

CCN and IN updates to the Thompson microphysics scheme

Trude Eidhammer and Greg Thompson

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NATIONAL CENTER FOR ATMOSPHERIC RESEARCH

Outline



Short introduction of the linking of aerosols to droplet and ice formation in the Thompson microphysics scheme in WRF.

Case study of dust – IN activation

Linking aerosols to CCN and IN in WRF

Crawl:

Constant cloud droplet number that influences precipitation. Ice initiation based on temperature or supersaturation alone

Walk:

Creating ice and droplet number based on simple, but realistic aerosols

Run:

Full integration with WRF-Chem and multiple aerosol species

CCN activation



Physical based droplet activation must be solved analytical and is not a simple computation.

Parameterization based on measurements N_{CCN} = cS^k (used in many models)

Parameterizations based on parcel model simulations (used in for example WRFchem)

Look up tables from parcel model simulations (used in for example RAMS and in the new Thompson scheme)

Droplet activation look-up table



- 1) aerosol concentration
- 2) updraft velocity
- 3) temperature
- 4) hygroscopicity (kappa)
- 5) aerosol mean radius
- 6) Standard deviation of size distribution
- 7) Condensation coefficient

Look-up table of activated fraction of aerosols created using parcel model by Feingold and Heymsfield (1992) and changes by T. Eidhammer & S. Kreidenweis

N _a (cm⁻³)	w (m/s)	T (°C)	k	r (µm)
10	0.01	-30	0.2	0.01
31.6	0.0316	-20	0.4	0.02
100	0.1	-10	0.6	0.04
316	0.316	0	0.8	0.08
1000	1	10		0.16
3160	3.16	20		
10000	10	30		
	31.6			
	100			

Ice hat leation peragesterization physical ets eme



$$N_{IN,T} = a(273.15-T)^{b}(n_{aer,0.5})^{(c(273.16-T)+d)}$$

a=0.0000594, b=3.33, c=0.0264, d=0.0033

Aerosol input



- Aerosol number concentration: Dust > 0.5 µm, and sulfate and sea-salt.
- Static file for input and boundary conditions: GOCART monthly average global mean concentration (2000-2007) 1x1.25 degree resolution
- Use a modified *mozbc* software to transport GOCART data to WRF grid. (see P.77 Pfister for information about *mozbc*)

Dust Module in the Thompson Scheme_{NCAR}

- Assume constant dust distribution shape (D = 0.8 μ m and σ = 2 (standard deviation))
- Emission of dust is included. Dependent on land surface, soil moisture and wind speed (need evaluation).
- Wet deposition included (main source of deposition for dust sizes important for ice nucleation).

Test case of dust/IN:

Precipitation in North and Western Colorado with heavy dust scavenging April 28-29, 2010



http://summitvoice.files.wordpress.com/2010/04/dust4.jpg

Summit county, Colorado

00Z 29 Apr 2010

University of Wyoming



McIDAS images of the GOES-13. Dust plume originate from White Sands, New Mexico

Dust Surface Concentration (ug/m**3) for 06:00Z 29 Apr 2010



NAAPS (Navy Aerosol Analysis and Prediction System) Global Aerosol Model

Dust > 0.5 micrometer (#/cm3)





Dust > 0.5 micrometer (#/cm3)

.2

.6



4

8

20

60 100

Dust Surface Concentration (ug/m**3) for 18:00Z 29 Apr 2010



Fri Apr 30 03:31:00 2010 UTC NRL/Monterey Aerosol Modeling

Hourly precipitation

Surface dust concentration

NCAR

REAL-TIME WRF

Init: 2010–04–28_12:00:00 Valid: 2010–04–29_21:00:00

Hourly preciptitation (mm)





REAL-TIME WRF

Init: 2010-04-28_12:00:00 Valid: 2010-04-29_21:00:00

Dust > 0.5 micrometer (#/cm3)





Dust concentration with height.



REAL-TIME WRF

Init: 2010-04-28_12:00:00 Valid: 2010-04-29_18:00:00

Dust > 0.5 micrometer (#/cm3)









Ice crystal concentration for different schemes and dust concentration





Snow, graupel and rain mixing ratio



Average 3rd domain precipitation





Difference between accumulated domain precipitation, "high – low" ice crystal concentration ("Cooper "– "0.1 cm⁻³ dust concentration")





Summary



We are working on a new version of the Thompson microphysics scheme that will be dependent on aerosols for droplet activation and ice initiation.

Aerosols initializations from global modeled climatology

Use Look-up Tables with 5 variables for droplet activation (N_{aer} , aerosol radius, chemical composition, temperature and updraft velocity).

Included new Ice nucleation parameterization by DeMott et al (2010).

Initial results for a simulated case with heavy dust loading do not indicate large difference in total domain precipitation, but the location of the precipitation have changed.