



Implementing a New Shallow Convection Scheme into WRF

Aijun Deng and Brian Gaudet

Penn State University

Jimy Dudhia

National Center for Atmospheric Research

Kiran Alapaty

United States Environmental Protection Agency

14th Annual WRF Users' Workshop

24 - 28 June 2013, Boulder, CO



Outline

- **Overview of the PSU Shallow Convection Scheme (SCP)**
- **WRF Implementation Plan**
- **Preliminary 1-D Results**
- **Summary and Conclusions**



Technique Details of the PSU Shallow Convection Parameterization (SCP) *(Deng et al. 2003a, 2003b, JAS)*

- **Triggering Function:**
 - 1) Parcel T', Q'
 - 2) Parcel vertical lifting determined by factors including TKE
 - **Cloud model:** K-F entraining/detraining cloud model
 - **Closure Assumption:** Amount of total cloud-base mass flux or number of updrafts is determined with a hybrid of TKE and CAPE, depending on the updraft depth.
- In addition -----
- Prognostic cloud scheme for cloud fraction and subgrid cloud water



PSU SCP description- Parcel initial vertical velocity

$$w_B = \bar{w} + w_T + w_{NH}$$

- 1) Resolvable-scale vertical velocity: \bar{w}
- 2) Boundary layer turbulence:

$$w_T = \sqrt{\frac{2}{3} \cdot TKE_{MAX}}$$

- 3) Vertical “pumping” due to non-hydrostatic pressure effects for deep convection:

$$w_{NH} = \max \{0.15 \cdot [w(z_{1000})^{\tau-1} - (\bar{w} + w_T)], 0\}$$



PSU SCP description- T and Q perturbation and updraft radius

Parcel perturbation in **temperature** and **moisture**,
averaged from the lowest 20% of the PBL \Rightarrow Cloud
base (LCL)

Updraft radius:

in a range of 150 to 1500 m, as a function of PBL depth
and cloud depth:

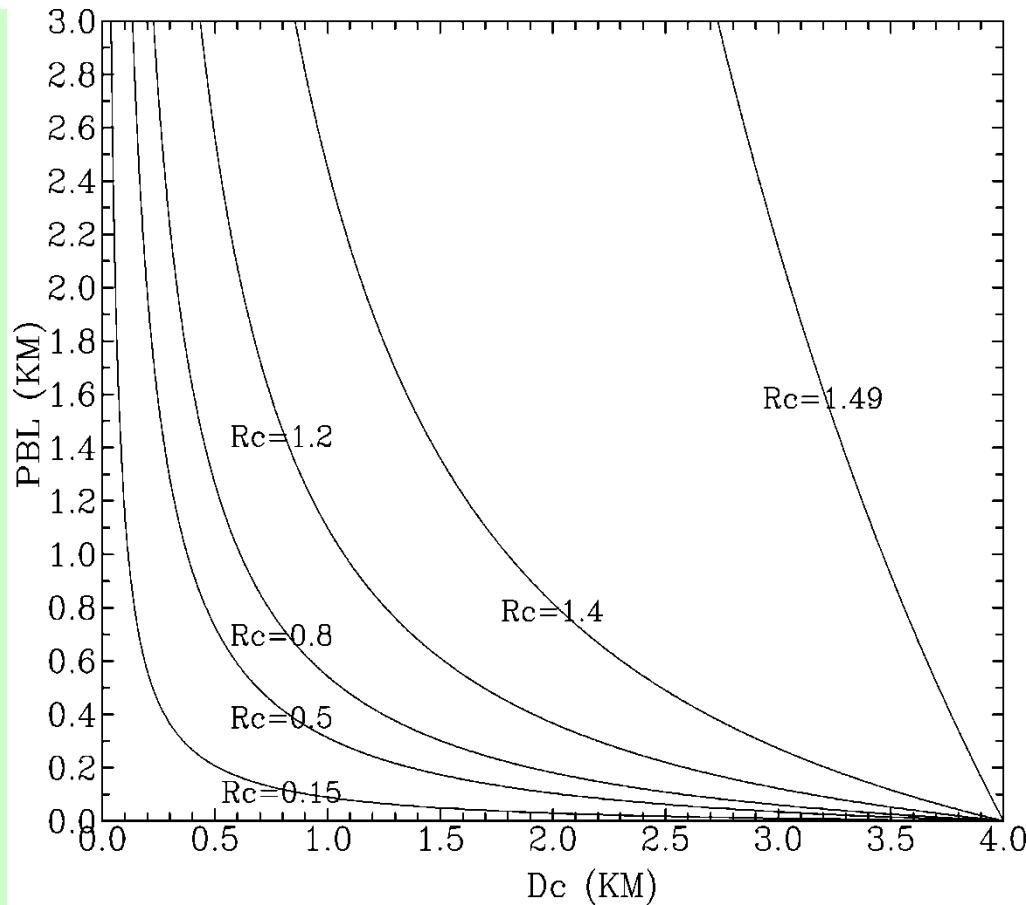
$$R_C = \frac{b - \sqrt{b^2 - 12\eta h_{PBL}}}{4}$$

where $b = \frac{7 + 2\eta h_{PBL}}{2}$

$$\eta = \frac{12D_C}{4 - D_C}$$

D_C : Cloud depth (km)

h_{PBL} : PBL (km)





PSU SCP description- Closure assumption

Cloud-base mass flux:

$$\mu_B = N\mu_{BS} \quad \text{where} \quad \mu_{BS} = \pi R_C^2 w_B \rho$$

where the number of updrafts

$$N = \begin{cases} N_4 & z_T \leq z_{LFC} \\ N_2 & \text{for } D_C \geq 4\text{km} \\ N_3 & z_T > z_{LFC} \text{ and } D_C < 4\text{km} \end{cases}$$

where $N_4 = \varepsilon \cdot TKE_{\max} \cdot N_T \rightarrow$ TKE-based closure

$$N_2 = \frac{\mu_B}{\mu_{BS}} \rightarrow \text{CAPE-removal closure}$$

$$N_3 = fN_2 + (1-f)N_4 \rightarrow \text{Hybrid closure}$$

$$\text{where } N_T = \frac{M_S}{\mu_{BS} \tau_C}$$

$$\varepsilon = 0.15 \text{ kg J}^{-1}$$

$$f = h_3/h_1 \quad h_3 = D_C - h_2$$

$$h_1 = 4 - h_2 \quad h_2 = z_{LFC} - z_B$$

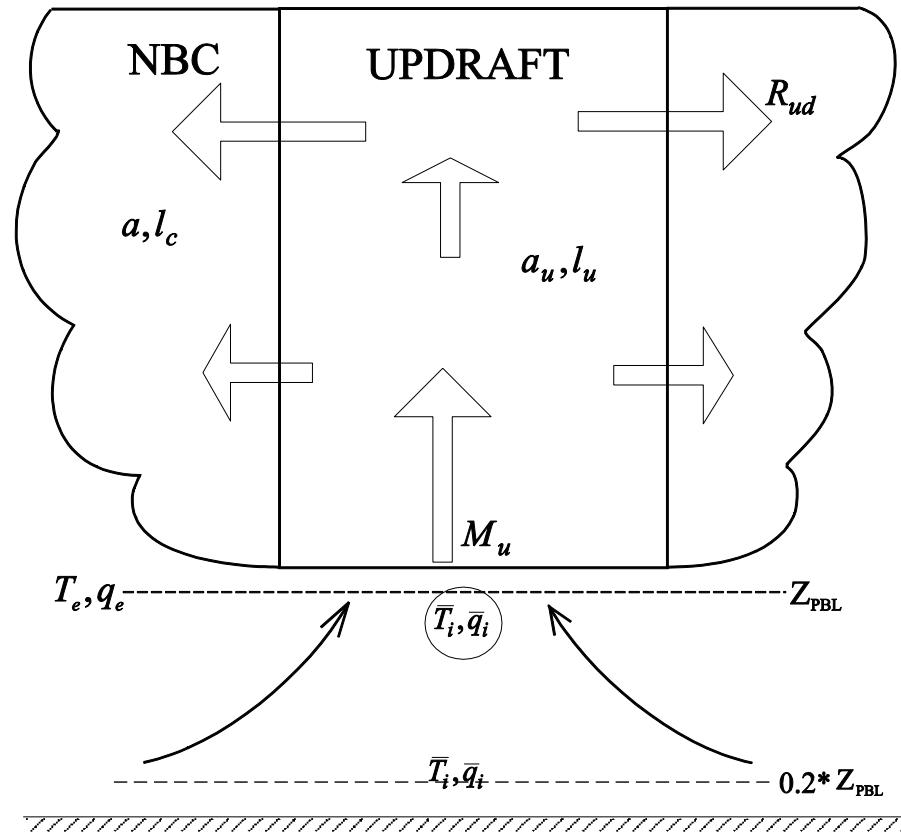
M_S : mass of air in the model layer.

TKE_{MAX} : maximum TKE below

the parcel-releasing level.



PSU SCP Schematic



Schematic of the prototype PSU SCP, where a is the neutrally-buoyant cloud (NBC) fraction, l_c is the NBC cloud water content, l_u is the cloud water content in the updraft denoted by subscript u , R_{ud} is the updraft detrainment rate; and Z_{pbl} is the depth of the PBL, etc.



$$\frac{\partial a}{\partial t} = S_a + D_a - \vec{v} \cdot \nabla a - w \frac{\partial a}{\partial z}$$

$$\frac{\partial l_c}{\partial t} = S_l + D_{mix} + D_{pre} + D_{ics} + D_{CTEI} - \vec{v} \cdot \nabla l_c - w \frac{\partial l_c}{\partial z}$$

a Cloud area

l_c Cloud water/ice content

S_a Source for cloud area due to updraft detrainment

D_a Dissipation of cloud area due to evaporation at cloud edge

S_l Source for cloud water due to updraft detrainment

D_{mix} Depletion of cloud water due to vertical mixing

D_{pre} Depletion of cloud water due to precipitation

D_{ics} Depletion of cloud ice due to ice settling

D_{CTEI} Depletion of cloud water due to CTEI

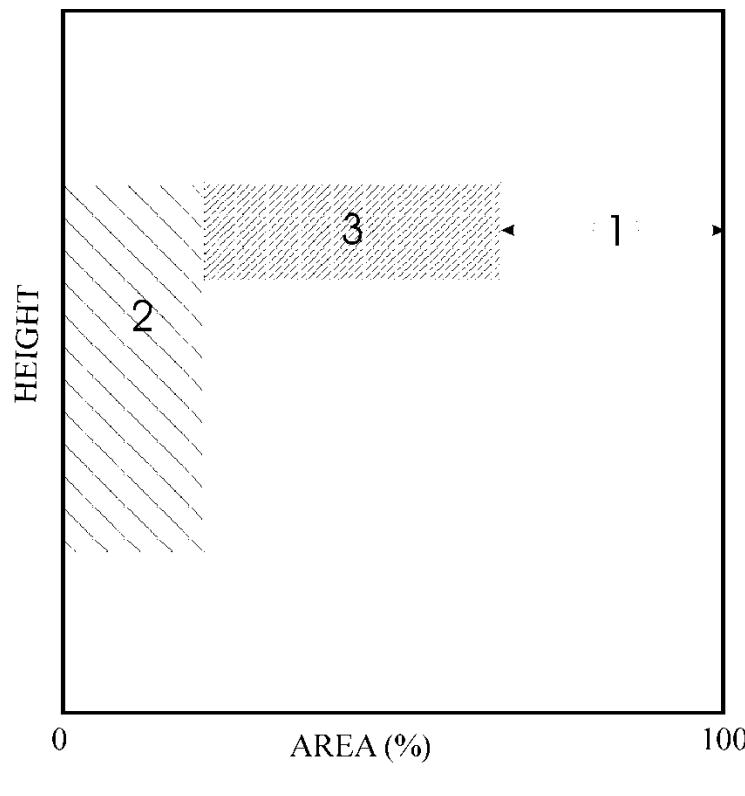
PSU SCP Interaction with Radiation Schemes

Effective cloud:

$$a_e = (1 - a)a_S + a$$

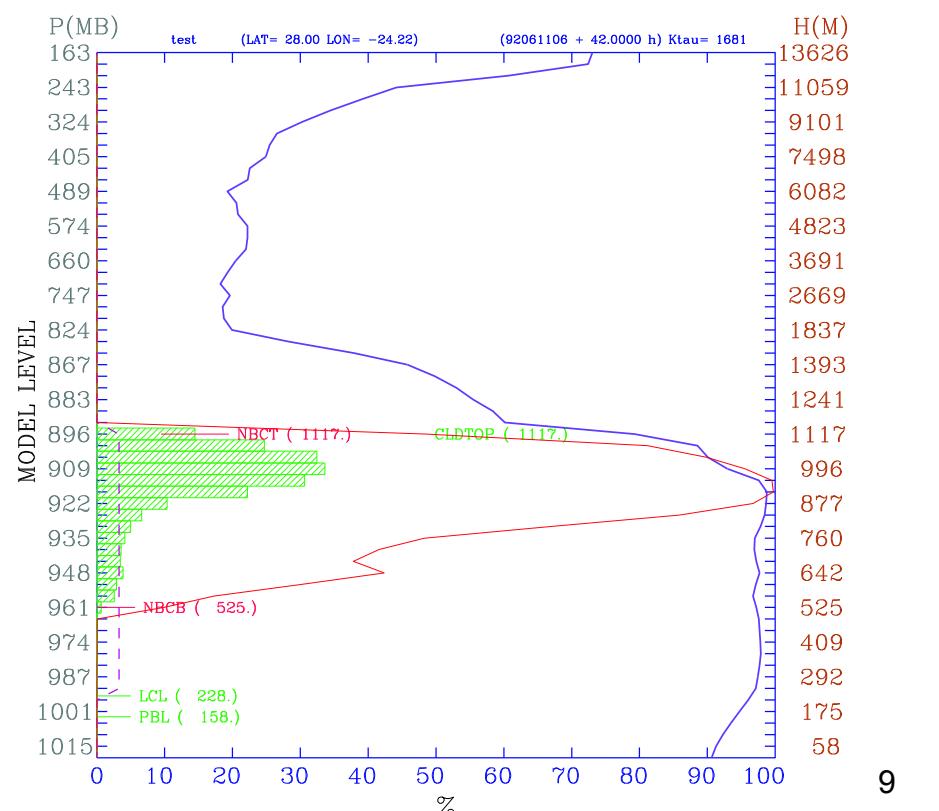
$$a_S = RH^{\alpha_1} [1 - \exp\left(\frac{-\alpha_3(l_c + q_c)}{[(1-RH)q_S]^{\alpha_2}}\right)] \quad \text{if } RH < 1$$

$\alpha_1, \alpha_2, \alpha_3$ Constants given by Xu and Randall (1996)



$$R_s = \bar{a}_{e2} R_2 + \bar{a}_{e3} R_3 + (1 - \bar{a}_{e2} - \bar{a}_{e3}) R_{clr}$$

$$\left. \frac{\partial T(z)}{\partial t} \right|_{rad} = [1 - a_e(z)] \left. \frac{\partial T(z)}{\partial t} \right|_{clr} + a_e(z) \left. \frac{\partial T(z)}{\partial t} \right|_{cld}$$

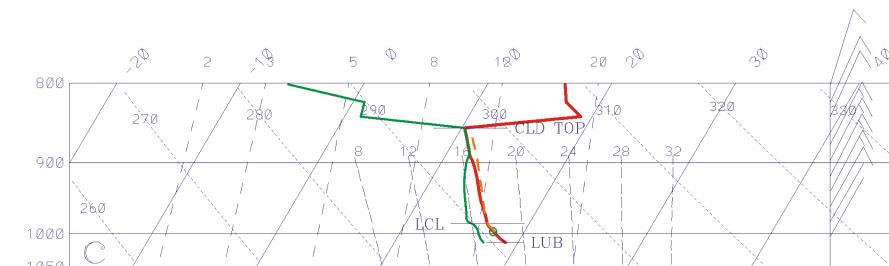
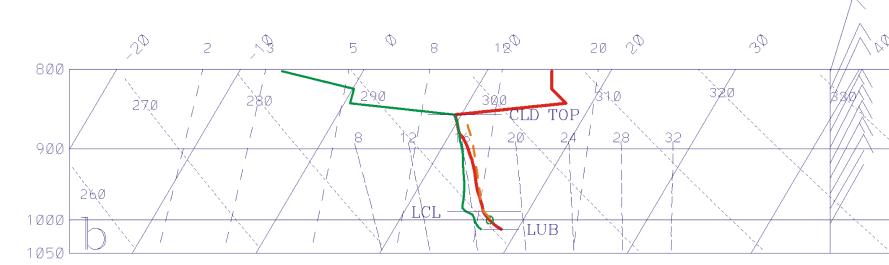
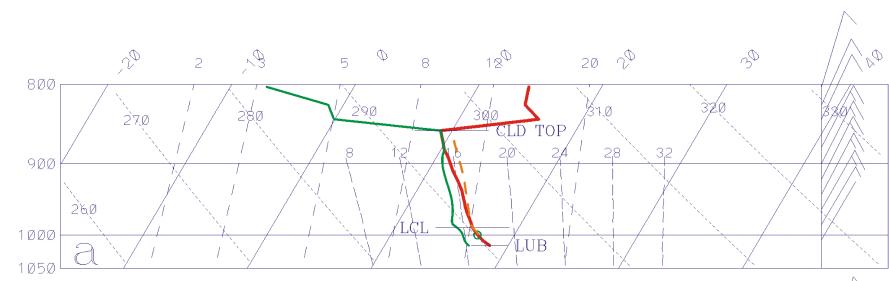
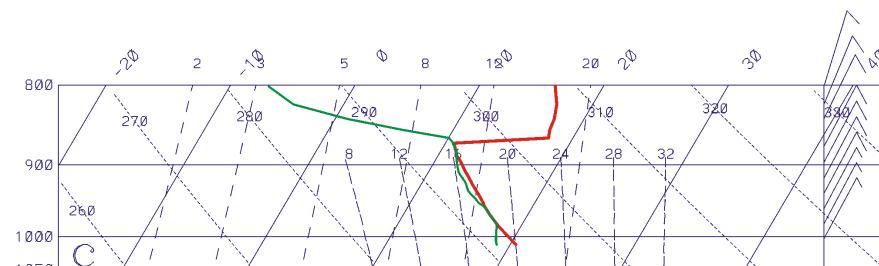
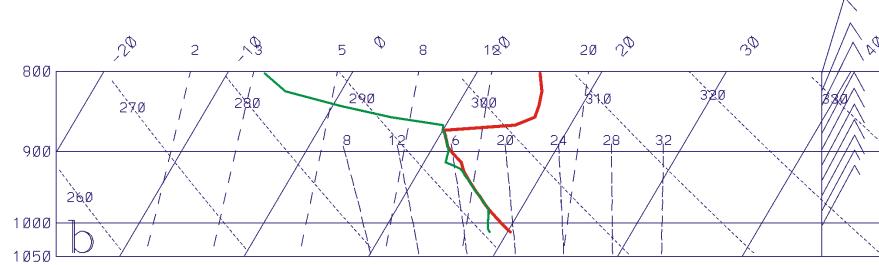
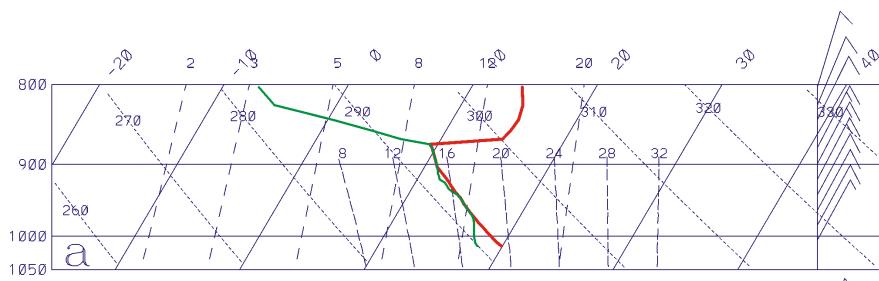


Effect of PSU SCP

Marine environment

Simulated sounding (20h, 40h, 60h)

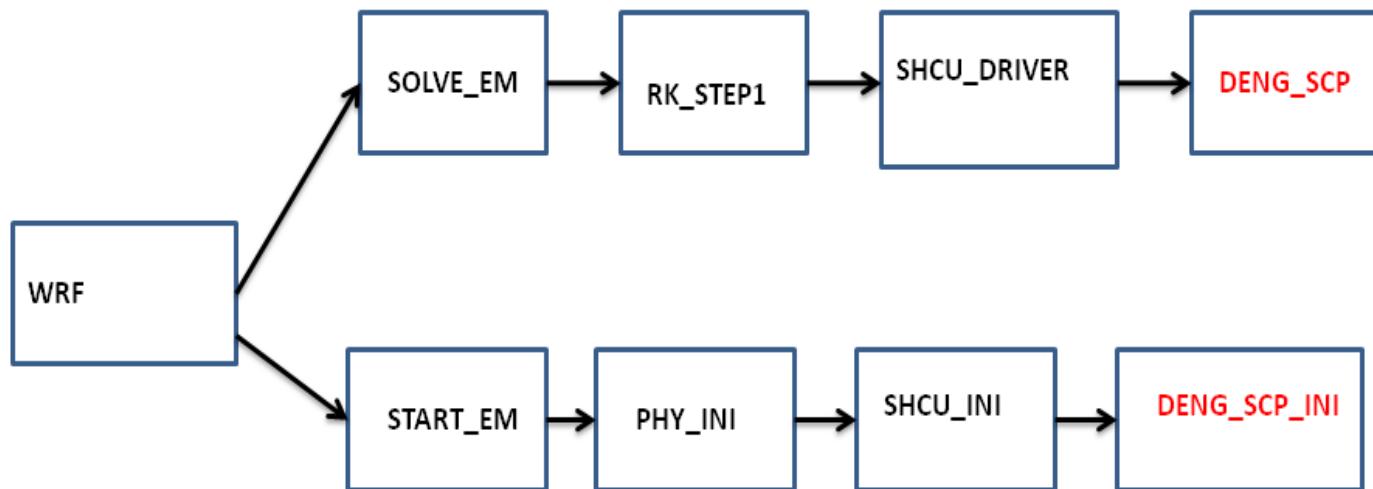
Without shallow convection With shallow convection





WRF Implementation Plan

WRF DENG-SCP FLOWCHART



In addition, MM5 GS PBL scheme is implemented



PSU SCP 1-D WRF Test Cases

- Case1 - Marine Shallow Stratus (Azores Islands, ASTEX)
 - 1992.06.11.06Z – 1992.06.14.06Z (72-hour simulation)
 - Location: (20.0N, 24.22W)
- Case2 - Continental Shallow Stratocumulus
 - 1998.06.08.12Z – 1998.06.09.12Z (24-hour simulation)
 - Location: Pittsburgh, Pennsylvania
- Case3 - Continental Deep Convection
 - 1997.07.10.12Z – 1997.07.11.12Z (24-hour simulation)
 - Location: SGP ARM CART Site



PSU SCP 1-D WRF Model Configuration

- DX=30 km, DT=90 S, 62 Eta layers
- WRF physics
 - Gayno-Seaman TKE-predicting PBL scheme (`bl_pbl_physics =11`) (newly implemented with the PSU SCP Scheme)
 - MM5 Monin-Obukhov scheme (`sf_sfclay_physics =1`)
 - 5-layer thermal diffusion scheme (`sf_surface_physics =1`)
 - PSU-Deng shallow convection scheme (`shcu_physics 3`)
 - RRTMG atmospheric radiation scheme (`ra_lw_physics=4, ra_sw_physics=4`)
 - WSM 3-class simple ice explicit moisture scheme (`mp_physics=3`)
- SCM initialization codes modified to allow pressure level sounding as in MM5



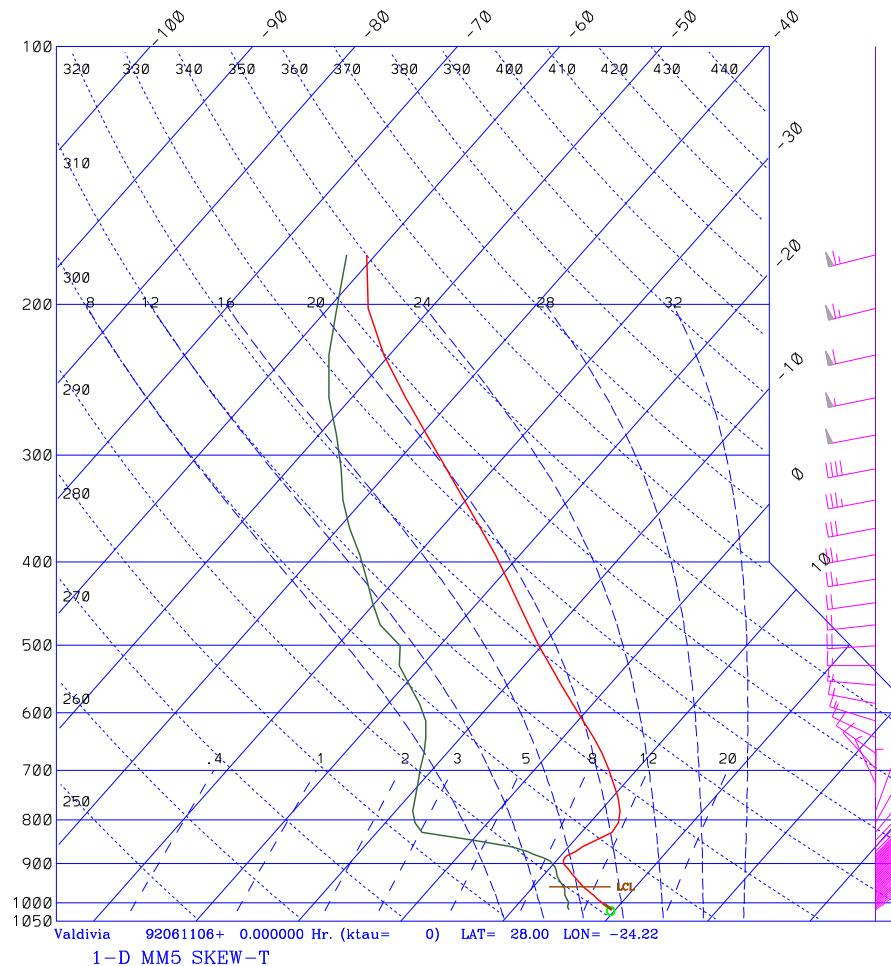
- **Case 1: Marine Shallow Stratus –Azores Islands, ASTEX, 06 UTC, 11-14 June 1992, 72-hour simulation**

Dominated by the persistent Bermuda High, this region exhibits deep tropospheric subsidence and a moist marine atmospheric boundary layer that is often capped by a strong inversion and stratocumulus clouds, with a prescribed subsidence profile.

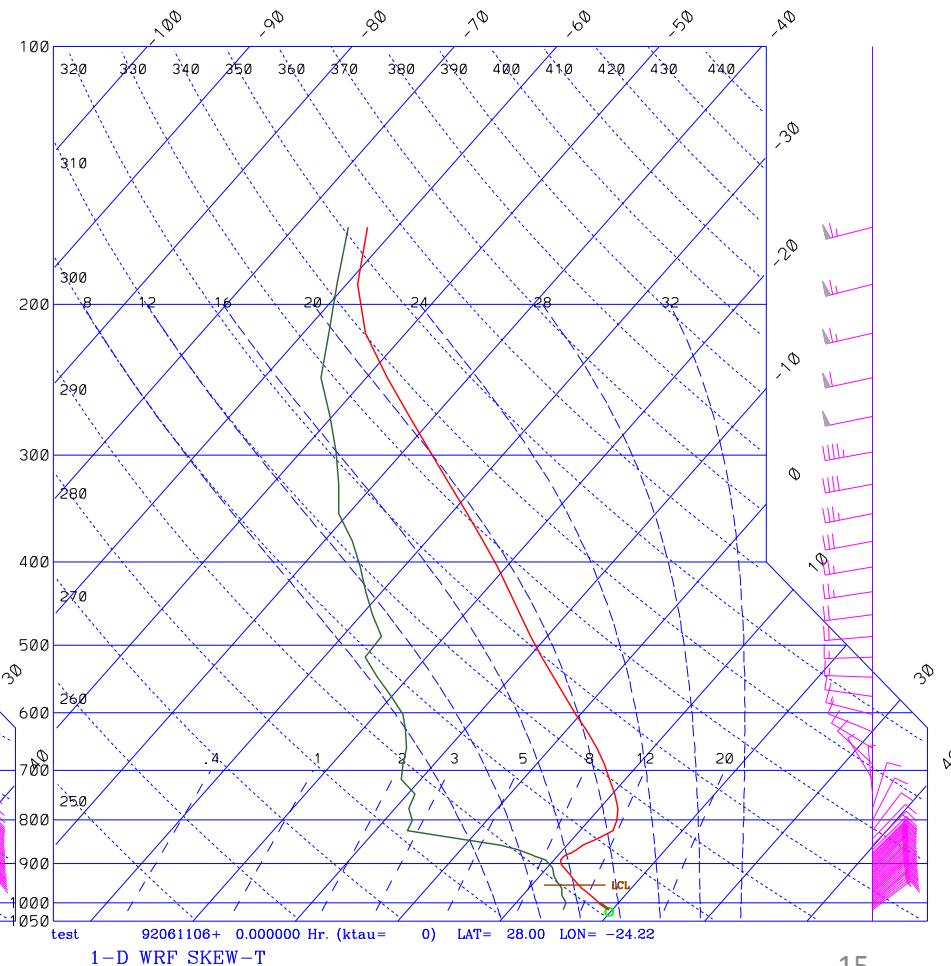


Initial Condition

MM5



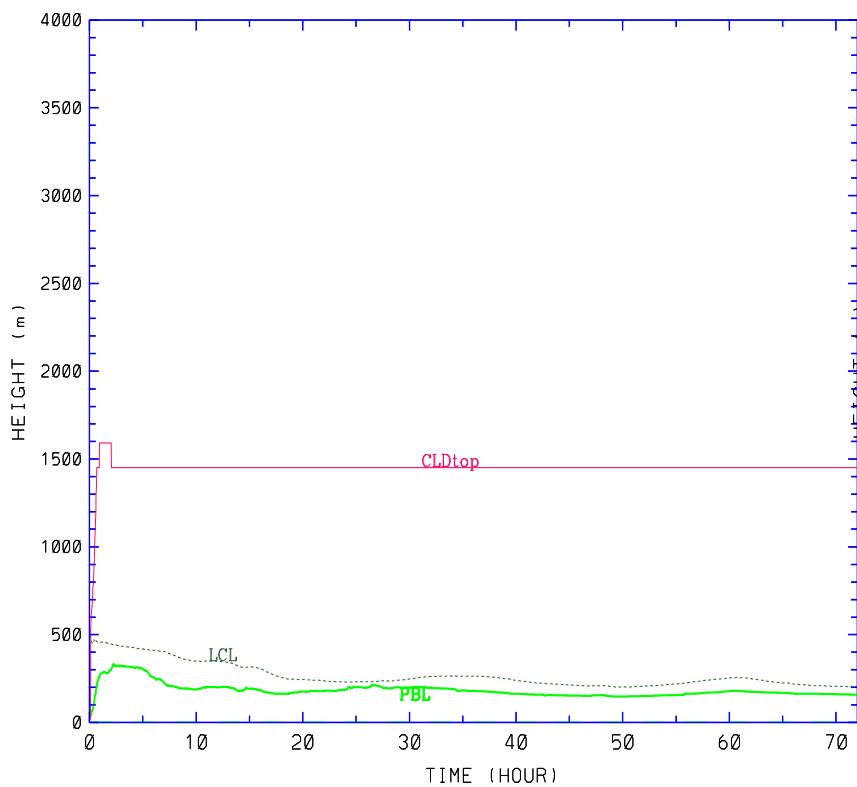
WRF



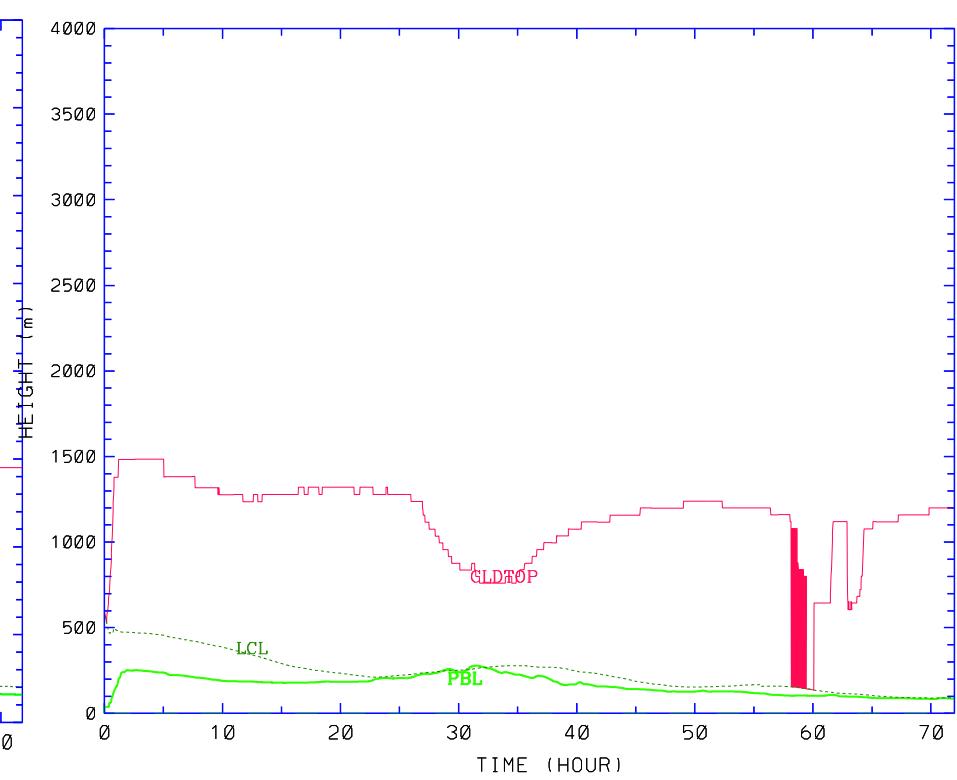


Model-predicted Cloud Updraft Top, LCL and PBL Depth

MM5



WRF

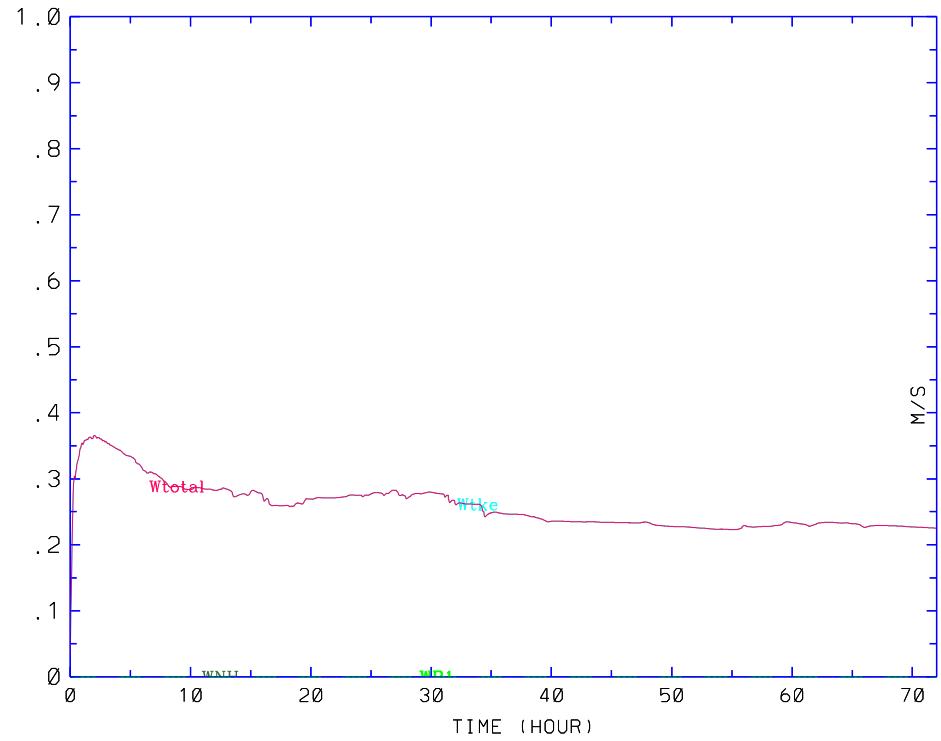




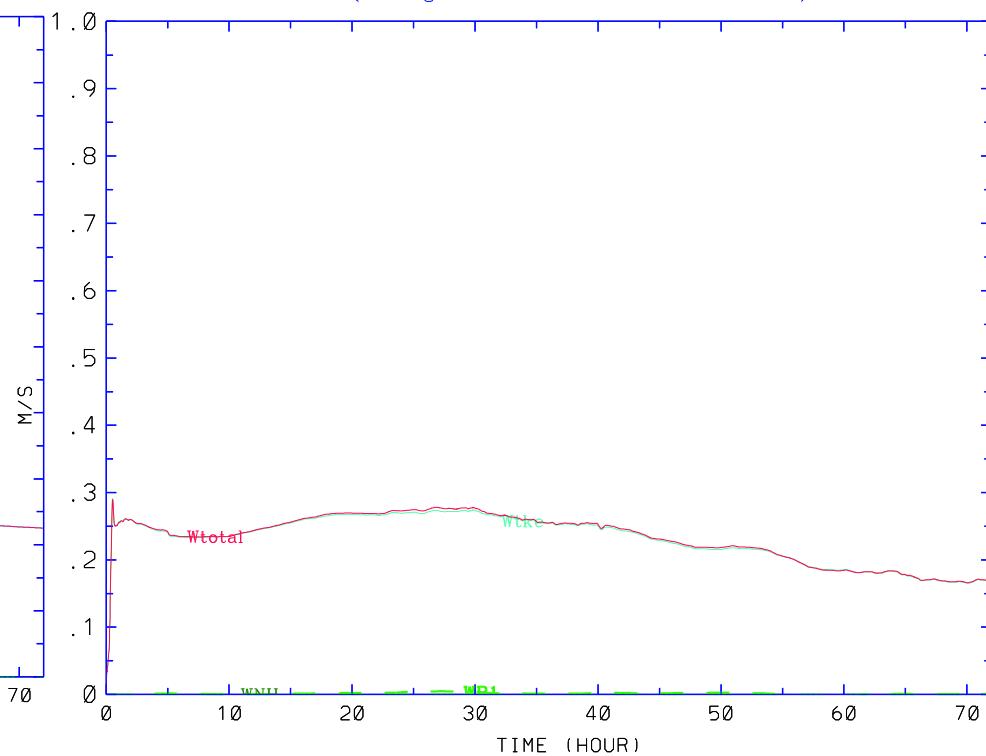
Cloud Parcel Initial Vertical Velocity

MM5**WRF**

PARCEL VELOCITIES AS A FUNCTION OF TIME
Valdivia (starting at 92061106 LAT= 28.00 LON= -24.22)

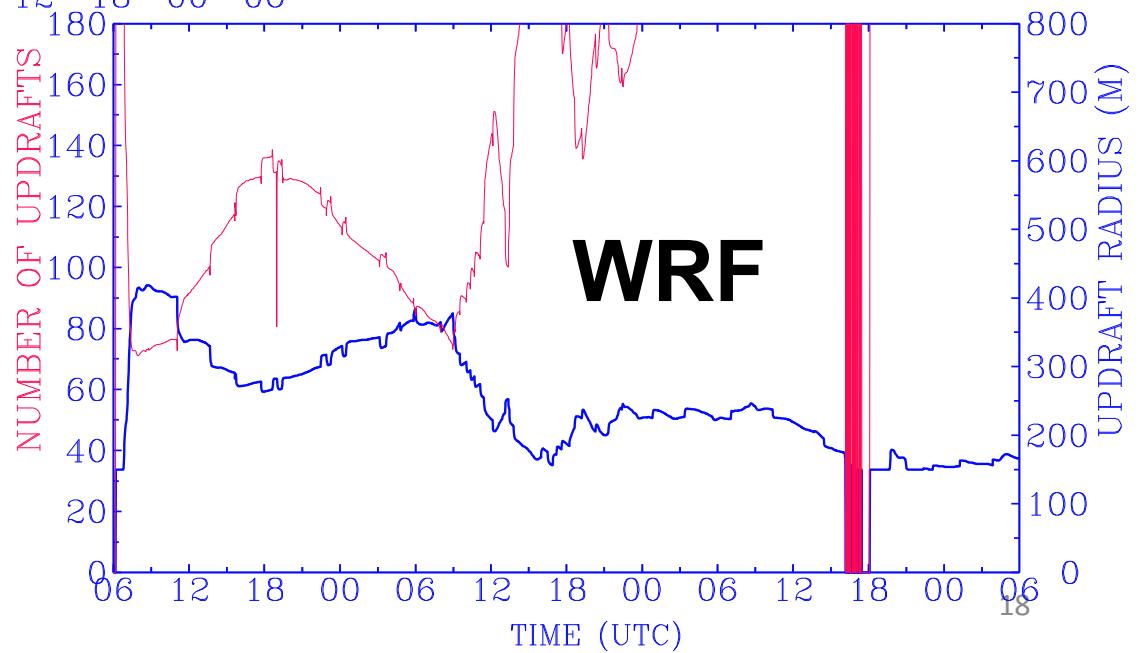
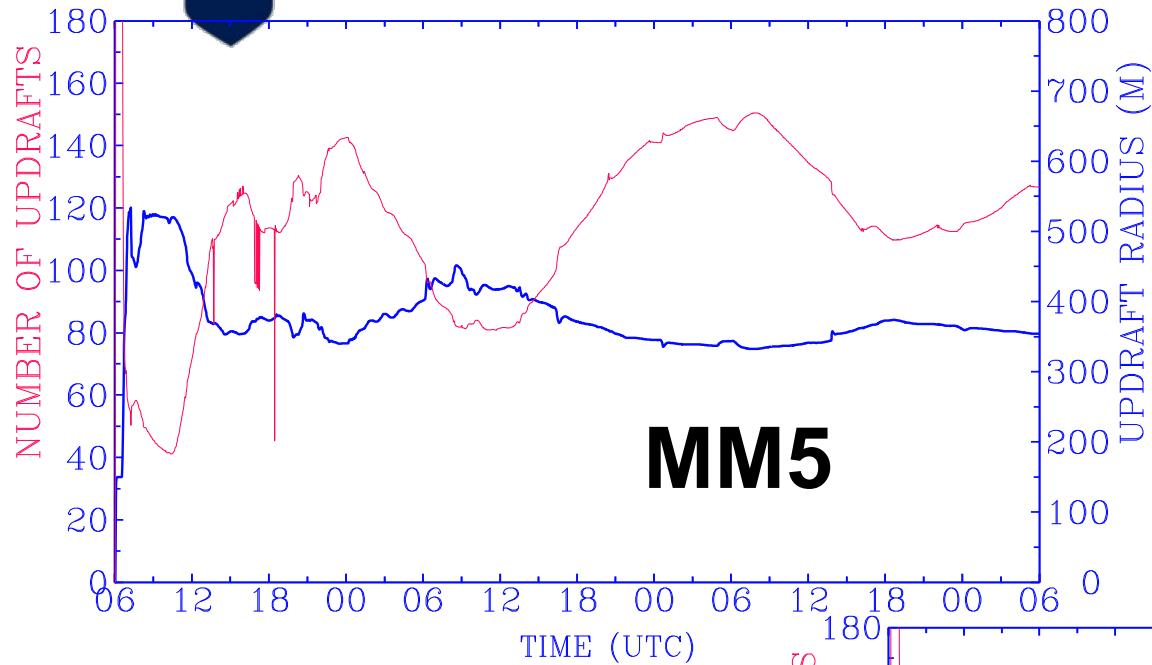


PARCEL VELOCITIES AS A FUNCTION OF TIME
test (starting at 92061106 LAT= 28.00 LON= -24.22)





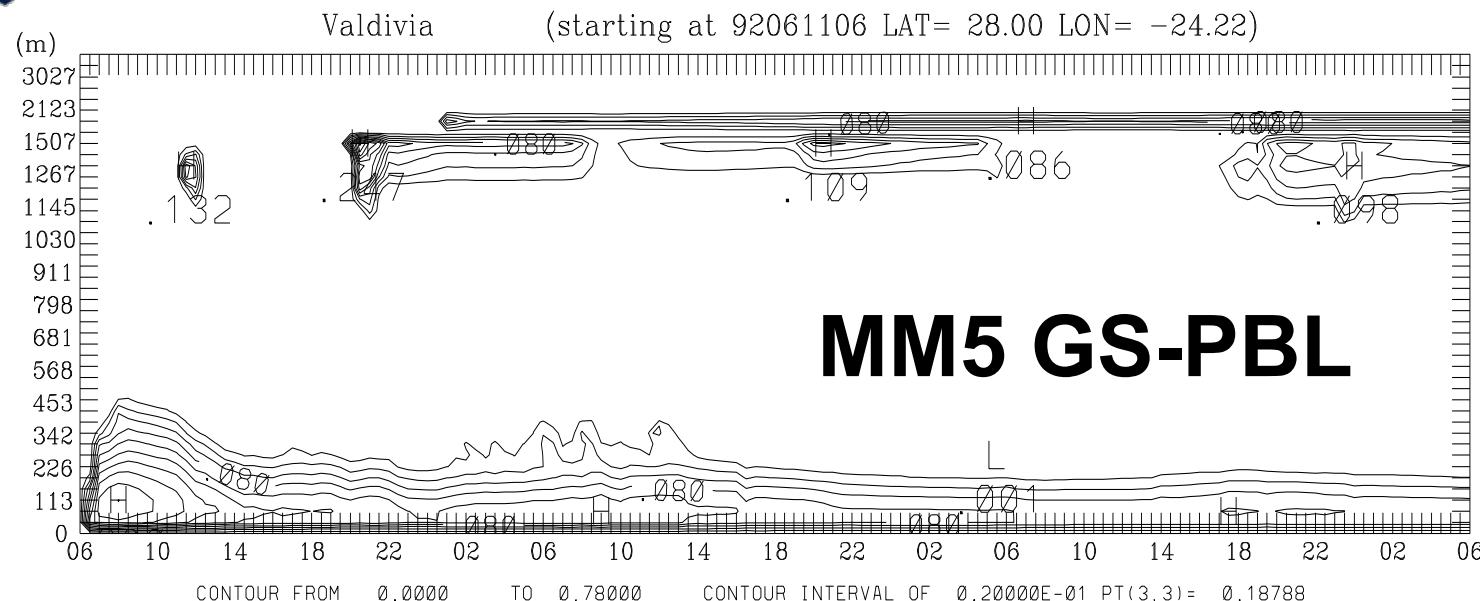
Model-predicted Number of Updrafts and Size



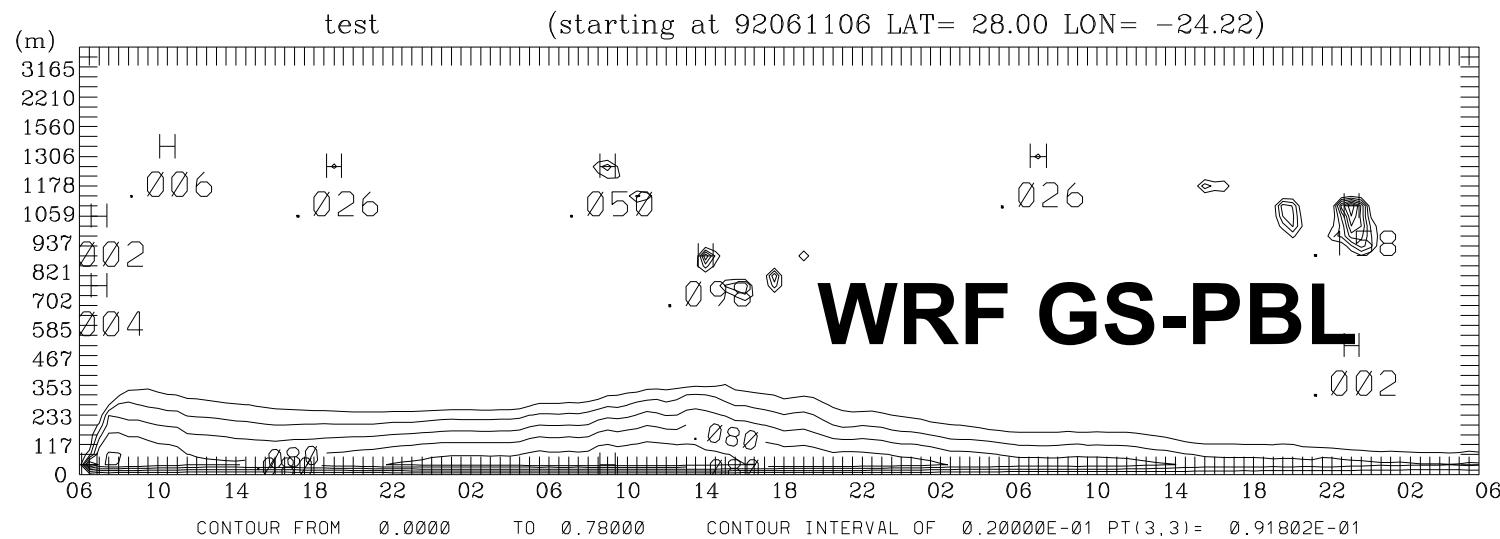


Model-predicted TKE

TIME-HEIGHT SECTION OF TKE



TIME-HEIGHT SECTION OF TKE

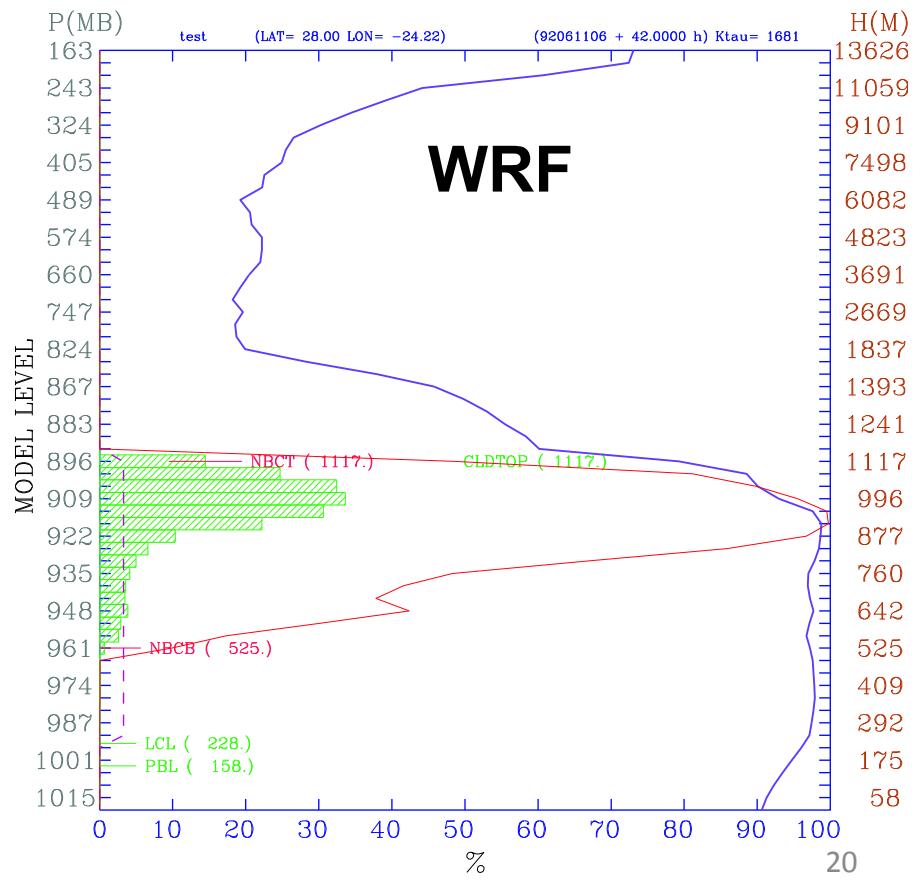
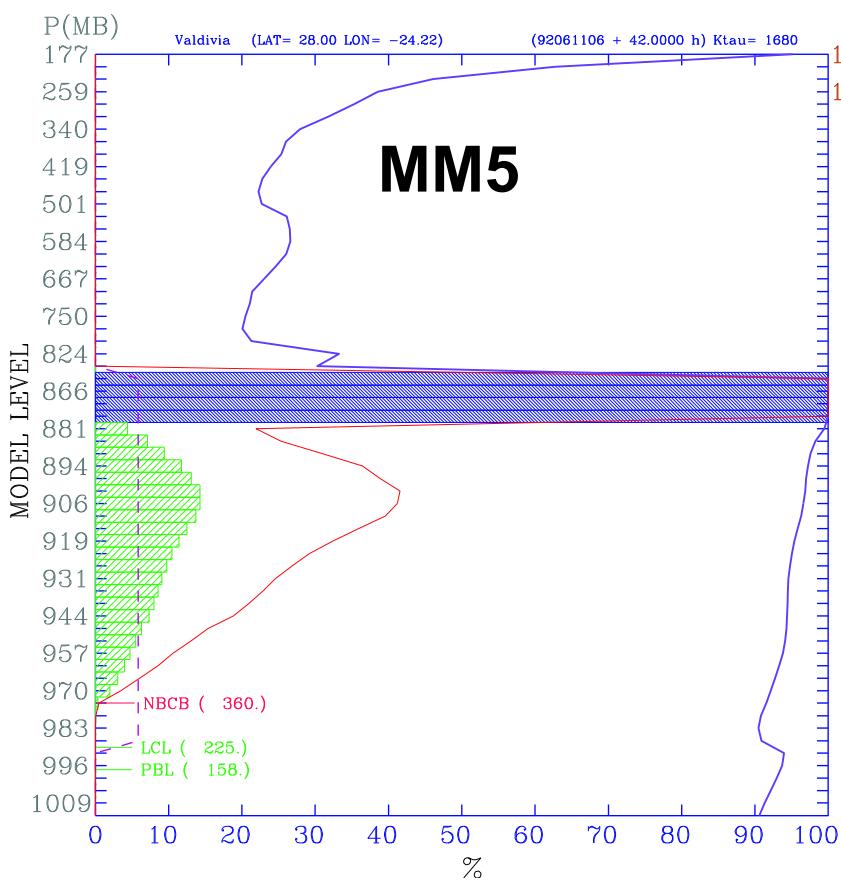




Model-predicted Cloud Fraction at 42 h

Dashed: Updraft
 Red Solid: Effective Cloud

Shaded: NBC
 Blue Solid: RH





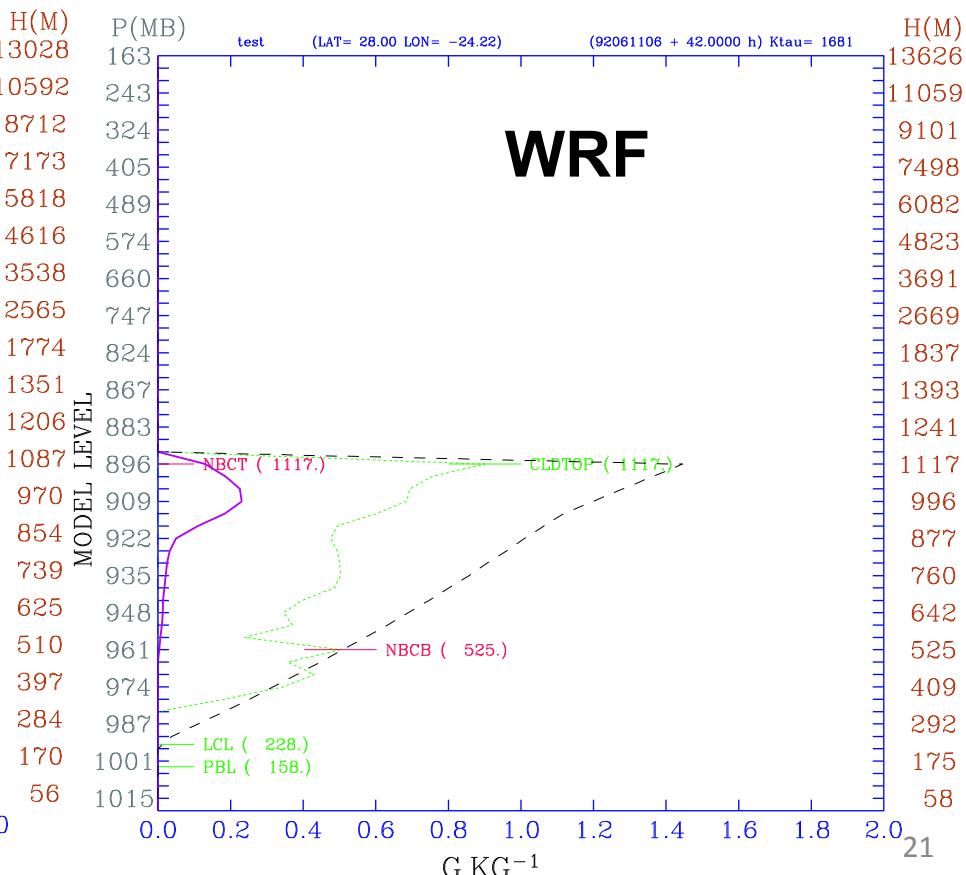
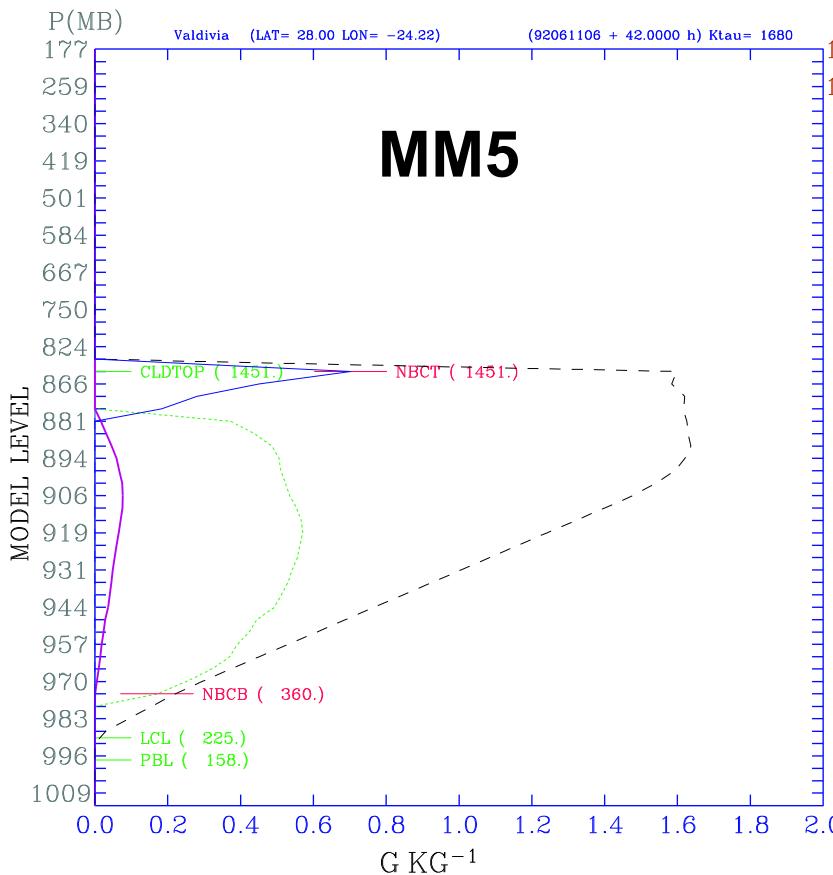
Model-predicted Cloud Water Content at 42 h

Dashed: Updraft

Heavy Solid: NBC avg. to grid

Green dotted: NBC

Thin Solid: Resolved





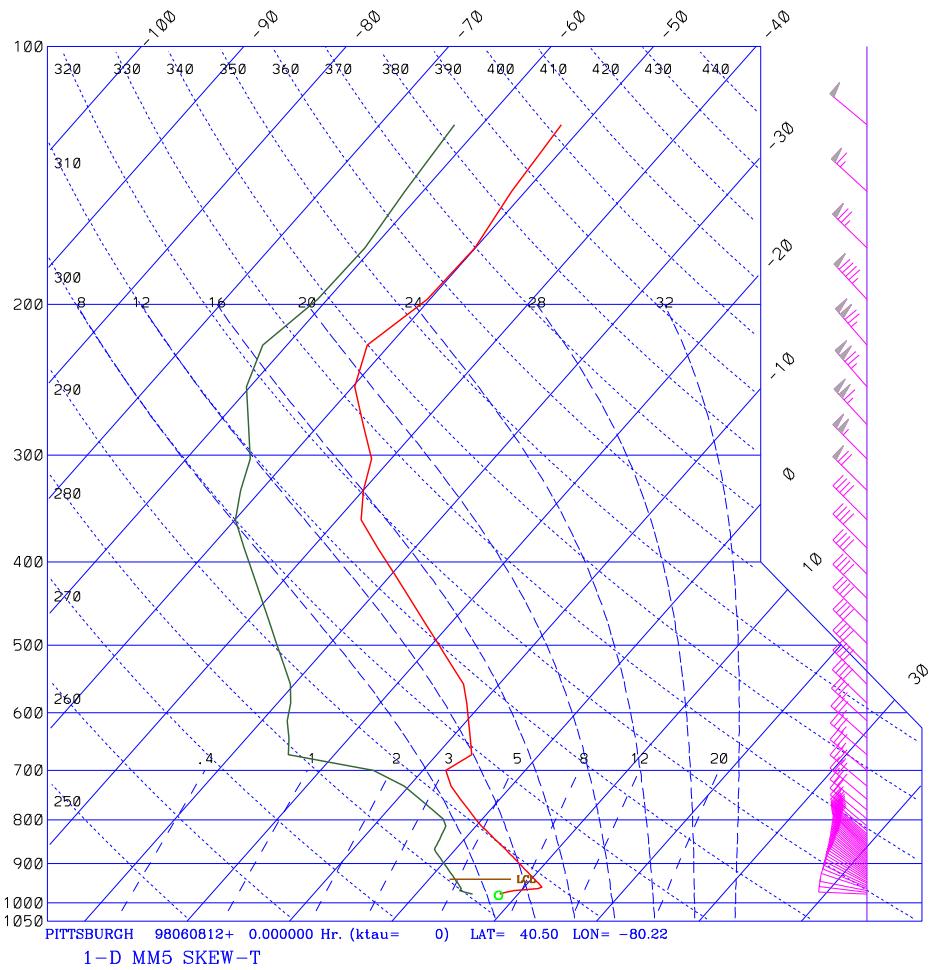
- **Case 2: Continental Shallow Stratocumulus - Pittsburgh, PA: 12 UTC, 8 June 1998 - 12 UTC, 9 June 1998, 24-hour simulation**

Cool northwesterly winds prevailed over western PA beneath the mid-level subsidence inversion associated with the ridge, shallow cloud generation during the daytime and evening hours was dominated by the local surface fluxes.

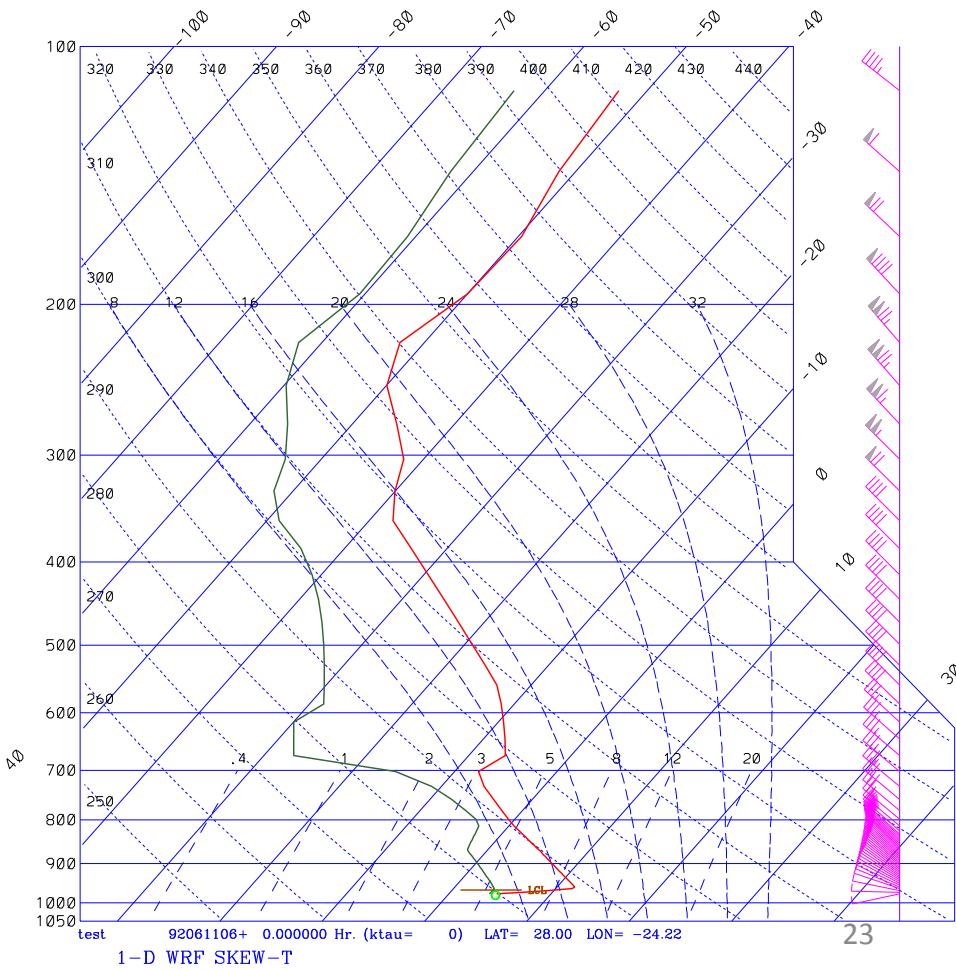


Initial Condition

MM5



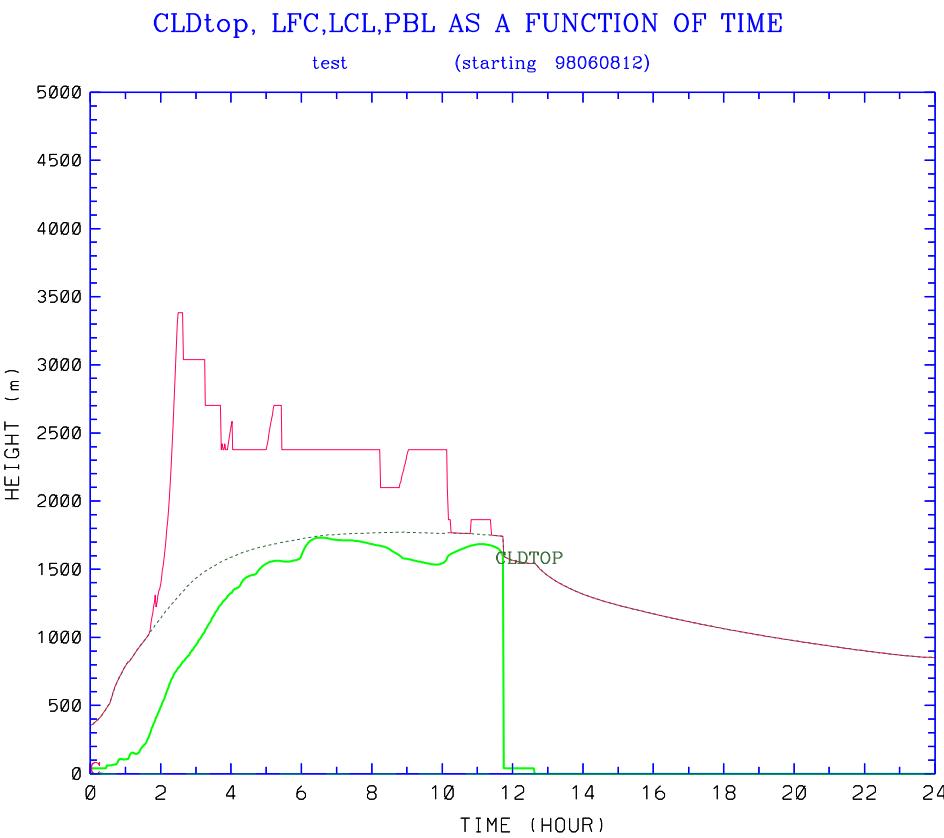
WRF



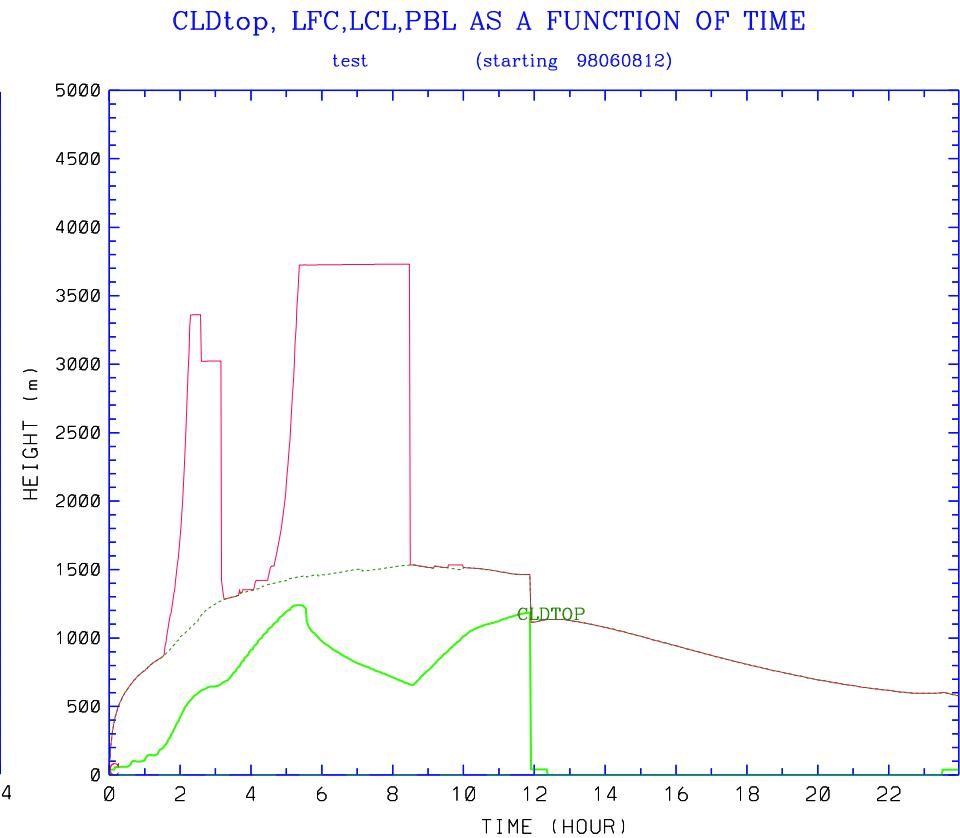


Model-predicted Cloud Updraft Top, LCL and PBL Depth

MM5



WRF

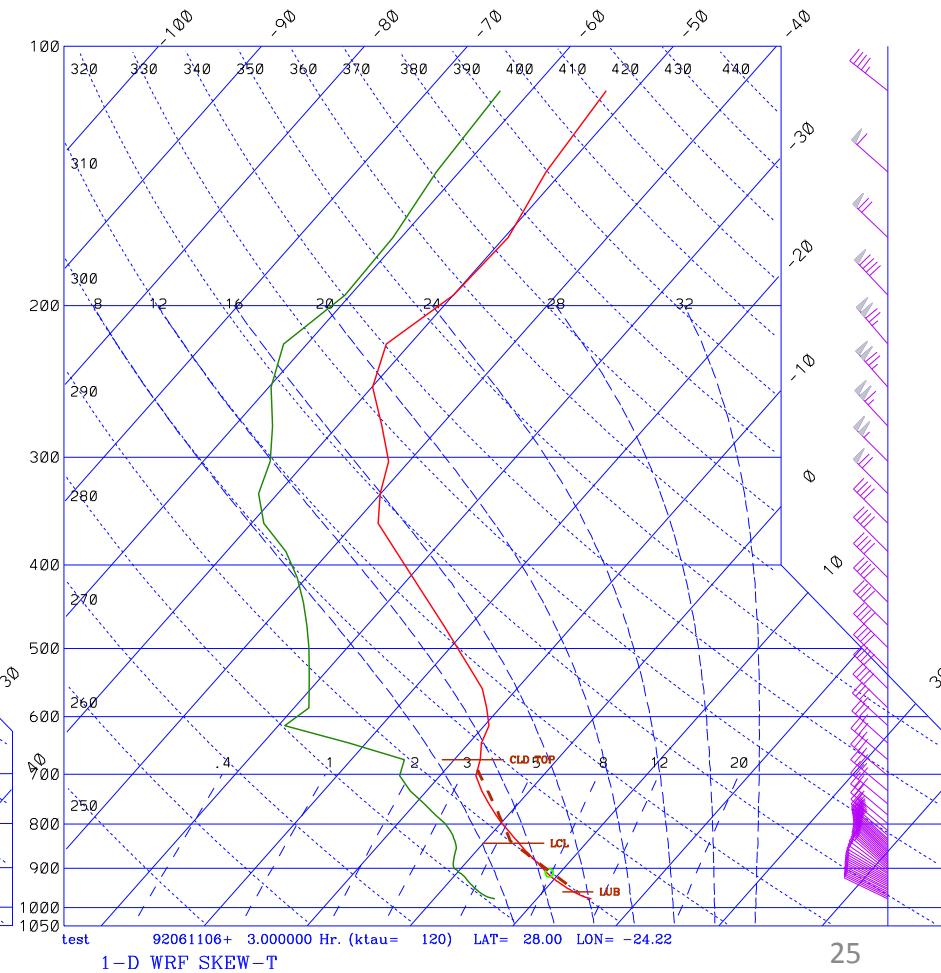
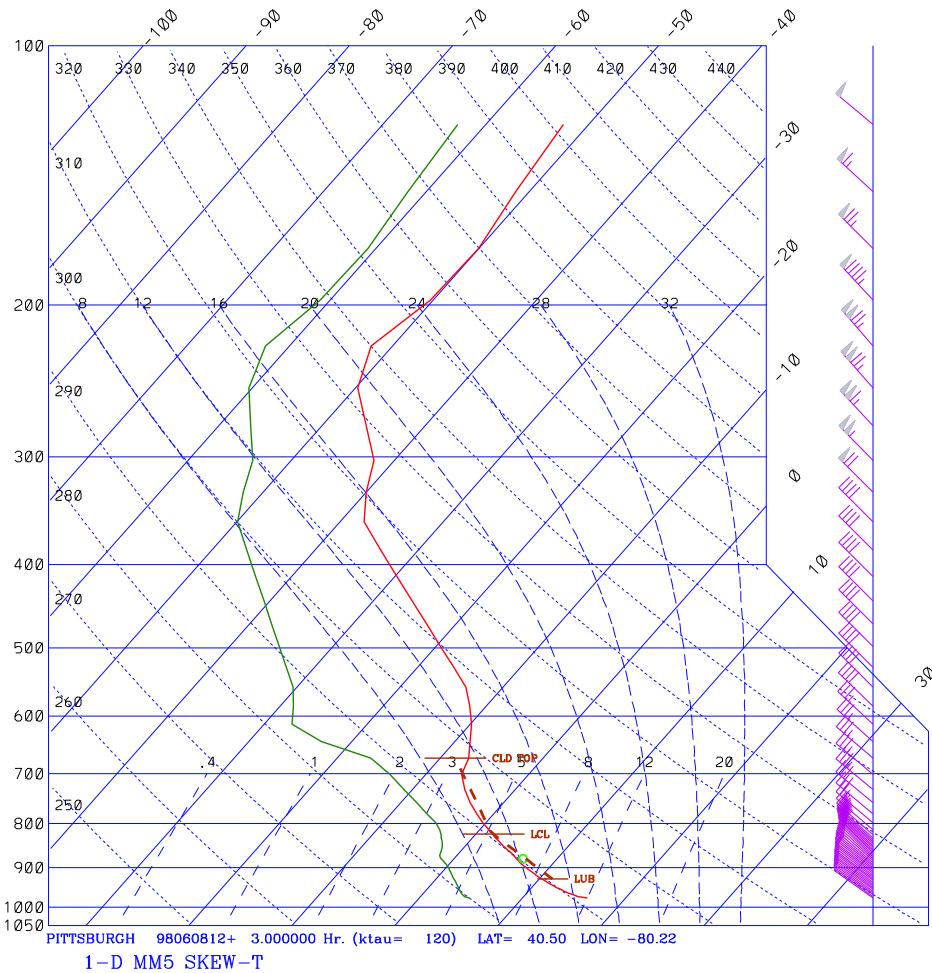




Model-predicted Condition at 3h

MM5

WRF

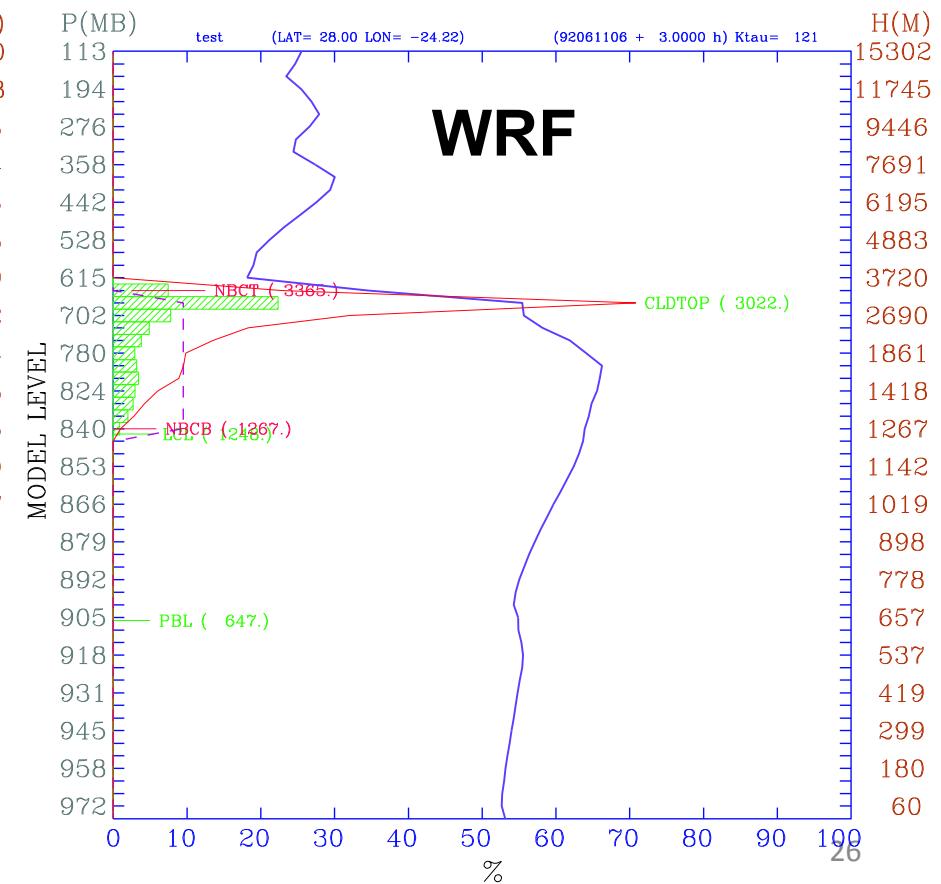
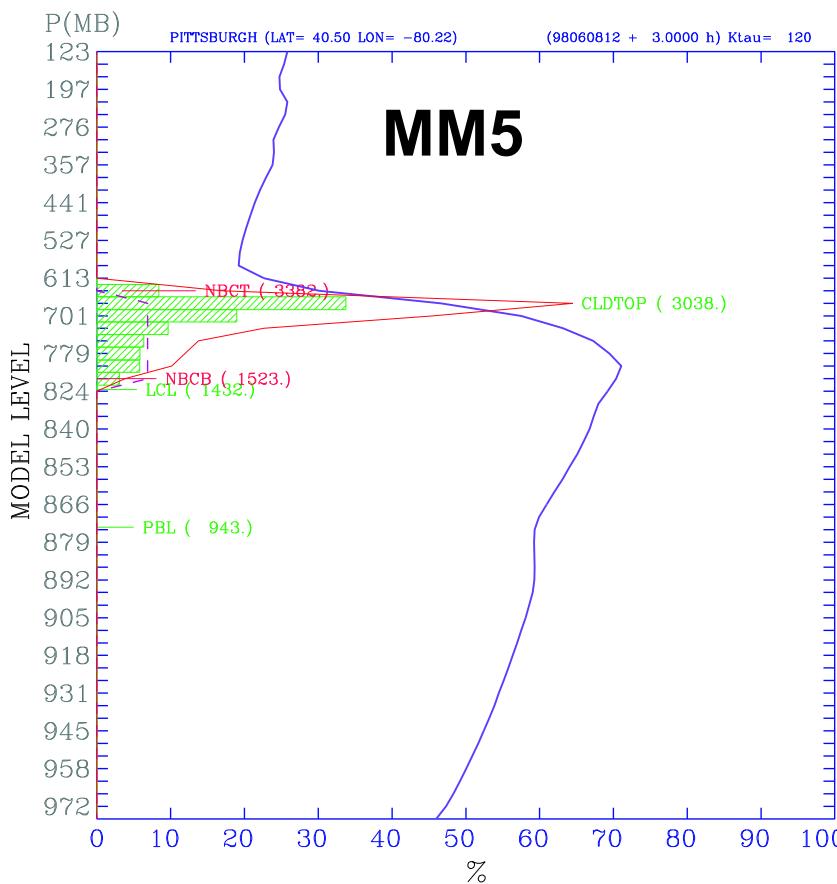




Model-predicted Cloud Fraction at 3 h

Dashed: Updraft
 Red Solid: Effective Cloud

Shaded: NBC
 Blue Solid: RH





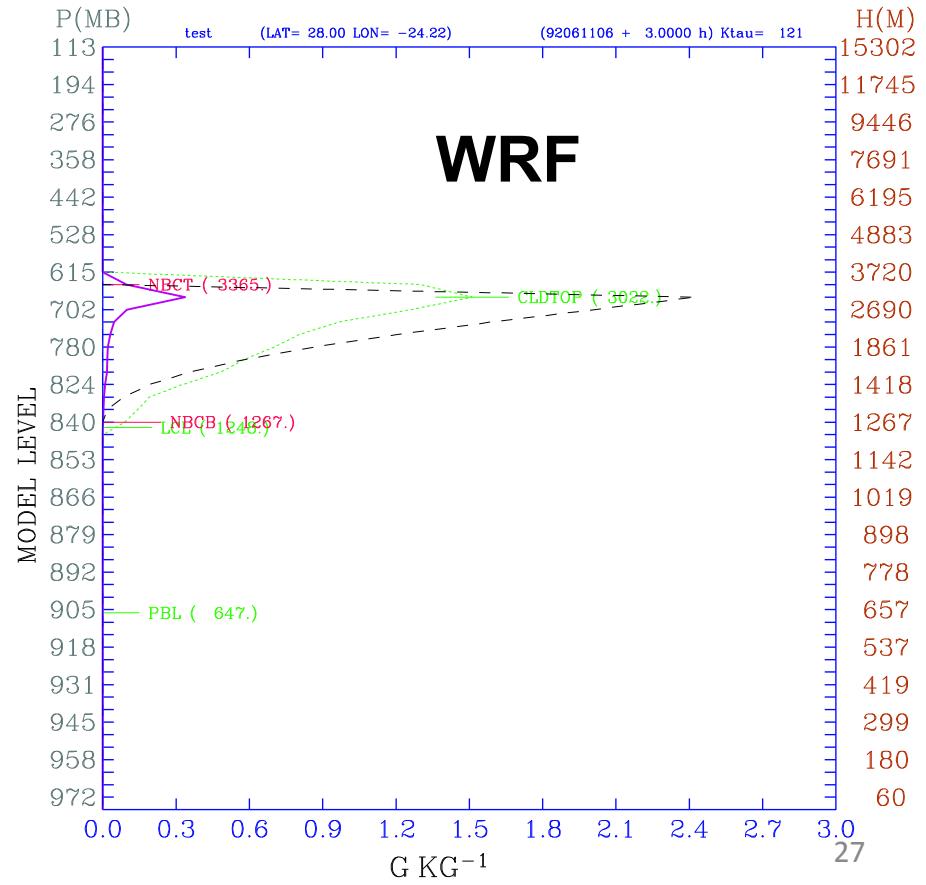
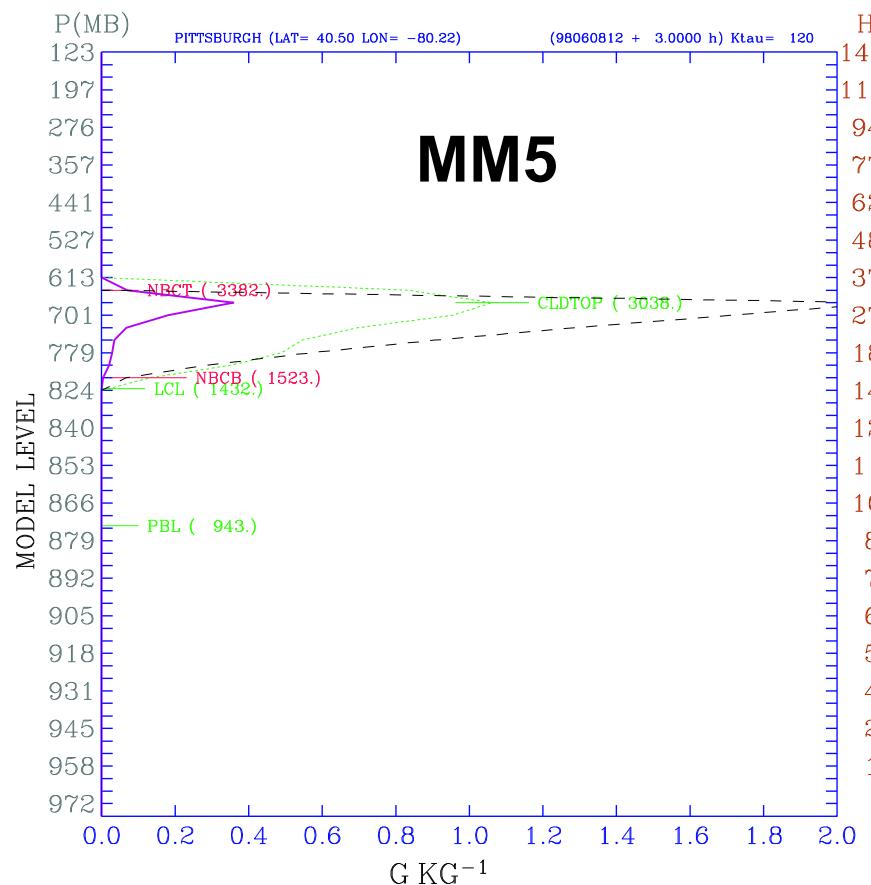
Model-predicted Cloud Water Content at 3 h

Dashed: Updraft

Heavy Solid: NBC avg. to grid

Green dotted: NBC

Thin Solid: Resolved





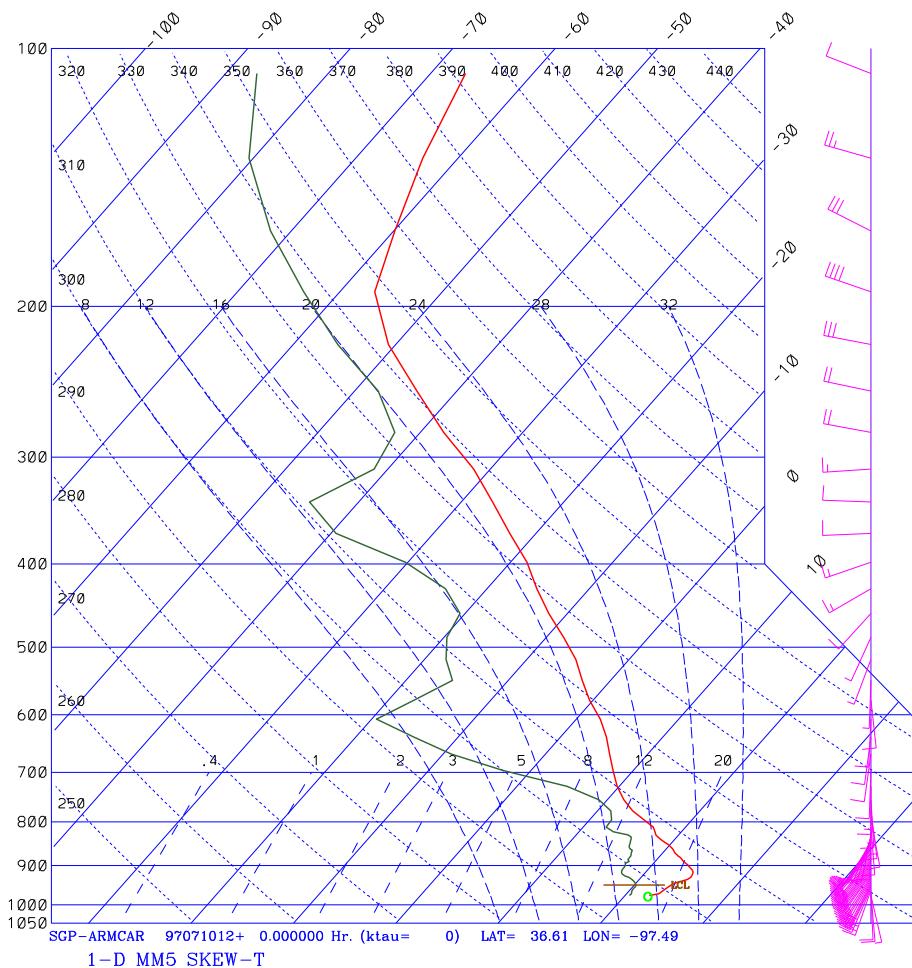
- Case 3: Continental Deep Convection - SGP
ARM CART Site: 12 UTC, 10 July - 12 UTC, 11 July 1997, 24-hour simulation

Warm south-southeasterly flow on the west side of the high advected warm moist air from the Gulf of Mexico to the region of interest, which supported deep convection later that day.



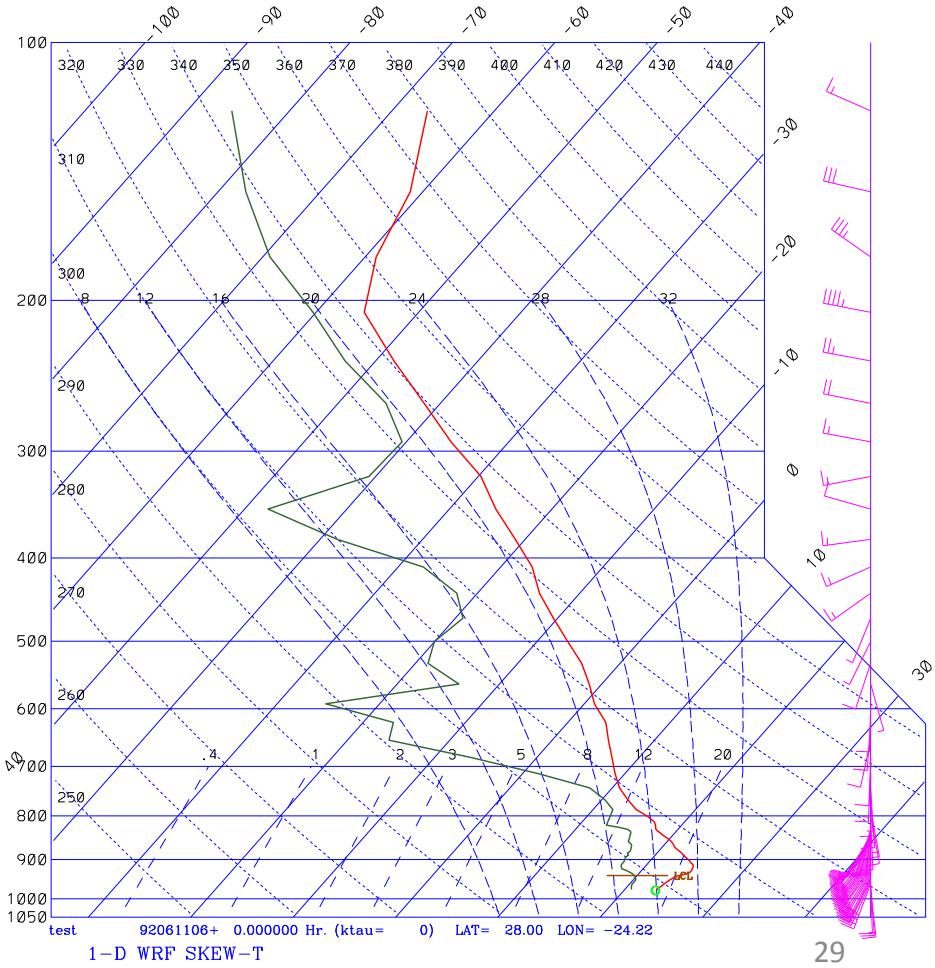
Initial Condition

MM5



1-D MM5 SKEW-T

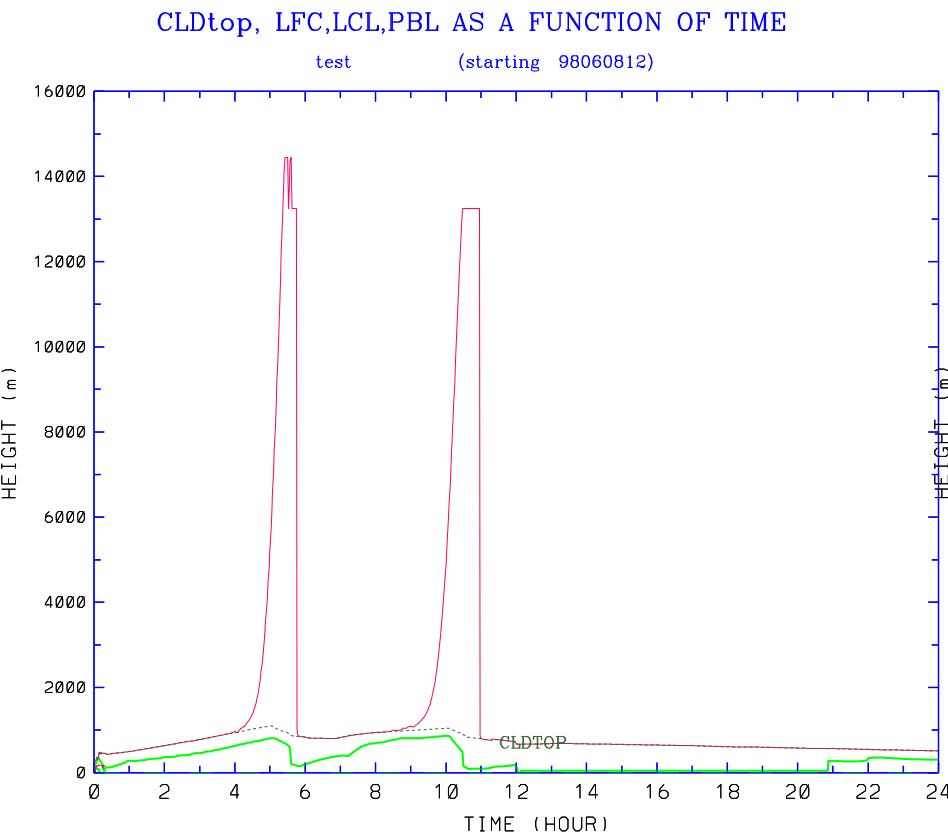
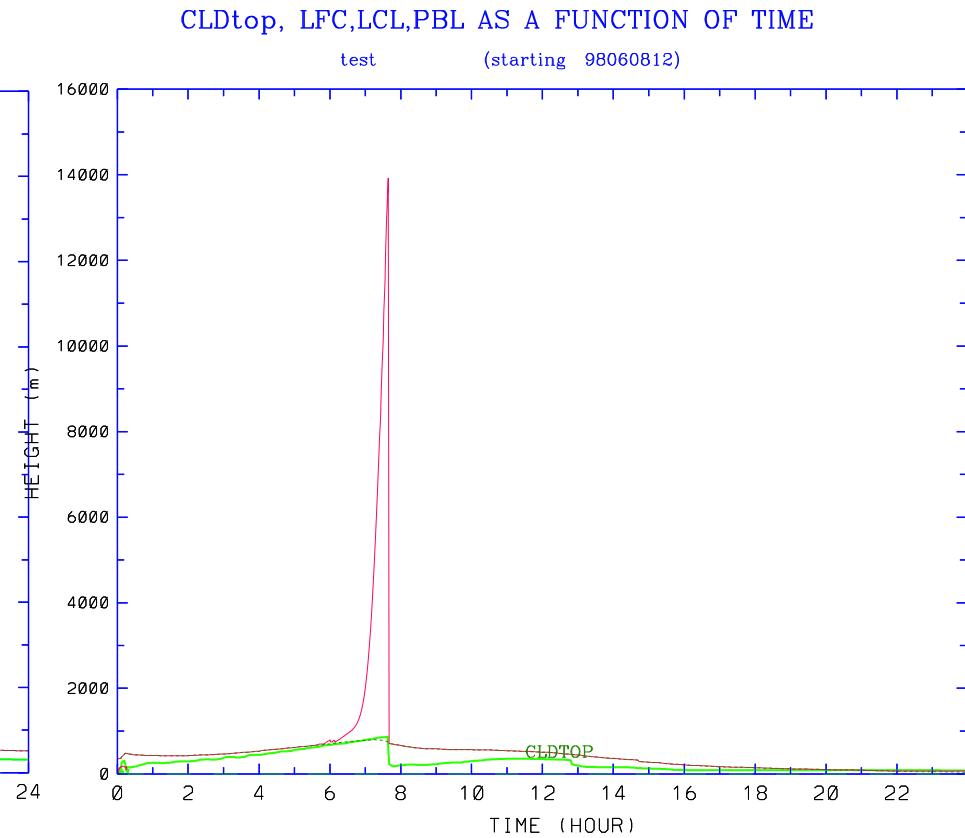
WRF



29



Model-predicted Cloud Updraft Top, LCL and PBL Depth

MM5**WRF**

Summary and Conclusions

- The PSU-Deng SCP scheme that was developed in MM5 is currently being implemented into the WRF modeling system, along with the PSU GS PBL scheme.
- Preliminary 1-D testing based on three convective cases showed that WRF with PSU-Deng SCP is able to reproduce the majority features of the MM5 results, including the cloud depth, fraction and cloud water.
- Future work include 1) addition of NBC advection terms, 2) extend to use with other PBL schemes, both TKE and none-TKE based; 3) 3-D evaluation; 4) comparison with existing shallow convection scheme in WRF; and 5) update and further improvement (e.g., allow more sophisticated microphysics, aerosol, tracer mixing, etc.).

Acknowledgement

- This research is supported by U. S. Environmental Protection Agency.



Supplementary Slides



Parcel buoyancy equation used in KF scheme, except

$$\frac{dz_T}{dt} = 0.2 \cdot W_{\max}$$

Shallow Conv. Scheme:	20-30 min.
KF scheme:	instantaneous

Source Layer:

Upper part of the PBL, a function of cloud width,
Min 10 mb to Max. 60 mb when transfer to K-F scheme



Shallow cloud model description- Formation of NBCs

$$S_a = -\tilde{w} \frac{\partial a}{\partial z} + \frac{R_{ud}}{M_L}$$

$$S_l = \left(\frac{\partial l}{\partial t} - l_c S_a \right) / a$$

where $\frac{\partial l}{\partial t} = -\tilde{w} \frac{\partial l}{\partial z} + \frac{R_{ud}}{M_L} l_u$

R_{ud} Updraft detrainment rate of (kg cloudy air/s)

M_L Mass of air in the grid cell at a specific layer

\tilde{w} Convective-induced subgrid-scale subsidence

l_u Cloud liquid/ice content in updrafts

$l = al_c$ Grid-averaged cloud water/ice

Shallow cloud model description- **Dissipation of NBCs****Dissipation by evaporation:**

$$D_a = -\frac{a}{l_c} K(q_s - q)$$

K : Diffusion coefficient ($10^{-5} s^{-1}$)

q : Water vapor mixing ratio

q_s : Saturation mixing ratio

Depletion by vertical mixing:

$$D_{mix} = \frac{1}{a} \frac{\partial}{\partial z} [(K_v + K_r) \frac{\partial (al_c)}{\partial z}]$$

K_v : Turbulent transport coefficient determined by TKE

K_r : Diffusion coefficient determined by cloud-top radiative forcing



Shallow cloud model description- Dissipation of NBCs

Depletion by precipitation:

$$D_{pre} = -C_1(l_c - l_{cr}) - C_2(\rho q_r)^{0.25} l_c$$

l_{cr} Critical value for liquid/ice water content
at which auto conversion is activated (0.5 g/kg)

ρ Air density

q_r Rain water mixing ratio

C_1, C_2 Dimensional constants

Depletion by ice settling:

$$D_{ics} = -C_3(\rho l_c)^{1.16} / a$$

Ice settling process when $T < 0^\circ C$

C_3 : Dimensional constant

PENNSTATE

 Shallow cloud model description- Dissipation of NBCs
 based on Randal 1980 and Del Genio et al. (1996)

Depletion by CTEI:

$$D_{CTEI} = -10^{-4} \frac{r - r_{\min}}{r_{\max} - r_{\min}} l_c$$

$$r = \Delta h / L \Delta (q + l_c)$$

$$r_{\min} = \frac{\kappa}{\beta} \quad (\approx 0.23)$$

$$r_{\max} = \frac{(1 + \gamma)[1 + \kappa(1 - \delta)]}{2 + \gamma[1 + \kappa(1 + \delta)]} \quad (\approx 0.7)$$

$$\kappa = \frac{C_P T}{L}$$

$$\gamma = \frac{L}{C_P} \left(\frac{\partial q_S}{\partial T} \right)_P$$

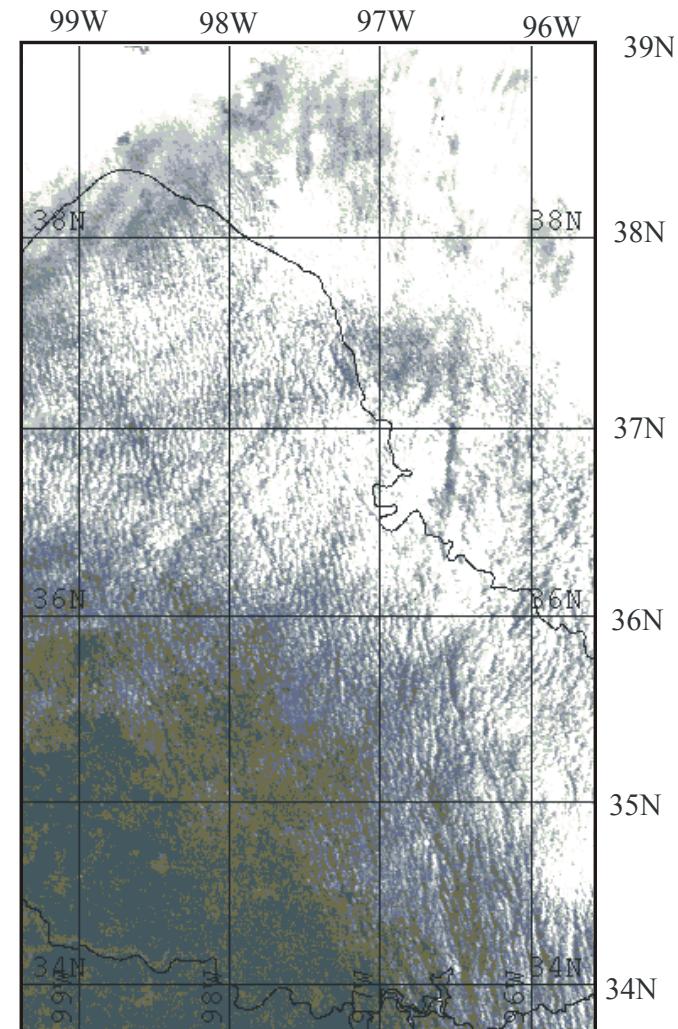
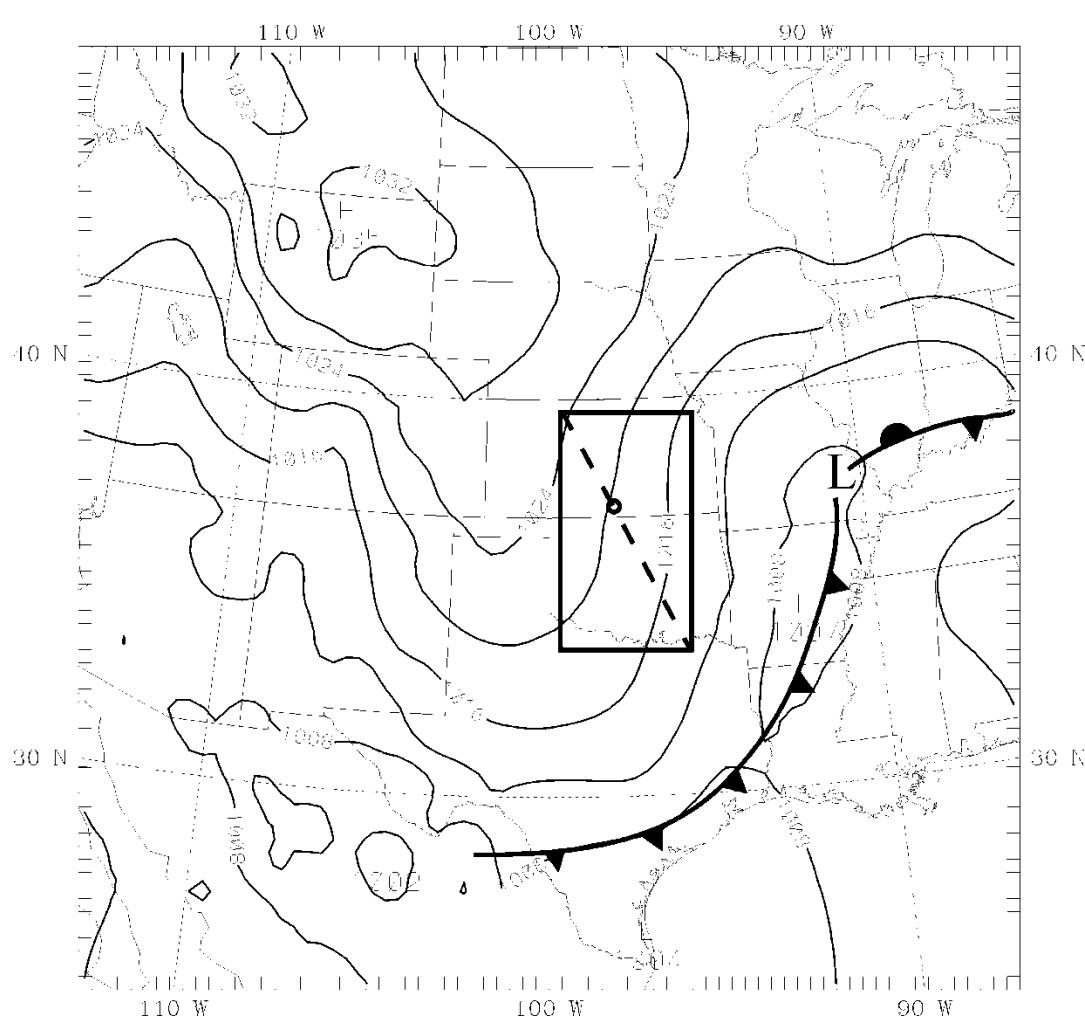
$$\beta = \frac{[1 + \gamma \kappa(1 + \delta)]}{1 + \gamma}$$

$$\delta = 0.608$$



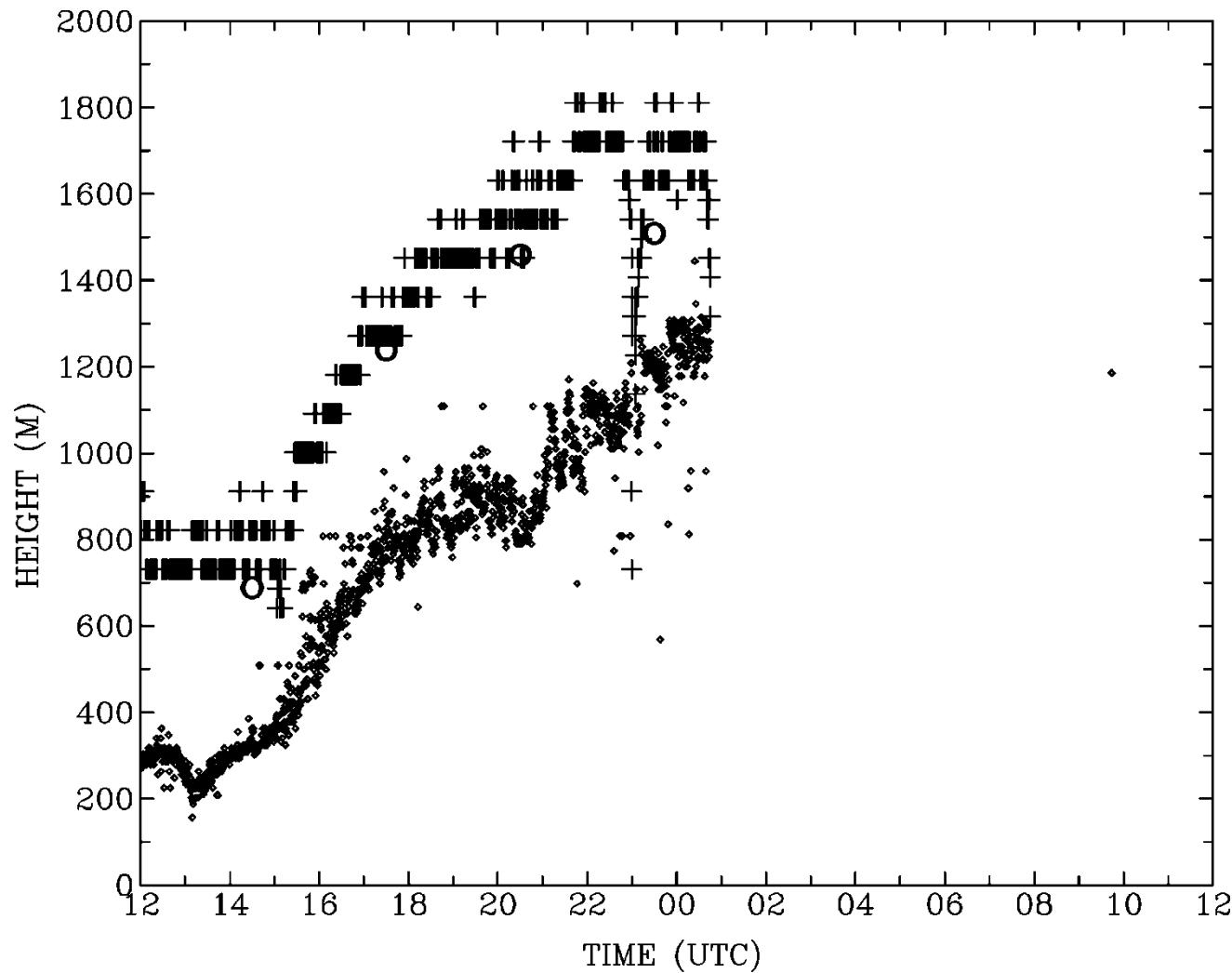
Effect of PSU SCP

Sea Level Pressure Analysis 1997.04.12.00Z +0h Satellite Image 1997.04.12.1709Z





Observed cloud top, base and PBL top

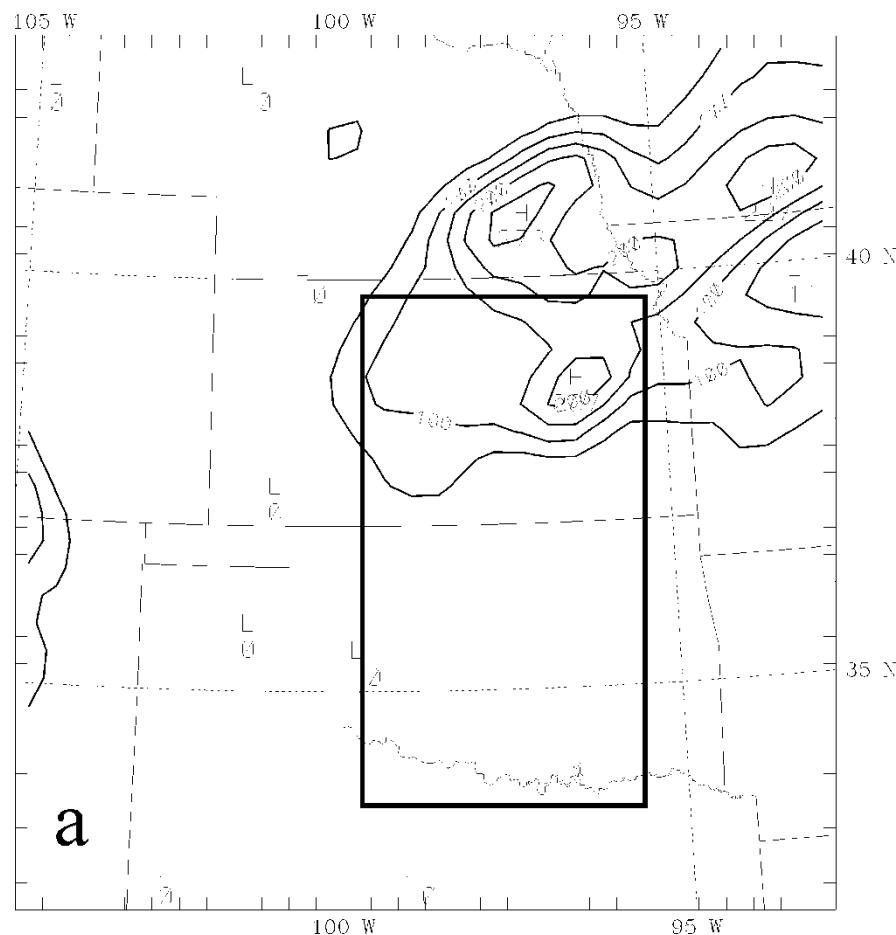




MM5 simulated cloud water content (No shallow convection)

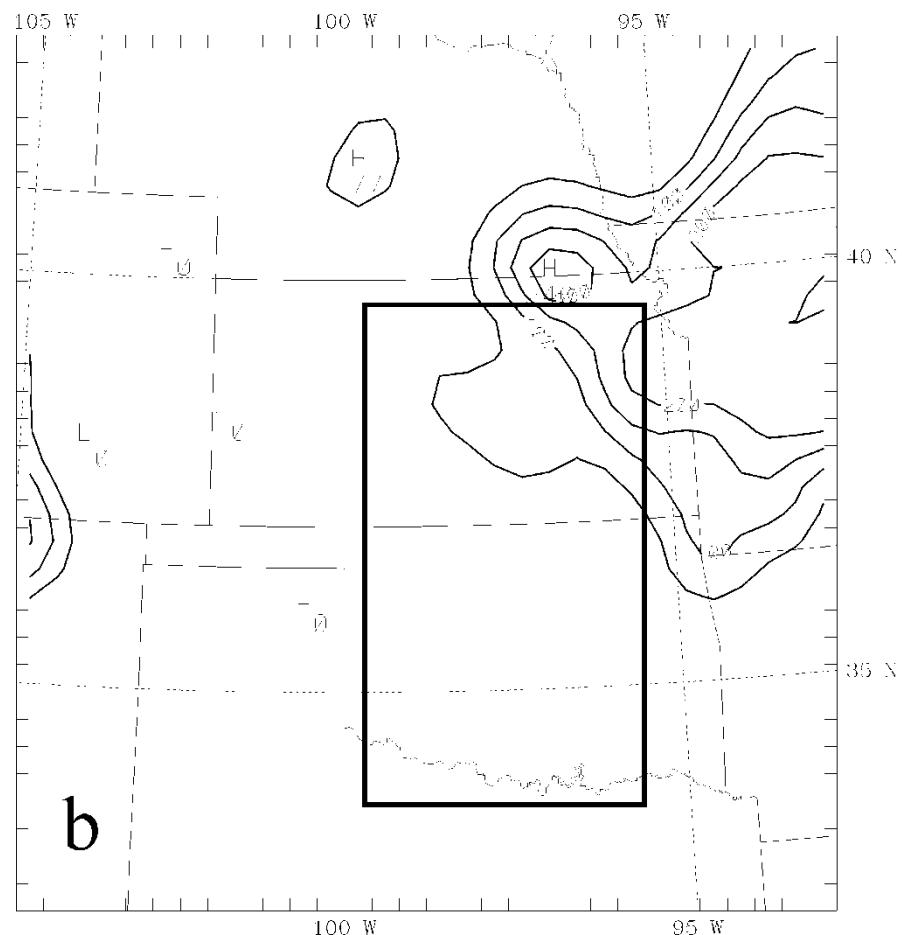
1997.04.12

18Z



a

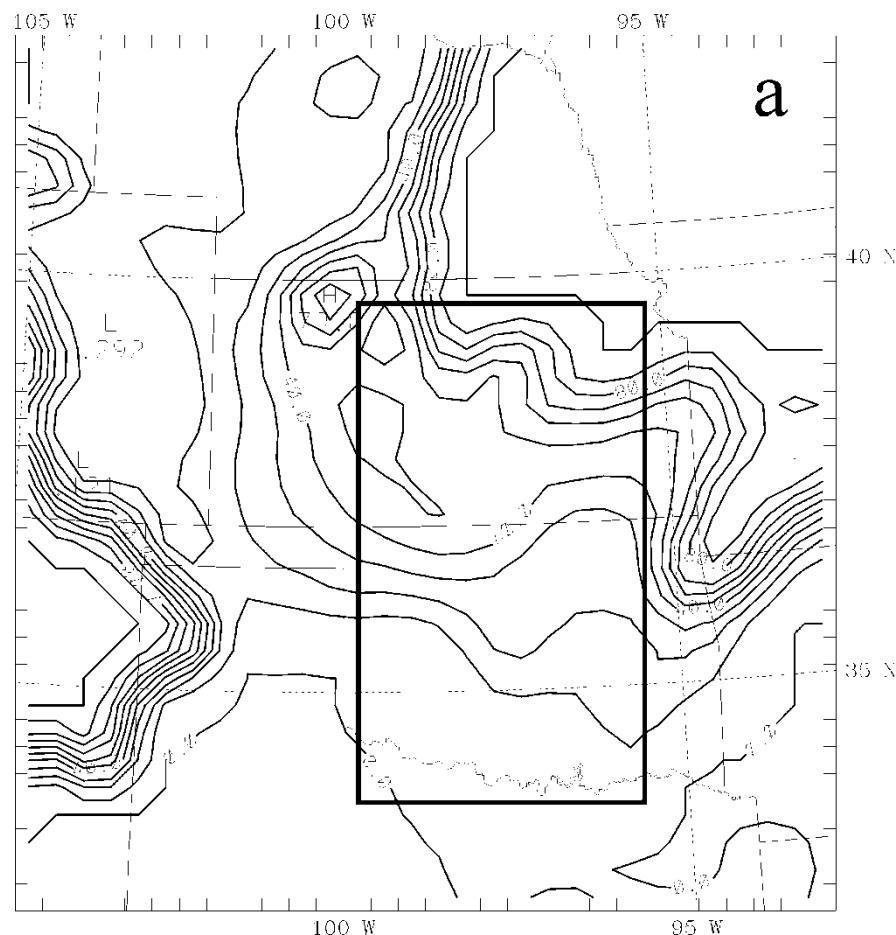
21Z



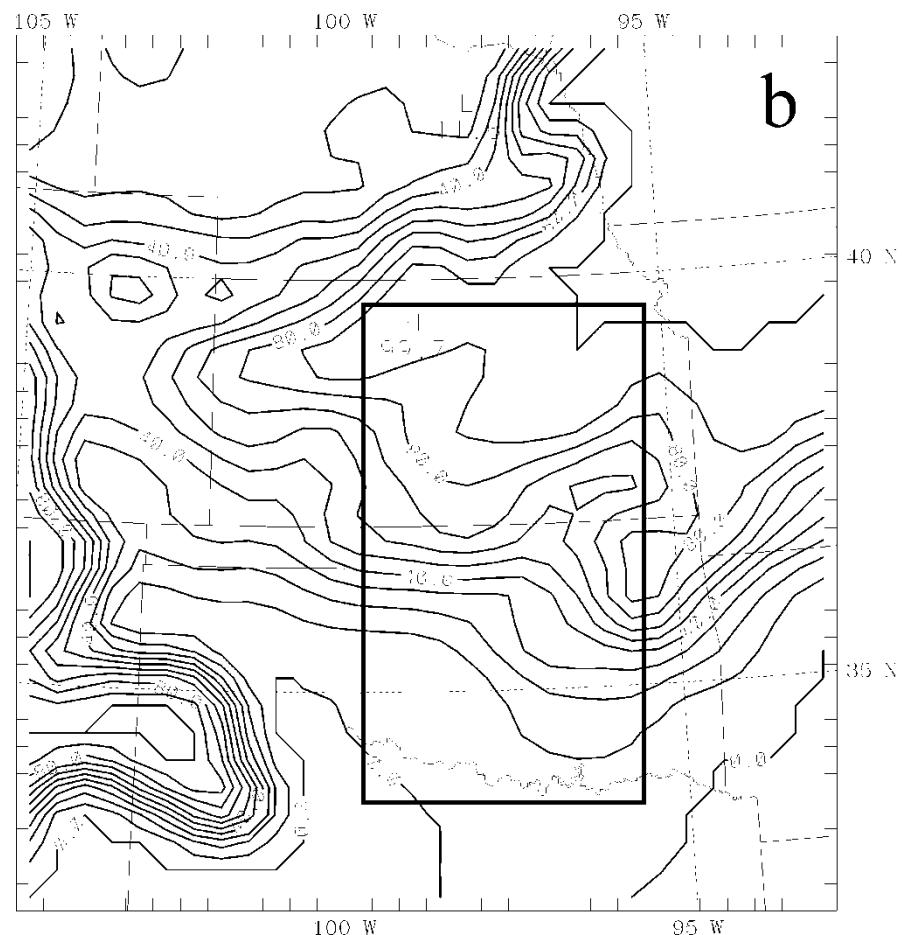
b

MM5 simulated cloud fraction (with shallow convection)
1997.04.12

18Z



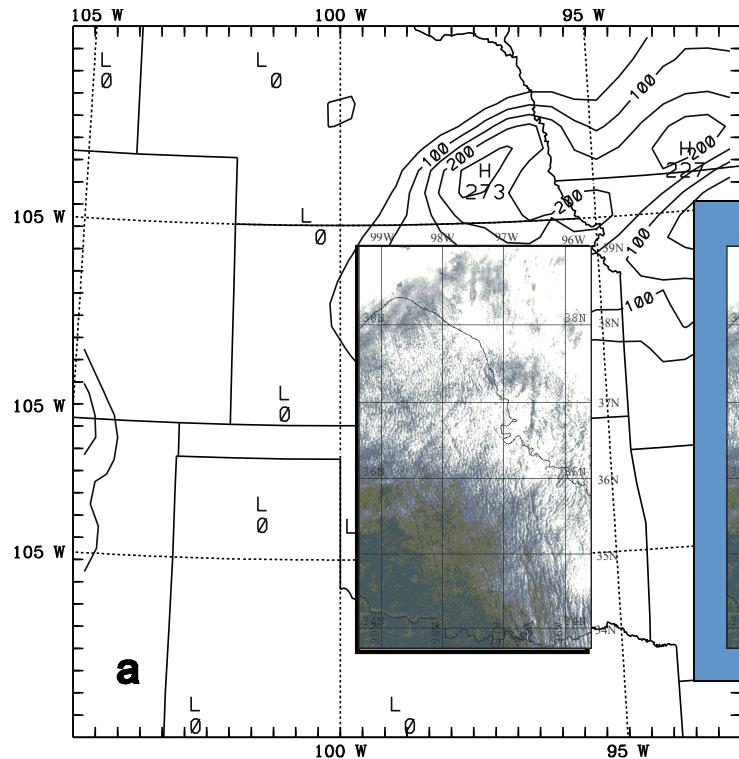
21Z



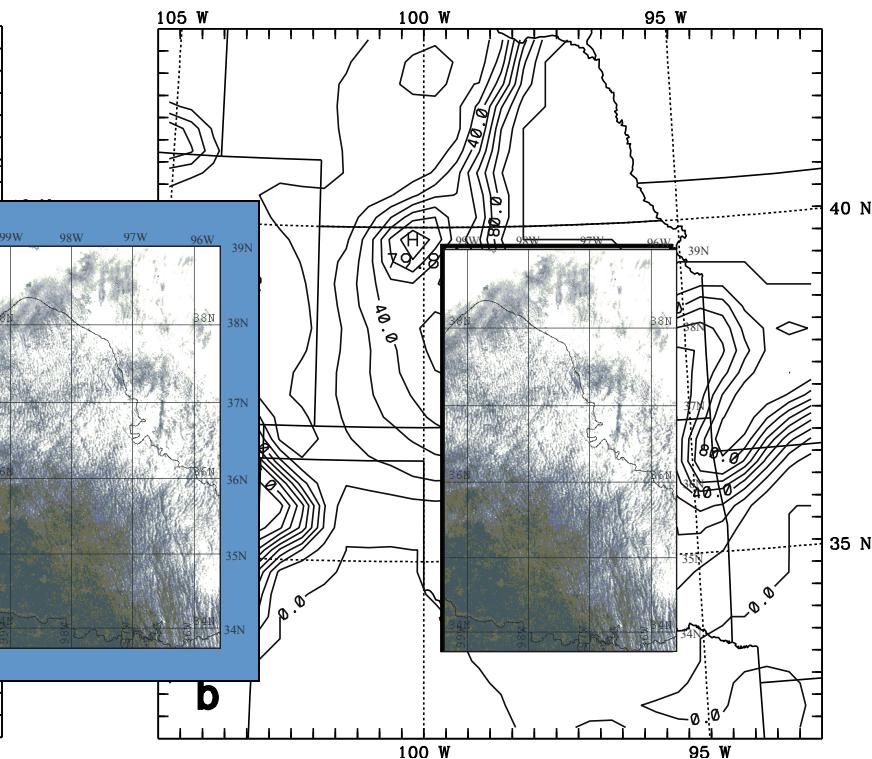


Effect of PSU SCP

MM5 Cloud Water Content



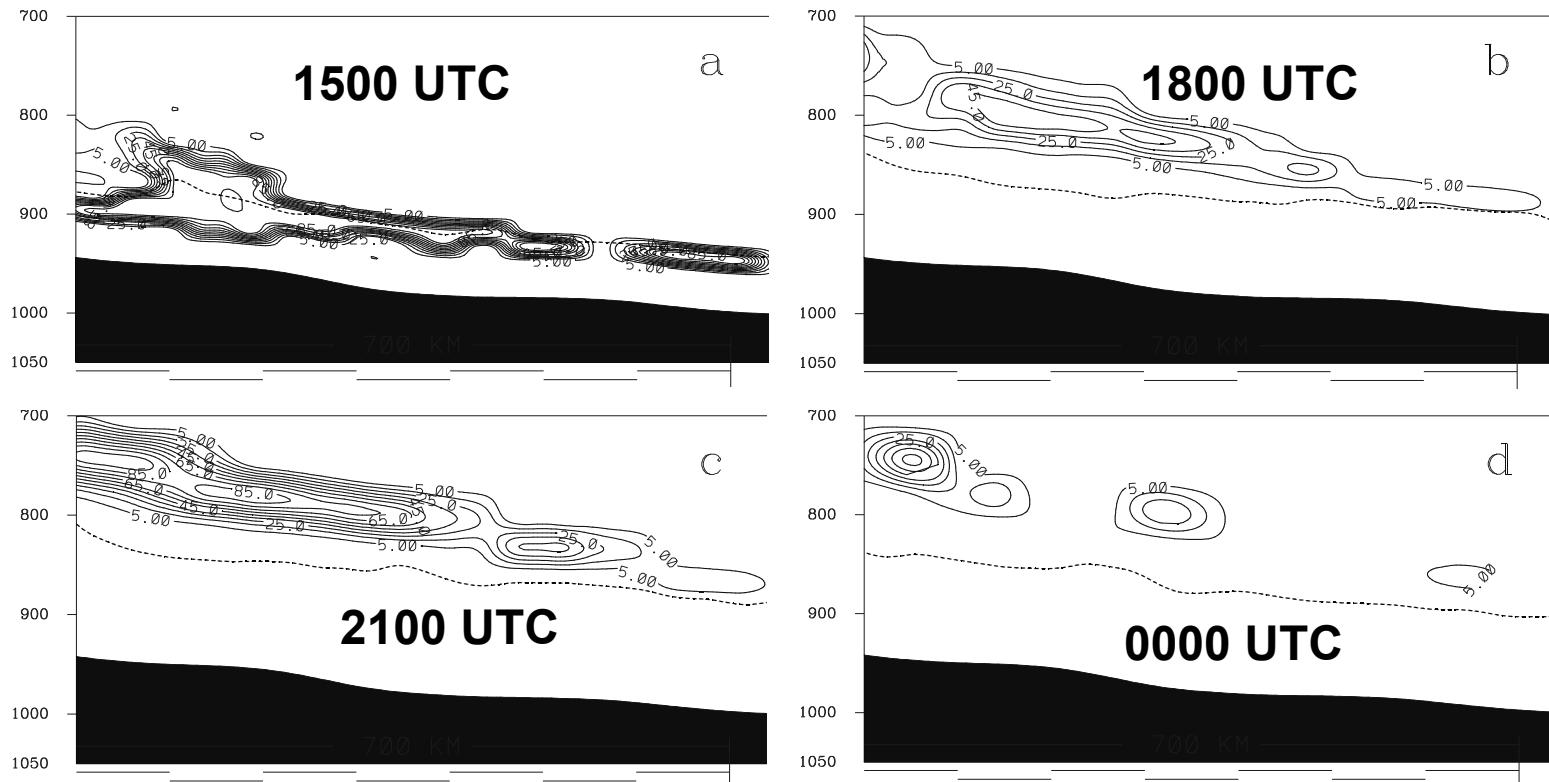
MM5 Cloud Fraction



The 3-D model sensitivity to SCP. a) MM5-predicted liquid/ice water content (g kg^{-1}) of explicit cloud (scaled by 10000), at 1 km AGL at 1800 UTC 12 April 1997 in the experiment without the PSU shallow convection scheme. Contour interval is 0.005 g kg^{-1} . b) MM5-predicted fractional cloud area (%) for low clouds predicted with the PSU shallow convection scheme at the same time and height level. Contour interval is 10%. The rectangular area corresponds to the area covered by the satellite image.



Effect of PSU SCP MM5-Predicted Cloud Fraction



MM5 north-south cross section of predicted fractional cloud area (%) versus pressure (hPa) with the PSU shallow convection scheme: a) 1500 UTC, b) 1800 UTC, c) 2100 UTC (this and the preceding from 12 April 1997), d) 0000 UTC 13 April 1997. Contour interval is 10%. Dashed curve is the MM5-predicted PBL depth.

