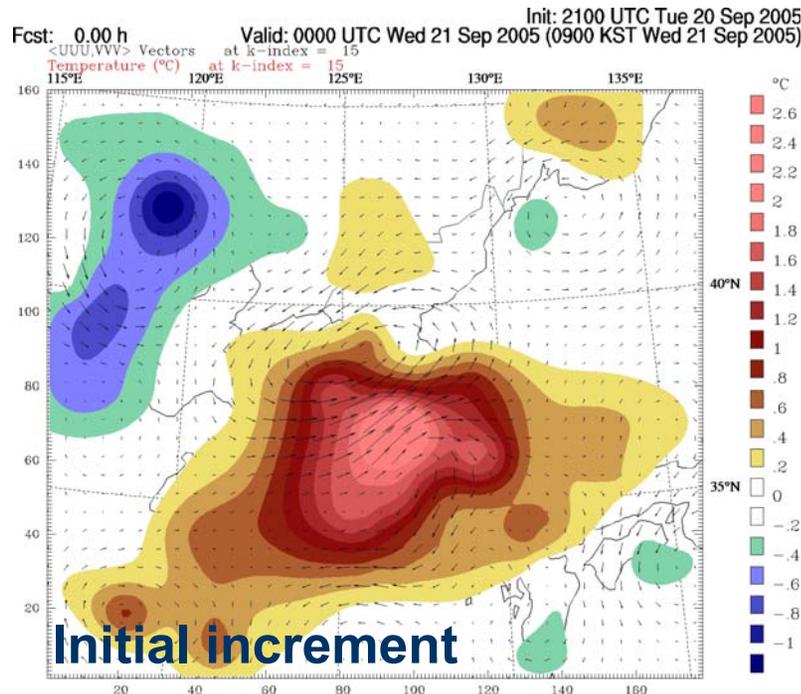
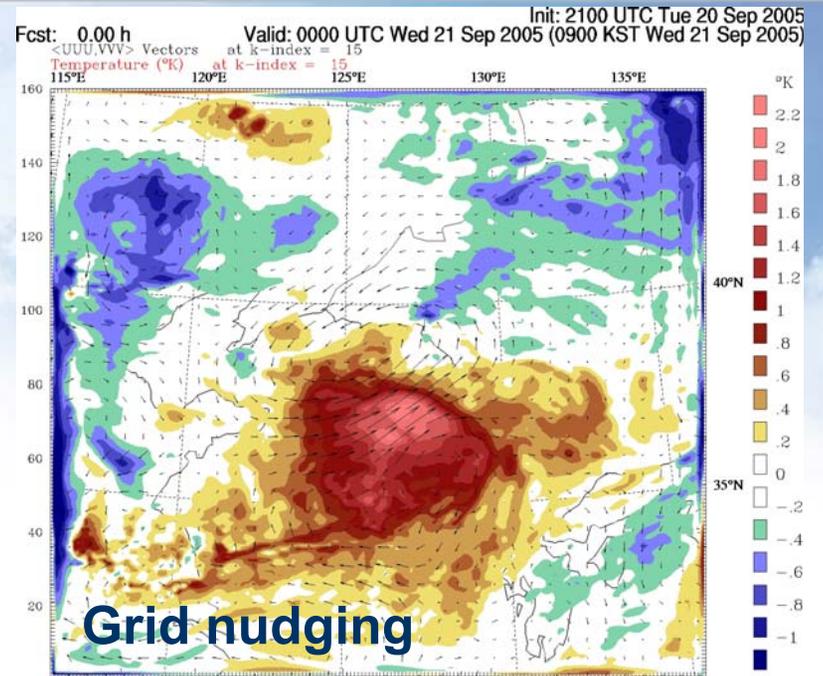
Grid nudging and IAU for WRF

$$\frac{\partial \mu \chi}{\partial t} = F_{\chi}(\mathbf{x}, t) + \left\{ \begin{array}{l} \text{Nudging} \\ G_{\chi} W(\mathbf{x}, t) \mu(\chi - \chi_{anal}), \end{array} \right. \left. \begin{array}{l} \text{IAU} \\ G_{\chi} \mu \chi'_{anal} \end{array} \right\}$$

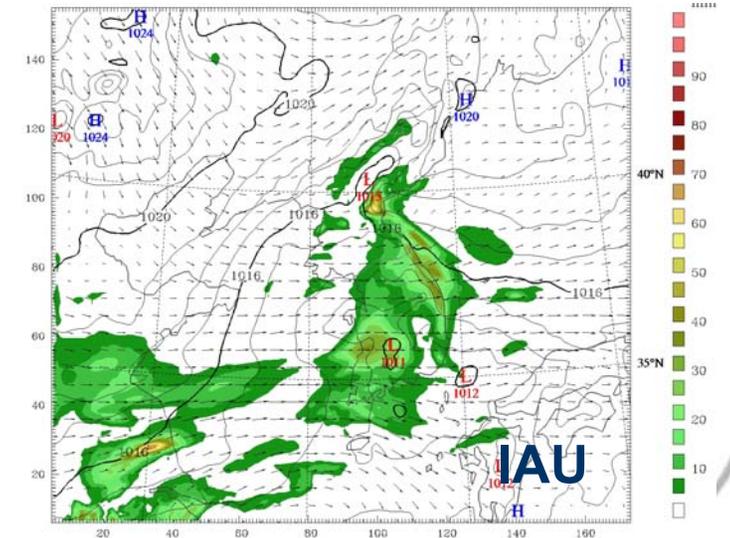
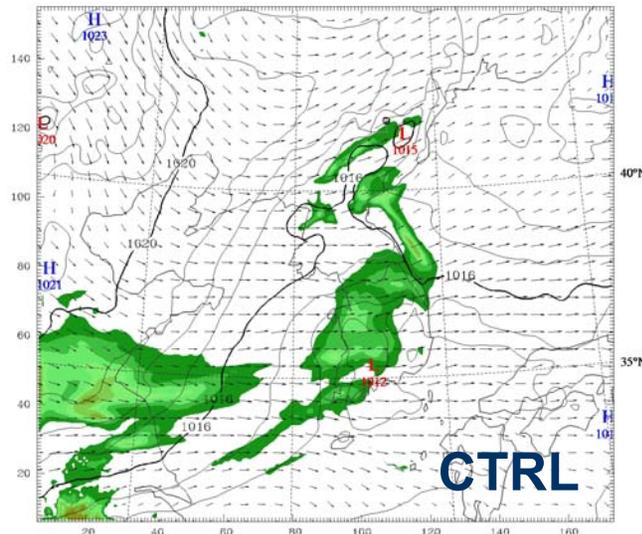
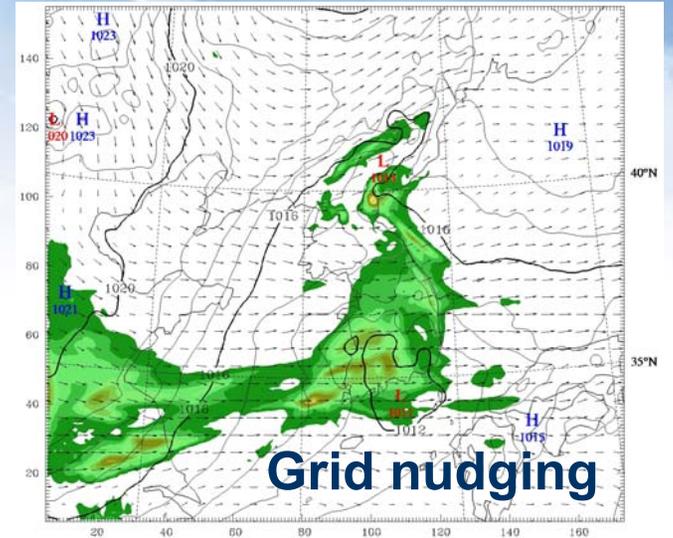
$W(x, t) = (w_{\eta} \cdot w_t)$, G_{χ} : nudging coefficient

- ❖ For the grid nudging, we have implemented time-variant nudging (=1) and target nudging (=3)
- ❖ Surface (in planetary boundary layer) nudging (or IAU) is not yet implemented.
- ❖ The nudging of U, V, and T is currently tested.

Verification of Grid nudging and IAU



6h accumulated rainfall amount (mm)



Initialization Experiments

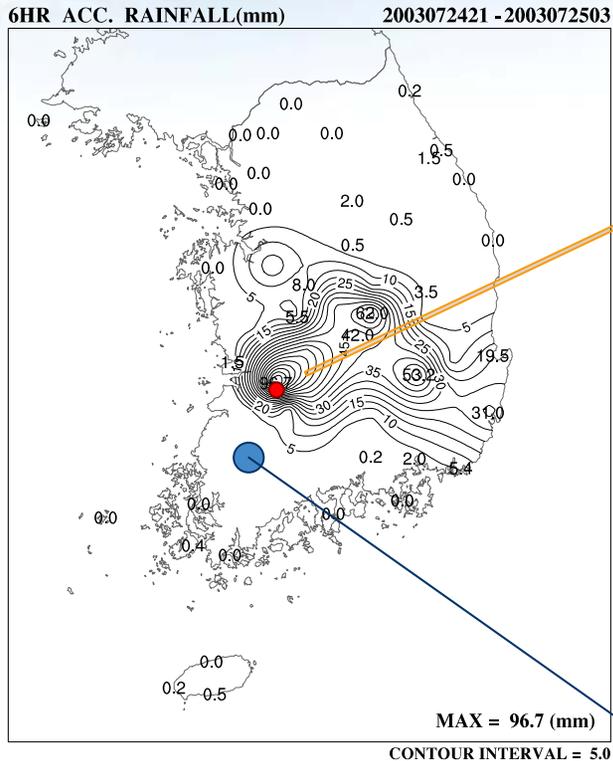
3-hr assimilation period with 1-hr update frequency and 6-hr forecast

Three initialization experiments are performed.

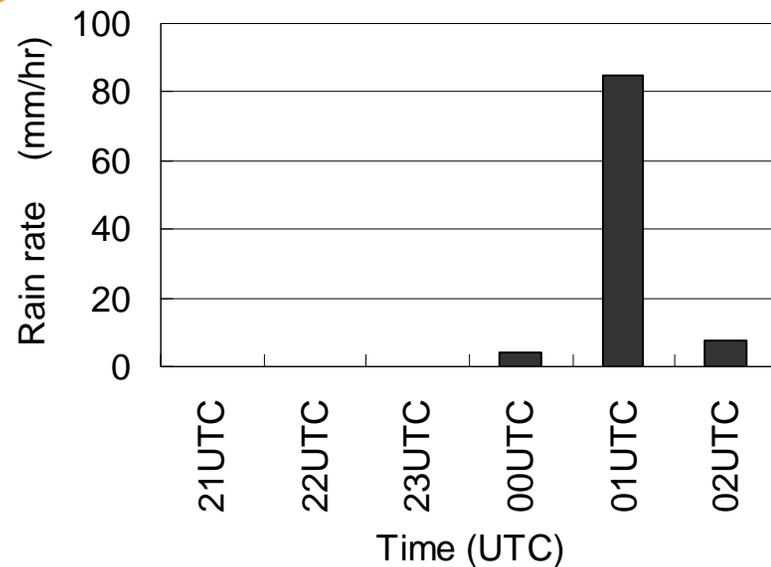
- UNTUNE : IAU with untuned 3DVAR increments**
- RUC : RUC with tuned 3DVAR**
- IAU : IAU with tuned 3DVAR increments**

Heavy Rainfall Case (25 July, 2003)

6-h accumulated rainfall and time series

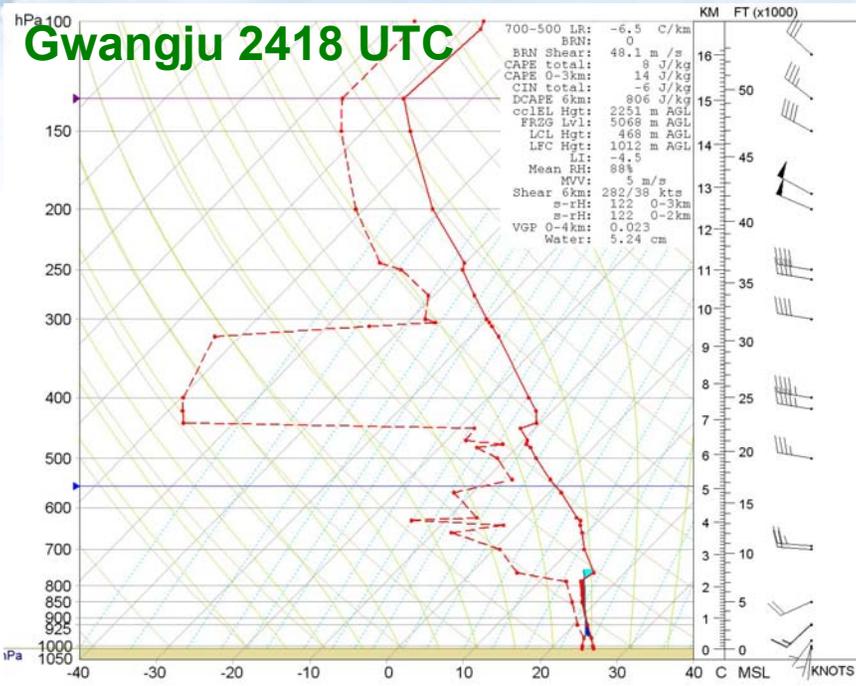


Jeonju (96.7 mm)

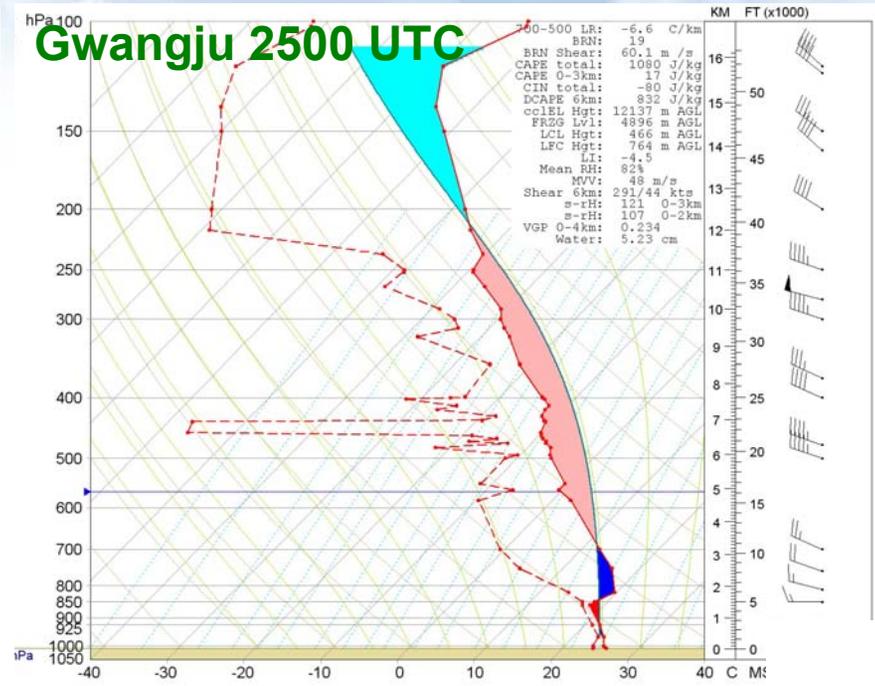


Upper-air station (Kwangju)

Soundings at Kwangju



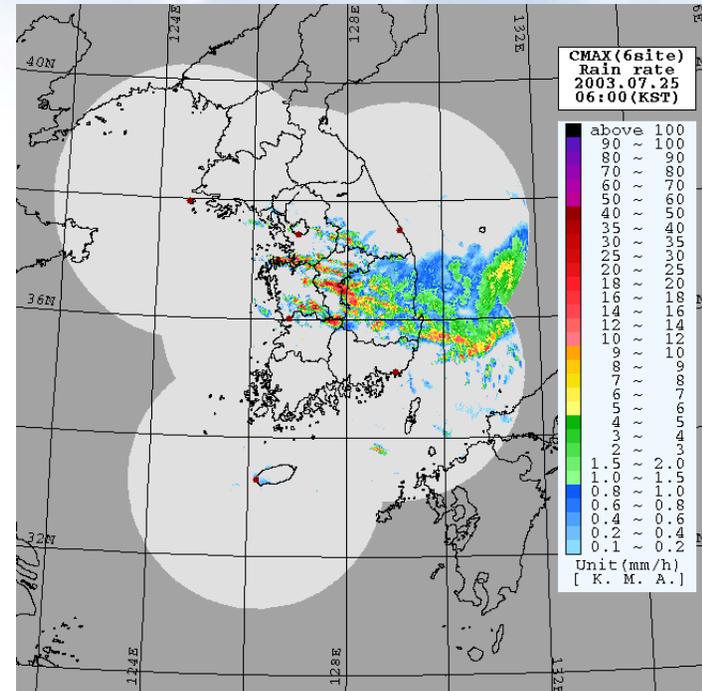
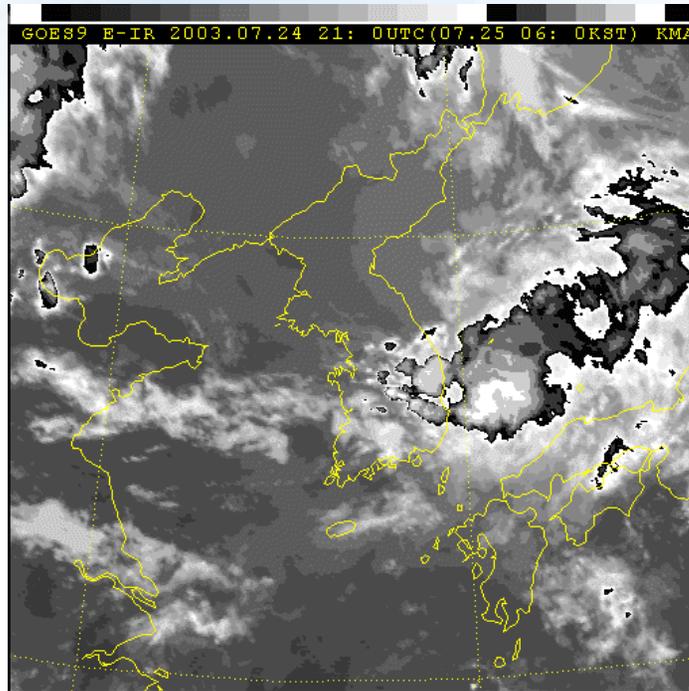
**Cape : 8 J/Kg
LFC : 1012 m AGL**



**Cape : 1080 J/Kg
LFC : 764 m AGL**

CAPE rapidly increased for 6 hours from 18 UTC 24 to 00 UTC 25 July. Mesoscale convective systems with 3 cells developed in the unstable environment.

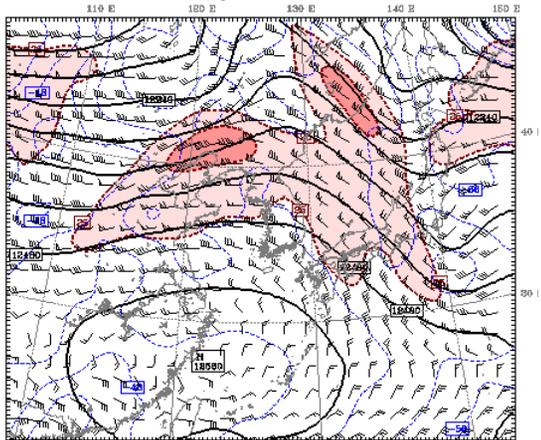
Enhanced IR and Radar images



The initialization experiments are applied to an intense rainfall event accompanying MCSs that has three convective cells.

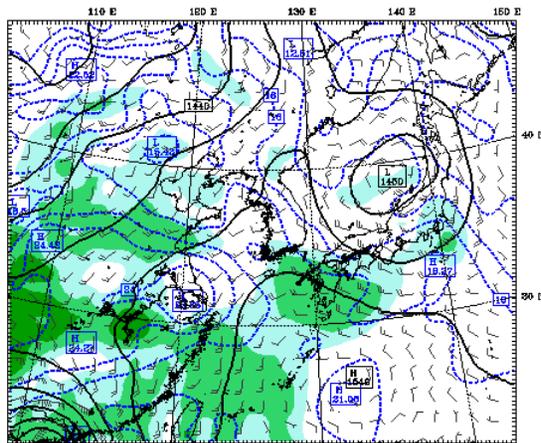
Synoptic Charts (00UTC 24-00UTC 25 July, 2003)

Dataset: CASE2003 4 RIP: 200 Init: 0000 UTC Thu 24 Jul 03
 Fcst: 0.00 h Valid: 0000 UTC Thu 24 Jul 03 (1800 MDT Wed 23 Jul 03)
 Horizontal wind speed at pressure = 200 hPa
 Horizontal wind speed at pressure = 200 hPa
 Temperature at pressure = 300 hPa
 Geopotential height at pressure = 200 hPa sm= 2
 Horizontal wind vectors at pressure = 300 hPa



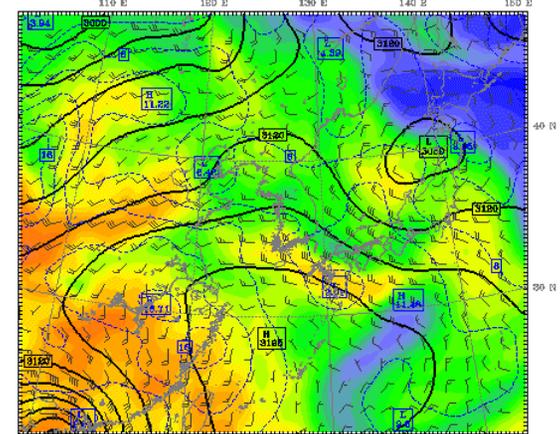
BARB VECTORS: FULL BARB = 5 $m s^{-1}$
 CONTOURS: UNITS=m LOF= 180.00 HIGH= 180.00 INTERVAL= 60.000
 CONTOURS: UNITS= $m^2 s^{-2}$ LOF= -28.500 HIGH= -48.500 INTERVAL= 2.0000
 CONTOURS: UNITS= $m s^{-2}$ LOF= 25.000 HIGH= 97.500 INTERVAL= 12.500

Dataset: CASE2003 4 RIP: 850 Init: 0000 UTC Thu 24 Jul 03
 Fcst: 0.00 h Valid: 0000 UTC Thu 24 Jul 03 (1800 MDT Wed 23 Jul 03)
 Taster vapor mixing ratio at pressure = 850 hPa
 Temperature at pressure = 850 hPa sm= 2
 Geopotential height at pressure = 660 hPa
 Horizontal wind vectors at pressure = 850 hPa



BARB VECTORS: FULL BARB = 5 $m s^{-1}$
 CONTOURS: UNITS=m LOF= 1320.0 HIGH= 1320.0 INTERVAL= 30.000
 CONTOURS: UNITS= $^{\circ}C$ LOF= 6.0000 HIGH= 24.0000 INTERVAL= 2.0000

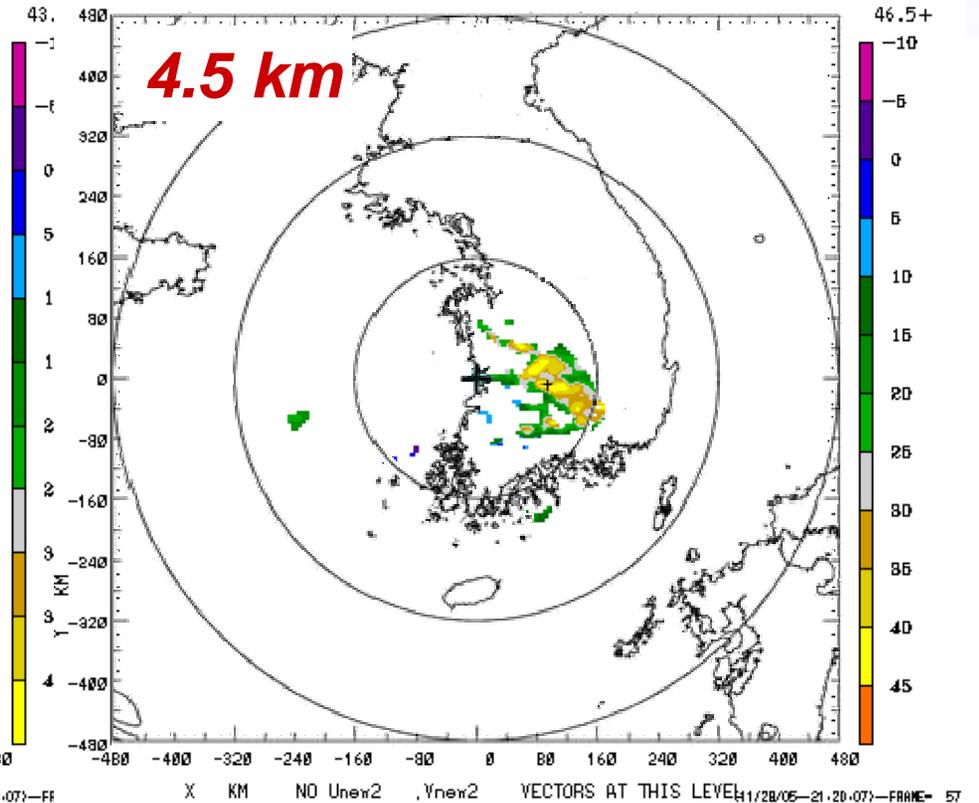
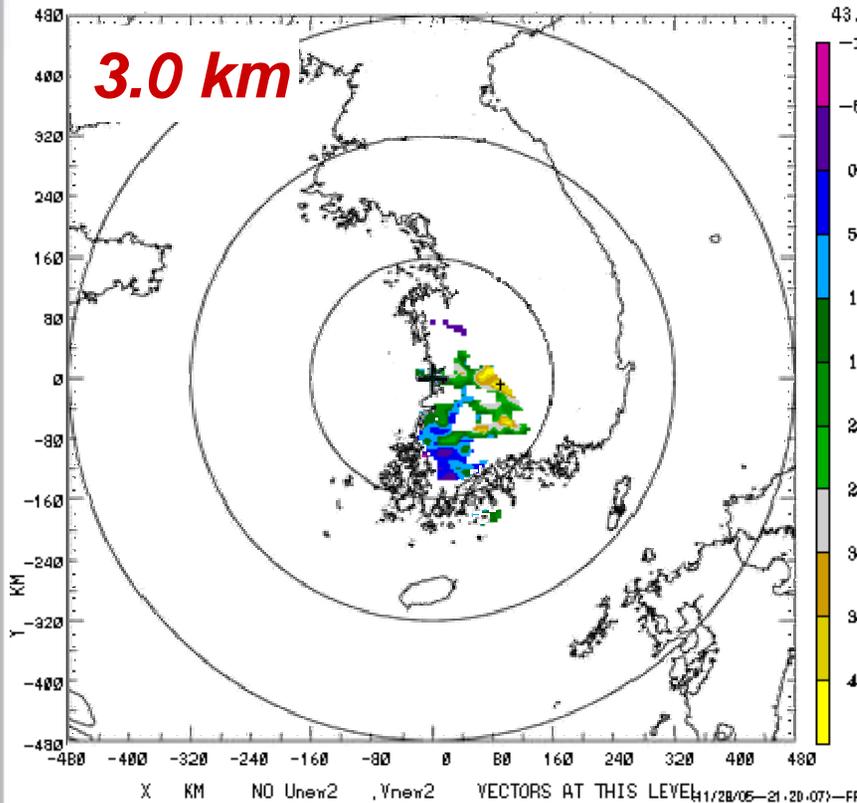
Dataset: CASE2003 4 RIP: 700 Init: 0000 UTC Thu 24 Jul 03
 Fcst: 0.00 h Valid: 0000 UTC Thu 24 Jul 03 (1800 MDT Wed 23 Jul 03)
 equivalent potential temperature at pressure = 700 hPa sm= 2
 geopotential height at pressure = 700 hPa
 horizontal wind vectors at pressure = 700 hPa
 temperature at pressure = 700 hPa



CONTOURS: UNITS= $^{\circ}C$ LOF= -6.0000 HIGH= 16.0000 INTERVAL= 8.0000
 BARB VECTORS: FULL BARB = 5 $m s^{-1}$
 CONTOURS: UNITS=m LOF= 2970.0 HIGH= 3190.0 INTERVAL= 30.000

Initial radar data (reflectivity and radial velocity)

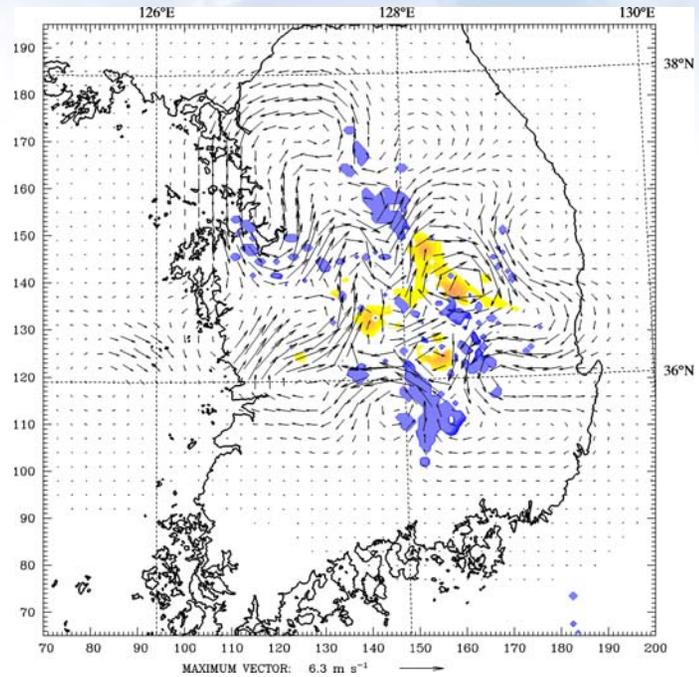
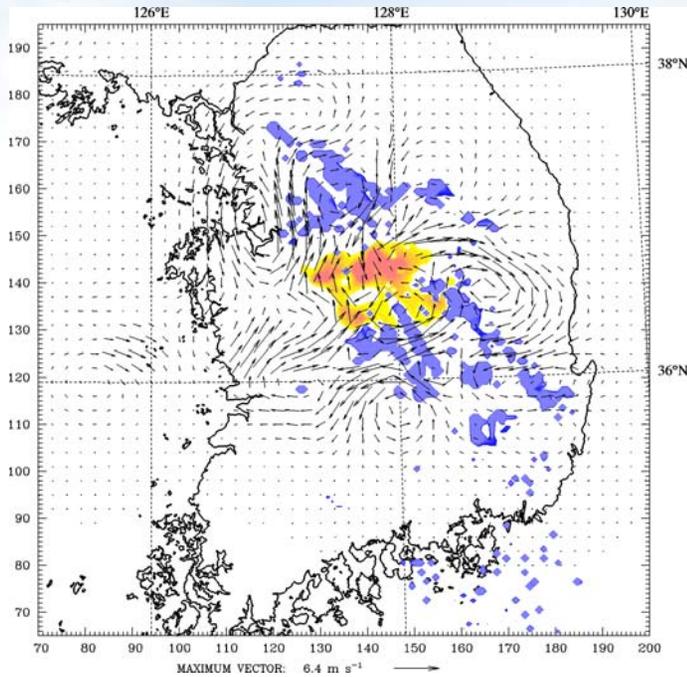
03/07/24 18:00:38-18:05:47 COMBIN Z = 3.00 KM SYN 03/07/24 18:00:38-18:05:47 COMBIN Z = 4.50 KM SYNDZ
 (AS OF 11/28/05) ORIGIN=(0.00, 0.00) KM X-AXIS= 90.0 D (AS OF 11/28/05) ORIGIN=(0.00, 0.00) KM X-AXIS= 90.0 DEG
 from RKSG & RKJK from RKSG & RKJK



Increment Fields

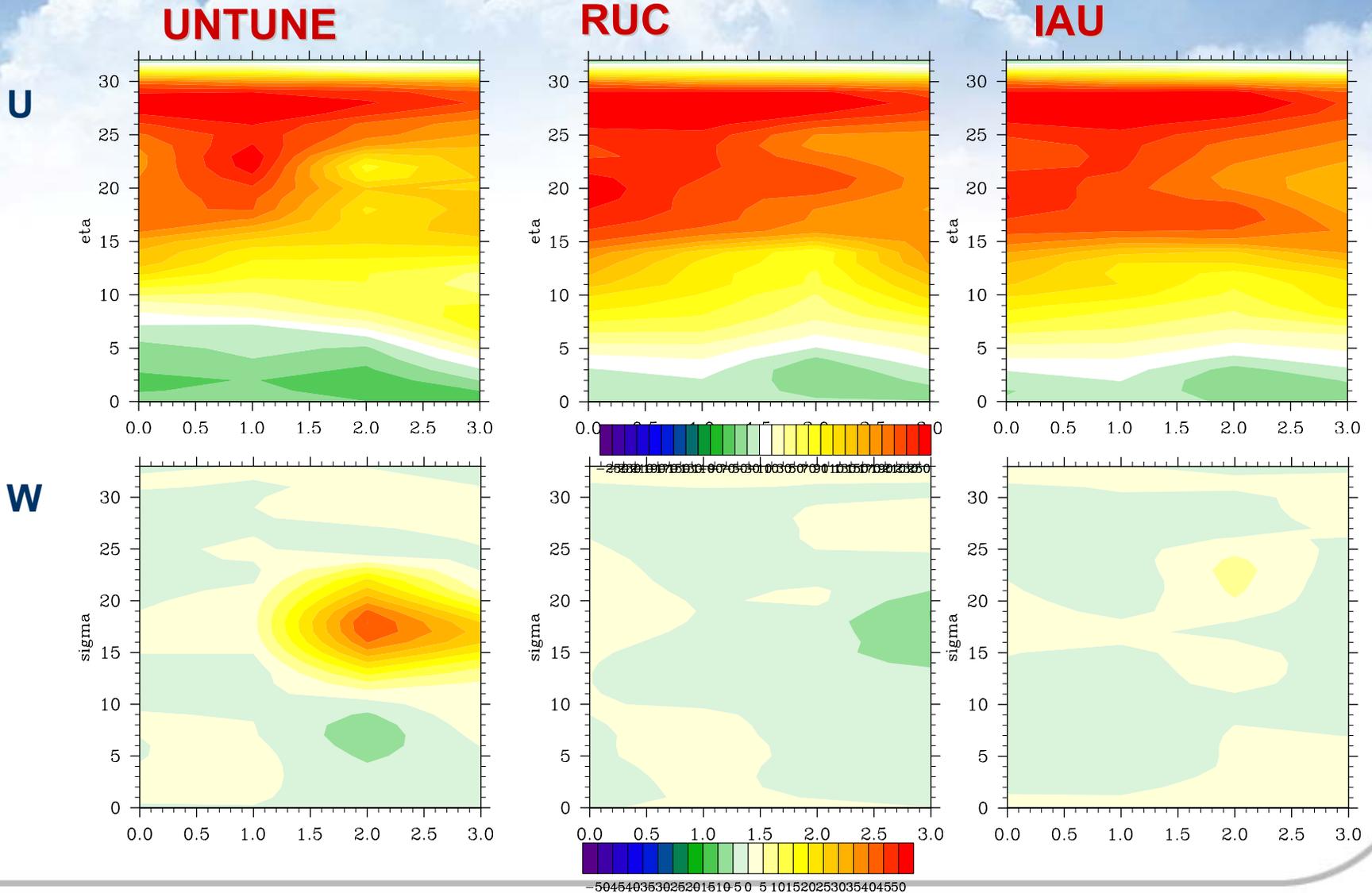
UNTUNE

IAU

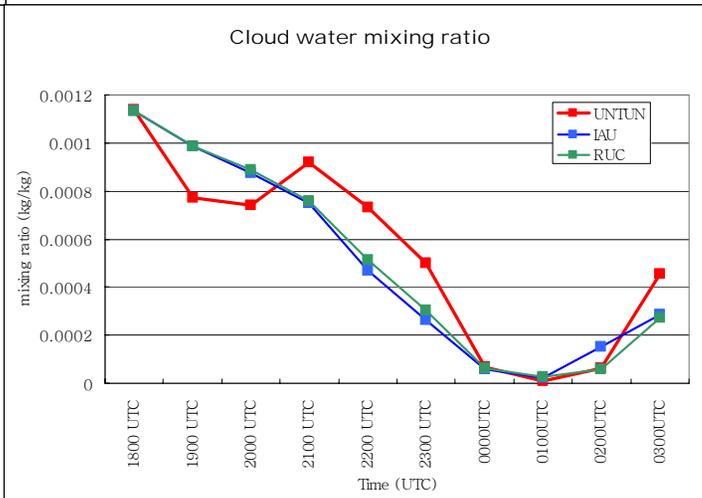
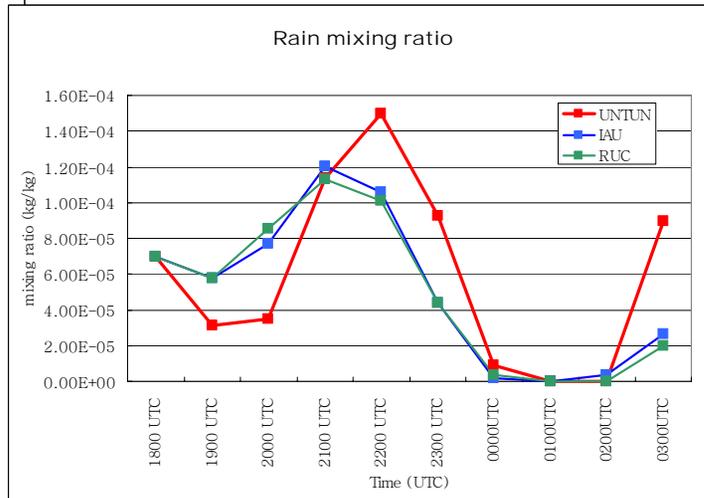
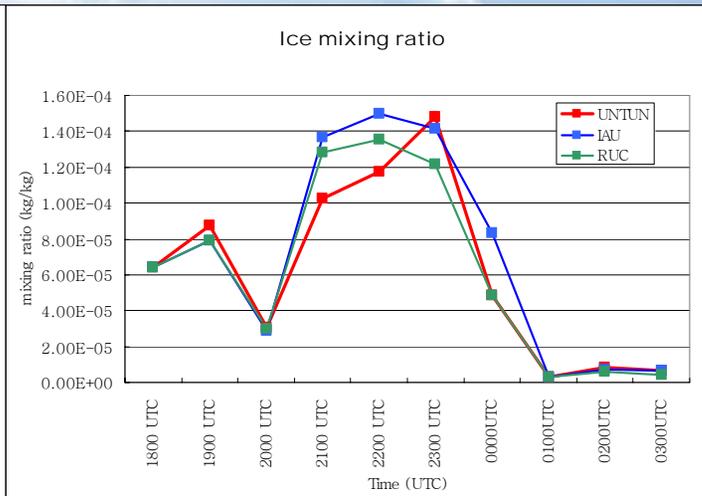
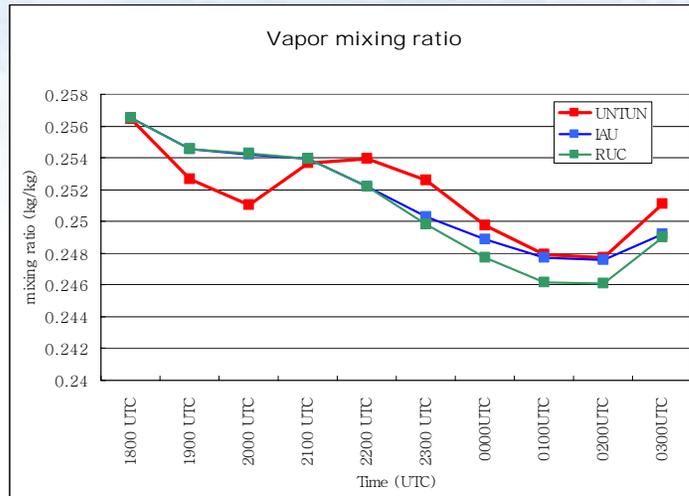


Water Vapor mixing ratio (g/kg)
Wind vector (m/sec)

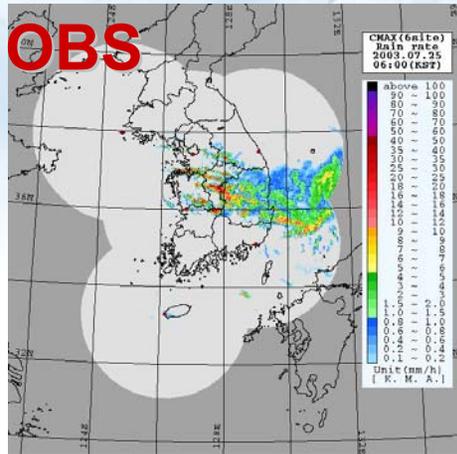
Time-height section of area-mean u- and w- wind



Time series of hydrometeors

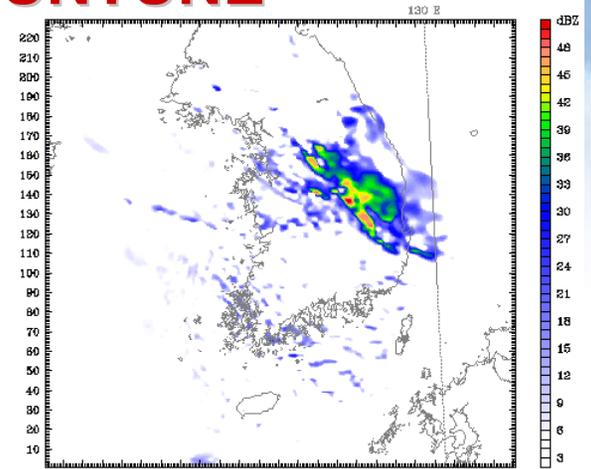


Radar images (dbz)



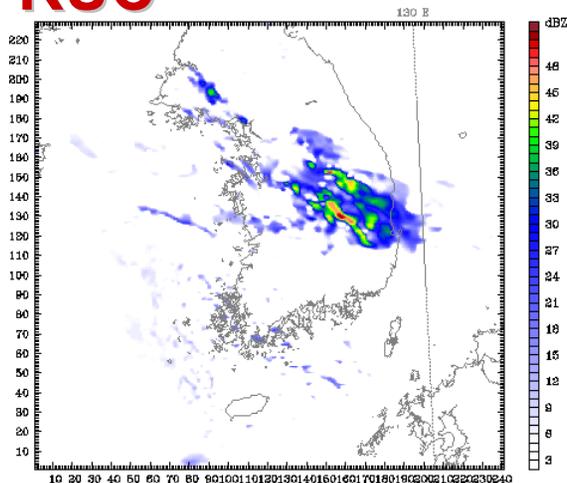
Dataset: IAU UNTUNED3H 1 RIP: dbz hor IAU UNTU
 Fcst: 2 h Valid: 22 UTC Thu 24 Jul 03 (07 LST Fri 25 Jul 03)
 Reflectivity at k-index = 33 sm= 1

UNTUNE



Dataset: RUC3H 1H 200307 RIP: dbz hor RUC 2003
 Fcst: 1 h Valid: 22 UTC Thu 24 Jul 03 (07 LST Fri 25 Jul 03)
 Reflectivity at k-index = 33 sm= 1

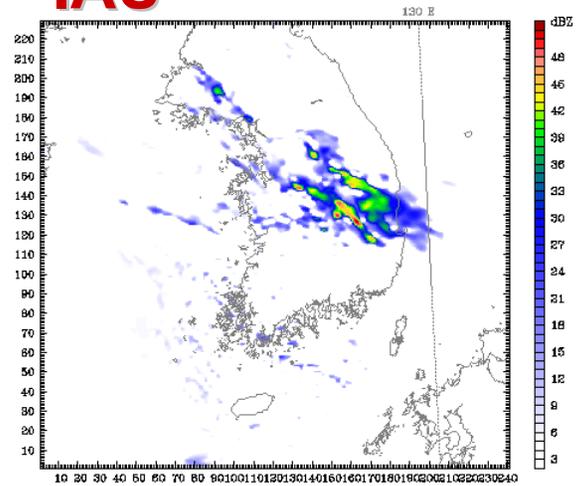
RUC



Model Infr: F2.9.3 No Cumulus YSU PBL Lin et al 3.3 km, 33 levels, 10 sec

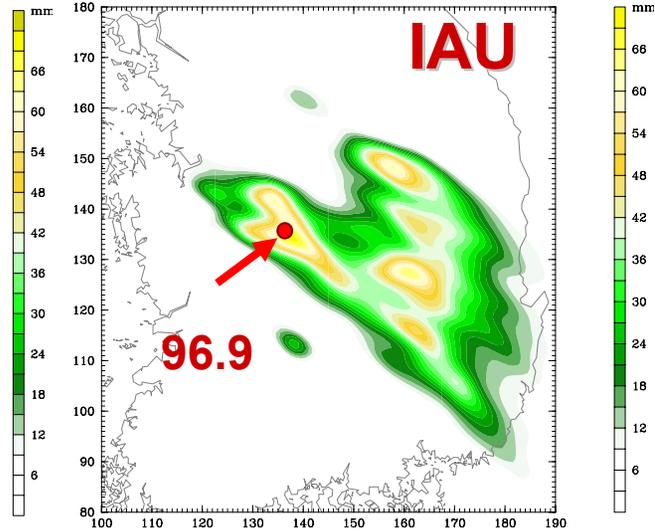
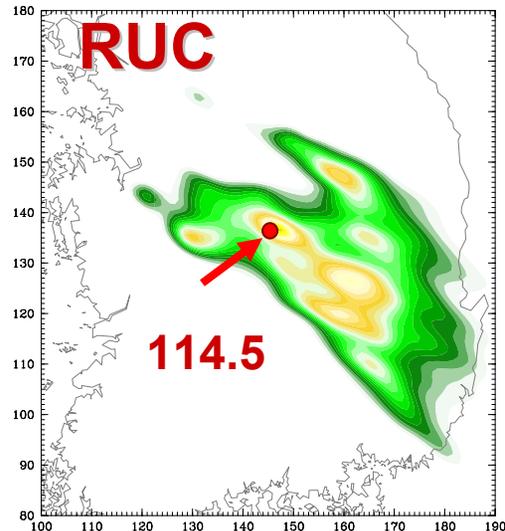
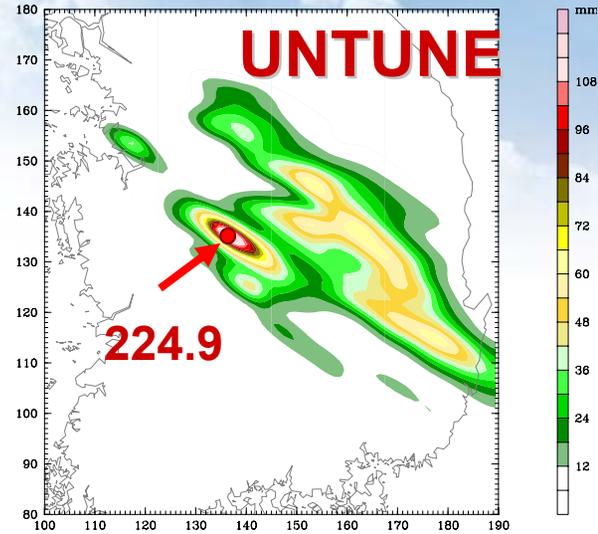
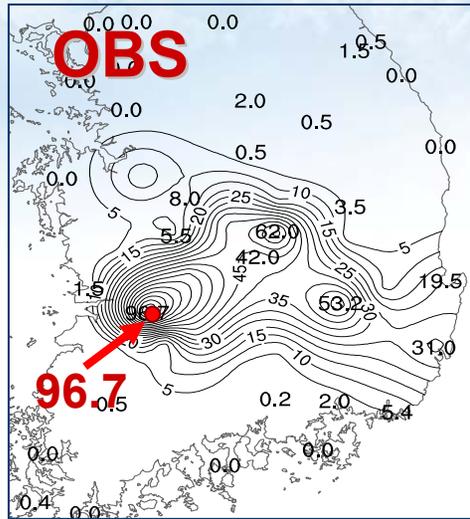
Dataset: IAU3H 1H 200307 RIP: dbz hor IAU3H 1H
 Fcst: 2 h Valid: 22 UTC Thu 24 Jul 03 (07 LST Fri 25 Jul 03)
 Reflectivity at k-index = 33 sm= 1

IAU

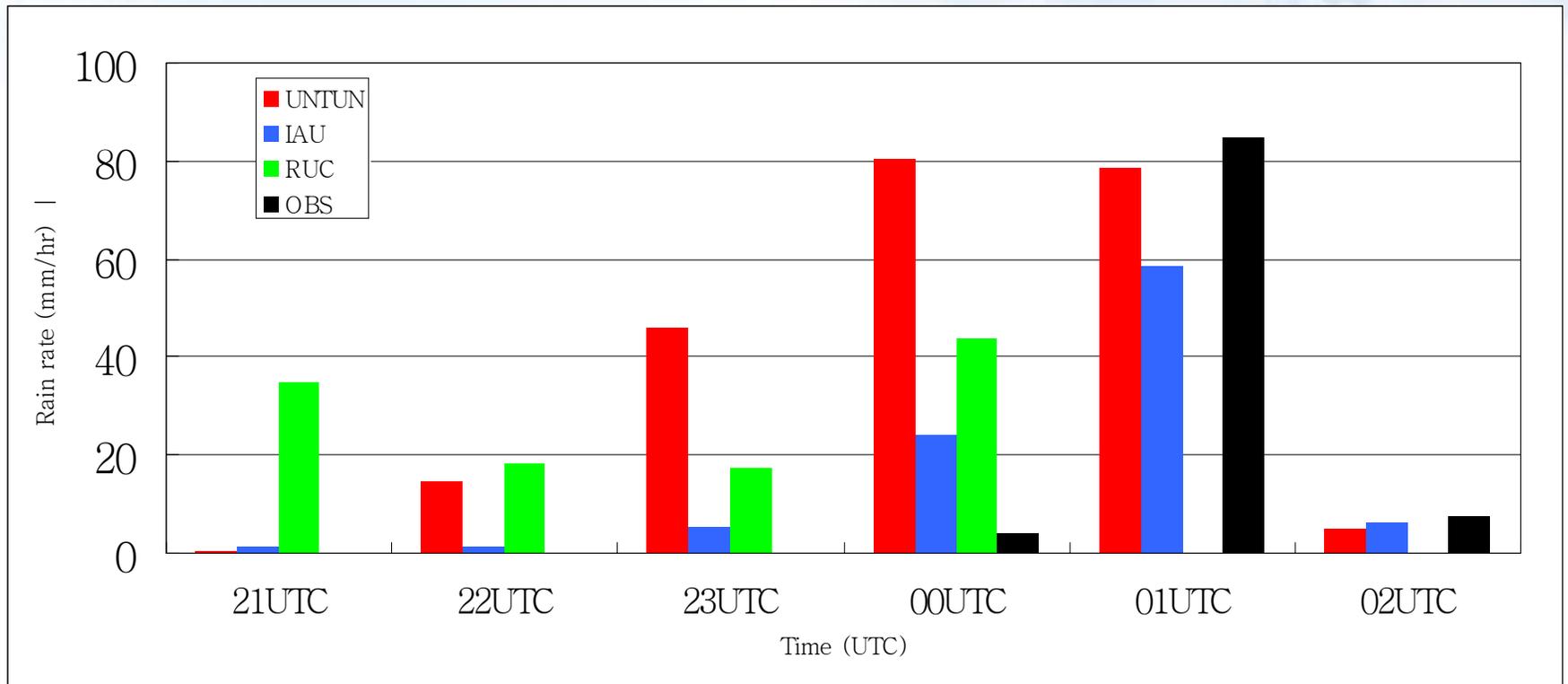


Model Infr: F2.9.3 No Cumulus YSU PBL Lin et al 3.3 km, 33 levels, 10 sec

6h accumulated rainfall amount (mm)

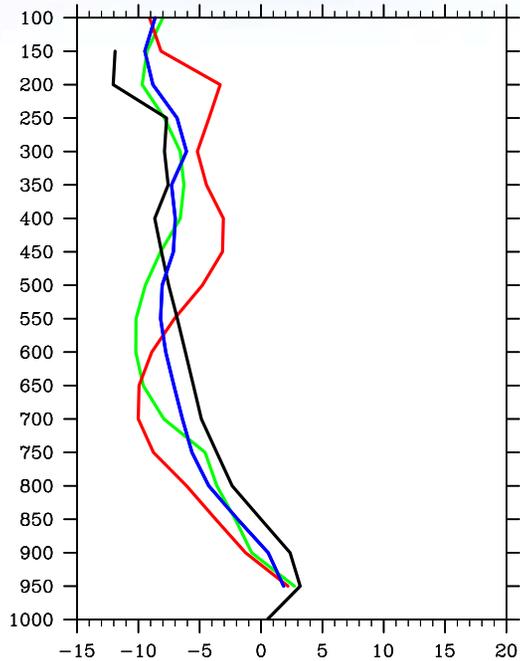


Time series of rainfall amount at maximum point

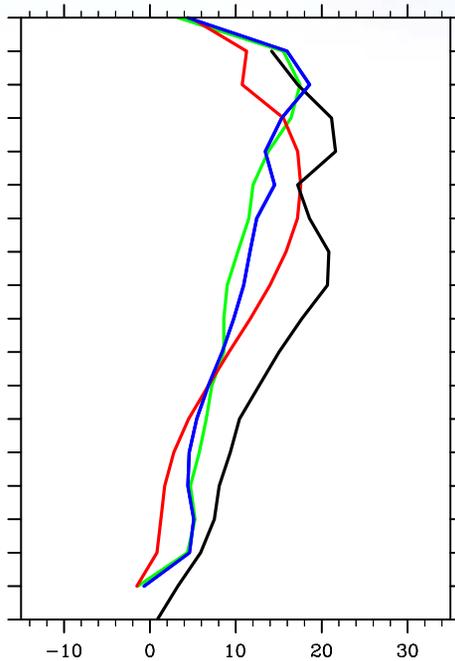


Profile of u , v , RH and T at 00 UTC 25 July

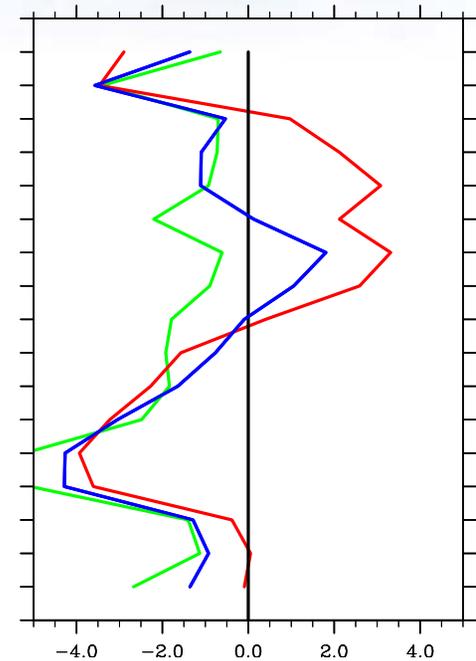
Vertical profile of v_a



Vertical profile of u_a



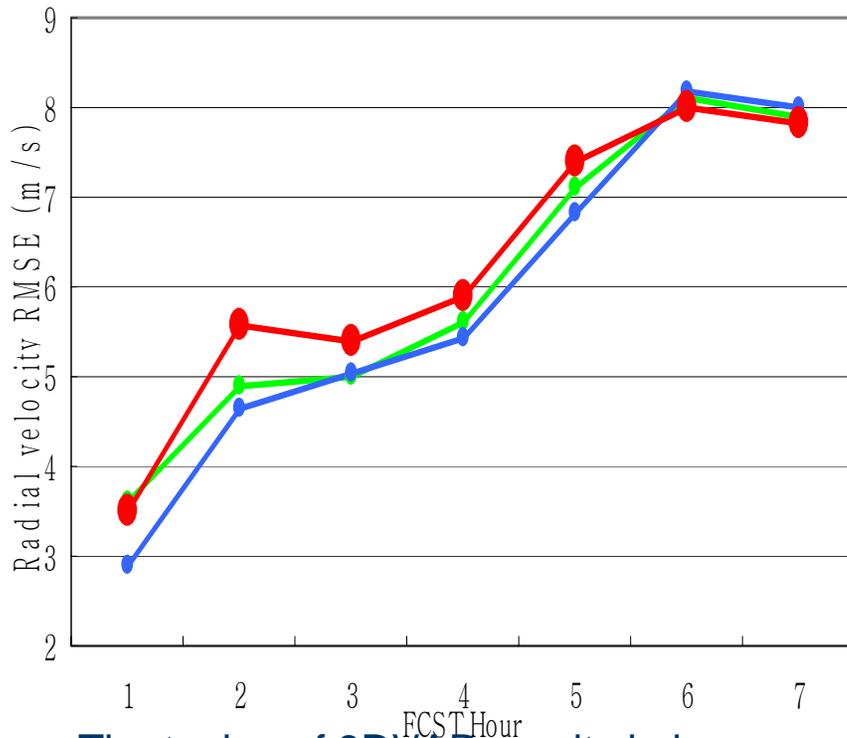
t_c difference from OBS



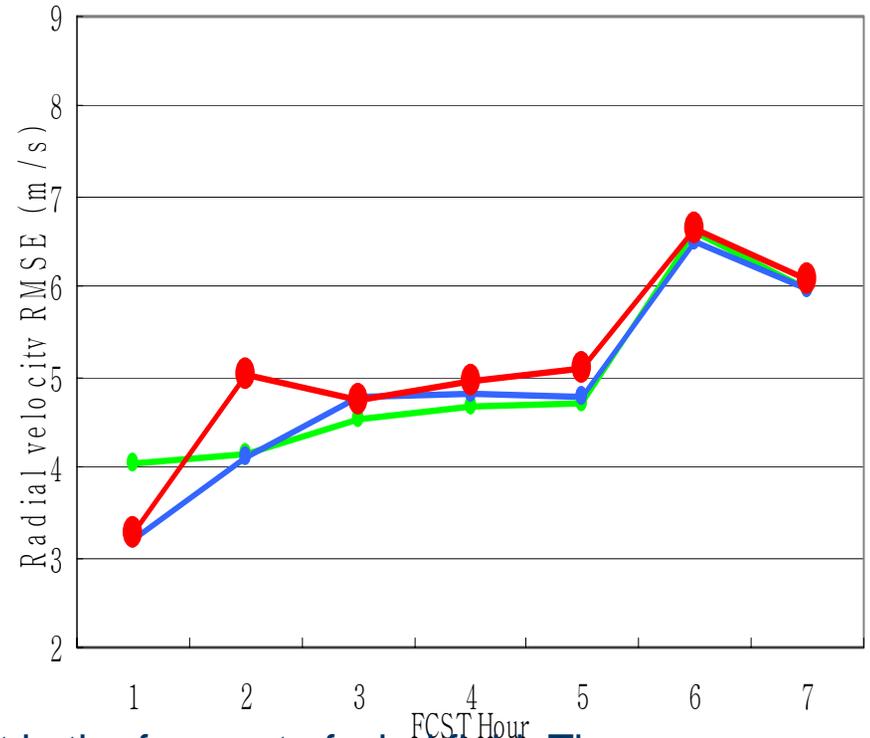
— : UNTUNE
— : IAU
— : RUC
— : OBS

RMSEs of radial velocity

RMSE of radial velocity at 3.0 km



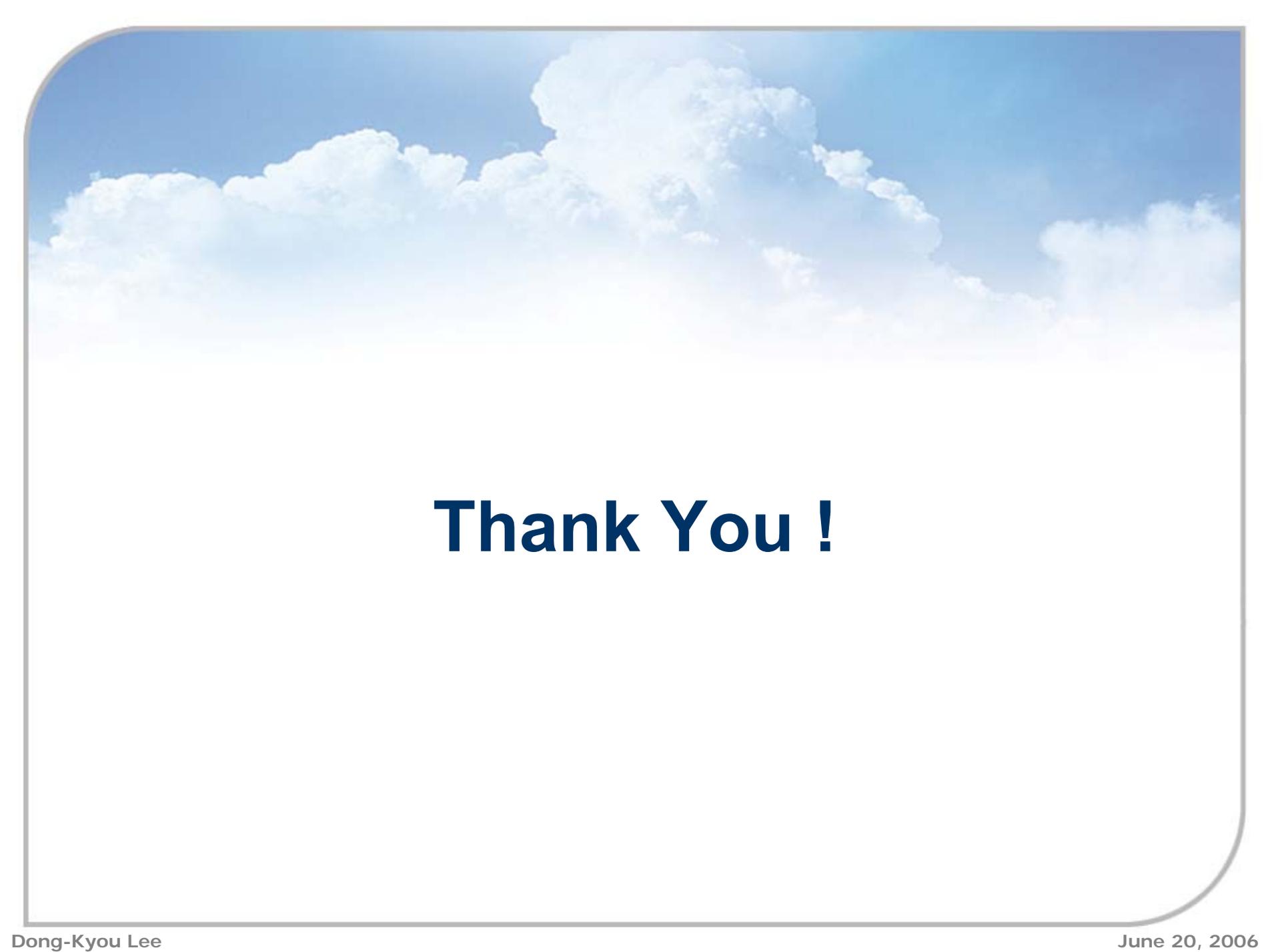
RMSE of radial velocity at 4.5 km



The tuning of 3DVAR results in improvement in the forecast of wind field. The improvement is effective up to 4-hour forecast.

Conclusions

- In this study, radar data has much shorter scale-lengths in 3DVAR compared to the typical synoptic observations. The scale length is 9 km for radial velocity, and 4 km for reflectivity.
- The error factors show that the error for radial velocity (2 m/s) is overestimated (70 % of the currently used error), and that for reflectivity (5 dBZ) is underestimated (190 % of the currently used error). Therefore, the error factors rapidly converge within one iteration.
- In the radar data assimilation, the tuned 3DVAR improves the maximum rainfall amount that is in better agreement with observation than that of the UNTUNE. The RMSEs of the tuned 3DVAR (RUC and IAU) are also smaller than that of the untuned 3DVAR. It is necessary to tune the scale lengths for the 3 DVAR assimilation of radar data.
- Effectiveness of model forecast by the assimilation of radar data appear to be within 3~4 hours.

A photograph of a bright blue sky filled with large, fluffy white cumulus clouds. The clouds are dense and appear to be rising or billowing, creating a sense of depth and volume. The sky is a clear, vibrant blue, and the overall scene is bright and cheerful.

Thank You !





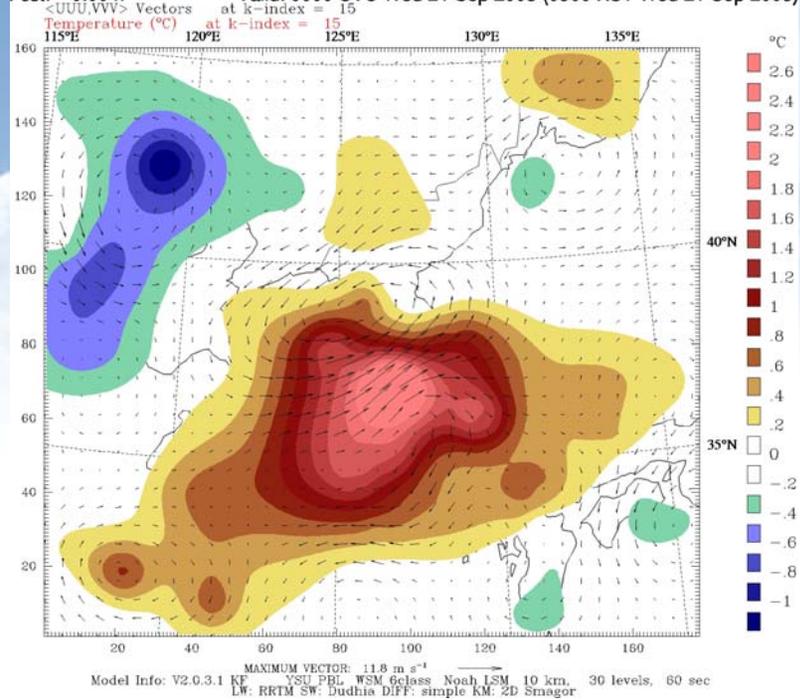
Grid nudging and IAU for WRF

$$\frac{\partial \mu \chi}{\partial t} = F_{\chi}(\mathbf{x}, t) + \left\{ \begin{array}{l} \text{Nudging} \\ G_{\chi} W(\mathbf{x}, t) \mu (\chi - \chi_{anal}), \end{array} \right. \left. \begin{array}{l} \text{IAU} \\ G_{\chi} \mu \chi'_{anal} \end{array} \right\}$$

$W(x, t) = (w_{\eta} \cdot w_t)$, G_{χ} : nudging coefficient

- ❖ For grid nudging, we have implemented time-variant nudging (=1) and target nudging (=3)
- ❖ Surface (in planetary boundary layer) nudging (or IAU) is not yet implemented
- ❖ The nudging of U, V, and T is currently tested

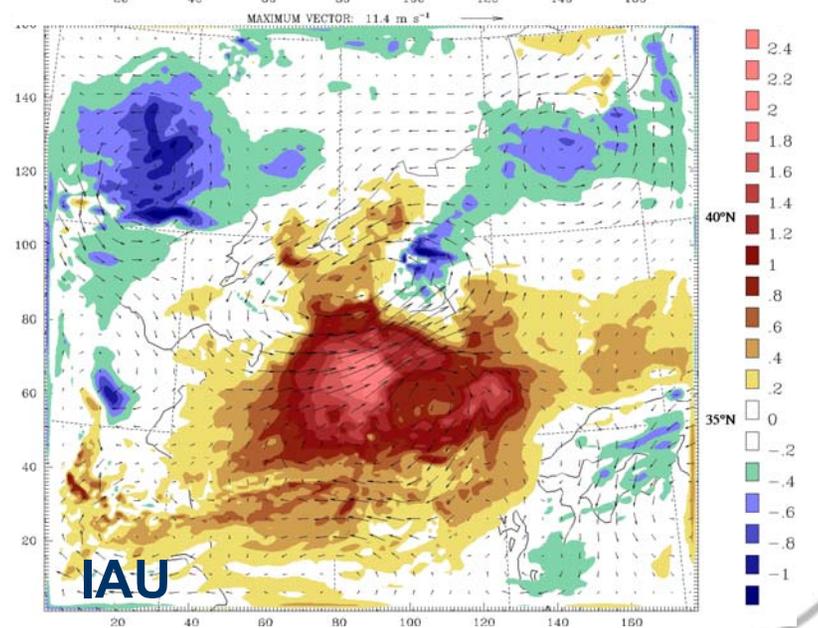
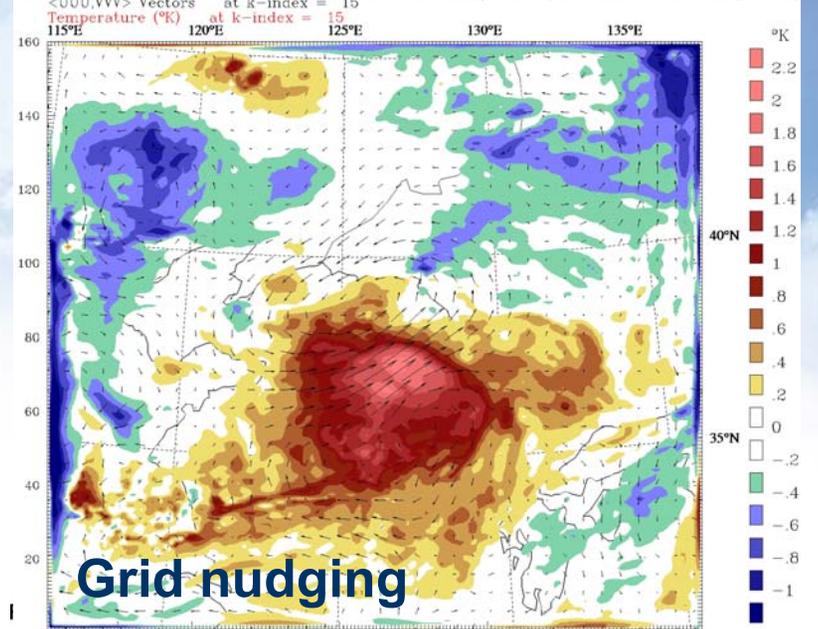
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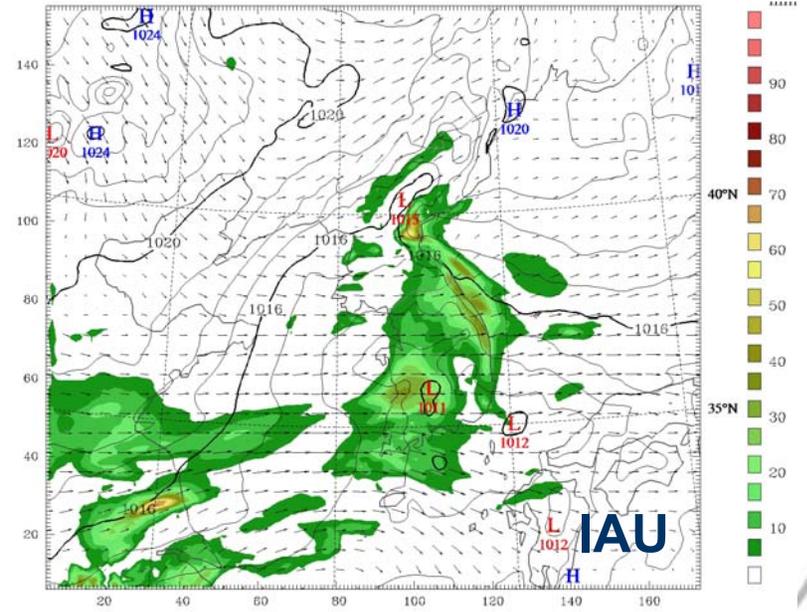
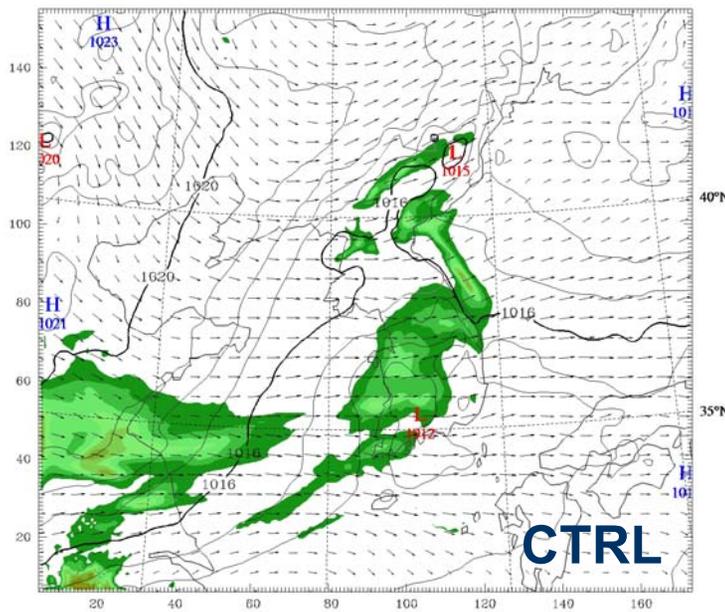
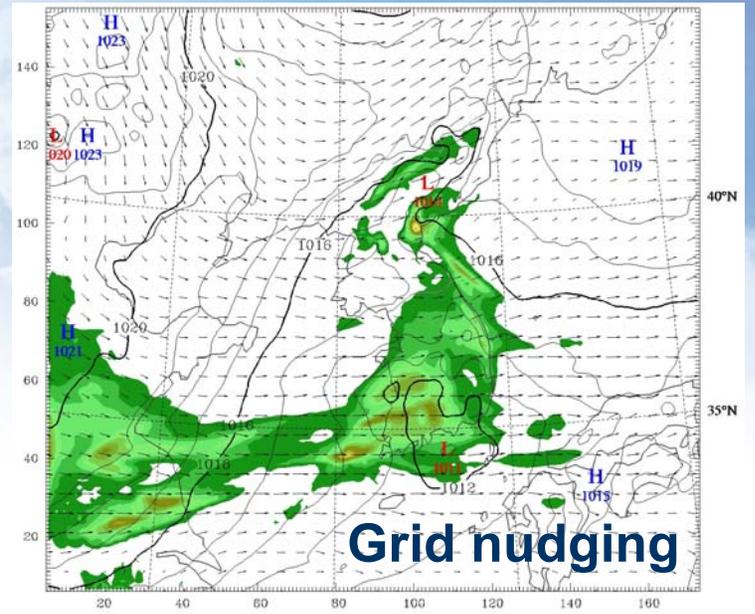
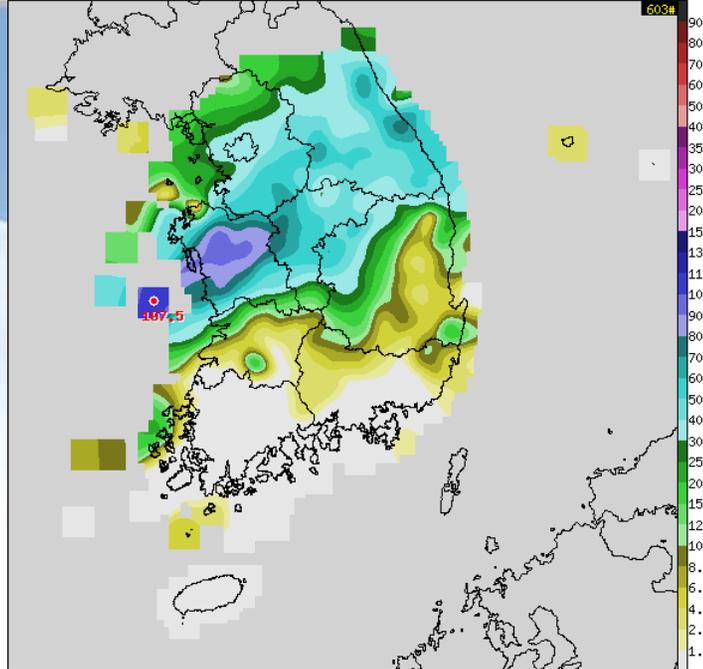
Initial increment

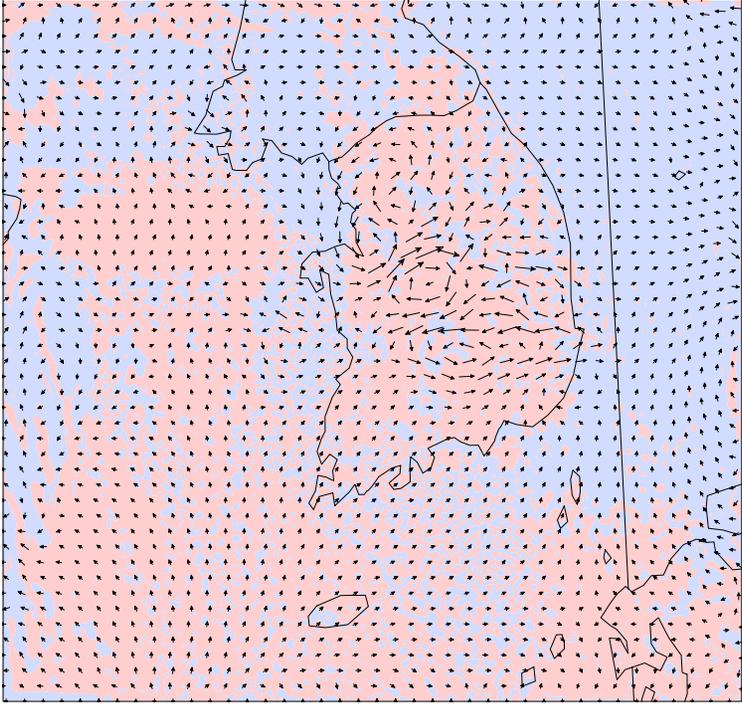
**Verification of Grid nudging
 and IAU**

Init: 2100 UTC Tue 20 Sep 2005
 Fcst: 0.00 h Valid: 0000 UTC Wed 21 Sep 2005 (0900 KST Wed 21 Sep 2005)



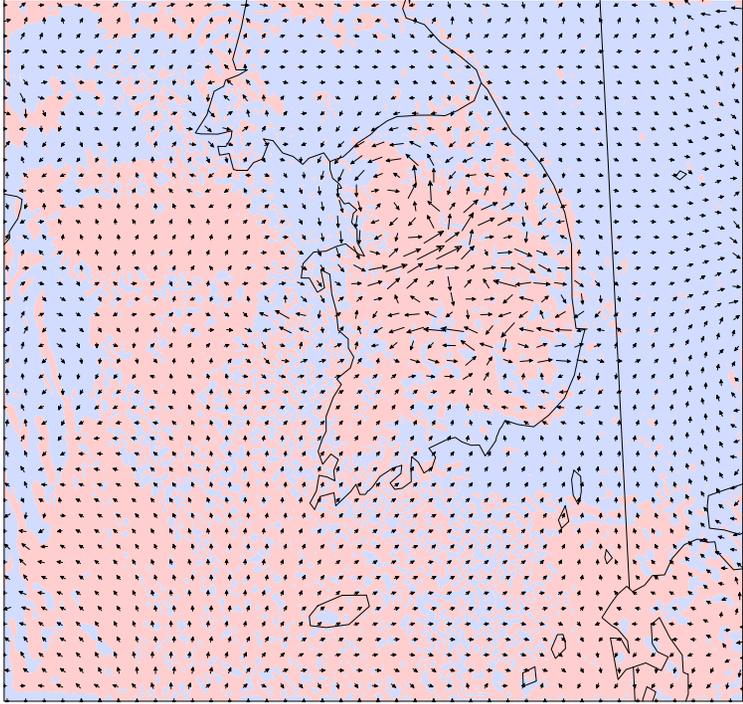
RAIN 2005.09.21.09:00-09.21.21:00





130E

T(C)



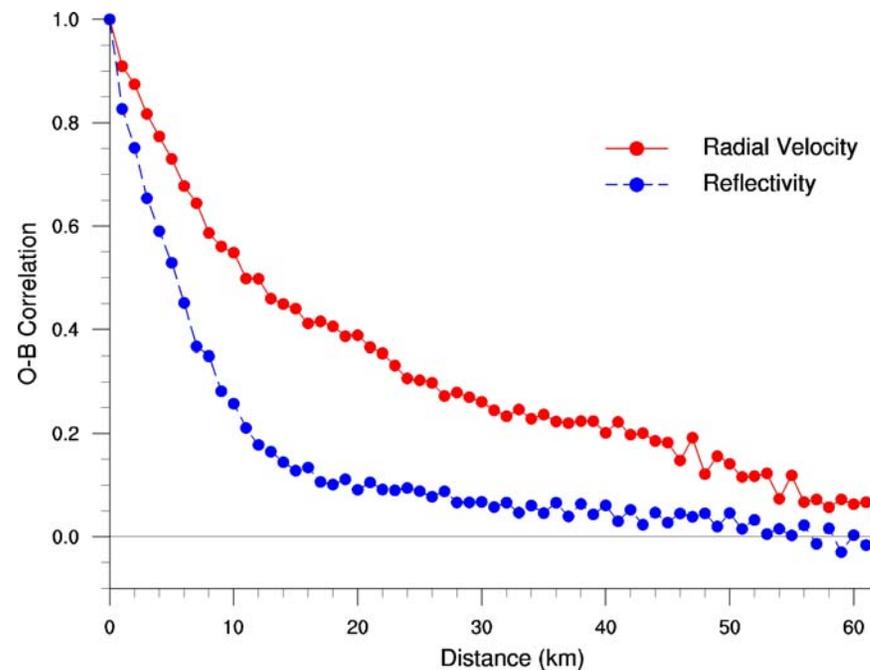
130E



Tuning of Background Scale-lengths

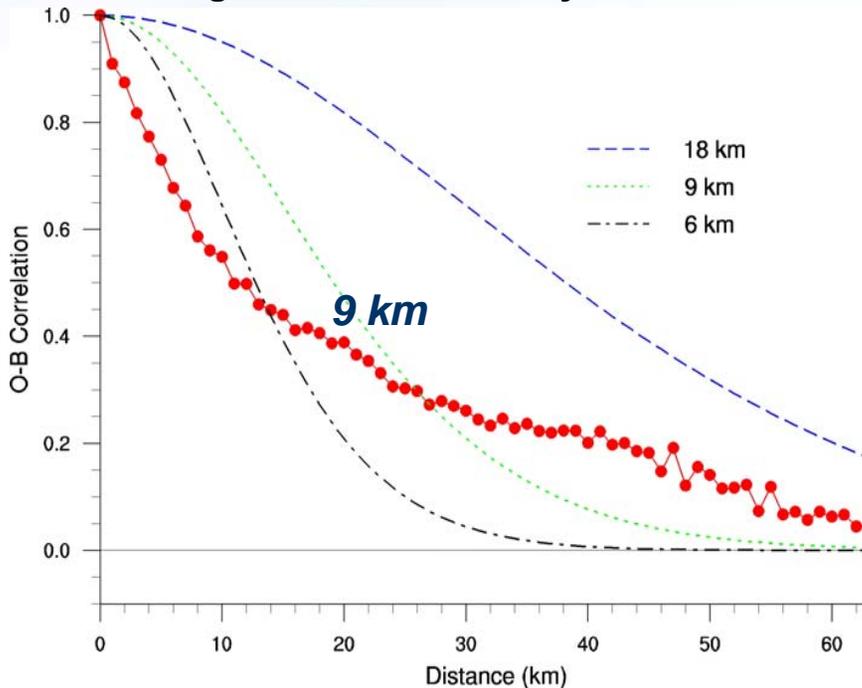
The O-B of radar data is calculated from two MCS cases and one frontal case. The O-B correlation decreases according to short distance. It means that the radar observation detects meso- as well as micro- scale phenomena. The locality of the radar reflectivity is higher than radical velocity

O-B Statistics

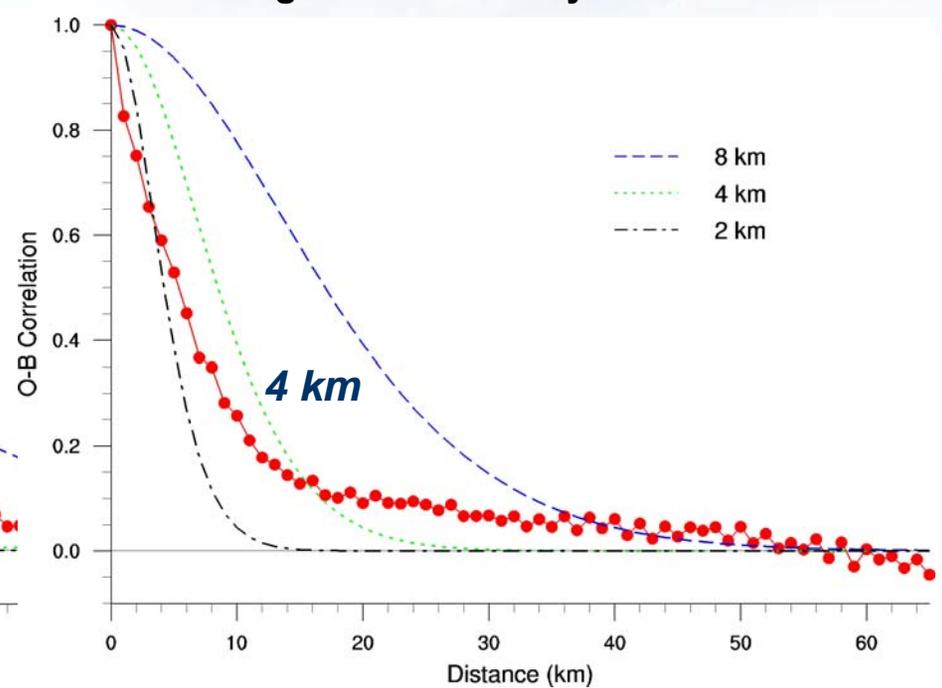


Scale-length

Scale-length for radial velocity



Scale-length for reflectivity



- The locality of the radar observation can be reflected by tuning the scale length of the background error using a recursive filter.
- 9 km and 4 km of scale lengths are proper for radial velocity and reflectivity, respectively.

3DVAR Error Tuning

- “a posteriori diagnosis on the minimized cost function to determine a priori parameters” (Desroziers and Ivanov, 2001).

Methodology of 3DVAR Error tuning

- ❖ The expectation values of the observation and background components of a cost function are given by

$$E(J^o) = \frac{1}{2}(p - \text{Tr}(\mathbf{HK})) \quad , \quad E(J^b) = \frac{1}{2} \text{Tr}(\mathbf{KH})$$

where p is the number of observations, H is the linearized observation operator, and K is the Kalman-gain matrix

$$\left(\mathbf{K} = \mathbf{BH}^T (\mathbf{HBH}^T + \mathbf{R})^{-1} = \mathbf{P}^a \mathbf{H}^T \mathbf{R}^{-1} \right).$$

Methodology of 3DVAR Error tuning

- ❖ The expectation value for the total cost function is given by a half of the number of effective observations (Desroziers and Ivanov, 2001).

$$E(J) = E(J^b) + E(J^o) = p / 2$$

- ❖ To adjust the real cost function to the expectation value, error factors will be applied to each cost function terms and to each observation types.

$$J = \frac{1}{s_b^2} J^b + \frac{1}{s_o^2} J^o$$

Methodology of 3DVAR Error tuning

- ❖ Error factors are determined iteratively,

$$s_o^2|^{i+1} = \left(\frac{J^o|^{i+1}}{E(J^o)} \right) \quad s_b^2|^{i+1} = \left(\frac{J^b|^{i+1}}{E(J^b)} \right)$$

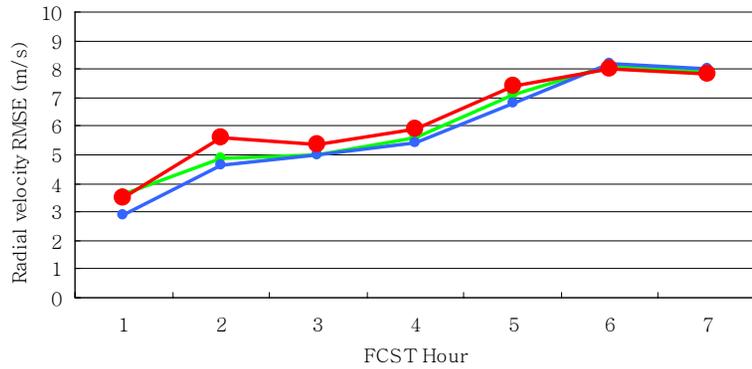
where i means iteration number. The expectation values are computed using a randomized estimation of the trace of matrix, HK and KH (Desroziers and Ivanov, 2001).

3DVAR Error Factors

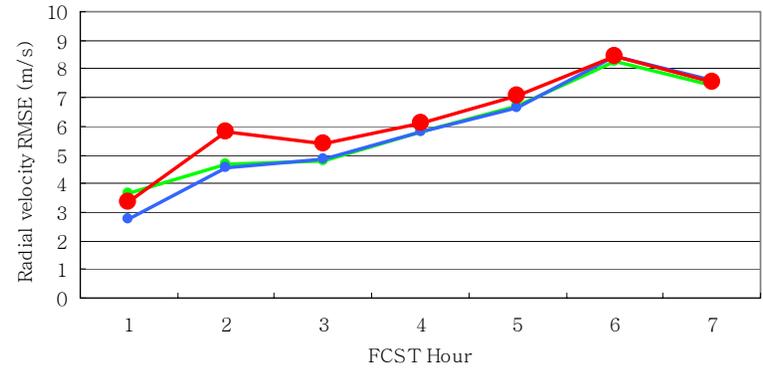
Observation	P	S(0)	S(1)	S(2)	S(3)
Radial Velocity	16820	1.00	0.72	0.69	0.68
Reflectivity	26977	1.00	1.95	1.95	1.95
J ^b		1.00	1.04	1.11	1.13

The number of effective radar radial velocity and reflectivity is 16,820 and 26,977, respectively, and the tuning parameters of the observation error converge.

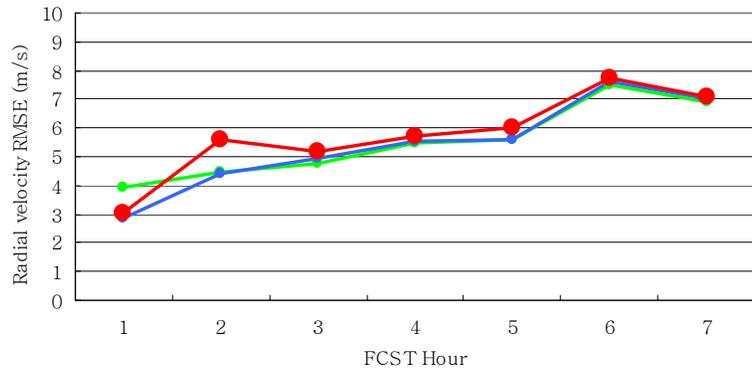
RMSE of radial velocity at 3.0 km



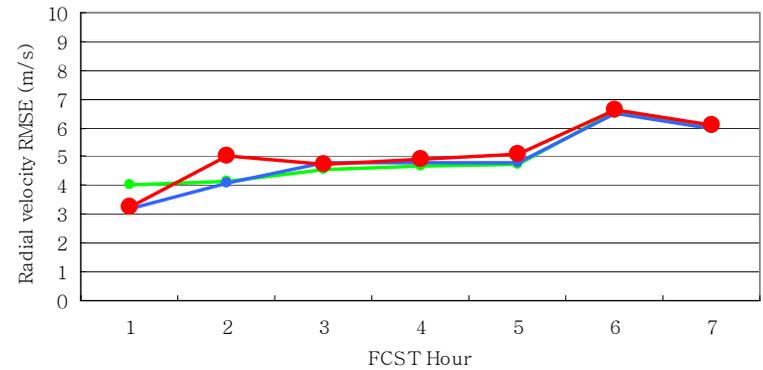
RMSE of radial velocity at 3.5 km



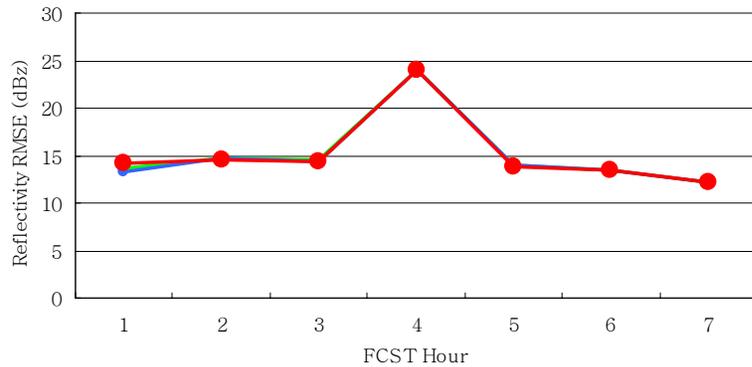
RMSE of radial velocity at 4.0 km



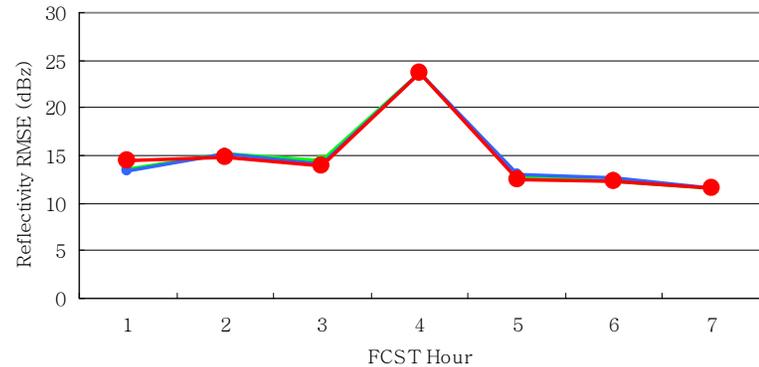
RMSE of radial velocity at 4.5 km



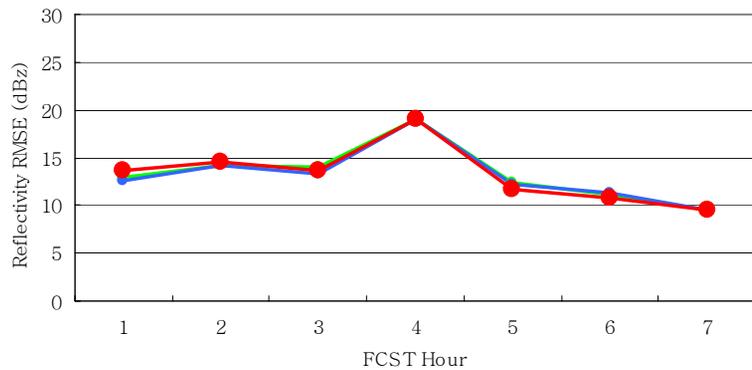
RMSE of reflectivity at 3.0 km



RMSE of reflectivity at 3.5 km



RMSE of reflectivity at 4.0 km



RMSE of reflectivity at 4.5 km

